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State of stress and stress rotations: Quantifying the role of surface topography and subsurface density contrasts in magmatic rift zones (Eastern Rift, Africa)

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Highlights

State of stress and stress rotations: Quantifying the role of surface topography and subsurface density contrasts in magmatic rift zones (Eastern Rift, Africa)

S. J. Oliva, C. J. Ebinger, E. Rivalta, C. A. Williams, C. Wauthier, C. Currie

- Surface topography influences shallow stress localization
- Subsurface density contrasts influences lower crustal stress localization
- Topography, subsurface, and extension can produce local stress rotation
- Subsurface stresses are important in complete spatial understanding of rift tectonics

State of stress and stress rotations: Quantifying the role of surface topography and subsurface density contrasts in magmatic rift zones (Eastern Rift, Africa)

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Abstract

In rift settings, the crustal stress field is dominated by extension, which leads to rift-parallel topography and basin alignments. However in some continental rift systems, some observables of the orientation of principal stresses show substantial deviations from these patterns. Such stress field rotations are currently poorly understood and could reflect the critical role of rift magmatism in the creation of topography, the plate state-of-stress, and volcanic and tectonic processes. Yet the role of magma intrusions, crustal thinning, and rift basin and flank topography on rift zone stress field rotations remain poorly quantified. The seismically- and volcanically-active Magadi-Natron-Manyara region of the East African Rift shows a 60° local stress field rotation with re-

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spect to regional extension. Here, we test the hypothesis that such rotation is due to the cumulative effects of surface and subsurface loads (lateral subsurface density contrasts). We use analytical and calibrated numerical models of magmatic rift zones to simulate lithospheric deformation in the presence of magma bodies, crustal thinning, and topography to quantify their effect on intrusions and fault kinematics in a rift setting. Our 3D static models of a weakly extended rift suggest that surface topography influences shallow stress localization, whereas subsurface density contrasts play a larger role in lower crustal stress localization. Both patterns suggest a preferred region for melt storage beneath the rift valley. We show that the interaction between topography, crustal thinning, extension, and a pressurized magma reservoir could generate principal stress orientations consistent with the local stress rotation observed from earthquake focal mechanisms. Our results demonstrate how rift topography and the geometry of crustal thinning can guide magmatism and strain localization, highlighting the need for a three-dimensional treatment of rift kinematics.

Keywords: rift, continental rifting, tectonics, stress, topography

1. Introduction

Continental rift zones are sites of intraplate stretching, commonly with adiabatic decompression melting accompanying lithospheric thinning. Large fault systems bound the rift and are segmented along the length of the rift, with strain linkage by rift-oblique transfer faults and relay ramps (e.g., Accella et al., 2005; Ebinger, 1989; Larsen, 1988; Tibaldi et al., 2020). Volcanoes and underlying crustal magma chambers cause fundamental changes

8 to the density and thermal structure, load the plates, and change the state
9 of stress within the crust and mantle lithosphere (e.g., Karlstrom et al.,
10 2009). In magmatic regions, the local stress field induced by volcanic edifice
11 loading/unloading (e.g., Kervyn et al., 2009; Maccaferri et al., 2014) and/or
12 weak zones (e.g., Corti et al., 2004) then influence subsequent magma migra-
13 tion and vent distribution (e.g., Karlstrom et al., 2009; Rivalta et al., 2019)
14 Generally, dike and fracture orientations are favored perpendicular to the
15 direction of the least compressive stress (e.g., Pollard, 1987; Sibson, 1985).
16 Although published two-dimensional models of loading on plates without in-
17 plane regional stresses offer insight (e.g., Karlstrom et al., 2009), a critical
18 gap in understanding remains on how surface (e.g., volcanoes, rift flanks) and
19 subsurface intrusions (sills, dikes, magma chambers) combine with regional
20 stresses in rift zones. In this study, we aim to address this gap and comple-
21 ment existing conceptual models of magmatic rift zone structural patterns
22 (e.g., Magee et al., 2016; Muirhead et al., 2015; Tibaldi et al., 2020).

23 We build analytical and numerical models of magmatic rift zones to simu-
24 late lithospheric deformation in the presence of magma bodies to quantify the
25 interactions between surface and subsurface loading on intrusions and fault
26 kinematics in a rift setting. We use the open-source finite element modeling
27 package PyLith v2.2.1 (Aagaard et al., 2013, 2017a,b) to build two- (2D) and
28 three-dimensional (3D) physics-based static models, calibrated to analytical
29 stress models, to quantify the effects of loading on the crustal state of stress.
30 We probe how inclusions (subsurface volumes of density different than host
31 material, e.g., magmatic intrusions) and loads/excavations of different spa-
32 tial scales guide local stress field rotation and strain localization to favor the

33 initiation of extensional faults, dikes, and eruptive centers oriented obliquely
34 to the regional extension direction (e.g., Ebinger, 1989; Acocella et al., 2005).
35 Our approach is two-fold: we first evaluate model sensitivities to topography
36 and subsurface density contrasts separately, and then we compare our models
37 with observed stress indicators (i.e., earthquake locations and focal mecha-
38 nisms). The static elastic modeling approach allows us to consider each load
39 separately, and summatively.

40 **2. Tectonic Background**

41 The Magadi-Natron-Manyara basins are located in the southernmost sec-
42 tor of the magmatic Eastern rift (Fig. 1). The Natron basin, which contains
43 lavas and sedimentary strata dating to ~ 3 Ma, developed in metasomatized
44 Archaean lithosphere, as confirmed by mantle xenoliths (Mana et al., 2015;
45 Aulbach et al., 2011; McHenry et al., 2011; Foster et al., 1997). The present-
46 day Natron basin is a N-S trending half-graben bounded by a monocline to
47 the east, that formed during a major phase of faulting at ~ 1.2 – 1 Ma when
48 the east-dipping Songo border fault developed within the original wider basin
49 (Foster et al., 1997; Muirhead et al., 2016). Gelai volcano formed at ~ 1 Ma
50 whereas carbonatite volcano Oldoinyo Lengai (~ 0.37 Ma, OL, Fig. 1) post-
51 dates the main phases of rift faulting and basaltic volcanism (Mana et al.,
52 2015; Foster et al., 1997; Muirhead et al., 2015, 2016). The Naibor Soito
53 monogenetic cone field (NS, Fig. 1) lies between Oldoinyo Lengai and Gelai
54 (e.g., Weinstein et al., 2017; Dawson, 1992).

55 The region is an ideal place to quantify the different loading contributions
56 on crustal state-of-stress because the 2013-2014 Continental Rift and Fluid-

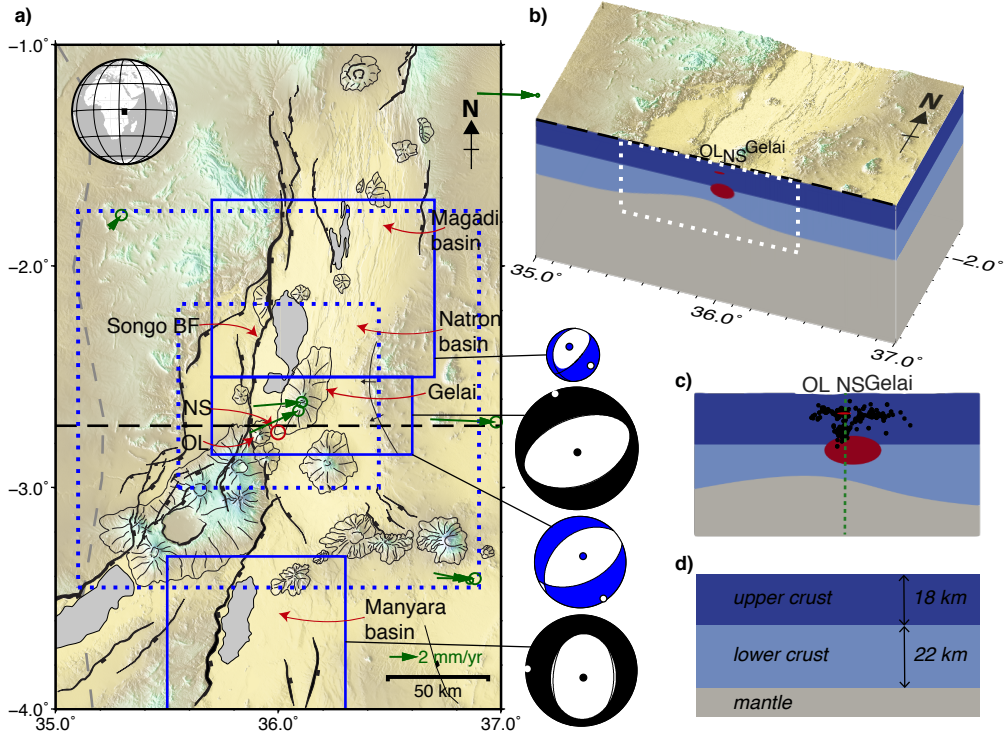


Figure 1: Study area and model geometry. a) Map of the Magadi-Natron-Manyara basins of the East African Rift system. Inset shows the location on the globe. Solid blue boxes enclose areas of seismicity analyzed by Weinstein et al. (2017) and the corresponding summed moment tensors (blue beachballs are from local focal mechanisms, black beachballs are from teleseismic focal mechanisms) that reveal along-axis local stress rotation. Green arrows are GNSS velocity vectors for a Nubia-fixed model with their error ellipses (King et al., 2019). Large dotted blue box describes the extent of the 3D model domain, with higher resolution topography (1 km) within the smaller dotted box (nodal spacing doubles at every step outward until the edge of the domain). The black dashed line represents the E-W transect at 2.72°S latitude which is used in the 2D models. Faults and edifices are from Muirhead et al. (2015). BF: border fault; NS: Naibor Soito; OL: Oldoinyo Lengai. b) Perspective figure looking northwest shows a section of the 3D model domain sliced along 2.72°S . The upper crust is dark blue, lower crust is light blue, mantle is gray, and the sill and magma chamber are red. c) Double-difference earthquake locations within 5 km north and south of the cross-section (dotted white box in b). Dashed green line ($x = 2750$ m, where $x = 0$ is 36° longitude) indicates the location of the 1D profile in Fig. 3. d) Reference model thicknesses.

57 Tectonic Interaction (CRAFTI) network (Weinstein et al., 2017), as well as
58 the recent Seismic and Infrasonic Networks to study the volcano Oldoinyo
59 Lengai (SEISVOL) project in 2019-2020 (Reiss et al., 2021), provide excel-
60 lent constraints for modeling. Nubia-fixed rigid plate models predict 2–5
61 mm/yr approximately E-W opening in the area (Saria et al., 2014; King
62 et al., 2019) (green arrows, Fig. 1). Receiver function studies show crustal
63 thinning from 40 km to 29 km in the middle of the rift valley (Plasman
64 et al., 2017). A velocity contrast at ~ 18 km depth is interpreted from a
65 clear negative peak in receiver functions around 3 s, which is also evident in
66 the tomography (Plasman et al., 2017; Roecker et al., 2017). Seismicity and
67 tomography studies (from the regional CRAFTI network) image a magma
68 chamber and sill complex in the subsurface (Weinstein et al., 2017; Roecker
69 et al., 2017; Reiss et al., 2021). The local SEISVOL network allowed fur-
70 ther detailed seismic studies and identified clusters beneath the Naibor Soito
71 volcanic field at 5–10 km depth interpreted as stress concentrations at the
72 edges of pressurized sills, and at 12–18 km depth representing the top of
73 the lower-crustal magma chamber, confirming that some seismicity patterns
74 observed in the earlier network are consistent over longer timescales (Reiss
75 et al., 2021). Focal mechanisms of both local and teleseismic earthquakes
76 reveal an along-axis local stress rotation of $\sim 60^\circ$ in the OL-NS-Gelai com-
77 plex ($\sim N150^\circ E$ extension), relative to the $\sim E-W$ regional extension indicated
78 by seismic and geodetic data (Weinstein et al., 2017) (Fig. 1). Reiss et al.
79 (2021) report a mean T-axis direction of NW-SE in the OL-NS-Gelai area,
80 noting that beneath Naibor Soito, there was no dominant mechanism above
81 the magma chamber (below 13 km) and in the sill complex (5–7 km; 9–11

82 km). Instead, the local stress rotation (\sim N-S T-axes) was mainly observed at
83 \sim 7–9 km depth. A detailed study of the variations in fault, dike, and cone
84 orientations in the East African Rift suggest that in different regions, dif-
85 ferent stresses dominate: the regional stress state, magma overpressure and
86 volcanic edifice load, pre-existing structures, or local stress rotations due to
87 mechanical interactions of rift basin segments (Muirhead et al., 2015). A
88 separate study of East African Rift volcanoes hypothesized that rift-oblique
89 dikes and fissures were either due to lithospheric heterogeneities or mag-
90 matic processes (Wadge et al., 2016). Since dikes are subvertical (e.g., Calais
91 et al., 2008; Muirhead et al., 2015; Weinstein et al., 2017), it is unlikely that
92 they reactivated Archaean and Pan-African strain fabrics that dip at an-
93 gles $\leq 45^\circ$ (e.g., Le Gall et al., 2008; Smith and Mosley, 1993), although the
94 pre-existing oblique fabric might influence later structures by lowering the
95 residual strength of the rock.

