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Semi-automatic detection of the overtopping waves and reconstruction of the overtopping flow characteristics at coastal structures

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2
3 **SEMI-AUTOMATIC DETECTION OF THE OVERTOPPING WAVES AND**
4 **RECONSTRUCTION OF THE OVERTOPPING FLOW CHARACTERISTICS**
5 **AT COASTAL STRUCTURES**

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7

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15 **Abstract**

16 This paper proposes a semi-automatic and customizable procedure for the identification of the
17 overtopping waves based on a threshold-down-crossing analysis of the sea surface elevation
18 signals. The procedure can be applied to 2D experimental and numerical signals, to emerged
19 and submerged structures, with the same accuracy of a human-supervised analysis. The
20 procedure includes an original and innovative algorithm to compare the water level signals at
21 consecutive gauges and couple the waves propagating in between. The coupling algorithm
22 implies a series of original applications of practical relevance, such as: i) the computation of the
23 wave celerity, which is a crucial parameter for the assessment of the structural stability and the
24 hydraulic vulnerability of the landward area; ii) the estimation of the wave overtopping discharge,
25 which can be obtained by integrating the wave celerities with the surface elevations; iii) the
26 description of the wave overtopping characteristics and their evolution over the structure crest;
27 iv) the evaluation of the volumes lost for percolation in permeable structures. The application to
28 new and literature data and the comparison with well-established formulae prove that the results
29 obtained from the identification and coupling procedures are accurate and reliable.

30
31 **Keywords:** semi-automatic procedure; wave coupling; wave celerity; flow velocity; wave
32 overtopping discharge; individual wave overtopping volumes; probability of overtopping.
33

34 **Highlights**

35 A new semi-automatic procedure for the identification and coupling of the overtopping waves is
36 presented

37 The accuracy of this new procedure is the same achieved by visual examination of the wave
38 signals time series

39 The procedure allows the calculation of the wave celerities and the estimation of the overtopping
40 discharge

41 The procedure allows to reconstruct the probability distribution of the overtopping volumes

42 The procedure allows for the analysis of the evolution of the overtopping flow characteristics
43 across the structure crest

44

45 **1. Introduction**

46 For design purposes, the identification of the single overtopping events and the analysis of their
47 evolution over the structure crest represent the key information for the assessment of the
48 hydraulic vulnerability of the structures and the stability of the armour layers. Specifically, the
49 accurate identification of the waves is essential for the reconstruction of the statistical distribution
50 of the individual overtopping volumes, which govern the hydrodynamic forces acting on the crests
51 and the landward-side slopes (Van der Meer et al., 2010; Hughes et al., 2012).

52 The more recent automatic procedures for the reconstruction of the individual overtopping
53 volumes are based on the zero up-crossing or down-crossing analysis of the sea surface
54 elevation signals at weigh cells (Victor, 2012, later modified by Platteeuw, 2015; Molines et al.
55 2019) or at resistant gauges (Nørgaard et al., 2014; Hughes, 2015; Hughes and Thornton, 2016).
56 Many of these procedures seem to be affected by unaffordable inaccuracy in comparison with
57 the manually supervised analysis of the discharge time series (Hughes, 2015).

58 None of these procedures are scoped to identify the waves at consecutive gauges and couple
59 the waves travelling along the structure crest. The computation of the wave time lags from the
60 off-shore to the in-shore gauge would allow the derivation of the wave front velocity propagation,
61 i.e. the wave celerity, which may be used as an estimator of the flow velocity (Schüttrumpf and
62 Oumeraci, 2005; Lykke Andersen et al., 2011) and can be integrated with the water levels to
63 derive the instantaneous and average overtopping discharges. Coupling the single overtopping
64 waves would also allow the reconstruction of the evolution of the wave shapes, which is
65 fundamental for characterizing the wave asymmetries (a.o., Peng et al., 2009; Chella et al., 2015)
66 or for estimating the volumes lost for percolation over the crest of permeable structures
67 (Zanuttigh and Lamberti, 2006).

68 Yet, as far as known, an automatic procedure for coupling the overtopping events, which is of
69 wide applicability and which can guarantee high standards of confidence, is missing. A first
70 attempt was made by Zanuttigh and Lamberti (2006), whose procedure was not applicable to all
71 crest level conditions and was strongly affected by the selection of appropriate threshold values.

72 This contribution describes a new, advanced and versatile procedure for the identification of the
73 overtopping waves specifically developed to overcome the problems of the automatic detection,
74 in order to reproduce the number of the overtopping waves as they would have been recognized
75 by a manual analysis of the wave signal (see Hughes et al., 2015; 2016). The procedure has
76 been recently presented in Formentin and Zanuttigh (2018, b) through a preliminary application
77 to a numerical database consisting of wave overtopping at smooth dikes. In the present work,
78 the procedure is fully validated against a new set of experiments recently carried out by the
79 authors (Zanuttigh and Formentin, 2018) and a set of original applications of practical interest
80 are presented with reference to a wide laboratory dataset.

81 The procedure is conceived to elaborate water level time signals and reconstruct the shape
82 parameters of the overtopping waves (crest and trough elevations and durations), but it can
83 process any kind of oscillatory signal in the time domain. The waves are detected through the

84 implementation of a threshold-down-crossing algorithm, where the 2 threshold values vary with
85 the wave characteristics, the structural parameters, and the scale of the data. The setting of the
86 thresholds can be customized by the user to achieve the desired level of accuracy in the
87 identification of the overtopping events. A second algorithm can couple the waves travelling over
88 the structure crest so that the wave celerity, the wave overtopping discharge, the overtopping
89 volumes can be estimated. The procedure can deal with both regular and irregular 2D waves
90 propagating perpendicularly towards the structures. The wave identification requires as input a
91 unique signal recorded at one gauge, while the wave coupling requires at least 2 signals
92 recorded at 2 consecutive gauges.

93 Section 2 presents a short overview of the datasets used to illustrate, calibrate and validate this
94 new procedure. In particular, the new set of overtopping experiments at dikes, used for the
95 validation of the procedure and the presentation of its results, is introduced. The algorithm and
96 the working principle of the procedure are described in Section 3 by steps, including the
97 validation of the procedure and the assessment of the validity of the adopted criteria. Sections 4
98 and 5 propose a set of applications to smooth and rubble mound structures, respectively. The
99 conclusions on the work are drawn in Section 6.

100

101 **2. Experimental and numerical datasets**

102 This Section presents and describes the 5 groups of tests that are used to describe, validate and
103 illustrate the application fields of the procedure. The ranges of variability of the hydraulic and
104 structural parameters relative to these datasets represent therefore the field of applicability on
105 which the procedure has been tested so far. Table 1 proposes a summary of the 5 datasets,
106 while their details are given in the Sub-section 2.1 to 2.5.