96 In June–August 2007, the Naibor Soito complex and the flanks of Gelai
97 were the site of $50+ 3.6 < m_b < 5.9$ earthquakes and diking with associ-
98 ated opening of 1.0–2.4 m, followed by eruption at Oldoinyo Lengai through
99 2008 (Baer et al., 2008; Biggs et al., 2009; Calais et al., 2008). During the
100 seismomagmatic sequence, 65% of the total geodetic moment release was at-
101 tributed to dike opening, highlighting the importance of magma intrusions
102 in accommodating rift opening (Calais et al., 2008). New geophysical results
103 reinterpret the 2007 diking event to have occurred above a shallow sill com-
104 plex underlain by a low-velocity mid-crustal magma mush (Oliva et al., 2019;
105 Roecker et al., 2017; Weinstein et al., 2017).

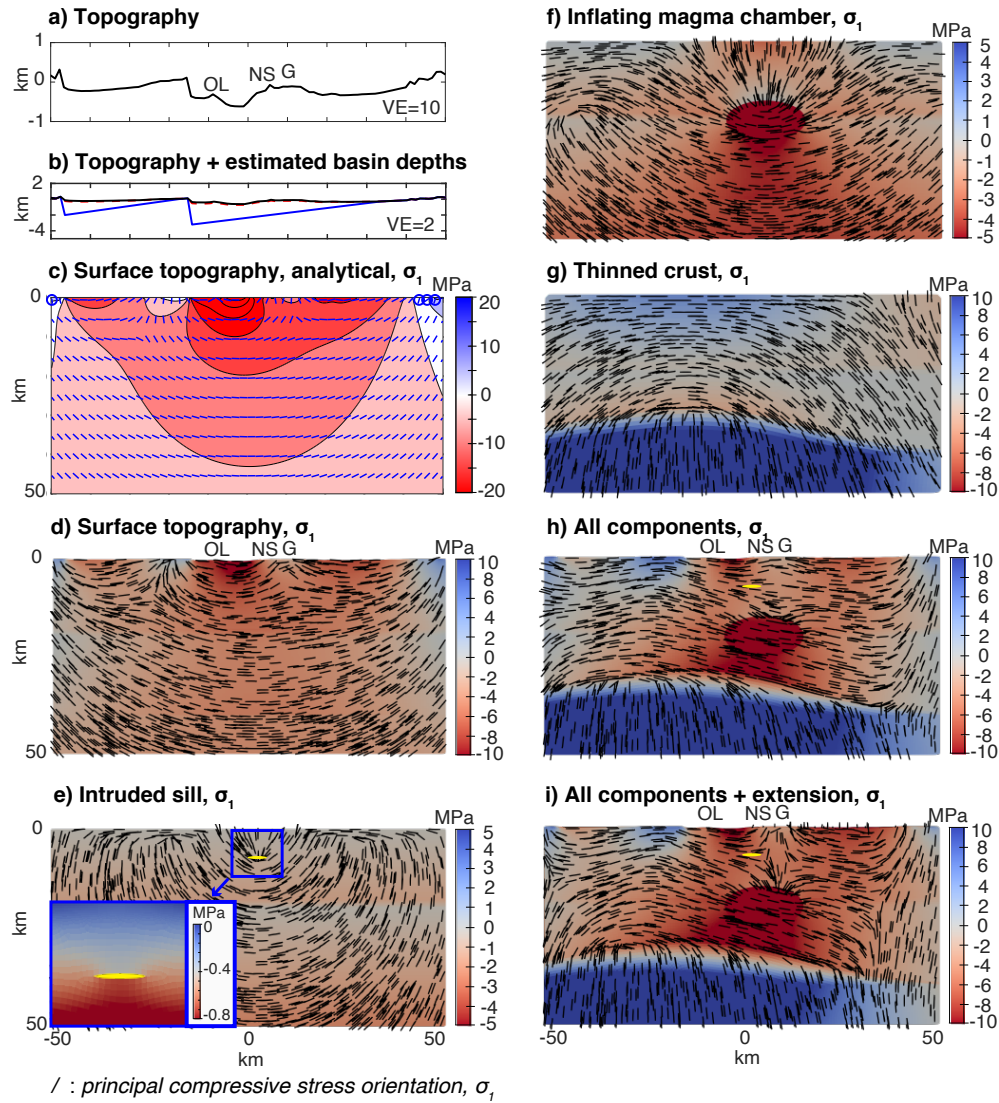


Figure 2: Two-dimensional static model results. Cross-sections are E-W transects at 2.72°S , centered at 36°E (black dashed line in Fig. 1). a) Topography with $10\times$ vertical exaggeration. b) Topography (black), estimated basin depths (blue) (Birt et al., 1997), and topography corrected for low-density basin fill (red dashed) with $2\times$ vertical exaggeration. c) Analytical model of topographic loading (as in b) colored by stress magnitude (positive stress values are compressive). Line segments (blue on c and black elsewhere) are orientations of the principal stress (greatest compressive), σ_1 . d-i) Numerical models of separate stress contributions of each rift component colored by stress magnitude. Stresses are deviations in gravitational stress ($g = 9.78 \text{ m/s}^2$) with respect to a reference model: a uniform-thickness three-layer model with no magmatic systems. Numerical models show the separate effects of d) surface topography, e) an intruded sill (density contrast, yellow ellipse), f) an overpressurized magma chamber (density contrast and outward pressure), g) the Moho topography of a thinned crust, h) the combined effects of the four above-mentioned components, and i) the combined effects of all components including a displacement boundary condition to represent regional extension. OL: Oldoinyo Lengai, NS: Naibor Soito, G: Gelai.

106 **3. Modeling approach and limitations**

107 The crustal stress field is determined by the combination of stress-generating
108 mechanisms (e.g., topographic loads, magma overpressure) and stress-relieving
109 mechanisms (e.g., inelastic deformation, earthquakes). Here we focus on the
110 former and treat the location and orientation of dikes and faults, as well as
111 Global Navigation Satellite System (GNSS)-derived plate motion directions,
112 as stress indicators. Crustal stresses control the predominant orientation of
113 faults and the slip direction during earthquakes. Elastic stresses cycle phases
114 of stress accumulation from the action of tectonic forces, punctuated by re-
115 lease events during major earthquakes. Constraining the pattern of elastic
116 stresses at the regional, “tectonic environment” scale can help us understand
117 these cycles and any stress rotation occurring.

118 To investigate first-order stress contributions of different components, we
119 use present-day constraints for model geometry and evaluate static elastic
120 stress models. With elastic models, we take advantage of linear superpo-
121 sition, wherein the net response to various components is the sum of the
122 individual responses. This enables us to investigate each component of inter-
123 est separately, while also being able to superimpose them to see the combined
124 effect. Note that because time is an important aspect of ductile behavior,
125 our static models cannot capture time-dependent rheological behavior.

126 In our models, we do not include faults as discontinuities and instead
127 only consider the influence of surface topography, magma bodies, and crustal
128 thinning to the stresses on a continuous crustal model. Faults both respond
129 to and influence stress orientations, and modeling this feedback is difficult to
130 capture in a static model and would be more appropriately done in a separate

131 dynamic modeling study. In addition, long-wavelength gravitational loads
132 (such as those considered in this study) would dominate the short-wavelength
133 effect of faults, except the largest faults (e.g., Maccaferri et al., 2014).

134 The abovementioned modeling decisions enable us to evaluate the state-
135 of-stress within the continental crust, specifically the stress contributions of
136 topography, crustal thinning, and magma bodies. Our goal is to quantify
137 the resulting spatial stress patterns, to summarize first-order effects, and
138 discuss the implications of our findings on the evolution of magmatic and
139 fault systems. Our goal is not to simulate the detailed state-of-stress within
140 the Natron-Magadi system. Hence, we avoid additional model complexities
141 to ensure that the first-order effects we are interested in are not obscured.
142 While viscous stresses are certainly important in the long-term evolution
143 of the system, in the short timescales of earthquakes and dikes, the elastic
144 component of the stresses are more important. The caveat is that our results
145 are generally limited to the strong part (\sim upper crust) of the lithosphere,
146 where viscous effects are less important.

147 **4. Methods**

148 *4.1. Model geometry*

149 Our 2D models are about 200 km across and 100 km deep, and the 3D
150 models are about 200 km \times 200 km \times 100 km (Fig. 1c,d). We only interpret
151 the results within a smaller 100 km (\times 100 km) \times 50 km box embedded
152 within the larger domain to avoid effects due to boundary conditions (see
153 dashed white box on Fig. 1b). The model geometry was built from the
154 Shuttle Radar Topography Mission (SRTM) 90m digital elevation database

Layer	Depths (km)	ρ (kg/m ³)	Vs (km/s)	Vp (km/s)	λ ($\times 10^{10}$)	μ ($\times 10^{10}$)
upper crust	0–18	2761	3.6	6.2	3.46	3.58
lower crust	18–22	2833	3.4	6.5	5.38	3.29
mantle	22–100	3300	4.5	8.1	8.29	6.68
sill	~ 7	2661	3.6	6.2	3.33	3.45
chamber	15–25	2661	3.3	5.5	2.25	2.90

Table 1: Material properties. First-order crustal densities (ρ) were estimated from seismic velocities (Vp, Vs) (Plasman et al., 2017; Roecker et al., 2017; Tiberi et al., 2018), using an empirical relation between crustal density and seismic velocity (Brocher, 2005). The mantle velocity and density were adapted from the velocity models of Plasman et al. (2017) and Tiberi et al. (2018). Sill and magma chamber velocities are taken from Roecker et al. (2017). Lamé parameters are calculated from density and seismic velocities: $\mu = \rho V s^2$; $\lambda = \rho V p^2 - 2\mu$.

155 v.4 (Jarvis et al., 2008), estimated basin depths from refraction and wide-
156 angle reflection data (Birt et al., 1997), and crustal tomography (Roecker
157 et al., 2017). The Moho structure was constrained from receiver functions
158 (Plasman et al., 2017) assuming a velocity contrast at 18 km (Plasman et al.,
159 2017; Roecker et al., 2017) and that unstretched crust is 40 km thick (Mansur
160 et al., 2014; Tugume et al., 2012).

161 4.2. Analytical models

162 We first modeled topography and extension using analytical 2D models.
163 Because our study area sits at the edge of the uplifted Tanzanian craton and
164 we do not include dynamic topography in the models, we removed a long-
165 wavelength trend from the SRTM topography. Detrended topography with

166 adjustments based on lower density basin fill (dotted red line, Fig. 2b) was
167 modeled analytically as a line load in an elastic 2D half-space (e.g., Dahm,
168 2000). Basin-fill adjustments were made such that the surface load of the
169 adjusted topography with uniform density ($\rho = 2.7 \text{ g/cm}^2$) is equivalent to
170 that of original topography with lower-density basins ($\rho = 2.55 \text{ g/cm}^2$, Birt
171 et al. (1997)).

172 *4.3. Numerical models*

173 The 2D and 3D finite-element (FE) models were built using the mesh-
174 ing software Trelis and modeled with regional extension and gravity using
175 the open-source lithospheric modeling package PyLith (Aagaard et al., 2013,
176 2017a,b). PyLith was specifically designed for the solution of tectonic prob-
177 lems, and is capable of solving static, quasi-static, and dynamic problems in
178 geophysics. We do not consider time and thermomechanical effects in our
179 modeling efforts, and we therefore restrict our models to static elastic simu-
180 lations without faults. Sample Trelis and PyLith input files are provided in
181 Supplementary materials SM3–SM4.

182 *4.3.1. Mesh generation*

183 In 2D models, we mesh with quadrilateral elements (paving) with a mesh
184 size of 0.5 km within the crust and the magma chamber, 0.2 km within the sill,
185 and 0.8 km in the mantle. In 3D models, we mesh with tetrahedral elements
186 and implement adaptive meshing to create a mesh with finer spacing along
187 areas of interest that gets coarser towards the lateral and bottom boundaries.
188 We use Trelis’ bias sizing function to produce gradients of mesh size that
189 respect the predefined discretization size along curves (see Supplementary

190 material SM2). The discretization size is set to 1 km along the surface
191 topography, 1 km along the magma chamber boundary, 0.2 km along the sill
192 boundary, 2.5 km along the interface at 18 km depth, 5 km along the Moho,
193 and 10 km at the bottom of the model.

194 *4.3.2. Material parameters*

195 Using an empirical relation between crustal density and seismic velocity
196 (Brocher, 2005), the first-order crustal densities were estimated from seismic
197 velocities (Plasman et al., 2017; Roecker et al., 2017; Tiberi et al., 2018).
198 The mantle velocity and density were adapted from the velocity models of
199 Plasman et al. (2017) and Tiberi et al. (2018). In the cases with the sill
200 and the magma chamber, magma density was taken to be slightly lower than
201 surrounding rock with seismic velocities taken from Roecker et al. (2017),
202 which imposes a buoyancy effect on the models. Although temperature,
203 pressure, composition, and hydration state also influence density, seismic
204 velocity, and rheology (e.g., Guerri et al., 2015), these are not considered
205 in this study; that is, the elastic parameters (Lamé constants) used in the
206 numerical model were calculated using density and seismic velocities only
207 (Table 1). Because the static models do not include the time component,
208 ductile behavior is not captured (e.g., Maccaferri et al., 2014).