107 All the datasets considered in this work consist of overtopping tests at structures (smooth dikes
108 or rubble mound breakwater) which are characterized by a similar, simple trapezoidal cross-
109 section. Figure 1 provides a schematic layout of the typical cross-section, with reference to the
110 main symbols adopted hereinafter to describe the geometrical parameters of the structures and
111 the working principle of the procedure:

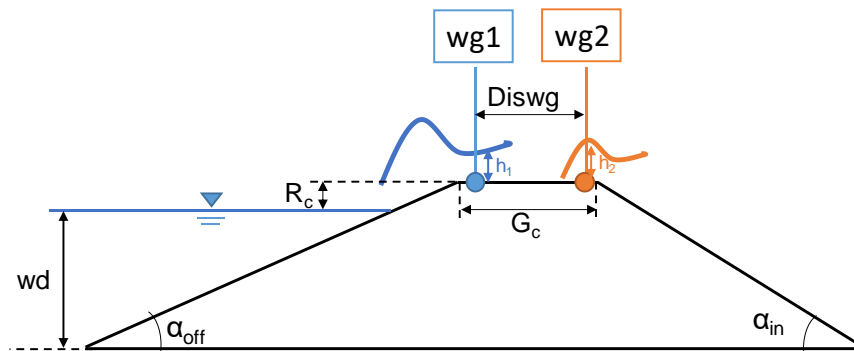
- 112 • the structure off-shore and in-shore slopes, α_{off} and α_{in} ;
- 113 • the water depth before the structure wd and structure crest freeboard R_c ;
- 114 • the structure crest width G_c ;
- 115 • the position of the 2 generic “wave gauges” (wgs, hereinafter) for the registration of the water
116 level signals at the structure crest to be processed with the procedure, and their distance
117 $diswg$; the term “wg” can refer to resistant wave gauges, pressure transducers, Acoustic
118 Doppler Velocimeters, etc., according to the instruments used in the different experiments;
- 119 • the water depth, or the free surface elevation, h measured at the 2 wgs, h_1 and h_2 ,
120 respectively.

121 Table 1. Summary and characteristics of the 5 datasets of tests on wave overtopping mentioned
 122 throughout the manuscript and involved in the application of the new procedure.

Dataset Label	Number of tests	Description of the tests	Freeboard range R_c/H_s	Slope $\cot(\alpha_{off})$	Employment of the dataset	Reference
UB-num	94	2D numerical tests at smooth dikes	+[-1.5; 1.5]	4; 6	Description of the procedure (Sec.s 3.1, 3.2)	Formentin and Zanuttigh (2018, a)
HT	8	2D and 3D experiments at smooth dikes	[0.319; 1.064]	3; 6	Validation of the identification step (Sec. 3.3.1)	Hughes and Thornton (2016)
HS	3	2D experiments on levees	[-0.430; -0.121]	4.25	Validation of coupling step (Sec. 3.3.1)	Hughes and Shaw (2011)
UB-exp	54	2D experiments at smooth dikes	0; 0.5; 1	2; 4	Results and applications of the procedure (Sec. 4)	Zanuttigh and Formentin (2018)
AAU	33	3D experiments tests at permeable breakwaters	[-1.59; 0.49]	2	Results and applications of the procedure (Sec. 5)	Kramer et al. (2005)

123

124



125

126 Figure 1 – Scheme of a dike for the application of the new procedure, including 2 wgs at the off-
 127 shore (wg1) and in-shore (wg2) edges of the crest.

128

129 2.1. Dataset “UB-num”

130 The numerical database of tests on wave overtopping against smooth dikes collected by
 131 Formentin et al. (2014) and extended by Formentin and Zanuttigh (2018, a), “UB-num”,
 132 hereinafter, includes 94 tests on structures at various crest freeboards, with R_c/H_s ranging from

133 -1.5 to +1.5, 2 off-shore slopes $\cot(\alpha_{\text{off}})=4$ and 6 and fixed in-shore slope $\cot(\alpha_{\text{in}})=3$ and fixed
 134 crest width $G_c=0.3$ m. The wave attacks included 2 target wave heights $H_s=0.1$ and 0.2 m and
 135 wave steepnesses $H_s/L_{m-1,0}$ in the range 0.02-0.05, where $L_{m-1,0}$ is the wave length computed on
 136 the spectral wave period $T_{m-1,0}$. The simulations were carried out with the IH-2VOF code
 137 developed by the University of Cantabria (Lara et al., 2011) and the summary of the tested
 138 conditions is given in Table 2.

139 For all the tests, 2 numerical wave gauges (wgs), namely wg1 and wg2, were placed in proximity
 140 of the off-shore and in-shore edges of the crest width, at the distance $diswg=0.27$ m. wg1 and
 141 wg2 provided the time records of the free-surface elevations (h , [m]) and the cross-shore flow
 142 velocities (u , [m/s]). These records can be processed by the procedure and were used indeed
 143 for the calibration of its parameters (Sub-sections 3.1 and 3.2) and in part for the validation of its
 144 results (Sub-section 3.3). For all the simulations, the adopted sampling frequency (sf) at wg1
 145 and wg2 was $sf=20$ Hz.

146
 147 Table 2. Summary of the tested conditions of the numerical database collected by Formentin and
 148 Zanuttigh (2018, a). For all the tests, $G_c=0.3$ m, $diswg=0.27$ and $\cot(\alpha_{\text{in}})=3$. The water depth
 149 $wd=0.85-R_c$. $sf=20$ Hz.

R_c/H_s	-1.5	-1	-0.5	-0.2	0	+0.5	+1	+1.5
$H_s/L_{m-1,0}$ [%]	2; 3	2; 3; 4	2; 3; 5	2	2; 3; 4; 5	2; 3; 4	2; 3; 4	3
H_s [m]	0.1; 0.2	0.1; 0.2	0.1; 0.2	0.2	0.1; 0.2	0.1; 0.2	0.2	0.2
$\cot(\alpha_{\text{off}})$	4; 6	4; 6	4; 6	4; 6	4; 6	4; 6	4; 6	4; 6
Tot. #	12	18	18	2	20	16	6	2

150

151 2.2. Dataset “HT”