209 *4.3.3. Initial and boundary conditions*

210 We first evaluated separate stress contributions of each rift structure,
211 then later combined them all together. Stress contributions were calculated
212 as deviations in gravitational stress (with acceleration due to gravity $g = 9.78$
213 m/s^2) with respect to the assumed lithostatic equilibrium state of a reference

214 model (Fig. 1d, see Supplementary material Fig. SM1): a uniform-thickness
215 three-layer model with no magmatic systems and no topographic variations
216 ($z = 0$ on Fig. 2a). Topography and crustal thinning were represented by
217 modifications to the model geometry. The upper-crustal sill was modeled as
218 an ellipsoid 500 m thick, circular in plan-view (major axis of 5 km), centered
219 at 7 km depth. The mid-crustal magma chamber, which likely consists of
220 geometrically complex zones of mush and melt (Cashman et al., 2017), was
221 modeled as a 10 km thick, 20 km wide spheroidal inclusion, centered at 20
222 km depth, with an outward-directed pressure boundary condition of 10 MPa
223 to represent inflation (Allmann and Shearer, 2009).

224 For all models, a roller boundary condition was imposed on the bottom
225 of the block and a free surface on the top. For models without extension,
226 the lateral boundaries had a zero-displacement boundary condition. For
227 extensional models, five meters of displacement was imposed on each lateral
228 side, 2.5×10^{-5} of strain, which corresponds to ~ 3 MPa horizontal upper
229 crustal stress. The choice of imposed stresses is guided by the assumption
230 that deviatoric stresses significantly larger than 10 MPa would be released
231 by earthquakes, since 10 MPa is the global median stress drop associated
232 with earthquakes (Allmann and Shearer, 2009).

233 In 2D models, plane strain is assumed. No discontinuity or weakness
234 is introduced to represent the known faults, since we treat faults as stress
235 indicators rather than stress-generating structures.

236 *4.3.4. Post-processing*

237 Principal stress orientations are calculated from total stress values whereas
238 stress magnitudes are relative to the assumed lithostatic equilibrium state of

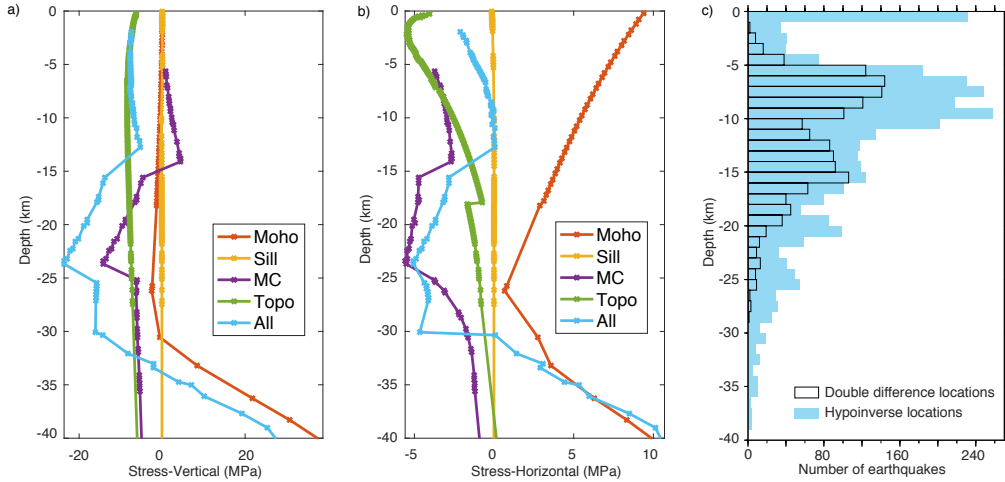


Figure 3: One-dimensional profiles from 2D static models. a) Vertical stresses along the 1D depth profile (green dashed line on Fig. 1c) with the stresses from a reference model subtracted. Initial stress state is assumed to be that of the reference model: a uniform-thickness three-layer model with no magmatic systems. Model values, marked with \times 's along the lines, are taken within 100 meters of the profile ($x = 2750$ m, where $x = 0$ is 36° longitude). Topo: topography model (Fig. 2d), Sill: sill model (Fig. 2e), MC: magma chamber model (Fig. 2f), Moho: thinned crust model (Fig. 2g), All: all-components model (Fig. 2h). b) Corresponding horizontal stresses along the 1D depth. c) Depth histogram of earthquakes in the study area from Weinstein et al. (2017).

239 the reference model (see Supplementary material SM5). The greatest com-
 240 pressive stress, σ_1 , is taken as a proxy for preferred orientation of magmatic
 241 intrusions which is perpendicular to the preferred opening direction. The
 242 orientation of the least compressive stress (σ_3) is comparable to the T-axis
 243 of earthquake focal mechanisms or the regional extensional GNSS direction.
 244 Model results are viewed using the open-source visualization software Par-
 245 aView 5.7.

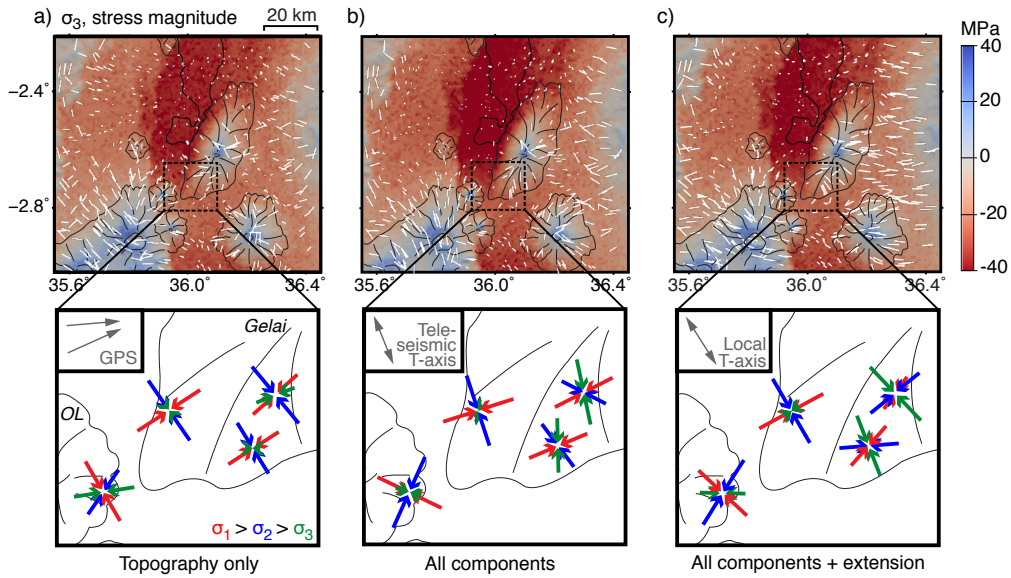


Figure 4: Depth slices of principal stresses in the upper crust (0–18 km depth). These are 3D model results of a) topography only, b) all components together, c) all components together with extension. The top row are colored by stress magnitudes with white lines representing the least compressive stress orientations (σ_3), where dots represent a line coming off the page or vertical. The bottom row shows 3D principal stress orientations (projected to 2D onto the map) in the Naibor Soito region along with other stress markers from Fig. 1 on the top left corner for comparison. σ_1 : greatest compressive stress; σ_2 : intermediate compressive stress.

246 5. Results

247 5.1. Topography and validation of finite-element model

248 The 2D analytical model of topographic loading (Fig. 2c) reveals stress
249 concentrations beneath the basin where sills or horizontal magma storage
250 may be favored (greatest compressive stresses are horizontal) (e.g., Maccaferri
251 et al., 2014). The corresponding numerical model (Fig. 2d) reveals similar
252 patterns of stress concentrations but with smaller stress amplitudes, due to
253 the material and domain constraints in the numerical model that do not exist
254 in the half-space of the analytical model. The similarities between Figs. 2c
255 and 2d help us validate the numerical model. In the analytical models, the
256 effect of the lower-density basin fill is minimal (black vs. dashed red line on
257 Fig. 2b are nearly coincident) hence the basin-fill density correction is not
258 considered in succeeding models.

259 The topography-driven preferred region (shallow red areas in Figs. 2c,d)
260 for melt storage beneath the basin is focused to a narrow region between the
261 Natron rift flank on the west and the Gelai volcanic edifice on the east. The
262 local topography creates a path of likely intrusion propagation (σ_1 , black
263 lines on Fig. 2) from beneath the basin towards the flanks of Gelai, similar
264 to the models of off-rift volcanism of Maccaferri et al. (2014); that is, if there
265 were magma at ~ 10 km depth between OL and NS, it might intrude to
266 shallower depths tracing along the σ_1 direction (black lines) towards Gelai
267 to the east or towards OL to the west. This topographic loading effect is
268 focused in the shallow crust, to a depth that scales with the length scale of
269 the loading/unloading. For example, in the cross-section of the analytical 2D
270 model (Fig. 2c), we see that the wide basin effect (~ 80 km wide) reaches 40

271 km depth, although the highest stresses due to the topographic depression
272 between the border fault and Gelai (~ 15 km wide) are limited to < 20 km
273 depth. On Fig. 3a, we observe that vertical strain due to topography peaks
274 at ~ 10 km depth. The lateral shift from zero of the entire topographic effect
275 profile is arbitrary and depends on the assumed height of the reference model,
276 and so for example, we can shift the profile to the right such that topographic
277 stress is zero at the base of unrifted crust (40 km thick) (e.g., Roman and
278 Jaupart, 2014).

279 *5.2. Subsurface magmatic system*

280 Both the sill and magma chamber produce a similar principal stress pat-
281 tern with near-vertical σ_1 above them, but the stresses are about an order of
282 magnitude apart, mainly due to the difference in size and magma pressure.
283 For the sill, the stress perturbation is only due to the density difference which
284 produces an upwards buoyancy push, but the stress magnitude is very small
285 compared to all the other components considered in the study (Fig. 2e). For
286 the magma chamber, in addition to the buoyancy effect, inflation adds an
287 outward-directed pressure (Fig. 2f). Vertical σ_1 above the magma bodies cre-
288 ates a possible pathway for magma to reach the surface, where diking may
289 be preferred.

290 *5.3. Thinned crust*

291 High-density mantle material replaces stretched continental crust that is
292 470 kg/m^2 lighter, imposing a downward pull on the system proportional to
293 the crustal thinning (Fig. 2g). Crustal thinning has a long wavelength (rift-
294 width scale) thereby affecting a larger region of the model, compared to the

295 high-frequency spatial variation of topography and smaller size of the magma
296 bodies. The vertical σ_1 beneath the crust and horizontal σ_1 at the bottom
297 of the crust suggest that stresses induced by crustal thinning may facilitate
298 ponding of mafic magma around the Moho (e.g., Thybo and Artemieva, 2013;
299 Maccaferri et al., 2014).

300 5.4. All components in 2D models

301 After testing each component separately, we create a model with all
302 the components – topography, magma chamber, sill. Without extension
303 (Fig. 2h), the 2D model has the recognizable stress patterns produced by the
304 thinned crust, the inflated magma chamber, and the topography. Adding
305 extension (Fig. 2i) noticeably changes the orientations of the greatest com-
306 pressive stress. At shallow depths just beneath the Naibor Soito monogenetic
307 cone complex and Gelai volcano, σ_1 is vertical, where it was near-horizontal
308 without extension, suggesting that the large-scale heterogeneities (Moho,
309 magma chamber, topography) facilitate localization of extensional stress and
310 indicating the importance of taking into consideration the regional exten-
311 sional setting in models.

312 On a one-dimensional profile of vertical stress (Fig. 3a), topography has
313 the largest effect close to the surface, showing negative stresses in our models
314 since topography at the location is below zero, hence lower than the refer-
315 ence model. The magma chamber inflation has a large effect near the magma
316 chamber but becomes negligible within a few km distance farther away (Ro-
317 man and Jaupart, 2014). The sill effect is negligible. The thinned crust
318 imposes a large stress at the bottom of the crust that gradually decreases
319 towards the surface. The horizontal stresses are smaller than the vertical

320 stresses, except for the rift-wide effect of the thinned crust which introduces
321 a horizontal stress that is comparable to the vertical stresses. However when
322 extension is added to the system, the effect of the thinned crust on the hor-
323 izontal stresses close to the surface is lessened, as also observed in Fig. 2i.

324 *5.5. Stress orientations in 3D*

325 The similarity between principal stress orientations due to topography
326 only (Fig. 4a) and all components (Fig. 4b) suggests that either 1) to-
327 pography exerts the dominant stress field in the shallow upper crust, or
328 2) the most dominant stress-generating mechanisms considered contribute
329 similarly-oriented stresses, especially away from southern Gelai and the Ngoron-
330 goro crater region. Beneath Lake Natron, σ_3 is near-vertical, and is probably
331 mainly controlled by basin topography and crustal thinning.