152 The *FlowDike 1* and *FlowDike 2* experiments on wave overtopping and run-up were conducted
 153 in the wave basin at the Danish Hydraulic Institute in Hørsholm, DK (Lorke et al., 2009 and 2010).
 154 These experiments involved 2D and 3D wave attacks against smooth dikes characterized by 2
 155 G_c values (0.6 and 0.7 m, model scale values) and 2 $\cot(\alpha_{\text{off}})$ values (3 and 6, respectively
 156 *FlowDike 1* and *FlowDike 2*). A subset of 8 irregular 2D tests at emerged freeboard ($R_c/H_s>0$)
 157 belonging to the *FlowDike* experiments were selected by Hughes and Thornton (2016) and
 158 elaborated to identify the individual wave overtopping waves based on a human-supervised
 159 analysis of the time series of the overtopping discharge. The identification step of the new
 160 procedure proposed in this contribution has been applied to the same 8 tests and the results are
 161 compared to the achievements of Hughes and Thornton (2016) for validation (Sub-section 3.3.1).
 162 The characteristics of the 8 tests are resumed in Table 3, where: the subscript “HT” means “Hughes
 163 and Thornton, 2016” and refers to the quantities obtained by the application of the procedure by
 164 Hughes and Thornton, 2016; N_w = number of waves; V = volumes; Pow = “probability of
 165 overtopping”; Man = “recognized by manual detection (human supervised)”; $Auto$ = “recognized
 166 by automatic detection”; $Total$ = $auto+man$. In Table 3 the symbol H_{m0} is used to refer to the
 167 spectral wave height calculated at the structure toe as reported in Hughes and Thornton (2018).

168 Table 3. Summary of the 8 selected experiments and corresponding wave volume determination
 169 performed by the supervised procedure by Hughes and Thornton (2016) and by the new
 170 procedure. The columns to be compared each other are shaded with the same colour.

Test ID	Tests parameters				Hughes and Thornton (2016) procedure					New procedure		
	H_{m0} [m]	$T_{m-1,0}$ [s]	R_c/H_s [-]	Nw [-]	Auto V_{HT}	Man V_{HT}	Total V_{HT}	(Auto V_{HT})/ (Total V_{HT})	Pow _{HT}	Auto V	(Auto V)/ (Total V_{HT})	Pow
0198	0.103	1.619	0.971	1180	486	177	663	73%	56%	636	96%	54%
0199	0.094	1.164	1.064	1102	211	235	446	47%	40%	356	80%	31%
0200	0.15	1.96	0.667	1276	453	361	814	56%	64%	839	103%	66%
0201	0.148	1.379	0.376	1150	254	493	747	34%	65%	680	91%	59%
0451	0.09	1.555	0.556	1097	535	46	581	92%	53%	621	107%	57%
0453	0.122	1.663	0.410	1120	726	142	868	84%	78%	931	107%	83%
0456	0.157	1.936	0.319	1093	617	275	892	69%	82%	904	101%	83%
0457	0.141	1.373	0.355	1116	521	267	788	66%	71%	794	101%	71%

171

172 2.3. Dataset “HS”

173 To validate the coupling step of the new procedure, 3 tests were selected from a set of
 174 experiments on wave overtopping carried out against a trapezoidal levee in 1:25 scale at the
 175 Coastal and Hydraulics Laboratory (CHL) in Vicksburg, MS (Hughes and Shaw, 2011). All the
 176 experiments (HS, hereinafter) were performed in submerged conditions (negative freeboard of
 177 the levee crest) to investigate the combined effects of the storm surge overflow and the wave
 178 overtopping. 2 wgs, placed over the crest of the levee at a distance of 62 cm ($diswg=0.62$ m),
 179 were used to measure the flow thickness h over the levee and the pressure, and a laser Doppler
 180 velocimeter (ADV) was installed in correspondence of the second wg to measure the horizontal
 181 component of the flow velocity u . The full description of the experiments is given in Hughes and
 182 Shaw (2011), while the main hydraulic parameters of the 3 selected tests are reported in model
 183 scale units in Table 4. Only 3 tests were used for the application of the procedure because these
 184 are the only ones for which the records of h at the 2 wgs were available. In Table 4, H_{m0} refers
 185 to the spectral wave height at the structure toe as reported by Hughes and Shaw (2018) while sf
 186 is the sample frequency adopted to record the free-surface elevations and velocities. $u_{2\%}$ and
 187 c_{mean} are, respectively, the values of the flow velocities measured exceeded by the 2% of the
 188 incoming waves and the mean wave celerities computed with the coupling step of the procedure.

189

190 Table 4. Comparison among the lab measurements of the horizontal flow velocity ($u_{2\%}$) and the
 191 wave celerities (c_{mean}) derived from the new procedure for 3 tests by Hughes and Shaw (2011).

Test ID	H_{m0} [m]	T_p [s]	R_c [m]	R_c/H_{m0}	H_{m0}/R_c	sf [Hz]	Lab data $u_{2\%}$ [m/s]	New procedure c_{mean} [m/s]
R14	0.071	2.09	-0.0110	-0.155	6.45	100	0.800	0.895
R18	0.100	2.77	-0.043	-0.430	2.33	100	1.113	1.268
R109	0.099	2.73	-0.012	-0.121	8.25	50	0.722	0.729

192 2.4. Dataset “UB-exp”

193 A set of new experimental tests of wave overtopping was recently carried out by the authors in
194 the wave flume of the Hydraulics Laboratory of the University of Bologna (Zanuttigh and
195 Formentin, 2018), UB-exp, hereinafter. The wave flume was equipped with an overtopping tank
196 for the storage of the overtopping volumes and a recirculation system consisting of a pump and
197 a flowmeter. The experiments consisted of irregular waves at dikes (54 tests) and dikes with a
198 crown wall and a parapet (91 tests) at zero or positive freeboards ($0 \leq R_o/H_s \leq 1$). The procedure
199 was applied to the 54 tests on dikes without walls exclusively. The selected structures –
200 schematized in Figure 2 – consist of 4 dike configurations obtained by combining 2 crest widths
201 ($G_c=0.3$ and 0.15 m) and 2 slopes ($\cot(\alpha_{off})=2$ and 4). The height of the dike crest was kept
202 constant and equal to 0.35 m (with respect to the bottom of the wave flume) and the different
203 freeboards were realized by varying the water depth wd from 0.29 to 0.35 m. For each dike, 6
204 irregular wave attacks were performed, combining 3 H_s (0.04 , 0.05 and 0.06 m) and 2 $H_s/L_{m-1,0}$
205 (3% and 4%). The matrix of the tested conditions is given in Table 5.

206 All the dikes were equipped with 3 Ultrasonic Doppler Velocity Profilers (UVPs), named “D4”,
207 “D5” and “D6”, characterized by $sf=50$ Hz and placed consecutively along the structure crest
208 from the off-shore edge to the in-shore edge, as displayed in Figure 2. These UVPs were used
209 to derive the time series of the vertical profiles of the flow velocities (u , [m/s]) and track the free
210 surface elevation (h , [m]) at D4, D5 and D6, based on the following methodologies.