332 On the southern flank of Gelai, topography alone without extension pro-
333 duces a stress field with σ_3 oriented NE–SW (Fig. 4a). Adding the subsurface
334 effects, the direction of σ_3 becomes roughly N–S in southern Gelai (Fig. 4b),
335 and with extension the σ_3 orients close to NW–SE (Fig. 4c), which is con-
336 sistent with the local extensional (T-axis) direction in previous seismicity
337 studies (Reiss et al., 2021; Weinstein et al., 2017). The agreement with local
338 stress rotation observed from the recent local seismicity studies as well as
339 teleseismically-detected earthquakes since 1964 in the Global Centroid Mo-
340 ment Tensor catalog suggest that subsurface effects and extension play an
341 important role in the crustal state-of-stress.

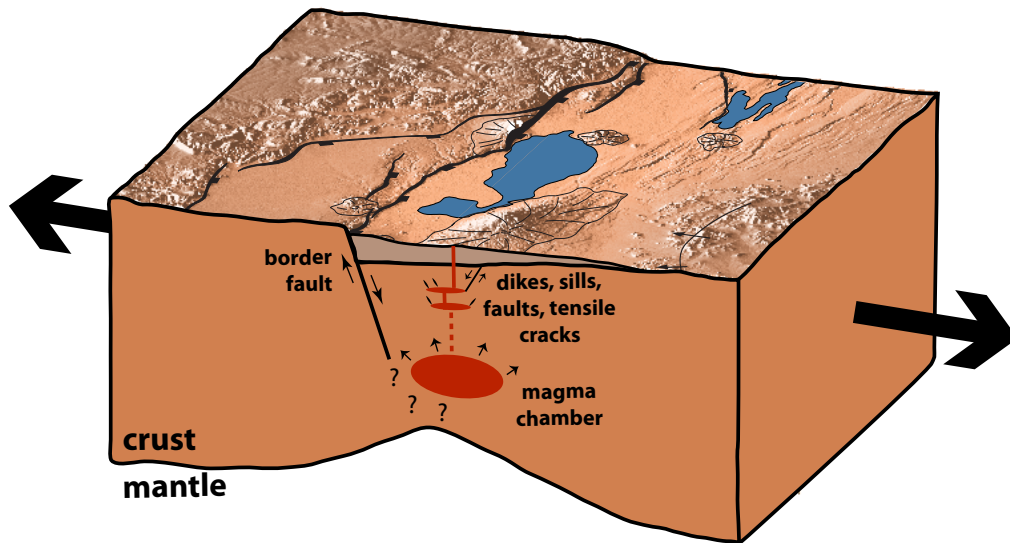


Figure 5: Schematic diagram of early-stage rifting. In addition to topographic loading, the subsurface density gradients from a thinned crust and the crustal magmatic complex also alter the stress field at depth. GNSS measurements constrain surface motions but do not represent strain at depth; they measure surface displacements. Thus, they are most sensitive to shallow stress fields and not deeper processes unless of greater magnitude, and generally are too sparse to provide information on deeper processes. Seismicity provides a better constraint on the stress field (strain field) at depth. Among the rift components considered in this study, surface topography and extension influences the shallow stress field the most (~upper crust), but at depth (~lower crust), subsurface density contrasts play a larger role. These large subsurface stresses may be overlooked when only considering surface-motion constraints but are key to a holistic three-dimensional understanding of rift tectonics.

342 **6. Discussion**

343 *6.1. Three-dimensional rift tectonics*

344 Stresses vary with depth, notably near-surface stresses differ from that
345 of the rest of the upper crust and the lower crust (Fig. 3a,b). We note that
346 topography has a significant influence on shallow vertical stresses, but subsur-
347 face components strongly affect deeper stresses and these cannot be ignored
348 (Fig. 5). Surface observations, such as GNSS and Interferometric Synthetic
349 Aperture Radar (InSAR) data, and surface lineaments, are most sensitive to
350 shallow stress fields and not deeper processes except those of greater magni-
351 tude. Hence, considering only surface constraints on plate motion paints a
352 two-dimensional picture of rifting that underestimates subsurface stress con-
353 trols on the system. In addition, at the shallow level there are sedimentary
354 basins that provide complexity through additional processes such as com-
355 paction, dewatering, and fluid overpressure. These processes influence near-
356 surface stress fields, which factor in strongly in near-surface observations, but
357 may not be representative of deeper, regional, tectonic-level processes. Most
358 of the earthquakes in the study region are in the upper crust, from 5–15 km
359 depths (Weinstein et al., 2017) (Fig. 3c), and hence provide stress constraints
360 at depth, which are highly complementary with surface observations.

361 *6.2. Rift evolution and strain transfer*

362 Border faults form very early during rift initiation and may reach a max-
363 imum length largely determined by the thickness of strong lithosphere (e.g.,
364 Jackson and Blenkinsop, 1997; Ebinger et al., 1999). New faults, called trans-
365 fer faults, may form to link initially discrete border faults and rift segments

366 into a contiguous system. Results of this study suggest that local changes in
367 the stress field caused by the topographic loads, presence of crustal magma
368 reservoirs, and evolution of crustal stretching within the system determine
369 the orientation and evolution of these transfer fault systems. The local stress
370 field influences where magmatism develops. For example, intrabasinal vol-
371 canic centers form in the high-strain accommodation zone. But it is im-
372 portant to note that once the magmatic systems are formed, they in turn
373 impose their own stresses on the region and provide a local stress control
374 where new faults and magmatic systems might develop, creating a feedback
375 loop. Hence, geometric rift segment linkage models that do not consider the
376 magmatic system (e.g., Acocella et al., 2005) may not be applicable in some
377 magmatic rift zones.

378 The Natron basin sits at a transition zone, between the Kenya rift to the
379 north which has central volcanoes along the rift axis, and the study region
380 where there are no central rift volcanoes. In the Magadi basin, intra-rift
381 faults accommodate 67–80% of regional extension, whereas in the Natron
382 basin, the border faults accommodate >69% of the extension (Muirhead
383 et al., 2016). The local changes to the stress field in the Natron area can
384 guide fluids that move up the plate and could be indicative of the mechanism
385 by which strain begins to localize intra-rift to the central depression during
386 early-stage rifting. The ~10-km crustal thinning imposes the largest stresses,
387 especially horizontal stresses, but the feature generally follows the rift axis
388 and so it cannot produce a local stress rotation (see Supplementary mate-
389 rial Fig. SM2). Instead, the local stress rotation could be produced by the
390 interplay between the volcanic edifices and the subsurface magmatic system.

391 The distribution and presence of magma within the crust contributes more to
392 stresses at depth, whereas the loading effect of volcanic edifices is largest at
393 shallow depths. The magmatic system (surface and subsurface) induces a lo-
394 cal stress field rotation that might allow extensional faults or dikes to form at
395 rift-oblique angles, which then later facilitate rift linkage between segmented
396 basins and strain migration towards intra-rift faults where magma intrusion
397 accommodates a larger percentage of extension.

398 **7. Conclusions**

399 Using physics-based numerical models, constrained by data and com-
400 pared to analytical models, we explore the effects of different rift loading
401 components on the crustal state-of-stress. Models suggest that a local stress
402 field rotation to NW–SE extension (T-axis) from the regional \sim E-W exten-
403 sion (from GNSS) can be due to the combined top loading from topography
404 and bottom loading from magma chamber inflation. We find that topo-
405 graphic loading governs shallow upper crustal stresses but subsurface com-
406 ponents (magma bodies, crustal thinning) can strongly influence lower crustal
407 stresses, which in turn may guide the location and upward path of magma
408 intrusion. The magmatic system (surface and subsurface) in an early-stage
409 rift provides a mechanism (local stress rotation) for later rift linkage between
410 segmented basins and strain migration towards the central rift zone. Studies
411 constrained only by surface observations underestimate stresses from crustal
412 thinning and lower crustal magmatic systems. Inclusion of the internal and
413 subsurface loads provides new insights into strain and magma migration dur-
414 ing rift evolution.

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427 **References**

- 428 Aagaard, B., Knepley, M., Williams, C., 2017a. PyLith User Manual, Version
429 2.2.1. Computational Infrastructure for Geodynamics (CIG). University
430 of California, Davis. URL: [https://geodynamics.org/cig/software/
431 github/pylith/v2.2.1/pylith-2.2.1_manual.pdf](https://geodynamics.org/cig/software/github/pylith/v2.2.1/pylith-2.2.1_manual.pdf).
- 432 Aagaard, B., Knepley, M., Williams, C., 2017b. PyLith v2.2.1. Compu-
433 tational Infrastructure for Geodynamics (CIG). University of California,
434 Davis. doi:<http://doi.org/10.5281/zenodo.886600>.
- 435 Aagaard, B.T., Knepley, M.G., Williams, C.A., 2013. A domain decomposi-
436 tion approach to implementing fault slip in finite-element models of quasi-

437 static and dynamic crustal deformation. *Journal of Geophysical Research:*
438 *Solid Earth* 118, 3059–3079. doi:<https://doi.org/10.1002/jgrb.50217>.

439 Acocella, V., Morvillo, P., Funicello, R., 2005. What controls relay ramps
440 and transfer faults within rift zones? insights from analogue models. *Jour-*
441 *nal of Structural Geology* 27, 397–408. doi:[https://doi.org/10.1016/](https://doi.org/10.1016/j.jsg.2004.11.006)
442 [j.jsg.2004.11.006](https://doi.org/10.1016/j.jsg.2004.11.006).

443 Allmann, B.P., Shearer, P.M., 2009. Global variations of stress drop for
444 moderate to large earthquakes. *Journal of Geophysical Research: Solid*
445 *Earth* 114, 1–22. doi:<https://doi.org/10.1029/2008JB005821>.

446 Aulbach, S., Rudnick, R.L., McDonough, W.F., 2011. Evolution of the litho-
447 spheric mantle beneath the east african rift in tanzania and its potential
448 signatures in rift magmas. *Geological Society of America Special Papers*
449 478, 105–125. doi:[http://doi.org/10.1130/2011.2478\(06\)](http://doi.org/10.1130/2011.2478(06)).

450 Baer, G., Hamiel, Y., Shamir, G., Nof, R., 2008. Evolution of a magma-driven
451 earthquake swarm and triggering of the nearby oldoinyo lengai eruption,
452 as resolved by insar, ground observations and elastic modeling, east african
453 rift, 2007. *Earth and Planetary Science Letters* 272, 339–352. doi:[https://doi.org/10.1016/j](https://doi.org/10.1016/j.epsl.2008.04.052)
454 [.epsl.2008.04.052](https://doi.org/10.1016/j.epsl.2008.04.052).

455 Biggs, J., Amelung, F., Gourmelen, N., Dixon, T.H., Kim, S.W., 2009. Insar
456 observations of 2007 tanzania rifting episode reveal mixed fault and dyke
457 extension in an immature continental rift. *Geophysical Journal Interna-*
458 *tional* 179, 549–558. doi:[https://doi.org/10.1111/j.1365-246X.2009.](https://doi.org/10.1111/j.1365-246X.2009.04262.x)
459 [04262.x](https://doi.org/10.1111/j.1365-246X.2009.04262.x).

- 460 Birt, C., Maguire, P., Khan, M., Thybo, H., Keller, G.R., Patel, J., 1997. The
461 influence of pre-existing structures on the evolution of the southern kenya
462 rift valley—evidence from seismic and gravity studies. *Tectonophysics* 278,
463 211–242. doi:[https://doi.org/10.1016/S0040-1951\(97\)00105-4](https://doi.org/10.1016/S0040-1951(97)00105-4).
- 464 Brocher, T.M., 2005. Empirical relations between elastic wavespeeds and
465 density in the earth’s crust. *Bulletin of the seismological Society of America*
466 95, 2081–2092. doi:<https://doi.org/10.1785/0120050077>.
- 467 Calais, E., d’Oreye, N., Albaric, J., Deschamps, A., Delvaux, D., Déverchère,
468 J., Ebinger, C., Ferdinand, R.W., Kervyn, F., Macheyeke, A.S., et al.,
469 2008. Strain accommodation by slow slip and dyking in a youthful conti-
470 nental rift, east africa. *Nature* 456, 783. doi:[https://doi.org/10.1038/](https://doi.org/10.1038/nature07478)
471 [nature07478](https://doi.org/10.1038/nature07478).
- 472 Cashman, K.V., Sparks, R.S.J., Blundy, J.D., 2017. Vertically extensive and
473 unstable magmatic systems: a unified view of igneous processes. *Science*
474 355, eaag3055. doi:<http://doi.org/10.1126/science.aag3055>.
- 475 Corti, G., Bonini, M., Sokoutis, D., Innocenti, F., Manetti, P., Cloetingh, S.,
476 Mulugeta, G., 2004. Continental rift architecture and patterns of magma
477 migration: A dynamic analysis based on centrifuge models. *Tectonics* 23.
478 doi:<https://doi.org/10.1029/2003TC001561>.
- 479 Dahm, T., 2000. Numerical simulations of the propagation path and the
480 arrest of fluid-filled fractures in the earth. *Geophysical Journal Interna-*
481 *tional* 141, 623–638. doi:[https://doi.org/10.1046/j.1365-246x.2000.](https://doi.org/10.1046/j.1365-246x.2000.00102.x)
482 [00102.x](https://doi.org/10.1046/j.1365-246x.2000.00102.x).

- 483 Dawson, J., 1992. Neogene tectonics and volcanicity in the north
484 tanzania sector of the gregory rift valley: contrasts with the kenya
485 sector. *Tectonophysics* 204, 81–92. doi:[https://doi.org/10.1016/
486 0040-1951\(92\)90271-7](https://doi.org/10.1016/0040-1951(92)90271-7).
- 487 Ebinger, C., Jackson, J., Foster, A., Hayward, N., 1999. Extensional basin ge-
488 ometry and the elastic lithosphere. *Philosophical Transactions of the Royal*
489 *Society of London. Series A: Mathematical, Physical and Engineering Sci-*
490 *ences* 357, 741–765. doi:<https://doi.org/10.1098/rsta.1999.0351>.
- 491 Ebinger, C.J., 1989. Geometric and kinematic development of border faults
492 and accommodation zones, kivu-rusizi rift, africa. *Tectonics* 8, 117–133.
493 doi:<https://doi.org/10.1029/TC008i001p00117>.
- 494 Foster, A., Ebinger, C., Mbede, E., Rex, D., 1997. Tectonic development
495 of the northern tanzanian sector of the east african rift system: *Journal*
496 *of the geological society of london*, v. 154. doi 10, 689–700. doi:[https:
497 //doi.org/10.1144/gsjgs.154.4.0689](https://doi.org/10.1144/gsjgs.154.4.0689).
- 498 Guerri, M., Cammarano, F., Connolly, J.A., 2015. Effects of chemical
499 composition, water and temperature on physical properties of continental
500 crust. *Geochemistry, Geophysics, Geosystems* 16, 2431–2449. doi:[https:
501 //doi.org/10.1002/2015GC005819](https://doi.org/10.1002/2015GC005819).
- 502 Jackson, J., Blenkinsop, T., 1997. The bilila-mtakataka fault in malaŵi:
503 An active, 100-km long, normal fault segment in thick seismogenic crust.
504 *Tectonics* 16, 137–150. doi:<https://doi.org/10.1029/96TC02494>.

- 505 Jarvis, A., Reuter, H.I., Nelson, A., Guevara, E., 2008. Hole-filled srtm for
506 the globe version 4, available from the cgiar-csi srtm 90m database. URL:
507 <http://srtm.csi.cgiar.org/>.
- 508 Karlstrom, L., Dufek, J., Manga, M., 2009. Organization of volcanic
509 plumbing through magmatic lensing by magma chambers and volcanic
510 loads. *Journal of Geophysical Research: Solid Earth* 114. doi:<https://doi.org/10.1029/2009JB006339>.
- 512 Kervyn, M., Ernst, G., van Wyk de Vries, B., Mathieu, L., Jacobs, P., 2009.
513 Volcano load control on dyke propagation and vent distribution: Insights
514 from analogue modeling. *Journal of Geophysical Research: Solid Earth*
515 114. doi:<https://doi.org/10.1029/2008JB005653>.
- 516 King, R., Floyd, M., Reilinger, R., Bendick, R., 2019. Gps velocity
517 field (mit 2019.0) for the east african rift system generated by king et
518 al. URL: <http://doi.iedadata.org/324785>, doi:[http://doi.org/10.](http://doi.org/10.1594/IEDA/324785)
519 [1594/IEDA/324785](http://doi.org/10.1594/IEDA/324785).
- 520 Larsen, H., 1988. A multiple and propagating rift model for the ne atlantic.
521 *Geological Society, London, Special Publications* 39, 157–158. doi:[http://doi.org/10.](http://doi.org/10.1144/GSL.SP.1988.039.01.15)
522 [1144/GSL.SP.1988.039.01.15](http://doi.org/10.1144/GSL.SP.1988.039.01.15).
- 523 Le Gall, B., Nonnotte, P., Rolet, J., Benoit, M., Guillou, H., Mousseau-
524 Nonnotte, M., Albaric, J., Déverchère, J., 2008. Rift propagation at craton
525 margin.: Distribution of faulting and volcanism in the north tanzanian
526 divergence (east africa) during neogene times. *Tectonophysics* 448, 1–19.
527 doi:<https://doi.org/10.1016/j.tecto.2007.11.005>.

- 528 Maccaferri, F., Rivalta, E., Keir, D., Acocella, V., 2014. Off-rift volcanism
529 in rift zones determined by crustal unloading. *Nature Geoscience* 7, 297.
530 doi:<https://doi.org/10.1038/ngeo2110>.
- 531 Magee, C., Muirhead, J.D., Karvelas, A., Holford, S.P., Jackson, C.A., Bas-
532 tow, I.D., Schofield, N., Stevenson, C.T., McLean, C., McCarthy, W.,
533 et al., 2016. Lateral magma flow in mafic sill complexes. *Geosphere* 12,
534 809–841. doi:<https://doi.org/10.1130/GES01256.1>.
- 535 Mana, S., Furman, T., Turrin, B.D., Feigenson, M.D., Swisher III, C.C.,
536 2015. Magmatic activity across the east african north tanzanian divergence
537 zone. *Journal of the Geological Society* 172, 368–389. doi:<https://doi.org/10.1144/jgs2014-072>.
- 539 Mansur, A.T., Manya, S., Timpa, S., Rudnick, R.L., 2014. Granulite-facies
540 xenoliths in rift basalts of northern tanzania: age, composition and origin
541 of archean lower crust. *Journal of Petrology* 55, 1243–1286. doi:<https://doi.org/10.1093/petrology/egu024>.
- 543 McHenry, L.J., Luque, L., Gómez, J.Á., Diez-Martín, F., 2011. Promise and
544 pitfalls for characterizing and correlating the zeolitically altered tephra of
545 the pleistocene peninj group, tanzania. *Quaternary Research* 75, 708–720.
546 doi:<https://doi.org/10.1016/j.yqres.2010.11.008>.
- 547 Muirhead, J., Kattenhorn, S., Lee, H., Mana, S., Turrin, B., Fischer, T.,
548 Kianji, G., Dindi, E., Stamps, D., 2016. Evolution of upper crustal faulting
549 assisted by magmatic volatile release during early-stage continental rift

550 development in the east african rift. *Geosphere* 12, 1670–1700. doi:<https://doi.org/10.1130/GES01375.1>.
551

552 Muirhead, J.D., Kattenhorn, S.A., Le Corvec, N., 2015. Varying styles of
553 magmatic strain accommodation across the east african rift. *Geochemistry,*
554 *Geophysics, Geosystems* 16, 2775–2795. doi:[https://doi.org/10.1002/](https://doi.org/10.1002/2015GC005918)
555 [2015GC005918](https://doi.org/10.1002/2015GC005918).

556 Oliva, S., Ebinger, C., Wauthier, C., Muirhead, J., Roecker, S., Rivalta, E.,
557 Heimann, S., 2019. Insights into fault-magma interactions in an early-stage
558 continental rift from source mechanisms and correlated volcano-tectonic
559 earthquakes. *Geophysical Research Letters* 46, 2065–2074. doi:<https://doi.org/10.1029/2018GL080866>.
560

561 Plasman, M., Tiberi, C., Ebinger, C., Gautier, S., Albaric, J., Peyrat, S.,
562 Deverchere, J., Le Gall, B., Tarits, P., Roecker, S., et al., 2017. Litho-
563 spheric low-velocity zones associated with a magmatic segment of the tan-
564 zanian rift, east africa. *Geophysical Journal International* 210, 465–481.
565 doi:<https://doi.org/10.1093/gji/ggx177>.

566 Pollard, D., 1987. Elementary fracture mechanics applied to the structural
567 interpretation of dykes, in: Halls, H.C., Fahrig, W.F. (Eds.), *Mafic Dyke*
568 *Swarms*. Geol. Assoc. of Can.. volume 34, pp. 112–128.

569 Reiss, M., Muirhead, J., Laizer, A., Link, F., Kazimoto, E., Ebinger, C.,
570 Rümpler, G., 2021. The impact of complex volcanic plumbing on the
571 nature of seismicity in the developing magmatic natron rift, tanzania.

572 front. Earth Sci 8, 609805. doi:[https://doi.org/10.3389/feart.2020.](https://doi.org/10.3389/feart.2020.609805)
573 609805.

574 Rivalta, E., Corbi, F., Passarelli, L., Acocella, V., Davis, T., Di Vito, M.A.,
575 2019. Stress inversions to forecast magma pathways and eruptive vent
576 location. Science advances 5, eaau9784. doi:[https://doi.org/10.1126/](https://doi.org/10.1126/sciadv.aau9784)
577 [sciadv.aau9784](https://doi.org/10.1126/sciadv.aau9784).

578 Roecker, S., Ebinger, C., Tiberi, C., Mulibo, G., Ferdinand-Wambura, R.,
579 Mtelela, K., Kianji, G., Muzuka, A., Gautier, S., Albaric, J., et al., 2017.
580 Subsurface images of the eastern rift, africa, from the joint inversion of
581 body waves, surface waves and gravity: investigating the role of fluids
582 in early-stage continental rifting. Geophysical Journal International 210,
583 931–950. doi:<https://doi.org/10.1093/gji/ggx220>.

584 Roman, A., Jaupart, C., 2014. The impact of a volcanic edifice on intru-
585 sive and eruptive activity. Earth and Planetary Science Letters 408, 1–8.
586 doi:<https://doi.org/10.1016/j.epsl.2014.09.016>.

587 Saria, E., Calais, E., Stamps, D., Delvaux, D., Hartnady, C., 2014. Present-
588 day kinematics of the east african rift. Journal of Geophysical Re-
589 search: Solid Earth 119, 3584–3600. doi:[https://doi.org/10.1002/](https://doi.org/10.1002/2013JB010901)
590 [2013JB010901](https://doi.org/10.1002/2013JB010901).

591 Sibson, R.H., 1985. A note on fault reactivation. Journal of Structural Geol-
592 ogy 7, 751–754. doi:[https://doi.org/10.1016/0191-8141\(85\)90150-6](https://doi.org/10.1016/0191-8141(85)90150-6).

593 Smith, M., Mosley, P., 1993. Crustal heterogeneity and basement influence

594 on the development of the kenya rift, east africa. *Tectonics* 12, 591–606.
595 doi:<https://doi.org/10.1029/92TC01710>.

596 Thybo, H., Artemieva, I.M., 2013. Moho and magmatic underplating in
597 continental lithosphere. *Tectonophysics* 609, 605–619. doi:<https://doi.org/10.1016/j.tecto.2013.05.032>.

599 Tibaldi, A., Bonali, F.L., Pasquaré Mariotto, F., Corti, N., Russo, E., Einars-
600 son, P., Hjartardóttir, Á.R., 2020. Rifting kinematics produced by mag-
601 matic and tectonic stresses in the north volcanic zone of iceland. *Frontiers*
602 *in Earth Science* 8, 174. doi:[https://doi.org/10.3389/feart.2020.](https://doi.org/10.3389/feart.2020.00174)
603 00174.

604 Tiberi, C., Gautier, S., Ebinger, C., Roecker, S., Plasman, M., Albaric, J.,
605 Déverchère, J., Peyrat, S., Perrot, J., Wambura, R.F., et al., 2018. Litho-
606 spheric modification by extension and magmatism at the craton-orogenic
607 boundary: North tanzania divergence, east africa. *Geophysical Journal In-*
608 *ternational* 216, 1693–1710. doi:<https://doi.org/10.1093/gji/ggy521>.

609 Tugume, F., Nyblade, A., Julià, J., 2012. Moho depths and poisson’s ratios
610 of precambrian crust in east africa: evidence for similarities in archean
611 and proterozoic crustal structure. *Earth and Planetary Science Letters*
612 355, 73–81. doi:<https://doi.org/10.1016/j.epsl.2012.08.041>.

613 Wadge, G., Biggs, J., Lloyd, R., Kendall, J.M., 2016. Historical volcanism
614 and the state of stress in the east african rift system. *Frontiers in Earth*
615 *Science* 4, 86. doi:<https://doi.org/10.3389/feart.2016.00086>.

616 Weinstein, A., Oliva, S.J., Ebinger, C., Roecker, S., Tiberi, C., Aman, M.,
617 Lambert, C., Witkin, E., Albaric, J., Gautier, S., et al., 2017. Fault-magma
618 interactions during early continental rifting: Seismicity of the magadi-n
619 atron-manyara basins, Africa. *Geochemistry, Geophysics, Geosystems* 18,
620 3662–3686. doi:<https://doi.org/10.1002/2017GC007027>.