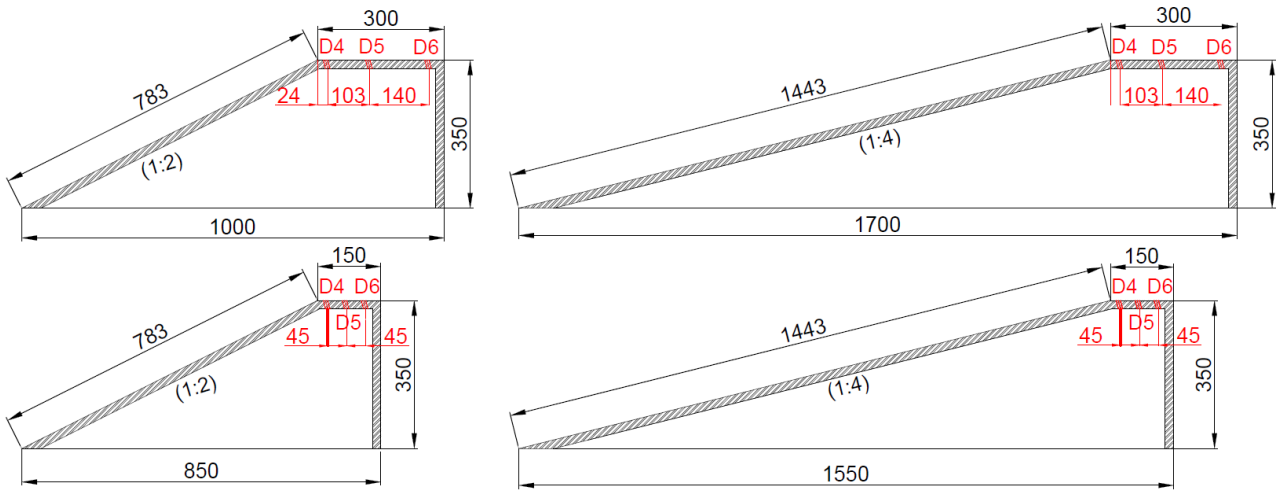
- 211 • u . At each time step, the UVPs measured the horizontal component of the flow velocity (u ,
212 [m/s]) at 40 gates linearly spaced along the direction of the acoustic impulse. The range of
213 the impulse was set approximately 10 cm for each UVP, i.e. a distance which should be
214 sufficient to sweep the whole water column from the dike crest to the free surface also in the
215 case of the largest waves ($H_s=0.06$ m). By dividing the range of the impulse (10 cm) by the
216 number of gates (40), the spatial resolution of ≈ 0.0025 m is obtained for the vertical profiles
217 of u .
- 218 • h . At each time step and at the same 40 gates along the direction of the sonic impulse, the
219 UVPs recorded also the values of the echo (dB). In correspondence of the free surface, the
220 acoustic impulse undergoes a strong reflection, which determines a sharp peak of the echo
221 value. The time series of the free-surfaces at each UVP have been reconstructed based on
222 the position of the peaks in the instantaneous vertical profiles of the echo.

223 The time series of h at D4 and D6 were elaborated with the new procedure to detect and couple
224 the individual overtopping volumes. In this application, D4 and D6 represent therefore “wg1” and
225 “wg2”, respectively. The distances between D4 and D6, $diswg$, are respectively equal to 0.09 m
226 and 0.243 m for $G_c=0.15$ m and $G_c=0.30$ m, see Table 5.

227 The time series of u were elaborated to get the statistics of the overtopping flow velocities at
228 the off-shore edge ($\approx D4$) and to validate the coupling step of the procedure (see Sub-section
229 3.4.1), but were not directly elaborated by the procedure itself.

230

231



232

233 Figure 2 – Scheme of the 4 dike configurations tested in the Laboratory of Hydraulics at the
 234 University of Bologna (Zanuttigh and Formentin, 2018). The positions of the 3 UVPs (D4, D5,
 235 D6) installed across the dike crest are marked with red colour. The measures are in mm.

236

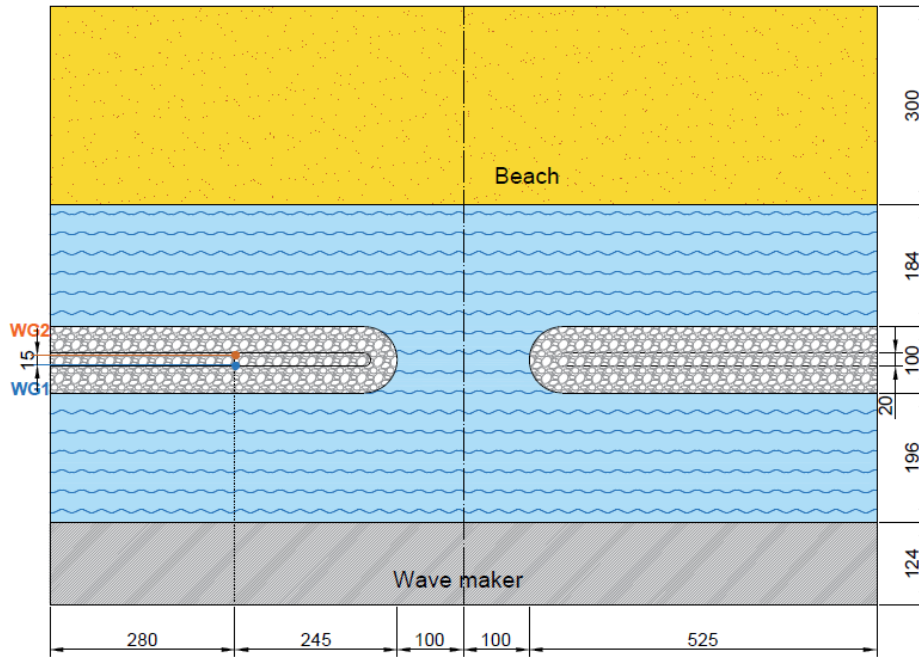
237 Table 5. Summary of the tested conditions characterizing the 54 experiments of wave
 238 overtopping at dikes selected from the new database by Zanuttigh and Formentin (2018). For all
 239 the tests, $sf=50$ Hz.

R_c/H_s	0	+0.5	+1
$H_s/L_{m-1,0}$ [%]	3; 4	3; 4	3; 4
H_s [m]	0.04; 0.05; 0.06	0.04; 0.05; 0.06	0.04; 0.05; 0.06
wd [m]	0.35	[0.32; 0.33]	[0.29; 0.31]
$\cot(\alpha_{off})$	2; 4	2; 4	2; 4
G_c [m]	0.15; 0.30	0.15; 0.30	0.15; 0.30
$diswg$ [m]	0.09 for $G_c=0.15$ m; 0.243 for $G_c=0.30$ m		
Tot. #	24	18	18

240

241 2.5.Dataset “AAU”

242 The dataset of experimental tests performed in the shallow water basin at the Aalborg University
 243 (Kramer et al., 2005; Zanuttigh and Lamberti, 2006) consists of 33 perpendicular, regular and
 244 irregular wave attacks against 2 permeable structures with a gap in between, was selected. The
 245 main hydraulic and structural parameters characterizing the 33 tests selected (dataset “AAU”)
 246 are summarized in Table 6, while the top view of the structures is here reported in Figure 3.



247
 248 Figure 3 – Top view of the permeable structures with a gap in between used for the experiments
 249 at the Aalborg University by Kramer et al., 2005 (dataset AAU). Narrow berm configuration. The
 250 positions of wg1 and wg2 are represented by filled-in circles. Measures in centimeters.

251

252 Table 6. Summary of the hydraulic and structural parameters characterizing the 33 tests selected
 253 from the database of Kramer et al. (2005) for the assessment of the accuracy of the new
 254 procedure. Scale 1:20.

Parameter	values or range	description of the parameter
R_c [m]	-0.07, 0 and 0.03	crest freeboard
wd [m]	0.27, 0.20 and 0.17	water depth at the structure toe
H_s [m]	[0.020; 0.129]	significant wave heights at the structure toe
T_p [s]	[0.74; 1.97]	peak wave period
G_c [m]	0.2 and 0.6	width of the structure crest
$cot(\alpha_{off})$	2	cotangent of the seaward slope
$cot(\alpha_{in})$	2	cotangent of the landward slope
L_b [m]	5.05	length of each barrier (2 barriers at all)
L_{gap} [m]	2.40	length of the gap between the two barriers
sf [Hz]	40	sample frequency adopted for the experiments
$diswg$ [m]	0.15 (for $G_c=0.2$) and 0.40 (for $G_c=0.6$)	distance between the experimental gauges (wg1 and wg2) placed over the structure crest

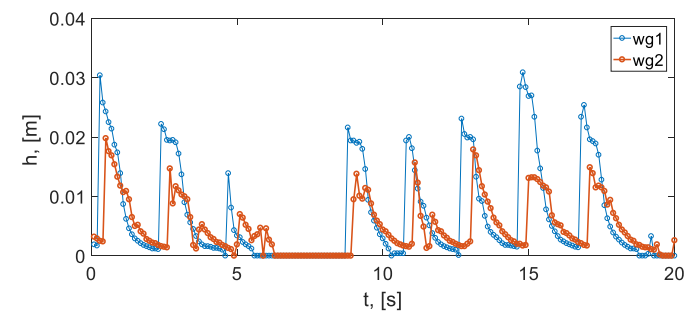
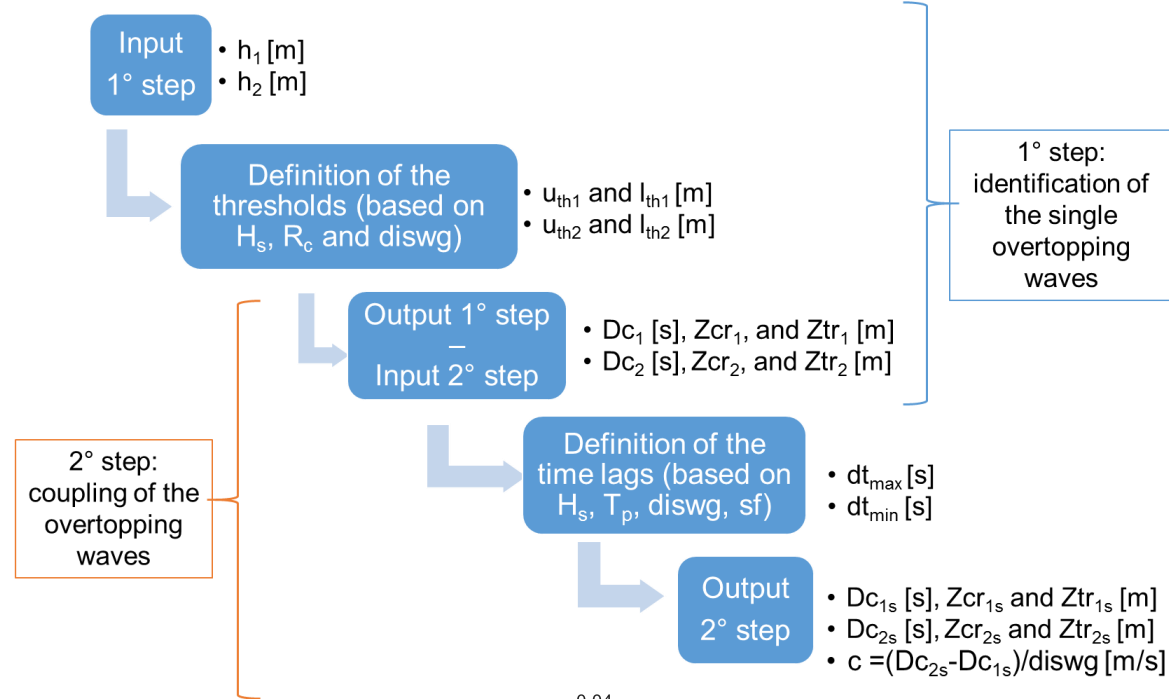
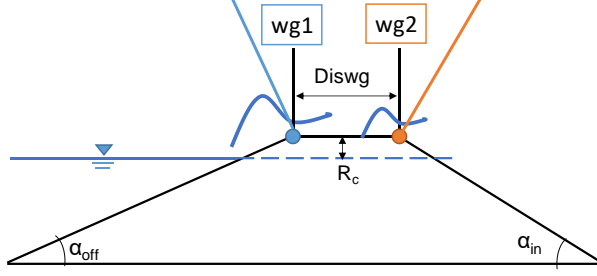
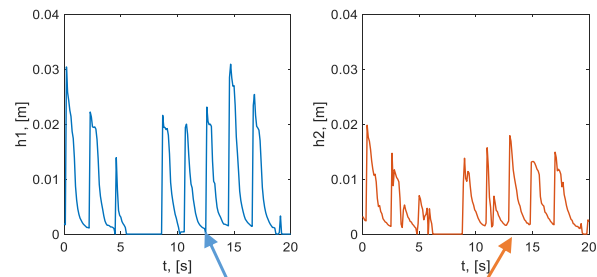
255 **3. Description and validation of the procedure**

256 The new procedure has been developed to i) identify the single overtopping waves from the
257 signals acquired at wave gauges (wgs) over the structure crest and ii) to couple 2 (or more) wave
258 signals registered at consecutive gauges, namely wg1 and wg2. The procedure can be applied
259 to both physical and numerical models, emerged and low-crested structures and to any structure
260 surface type (i.e. smooth or rough). It can process any kind of oscillatory signal in the time
261 domain, i.e. not exclusively water surface (h , hereinafter) signals but also, e.g., wave overtopping
262 discharge (q , hereinafter) time series.