Supplementary material of

State of stress and stress rotations: Quantifying the role of surface topography and subsurface density contrasts in magmatic rift zones (Eastern Rift, Africa)

S. J. Oliva, C. J. Ebinger, E. Rivalta, C. A. Williams, C. Wauthier, C. Currie

Contents of this file:

- SM1 Figure: Gravitational stresses due to reference model
- SM2 Figure: Map view of stresses due to crustal thinning
- SM3 Trellis input for 2D and 3D model geometry and meshing
- SM4 General PyLith parameters (pylithapp.cfg)
- SM5 Post-processing calculations

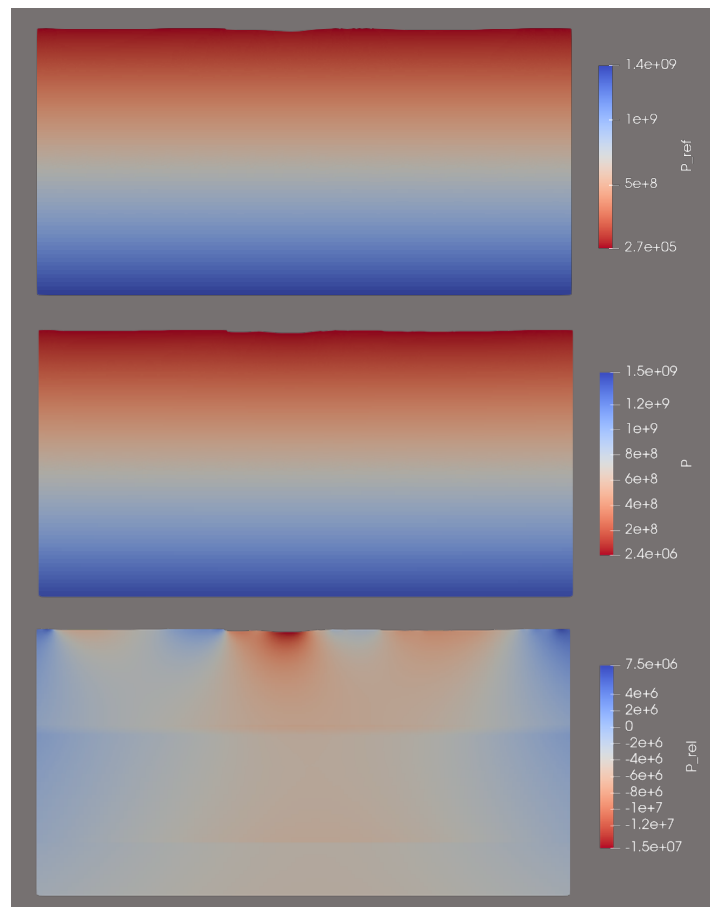


Figure SM1. The top figure (P_{ref}) shows the gravitational stresses of the reference model, a uniform-thickness three-layer model with no magmatic systems. The middle figure (P) shows the stresses due to topography – note that the gravitational stresses dominate and obscure any stress patterns due to topography. Hence, we subtract the assumed initial stress state corresponding to the reference model (P_{ref}) from our models of different components (P) to obtain the stress

patterns that we interpret in the figure (P_rel, bottom figure). All three figures have color bars scaled to the data range. The bottom figure is equivalent to Fig. 2d, but in the main text all color bars are centered at zero.

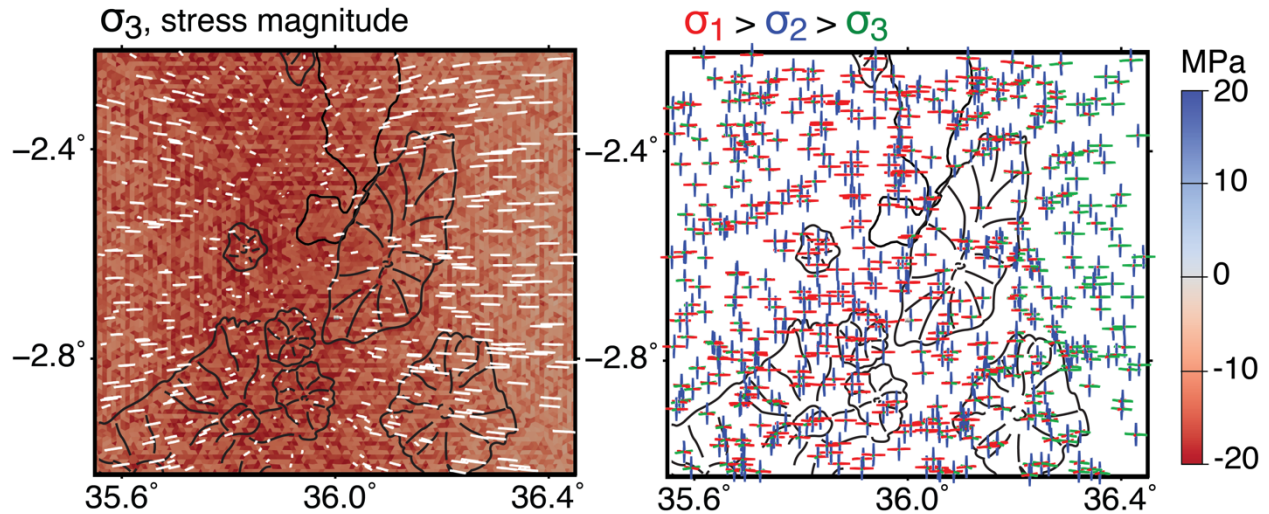


Figure SM2. 3D modeling of crustal thinning: Left panel is colored by stress magnitude with white lines representing the least compressive stress orientations (sigma 3), where dots represent a line coming off the page or vertical. Right panel shows the 3D principal stress orientations, projected to 2D onto the map. Note that sigma 1 is oriented E-W, sigma 2 is N-S, and sigma 3 is near-vertical along the rift basin. With the principal stress orientations generally aligned rift-parallel, rift-perpendicular, and vertical, we conclude that the crustal thinning cannot produce the map-view 60-degree local stress rotation observed from focal mechanisms.

SM3. Geometry and Mesh

Below we include the Trelis codes (.jou) to create the geometry and build the mesh of the complete model that includes all components (topography, moho, sill, magma chamber) in 2D and 3D.

```
# 2D Geometry
# -----
# -*- Python -*- (syntax highlighting)
# -----
reset
${Units('si')}

# -----
# Import Topo and Moho curves
# -----
import Acis "curve_moho.sat"
import Acis "curve_topo_272_100m_detrendzero.sat"

# -----
# Create surface using vertices
# -----

${domainW=300.0*km}
${domainH=110.0*km}
```

```

create vertex {-0.5*domainW} {-domainH+10.0*km} 0.0
${idVsbegin=Id("vertex")}
create vertex {-0.5*domainW} {10.0*km} 0.0
create vertex {+0.5*domainW} {10.0*km} 0.0
create vertex {+0.5*domainW} {-domainH+10.0*km} 0.0
${idVsend=Id("vertex")}
create surface vertex {idVsbegin} to {idVsend}

# -----
# Create material interface surfaces
# -----

# Magma chamber
create vertex {5.5*km} {-20.0*km} 0.0
${idVCirc=Id("vertex")}
create curve arc center vertex {idVCirc} radius {10.0*km} full
curve 7 name "chamber"
surface chamber scale x 1 y 0.5 about {5.5*km} {-20*km}
imprint surface 1 with curve chamber

# Constant thickness crust
${ucrustD=18.0*km}
create vertex {-0.5*domainW} {-ucrustD} 0.0
${idVM1=Id("vertex")}
create vertex {+0.5*domainW} {-ucrustD} 0.0
${idVM2=Id("vertex")}
create curve vertex {idVM1} vertex {idVM2} on surface 2
imprint surface 2 with curve 9

# Split across Moho curve
split curve 17 crossing curve moho
split curve 18 crossing curve moho
imprint surface 5 with curve moho

# Split across topo curve
split curve 11 crossing curve topo
split curve 12 crossing curve topo
imprint surface 4 with curve topo
delete surface 8

# Sill complex
create vertex {0.0*km} {-7.0*km} 0.0
${idVertex1=Id("vertex")}
create vertex {2.5*km} {-6.75*km} 0.0
${idVertex2=Id("vertex")}
create vertex {2.5*km} {-7.0*km} 0.0
${idVertex3=Id("vertex")}
create curve vertex {idVertex1} vertex {idVertex2} vertex {idVertex3} ellipse first
angle 0 last angle 360
curve 29 name "sill"
imprint surface 9 with curve sill

surface 11 name "s_ucrust"
surface 6 name "s_lcrust"
surface 7 name "s_mantle"
surface 10 name "s_sill"
surface 3 name "s_chamber"

# -----
# Name geometric entities for use in meshing.
# -----

curve 28 name "c_topo"
#curve 6 name "c_conrad"
#curve 15 name "c_moho"

```



```
curve 6 name "c_bottom"

curve 30 name "c_sill"
curve 14 name "c_chambertop"
curve 15 name "c_chamberbottom"
curve 16 name "c_chamberright"

curve 24 name "c_ucrust_xneg"
curve 27 name "c_ucrust_xpos"
curve 22 name "c_lcrust_xneg"
curve 19 name "c_lcrust_xpos"
curve 21 name "c_mantle_xneg"
curve 20 name "c_mantle_xpos"

delete vertex all

# End of file
```

```
# 2D Mesh
```

```
# -----
# -*- Python -*- (syntax highlighting)
# -----

# -----
# Set discretization size
# -----
surface all size {1.0*km}
surface s_sill size {0.2*km}
surface s_chamber size {0.5*km}
surface s_ucrust size {0.5*km}
surface s_lcrust size {0.5*km}
surface s_mantle size {0.8*km}

# -----
# Generate the mesh
# -----
delete mesh
surface all scheme pave
mesh surface all

# -----
# Smooth mesh to improve quality
# -----
surface all smooth scheme condition number beta 1.3 cpu 10
smooth surface all

# -----
# Create blocks for materials
# -----
block 1 surface s_ucrust
block 1 name "ucrust"

block 2 surface s_lcrust
block 2 name "lcrust"

block 3 surface s_mantle
block 3 name "mantle"

block 4 surface s_sill
block 4 name "sill"

block 5 surface s_chamber
block 5 name "chamber"
```

```

# -----
# Create nodeset for +x edge
# -----
group "boundary_xpos" add node in curve c_ucrust_xpos
group "boundary_xpos" add node in curve c_lcrust_xpos
group "boundary_xpos" add node in curve c_mantle_xpos
nodeset 20 group boundary_xpos
nodeset 20 name "boundary xpos"

# -----
# Create nodeset for -x edge
# -----
group "boundary_xneg" add node in curve c_ucrust_xneg
group "boundary_xneg" add node in curve c_lcrust_xneg
group "boundary_xneg" add node in curve c_mantle_xneg
nodeset 21 group boundary_xneg
nodeset 21 name "boundary xneg"

# -----
# Create nodeset for +y edge
# -----
group "boundary_ypos" add node in curve c_topo
nodeset 22 group boundary_ypos
nodeset 22 name "boundary ypos"

# -----
# Create nodeset for -y edge
# -----
group "boundary_yneg" add node in curve c_bottom
nodeset 23 group boundary_yneg
nodeset 23 name "boundary yneg"

# -----
# Create nodeset for sill complex
# -----
group "boundary_sill" add node in curve c_sill
nodeset 24 group boundary_sill
nodeset 24 name "boundary sill"

# -----
# Create nodeset for magma chamber
# -----
group "boundary_chamber" add node in curve c_chambertop
group "boundary_chamber" add node in curve c_chamberbottom
group "boundary_chamber" add node in curve c_chamberright
nodeset 25 group boundary_chamber
nodeset 25 name "boundary chamber"

# -----
# Export exodus file
# -----
export mesh "mesh_2d_mod6.exo" dimension 2 overwrite

# End of file

```

```

# 3D Geometry
# -----
# -*- Python -*- (syntax highlighting)
# -----
reset
${Units('si')}

# -----
# Create block

```

```

# -----
# Block is 250 km x 500 km x 62 km
${blockLength=200.0*km}
${blockWidth=190.0*km}
${blockHeight=110.0*km}

brick x {blockLength} y {blockWidth} z {blockHeight}
${idVol=Id("volume")}

# Translate block so top can be chopped with topography and so faults
# intersect upper surface rather than side.
${moveX=0.0*km}
${moveY=13.0*km} #topo centered at 2.72 S
#volume {idVol} move x {moveX} y {moveY} z {0.0*km}
volume {idVol} move x {moveX} y {moveY} z {-0.5*blockHeight+10.0*km}

# -----
# Import topography.
# -----
import Acis "srtm_1km_detrend/crafti_topo_detrend.sat"

# Scale topo, 1 km = 1000 m
surface 7 scale 1000
#save
# -----
# Webcut block with topography and delete volume lying above topography.
# -----
webcut volume {idVol} with sheet surface 7
delete volume 3
delete body 2
#save
# -----
# Divide upper crust, lower crust, and mantle
# -----
create planar surface with plane zplane offset {-18.0*km}
webcut volume 1 with sheet surface 18

# Magma chamber
create sphere radius {5*km}
${idSphere=Id("volume")}
volume {idSphere} move x {5.5*km} y 0 z {-20*km}
volume {idSphere} scale x 2 y 2 z 1 about {5.5*km} 0 {-20*km}

subtract volume 6 from volume 1 imprint keep
subtract volume 6 from volume 5 imprint keep
delete volume 1 5

# Sill
create sphere radius {0.5*km}
${idSphere=Id("volume")}
volume {idSphere} move x {2.5*km} y 0 z {-7*km}
volume {idSphere} scale x 10 y 10 z 1 about {2.5*km} 0 {-7*km}

subtract volume 9 from volume 7 imprint keep
delete volume 7

# Moho
import Acis "curve_moho_3d.sat"
curve moho move x 0 y {-90*km} z 0
webcut volume 8 sweep curve moho vector 0 1 0 distance {220*km}

delete body 4
delete curve moho