263 The whole procedure is structured into 2 sequential steps, i.e. the identification of the waves and
264 the coupling. The first step can be applied to one or more h -signals (or q -signals) registered at
265 one or more wgs and it is independent of the second one. The second step instead requires at
266 least two h -signals at two wgs, which must be previously processed by the first step. The
267 conceptual layout of the full procedure is schematized in Figure 4.

268 The description of the wave identification and wave coupling steps is given in the Sub-sections
269 3.1 and 3.2 respectively. The validation of the procedure and the assessment of the adequacy
270 of the adopted criteria are respectively provided in Sub-sections 3.3 and 3.4.

271



272

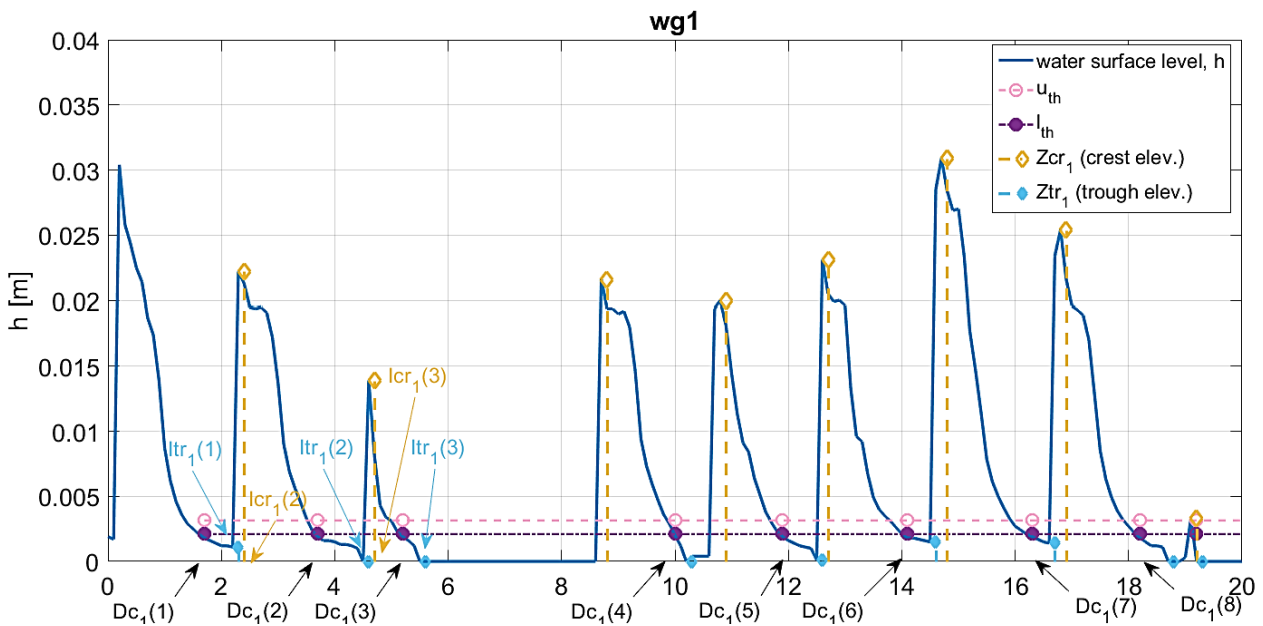
273 Figure 4 – Conceptual layout of the new procedure for the identification (1st step) and the coupling (2nd step) of the single wave overtopping
 274 waves.

275

276 **3.1. Wave identification**

277 The first step of the procedure is the wave identification algorithm, that is based on a time-domain
 278 threshold-down-crossing analysis (tdc, hereinafter) of the sea surface elevation h . The algorithm
 279 takes as input the time series of h -signals registered at one or more wgs and provides as outputs
 280 the time-ordered sequence of specific wave overtopping events. Each event can be described
 281 by the crest and trough wave heights (Z_{cr} and Z_{tr}) and the instants of zero-down-crossing (Dc).
 282 2 consecutive instants of tdc, namely $Dc(i-1)$ and $Dc(i)$, define the period $[Dc(i)-Dc(i-1)]$, the crest
 283 $Z_{cr}(i)$ and the trough $Z_{tr}(i)$ elevations of the i -th overtopping event. The procedure provides also
 284 the records of the instants of occurrence of Z_{cr} and Z_{tr} , named respectively l_{cr} and l_{tr} .

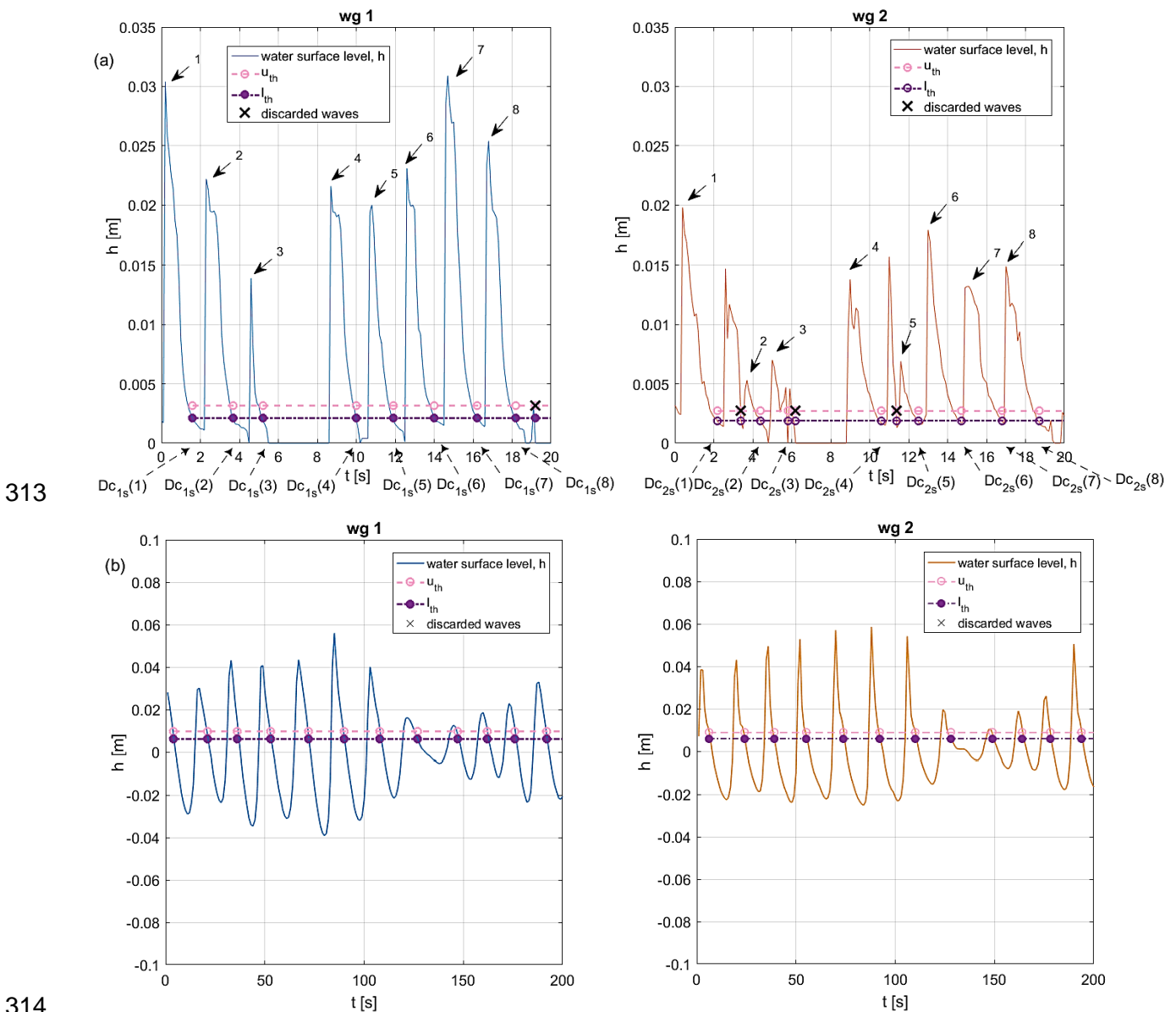
285 As a reference to the mentioned symbols, Figure 5 provides an example of a sea surface
 286 elevation signal (h) registered at wg1 and processed with the new procedure. The signal in this
 287 Figure refers to a numerical test (dataset UB-num) of wave overtopping at a smooth dike
 288 ($G_c=0.3$ m, $\cot(\alpha_{off})=4$, $\cot(\alpha_{in})=3$) in emerged conditions ($R_c/H_s=0.5$) and subjected to a
 289 perpendicular wave attack ($H_s=0.1$ m, $T_{m-1,0}=2.2$ s).
 290



291
 292 Figure 5 – Example of a sea surface elevation signal (h) registered at wg1 and processed with
 293 the new procedure. Data belonging to the example numerical test (UB-num dataset).

294
 295 The tdc algorithm for the wave identification depends on the definition of a threshold value of h
 296 representing the zero-control-value of the sea surface level. When h down-crosses this threshold
 297 value, a single event is identified. The value of this threshold, which is going to be referred as
 298 “lower threshold” l_{th} hereinafter, may be 0 (i.e. $l_{th}=0$ m) or a different – positive or negative –
 299 value, e.g. the still water level, depending on the nature of the signal itself. In Fig. 5, l_{th} is
 300 represented by the filled-in circles and is >0 to account for the water layer over the crest.

301 To better understand the procedure and its parameters, the different characteristics of the flow depth
 302 over the structure crest depending on the frequent or rare overtopping conditions should be
 303 observed. In case of rare overtopping (see the example at $R_o/H_s=1$ in Fig. 6-a), the sea surface
 304 elevation shows usually some bursts of small amplitude and almost instantaneous duration, which
 305 are signal noises or irregularities that typically precede or follow the waves and are due to the crest
 306 friction and/or wave breaking. The presence of these bursts increases with increasing R_o/H_s and with
 307 the structure roughness. Especially for permeable structures, the bursts are more frequent with
 308 increasing *diswg* (i.e. at wg2), since the water percolates in the mound (see Section 5.2). These
 309 irregularities should be discarded in the identification step, as they will also lead to a spurious
 310 coupling (see Sub-section 3.2). In case of frequent overtopping, the flow depths over the crest are
 311 larger and the wave events are instead well-defined (Fig. 6-b), resulting in a more regular signal.
 312



314 Figure 6 – h -signals measured at wg1 (left) and wg2 (right) at a dike at $R_o/H_s=1$ (panel a) and
 315 $R_o/H_s=-0.5$ (panel b). The coupled events and the corresponding instants tdc are indicated and
 316 numbered (panel a). The discarded uncoupled events are marked with black crosses.
 317

318 To ensure a correct identification (and coupling) of the waves also in case of extremely rare
319 overtopping, another threshold of the h -signals has been defined in the tdc algorithm. This
320 parameter is an upper threshold u_{th} , which is set greater than a certain percentage of l_{th} and it is
321 used to discard the “small” oscillations of h -signal whose amplitude is lower than the thresholds
322 difference ($u_{th}-l_{th}$).

323 It is not possible to univocally define the values of l_{th} and u_{th} , as the definition depends on the
324 nature (numerical, experimental) and the level of noise of the input signals and on the
325 characteristic of the single case to be processed. Based on the analyses carried out on the
326 available laboratory (datasets HT, HS, AAU and UB-exp) and numerical (UB-num) records, the
327 determination of l_{th} and u_{th} may vary according to:

- 328 • the emergence or submergence of the structure crest, R_c . The lower the R_c , the higher the
329 wave run-up and the lower the wave energy dissipation. Hence, for a given wave attack
330 (same H_s and $T_{m-1,0}$), the amplitude of the waves is generally lower over emerged structures
331 ($R_c>0$) than over submerged or zero-freeboard structures ($R_c\leq 0$). In conclusion, the
332 difference ($u_{th} - l_{th}$) should be lower for $R_c>0$ than for $R_c\leq 0$;
- 333 • the significant wave height, H_s . The difference ($u_{th} - l_{th}$) should be a function of H_s , as the
334 amplitude of the bursts of the h -signals to be discarded depends on the amplitude of the
335 incident waves;
- 336 • the distance between the wg1 and wg2, $diswg$, in case of 2 (or more) wgs. The greater
337 $diswg$, the greater the wave energy dissipation and the reduction of the amplitude of the
338 wave at wg2 with respect to wg1.