```

```

imprint all with volume all
merge all
delete vertex all

# -----
# Name geometric entities for use in meshing.
# -----

surface 49 name "topo"
surface 38 name "conrad"
surface 53 name "moho"
surface 43 name "sbottom"
surface 45 name "chambertop"
surface 37 name "chamberbottom"
surface 44 name "sill_surf"

volume 10 name "ucrust"
volume 8 name "lcrust"
volume 11 name "mantle"
volume 6 name "chamber"
volume 9 name "sill"

surface 50 name "ucrust_ypos"
surface 51 name "ucrust_xneg"
surface 47 name "ucrust_yneg"
surface 48 name "ucrust_xpos"

surface 56 name "lcrust_yneg"
surface 57 name "lcrust_xpos"
surface 54 name "lcrust_ypos"
surface 55 name "lcrust_xneg"

surface 61 name "mantle_yneg"
surface 62 name "mantle_xneg"
surface 59 name "mantle_ypos"
surface 60 name "mantle_xpos"

# End of file

# 3D Mesh
# -----
# -*- Python -*- (syntax highlighting)
# -----

# -----
# Set discretization size
# -----
surface topo size {1.0*km}
surface conrad size {2.5*km}
surface moho size {5.0*km}
surface sbottom size {10.0*km}
surface chambertop size {1.0*km}
surface chamberbottom size {1.0*km}
surface sill_surf size {0.2*km}

# Sizing functions for faces surrounding upper crust.
surface ucrust_xneg sizing function type bias start curve 95 finish curve 71
surface ucrust_xpos sizing function type bias start curve 93 finish curve 73
surface ucrust_yneg sizing function type bias start curve 90 finish curve 72
surface ucrust_ypos sizing function type bias start curve 94 finish curve 70

# Sizing functions for faces surrounding lower crust.
surface lcrust_xneg sizing function type bias start curve 71 finish curve 100
surface lcrust_xpos sizing function type bias start curve 73 finish curve 102

```

```

surface lcrust_yneg sizing function type bias start curve 72 finish curve 99
surface lcrust_ypos sizing function type bias start curve 70 finish curve 101

# Sizing functions for faces surrounding mantle.
surface mantle_xneg sizing function type bias start curve 100 finish curve 78
surface mantle_xpos sizing function type bias start curve 102 finish curve 81
surface mantle_yneg sizing function type bias start curve 99 finish curve 75
surface mantle_ypos sizing function type bias start curve 101 finish curve 80

# -----
# Generate the mesh
# -----
delete mesh
volume all scheme tetmesh
mesh surface topo
mesh surface all except topo
mesh volume all

# -----
# Smooth mesh to improve quality
# -----
volume all smooth scheme condition number beta 2.0 cpu 3
smooth volume all

# -----
# Create blocks for materials
# -----
block 1 volume ucrust
block 1 name "ucrust"

block 2 volume lcrust
block 2 name "lcrust"

block 3 volume mantle
block 3 name "mantle"

block 4 volume sill
block 4 name "sill"

block 5 volume chamber
block 5 name "chamber"

# -----
# Create nodeset for +x edge
# -----
group "boundary_xpos" add node in surface ucrust_xpos
group "boundary_xpos" add node in surface lcrust_xpos
group "boundary_xpos" add node in surface mantle_xpos
nodeset 30 group boundary_xpos
nodeset 30 name "boundary_xpos"

# -----
# Create nodeset for -x edge
# -----
group "boundary_xneg" add node in surface ucrust_xneg
group "boundary_xneg" add node in surface lcrust_xneg
group "boundary_xneg" add node in surface mantle_xneg
nodeset 31 group boundary_xneg
nodeset 31 name "boundary_xneg"

# -----
# Create nodeset for +y edge
# -----
group "boundary_ypos" add node in surface ucrust_ypos
group "boundary_ypos" add node in surface lcrust_ypos

```

```

group "boundary_ypos" add node in surface mantle_ypos
nodeset 32 group boundary_ypos
nodeset 32 name "boundary_ypos"

# -----
# Create nodeset for -y edge
# -----
group "boundary_yneg" add node in surface ucrust_yneg
group "boundary_yneg" add node in surface lcrust_yneg
group "boundary_yneg" add node in surface mantle_yneg
nodeset 33 group boundary_yneg
nodeset 33 name "boundary_yneg"

# -----
# Create nodeset for +z edge
# -----
group "boundary_zpos" add node in surface topo
nodeset 34 group boundary_zpos
nodeset 34 name "boundary_zpos"

# -----
# Create nodeset for -z edge
# -----
group "boundary_zneg" add node in surface sbottom
nodeset 35 group boundary_zneg
nodeset 35 name "boundary_zneg"

# -----
# Create nodeset for sill
# -----
group "boundary_sill" add node in surface sill_surf
nodeset 36 group boundary_sill
nodeset 36 name "boundary_sill"

# End of file

```

SM4. General PyLith parameters

Below is the pylithapp.cfg file used in 2D and 3D models.

```

# PyLith parameters for 2D models

# The settings in this file (pylithapp.cfg) will be read automatically
# by pylith, as long as the file is placed in the run directory.

# The settings in this file will override any settings in:
# PREFIX/etc/pylithapp.cfg
# $HOME/.pyre/pylithapp/pylithapp.cfg

# The settings in this file will be overridden by any .cfg file given
# on the command line or by any command line settings.

[pylithapp]

# -----
# journal
# -----
# The settings below turn on journal info for the specified components.
# If you want less output to stdout, you can turn these off.
[pylithapp.journal.info]
timedependent = 1

```

```

implicit = 1
petsc = 1
solverlinear = 1
meshiocubit = 1
homogeneous = 1
elasticityimplicit = 1
fiatlagrange = 1
pylithapp = 1
materials = 1

# -----
# mesh_generator
# -----
# The settings below control the mesh generation (importing mesh info).
# Turn on debugging output for mesh generation.
[pylithapp.mesh_generator]
debug = 1
reader = pylith.meshio.MeshIOcubit

# This component specification means we are using PyLith ASCII format,
# and we then specify the filename and number of space dimensions for
# the mesh.
[pylithapp.mesh_generator.reader]
filename = mesh/mesh_2d_mod1.exo
coordsys.space_dim = 2

# -----
# problem
# -----
# Specify the problem settings.
# This is a time-dependent problem, so we select this as our problem type.
# We select a total time of 0 sec, and a time step size of 1 sec, so we
# are performing a single time step.
# The spatial dimension for this problem is 2.
# For an implicit formulation (using implicit.cfg), we will perform 1
# implicit time step from t = -1.0 to t = 0.0 (elastic solution step).
[pylithapp.timedependent]
dimension = 2
formulation = pylith.problems.Implicit
formulation.solver = pylith.problems.SolverNonlinear

[pylithapp.timedependent.formulation.time_step]
total_time = 1.0*year
dt = 1.0*year

# -----
# materials
# -----
# Specify the material information for the problem.
# The material type is isotropic elastic formulated for plane strain.

[pylithapp.timedependent]
materials = [ucrust,lcrust,mantle]

[pylithapp.timedependent.materials]
ucrust = pylith.materials.ElasticPlaneStrain
lcrust = pylith.materials.ElasticPlaneStrain
mantle = pylith.materials.ElasticPlaneStrain

#[pylithapp.timedependent.materials.material]

# Upper Crust
[pylithapp.timedependent.materials.ucrust]
label = Elastic ucrust

```

```

# The id corresponds to the block number from CUBIT.
id = 1

db_properties = spatialdata.spatialdb.SimpleDB
db_properties.label = Elastic ucrust properties
db_properties.iohandler.filename = spatialdb/mat_ucrust_elastic.spatialdb
db_properties.query_type = nearest

# We are doing 2D quadrature for a triangle.
quadrature.cell = pylith.feassemble.FIATLagrange
quadrature.cell.dimension = 2

# Lower crust
[pylithapp.timedependent.materials.lcrust]
label = Elastic lower crust

# The id corresponds to the block number from CUBIT.
id = 2

db_properties = spatialdata.spatialdb.SimpleDB
db_properties.label = Elastic lcrust properties
db_properties.iohandler.filename = spatialdb/mat_lcrust_elastic.spatialdb
#db_properties.iohandler.filename = spatialdb/mat_lcrust_Maxwell.spatialdb
db_properties.query_type = nearest

# We are doing 2D quadrature for a triangle.
quadrature.cell = pylith.feassemble.FIATLagrange
quadrature.cell.dimension = 2

# Mantle
[pylithapp.timedependent.materials.mantle]
label = Elastic mantle

# The id corresponds to the block number from CUBIT.
id = 3

db_properties = spatialdata.spatialdb.SimpleDB
db_properties.label = Elastic mantle properties
db_properties.iohandler.filename = spatialdb/mat_mantle_elastic.spatialdb
#db_properties.iohandler.filename = spatialdb/mat_mantle_Maxwell.spatialdb
db_properties.query_type = nearest

# We are doing 2D quadrature for a triangle.
quadrature.cell = pylith.feassemble.FIATLagrange
quadrature.cell.dimension = 2

# -----
# output
# -----
# Names of output files are set in stepXX.cfg. We consolidate all of the
# output settings that are common to all of the simulations here.

[pylithapp.timedependent.formulation]
# Set the output to an array of 2 output managers.
# We will output the solution over the domain and the ground surface.
output = [domain]

# Domain
[pylithapp.problem.formulation.output.domain]
vertex_data_fields = [displacement]
writer = pylith.meshio.DataWriterHDF5

# Materials
[pylithapp.timedependent.materials.ucrust.output]
writer = pylith.meshio.DataWriterHDF5

```



```

cell_data_fields = [stress,cauchy_stress,total_strain]

# Comment this out for if rerunning Step 3 to get initial conditions for Step 8.
cell_filter = pylith.meshio.CellFilterAvg

[pylithapp.timedependent.materials.lcrust.output]
writer = pylith.meshio.DataWriterHDF5
cell_data_fields = [stress,cauchy_stress,total_strain]

# Comment this out for if rerunning Step 3 to get initial conditions for Step 8.
cell_filter = pylith.meshio.CellFilterAvg

[pylithapp.timedependent.materials.mantle.output]
writer = pylith.meshio.DataWriterHDF5
cell_data_fields = [stress,cauchy_stress,total_strain]

# Comment this out for if rerunning Step 3 to get initial conditions for Step 8.
cell_filter = pylith.meshio.CellFilterAvg

# -----
# PETSc
# -----
# We are using all of the default settings for PETSc except for
# specifying the block Jacobi preconditioner.  Additional PETSc
# command-line arguments may be found in the PETSc documentation.
[pylithapp.petsc]

# Preconditioner settings.
pc_type = bjacobi

# Convergence parameters.
ksp_rtol = 1.0e-50
ksp_atol = 1.0e-12

# Linear solver monitoring options.
ksp_monitor = true
ksp_converged_reason = true
ksp_error_if_not_converged = true

# Nonlinear solver monitoring options.
snes_rtol = 1.0e-50
snes_atol = 1.0e-10
snes_max_it = 100
snes_monitor = true
snes_converged_reason = true
snes_linesearch_monitor = true
snes_error_if_not_converged = true

# start_in_debugger = true
# debugger_timeout = 100

# End of file

```

PyLith parameters for 3D MODELS

```

# The settings in this file (pylithapp.cfg) will be read automatically
# by pylith, as long as the file is placed in the run directory.

# The settings in this file will override any settings in:
# PREFIX/etc/pylithapp.cfg
# $HOME/.pyre/pylithapp/pylithapp.cfg

# The settings in this file will be overridden by any .cfg file given
# on the command line or by any command line settings.

```

```

[pylithapp]

# -----
# journal
# -----
# The settings below turn on journal info for the specified components.
# If you want less output to stdout, you can turn these off.
[pylithapp.journal.info]
timedependent = 1
implicit = 1
petsc = 1
solverlinear = 1
meshiocubit = 1
homogeneous = 1
elasticityimplicit = 1
fiatsimplex = 1
pylithapp = 1
materials = 1
mesh_distributor = 1

# -----
# mesh_generator
# -----
# The settings below control the mesh generation (importing mesh info).
# Turn on debugging output for mesh generation.
[pylithapp.mesh_generator]
debug = 1
reader = pylith.meshio.MeshIOcubit

# This component specification means we are using PyLith ASCII format,
# and we then specify the filename and number of space dimensions for
# the mesh.
[pylithapp.mesh_generator.reader]
filename = ../mesh/mesh_3d_mod6.exo

# -----
# problem
# -----
# Specify the problem settings.
# This is a time-dependent problem, so we select this as our problem type.
# We select a total time of 0 sec, and a time step size of 1 sec, so we
# are performing a single time step.
# The spatial dimension for this problem is 2.
# For an implicit formulation (using implicit.cfg), we will perform 1
# implicit time step from t = -1.0 to t = 0.0 (elastic solution step).
[pylithapp.timedependent]
dimension = 3
formulation = pylith.problems.Implicit
formulation.solver = pylith.problems.SolverNonlinear

[pylithapp.timedependent.formulation.time_step]
total_time = 1.0*year
dt = 1.0*year

# -----
# materials
# -----
# Specify the material information for the problem.
# The material type is isotropic elastic formulated for plane strain.

[pylithapp.timedependent]
materials = [ucrust,lcrust,mantle,sill,chamber]

# Upper Crust

```

```

[pylithapp.timedependent.materials.ucrust]
label = Elastic ucrust

# The id corresponds to the block number from CUBIT.
id = 1

db_properties = spatialdata.spatialdb.SimpleDB
db_properties.label = Elastic ucrust properties
db_properties.iohandler.filename = ../spatialdb/mat_ucrust_elastic.spatialdb
db_properties.query_type = nearest

# We are doing 2D quadrature for a triangle.
quadrature.cell = pylith.feassemble.FIATSimplex
quadrature.cell.dimension = 3

# Lower crust
[pylithapp.timedependent.materials.lcrust]
label = Elastic lower crust

# The id corresponds to the block number from CUBIT.
id = 2

db_properties = spatialdata.spatialdb.SimpleDB
db_properties.label = Elastic lcrust properties
db_properties.iohandler.filename = ../spatialdb/mat_lcrust_elastic.spatialdb
db_properties.query_type = nearest

# We are doing 2D quadrature for a triangle.
quadrature.cell = pylith.feassemble.FIATSimplex
quadrature.cell.dimension = 3

# Mantle
[pylithapp.timedependent.materials.mantle]
label = Elastic mantle

# The id corresponds to the block number from CUBIT.
id = 3

db_properties = spatialdata.spatialdb.SimpleDB
db_properties.label = Elastic mantle properties
db_properties.iohandler.filename = ../spatialdb/mat_mantle_elastic.spatialdb
db_properties.query_type = nearest

# We are doing 2D quadrature for a triangle.
quadrature.cell = pylith.feassemble.FIATSimplex
quadrature.cell.dimension = 3

# Sill
[pylithapp.timedependent.materials.sill]
label = Elastic sill

# The id corresponds to the block number from CUBIT.
id = 4

db_properties = spatialdata.spatialdb.SimpleDB
db_properties.label = Elastic sill properties
db_properties.iohandler.filename = ../spatialdb/mat_magmanew_elastic.spatialdb
db_properties.query_type = nearest

# We are doing 2D quadrature for a triangle.
quadrature.cell = pylith.feassemble.FIATSimplex
quadrature.cell.dimension = 3

# Chamber
[pylithapp.timedependent.materials.chamber]

```

```

label = Elastic chamber

# The id corresponds to the block number from CUBIT.
id = 5

db_properties = spatialdata.spatialdb.SimpleDB
db_properties.label = Elastic chamber properties
db_properties.iohandler.filename = ../spatialdb/mat_mush_elastic.spatialdb
db_properties.query_type = nearest

# We are doing 2D quadrature for a triangle.
quadrature.cell = pylith.feassemble.FIATSimplex
quadrature.cell.dimension = 3

# -----
# output
# -----
# Names of output files are set in stepXX.cfg. We consolidate all of the
# output settings that are common to all of the simulations here.

[pylithapp.timedependent.formulation]
# Set the output to an array of 2 output managers.
# We will output the solution over the domain and the ground surface.
output = [domain]

# Domain
[pylithapp.problem.formulation.output.domain]
vertex_data_fields = [displacement]
writer = pylith.meshio.DataWriterHDF5

# Materials
[pylithapp.timedependent.materials.ucrust.output]
writer = pylith.meshio.DataWriterHDF5
cell_data_fields = [stress,cauchy_stress,total_strain]

# Comment this out for if rerunning Step 3 to get initial conditions for Step 8.
cell_filter = pylith.meshio.CellFilterAvg

[pylithapp.timedependent.materials.lcrust.output]
writer = pylith.meshio.DataWriterHDF5
cell_data_fields = [stress,cauchy_stress,total_strain]

# Comment this out for if rerunning Step 3 to get initial conditions for Step 8.
cell_filter = pylith.meshio.CellFilterAvg

[pylithapp.timedependent.materials.mantle.output]
writer = pylith.meshio.DataWriterHDF5
cell_data_fields = [stress,cauchy_stress,total_strain]

# Comment this out for if rerunning Step 3 to get initial conditions for Step 8.
cell_filter = pylith.meshio.CellFilterAvg

[pylithapp.timedependent.materials.sill.output]
writer = pylith.meshio.DataWriterHDF5
cell_data_fields = [stress,cauchy_stress,total_strain]

# Comment this out for if rerunning Step 3 to get initial conditions for Step 8.
cell_filter = pylith.meshio.CellFilterAvg

[pylithapp.timedependent.materials.chamber.output]
writer = pylith.meshio.DataWriterHDF5
cell_data_fields = [stress,cauchy_stress,total_strain]

# Comment this out for if rerunning Step 3 to get initial conditions for Step 8.
cell_filter = pylith.meshio.CellFilterAvg

```

```

# -----
# PETSc
# -----
# We are using all of the default settings for PETSc except for
# specifying the block Jacobi preconditioner.  Additional PETSc
# command-line arguments may be found in the PETSc documentation.
[pylithapp.petsc]
malloc_dump =

# Preconditioner settings.
pc_type = asm
sub_pc_factor_shift_type = nonzero

# Convergence parameters.
ksp_rtol = 1.0e-10
ksp_atol = 1.0e-20
ksp_max_it = 1000
ksp_gmres_restart = 200

# Linear solver monitoring options.
ksp_monitor = true
ksp_converged_reason = true
ksp_error_if_not_converged = true

# Nonlinear solver monitoring options.
snes_rtol = 1.0e-10
snes_atol = 1.0e-9
snes_max_it = 100
snes_monitor = true
snes_converged_reason = true
snes_linesearch_monitor = true
snes_error_if_not_converged = true

# PETSc summary -- useful for performance information.
log_view = true

# End of file

```

SM5. Post-processing.

We subtract the gravitational stresses due to the reference model (SM1), and calculate principal stresses and pressure = $(\sigma_1 + \sigma_2 + \sigma_3) / 3$ using the python script shown below for the 3D case:

```

#!/usr/bin/env nemesis
# -*- Python -*- (syntax highlighting)
# -----

import numpy
import h5py
import sys
from shutil import copyfile

from spatialdata.spatialdb.SimpleIOAscii import SimpleIOAscii
from coordsys import cs_mesh
cs = cs_mesh()

# Constants for computing reference model.
rhoUC = 2761.0

```

```

rhoLC = 2833.0
rhoM = 3300.0
vsUC = 3600.0
vsLC = 3400.0
vsM = 4500.0
vpUC = 6200.0
vpLC = 6500.0
vpM = 8100.0
g = 9.80665
z1 = -20000.0
z2 = -40000.0
stressUC = rhoUC*g*z1
stressLC = rhoLC*g*(z2 - z1)

# Compute material properties.
muUC = rhoUC*vsUC*vsUC
muLC = rhoLC*vsLC*vsLC
muM = rhoM*vsM*vsM
lamUC = rhoUC*vpUC*vpUC - 2.0*muUC
lamLC = rhoLC*vpLC*vpLC - 2.0*muLC
lamM = rhoM*vpM*vpM - 2.0*muM

def getCellCenters(vertices, cells):
    """
    Function to compute cell centers.
    """
    cellCoords = vertices[cells, :]
    cellCenters = numpy.mean(cellCoords, axis=1)
    return cellCenters

def generate(sim, fileRoot, materials):

    for material in materials:

        filenameH5 = "%s-%s.h5" % (sim, material)

        # Open HDF5 file and get coordinates, cells, and stress.
        h5 = h5py.File(filenameH5, "r")
        vertices = h5['geometry/vertices'][:]
        cells = numpy.array(h5['topology/cells'][:, :], dtype=numpy.int)

        filenewH5 = "%s-%s-pstresses.h5" % (sim, material)
        copyfile(filenameH5, filenewH5)
        h5new = h5py.File(filenewH5, "r+")

        # Get stresses from final time step.
        stress = h5['cell_fields/stress'][-1, :, :]
        refStress = numpy.zeros_like(stress)
        h5.close()

        # Compute coordinates of quadrature points.
        quadCoords = getCellCenters(vertices, cells)

        # Compute reference stress.
        z = quadCoords[:, 2]
        coordsUC = numpy.where(z >= z1)
        coordsLC = numpy.where(numpy.logical_and(z >= z2, z < z1))
        coordsM = numpy.where(z < z2)
        if (coordsUC[0].shape[0] != 0):
            refStress[coordsUC, 0] = rhoUC*g*z[coordsUC]
        if (coordsLC[0].shape[0] != 0):
            refStress[coordsLC, 0] = stressUC + rhoLC*g*(z[coordsLC]-z1)
        if (coordsM[0].shape[0] != 0):
            refStress[coordsM, 0] = stressUC + stressLC + rhoM*g*(z[coordsM]-z2)
        refStress[:, 1] = refStress[:, 0]

```

```

refStress[:,2] = refStress[:,0]
refStress *= -1.0
pressureRef = (refStress[:,0] + refStress[:,1] + refStress[:,2])/3

# Define square 3x3 array

sigma1 = [] # principal compressive stress, negative stress in PyLith
sigma2 = []
sigma3 = [] # principal tensile stress, positive stress in PyLith
theta1 = []
theta2 = []
theta3 = []
pressure = []
for cell in stress:
    stresstensor =
numpy.array([(cell[0],cell[3],cell[5]),(cell[3],cell[1],cell[4]),(cell[5],cell[4],ce
ll[2])])

# Calculate principal stresses
[w, v] = numpy.linalg.eigh(stresstensor)
sigma1.append(w[0])
sigma2.append(w[1])
sigma3.append(w[2])
theta1.append(v[:,0])
theta2.append(v[:,1])
theta3.append(v[:,2])
pressure.append((w[0]+w[1]+w[2])/3)

# Create pressure array as well as pressure minus reference model.
pressure = numpy.array(pressure)
pressureRel = pressure - pressureRef
pressureRelneg = pressureRel*(-1)

# change row vector to column vector
sigma1 = numpy.array(sigma1).reshape(-1,1)
sigma2 = numpy.array(sigma2).reshape(-1,1)
sigma3 = numpy.array(sigma3).reshape(-1,1)

# add zero for z column
theta1 = numpy.array(theta1)
theta3 = numpy.array(theta3)
theta2 = numpy.array(theta2)

nrows = stress.shape[0]
ncols = stress.shape[1]

#dataset_sigma = h5new.create_dataset('cell_fields/sigma', (1,nrows,3),
maxshape=(None,nrows,3), data=sigma)
dataset_sigma1 = h5new.create_dataset('cell_fields/s1', (1,nrows,1),
maxshape=(None,nrows,1), data=sigma1)
dataset_sigma2 = h5new.create_dataset('cell_fields/s2', (1,nrows,1),
maxshape=(None,nrows,1), data=sigma2)
dataset_sigma3 = h5new.create_dataset('cell_fields/s3', (1,nrows,1),
maxshape=(None,nrows,1), data=sigma3)
dataset_theta1 = h5new.create_dataset('cell_fields/t1', (1,nrows,3),
maxshape=(None,nrows,3), data=theta1)
dataset_theta2 = h5new.create_dataset('cell_fields/t2', (1,nrows,3),
maxshape=(None,nrows,3), data=theta2)
dataset_theta3 = h5new.create_dataset('cell_fields/t3', (1,nrows,3),
maxshape=(None,nrows,3), data=theta3)
dataset_pressure = h5new.create_dataset('cell_fields/P', (1,nrows,1),
maxshape=(None,nrows,1), data=pressure)
dataset_pressure_ref = h5new.create_dataset('cell_fields/P_ref',
(1,nrows,1), maxshape=(None,nrows,1), data=pressureRef)

```

```

    dataset_pressure_rel = h5new.create_dataset('cell_fields/P_rel',
(1,nrows,1), maxshape=(None,nrows,1), data=pressureRel)
    dataset_pressure_relneg = h5new.create_dataset('cell_fields/P_relneg',
(1,nrows,1), maxshape=(None,nrows,1), data=pressureRelneg)

    #dataset_sigma.attrs.create(name='vector_field_type',data="tensor")
    dataset_sigma1.attrs.create(name='vector_field_type',data="scalar")
    dataset_sigma2.attrs.create(name='vector_field_type',data="scalar")
    dataset_sigma3.attrs.create(name='vector_field_type',data="scalar")
    dataset_theta1.attrs.create(name='vector_field_type',data="vector")
    dataset_theta2.attrs.create(name='vector_field_type',data="vector")
    dataset_theta3.attrs.create(name='vector_field_type',data="vector")
    dataset_pressure.attrs.create(name='vector_field_type',data="scalar")
    dataset_pressure_ref.attrs.create(name='vector_field_type',data="scalar")
    dataset_pressure_rel.attrs.create(name='vector_field_type',data="scalar")
    dataset_pressure_relneg.attrs.create(name='vector_field_type',data="scalar")

    h5new.close()

    return

# -----
if __name__ == "__main__":
    import argparse
    parser = argparse.ArgumentParser()
    parser.add_argument("--sim", action="store", dest="sim",
default="novel_grav_mod0")
    parser.add_argument("--file-root", action="store", dest="fileRoot",
default="mat_principal_stress")
    parser.add_argument("--materials", action="store", dest="materials",
default="ucrust,lcrust,mantle")
    args = parser.parse_args()

    materials = args.materials.split(",")
    generate(args.sim, args.fileRoot, materials)

# End of file

```