339 Though the parameters u_{th} and l_{th} can be customized by the user of the procedure upon
340 necessity, it is suggested to set the values of u_{th} and l_{th} accounting for the following
341 recommendations:

- 342 • in case of $R_c\geq 0$, l_{th} should be set equal to the level of the structure crest (typically, 0) or to
343 the minimum of the h -values if a small layer of water is present over the crest (see Fig. 6-a);
- 344 • in case of $R_c<0$, l_{th} should be set equal to the mean of the h -signal or to the still water level;
- 345 • based on the number of bursts to be discarded, u_{th} can be increased of a certain percentage
346 p of H_s with respect to l_{th} , i.e. $u_{th}=l_{th}+H_s/p$;
- 347 • to account for the dissipation over the crest, the difference ($u_{th} - l_{th}$) should be greater for
348 wg1 than for wg2, especially in case of $R_c>0$;
- 349 • to keep all the bursts, it is possible to simply set $u_{th}=l_{th}$.

350 **3.2. Wave coupling**

351 A “coupled event” (or coupled wave) is an event that is firstly identified at wg1 and consecutively
352 at wg2 after a certain time lag necessary to the wave propagation from wg1 to wg2. The coupling
353 occurs if a series of criteria are satisfied to ensure that the events recognized at the 2 wgs are
354 the same event that has propagated from wg1 to wg2.

355 Figure 6-a compares the results of the tdc applied to the signal h at wg1 and wg2 of the example
356 numerical test of Fig. 1. In Figure 6-a, h shows a non-negligible damping and a significant change
357 of the wave shape occurred in the propagation between wg1 and wg2. Due to the modification

358 of the signal, the tdc has identified a different number of events, and specifically: 9 events at
 359 wg1 and 11 waves at wg2. Figure 6-b provides for comparison the results of the tdc applied to
 360 the same dike in submerged conditions ($R_d/H_s=0.5$, $H_s=0.1$ m, $T_{m-1,0}=1.8$ s). In the case of Figure
 361 6-b, the h -signals at wg1 and wg2 are very similar, no significant shape irregularity appears and
 362 the same number of waves is recognized at wg1 and wg2.

363 The coupling step of the procedure is the algorithm that:

- 364 • processes the first-step-outputs, i.e. Zcr_1 , Ztr_1 , Dc_1 at wg1 and Zcr_2 , Ztr_2 , Dc_2 at wg2;
- 365 • checks the satisfaction of a set of criteria to couple the corresponding waves at the 2 wgs;
- 366 • provides as final output the time-ordered sequence of the “coupled events”, i.e. Zcr_{1s} , Ztr_{1s} ,
 367 Dc_{1s} and Zcr_{2s} , Ztr_{2s} , Dc_2 , see Figure 6-a. From now on, the subscript “s” is used to refer to
 368 the “coupled events”.

369 While the number of the elements of the sequences of Dc_1 and Dc_2 may differ depending on nw1
 370 and nw2, the lengths of Dc_{1s} and Dc_{2s} are equal. Each element of Dc_{1s} does have a
 371 corresponding element in Dc_{2s} . For each couple of elements [$Dc_{1s}(i)$, $Dc_{2s}(i)$], the difference
 372 $Dc_{1s}(i)-Dc_{2s}(i)$ represents the time of propagation of the single wave from wg1 to wg2.

373 The coupling criteria to be satisfied are based on the definition of the minimum and maximum
 374 time lags (dt_{min} and dt_{max}) that may occur for the wave propagation from wg1 to wg2. These time
 375 lags depend in turn on:

- 376 • $diswg$, see Fig. 1-a;
- 377 • the celerity c of the single waves, which depends also on the thickness of the water layer
 378 over the structure crest h ($c \propto \sqrt{gh}$). Therefore, c – and the time lags – vary also with R_c ;

379 In the coupling algorithm, dt_{min} and dt_{max} vary with the peak wave period T_p and with the sampling
 380 frequency sf (or sampling time interval) of the sea surface signal. Their values are defined as
 381 follows:

$$382 \begin{cases} dt_{min} = \max\left(\frac{diswg}{c_{dw}}, \frac{1}{sf}\right), & \text{with } c_{dw} = \frac{L_{0,p}}{T_p} \\ dt_{max} = \frac{diswg}{c_{sw}}, & \text{with } c_{sw} = \min(\sqrt{gh_1}) \end{cases}, \quad (1)$$

383 where c_{dw} and c_{sw} represent the celerity in deep water and in shallow water, respectively.

384 It is assumed that c_{dw} is the maximum possible wave celerity which therefore determines dt_{min} .
 385 In eq. (1), the peak wave period and length, T_p and $L_{0,p}=g \cdot T_p^2/(2\pi)$, are the deep water values
 386 that are used as theoretical upper limit for the wave celerity. Note that dt_{min} might not correspond
 387 to c_{dw} but to the minimum sampling time step $1/sf$ because of the inherent constraint imposed by
 388 sf . Actually, sf does affect the accuracy of the celerity assessment. A too-low value of sf may be
 389 insufficient to catch the wave propagation from wg1 to wg2 and the tdc analysis may register the
 390 i -th wave passage at the same time, i.e. $Dc_1(i)=Dc_2(i)$. When this happens, the coupling
 391 procedure is forced to discard the event, resulting into a loss of data and an underestimation of
 392 the maximum and mean value of c . The sufficiency/insufficiency of sf depends on several factors,
 393 such as: $diswg$, the actual celerity of the single waves, which depends in turn on the structure
 394 emergence/submergence, wave characteristics, etc. Lykke Andersen et al. (2011) suggested to

395 consider a minimum distance among the gauges that equals 5 samples, i.e. $5/sf$, assuming a
 396 maximum celerity of 2 m/s. In other words, $diswg$ should be $\geq 2 \cdot (5/sf) = 10/sf$. In this work, $diswg$
 397 was in the range $5/sf - 12/sf$. We can recommend that $diswg \geq c_{dw} / sf$, or at the contrary, that
 398 $sf \geq c_{dw} / diswg$.

399 On the contrary, c_{sw} is computed based on $\min(\sqrt{gh_1})$, i.e. on the minimum of the values h_1 of
 400 the water surface elevation recorded at wg1 at the instants lcr_1 . In other words, h_1 is the record
 401 of the values extracted from $h(wg_1, t = lcr_1(i), i=1, \dots, nw1)$ and it is therefore assumed the
 402 minimum possible c determining in turn dt_{max} . Clearly, the value of c_{sw} depends on the sensitivity
 403 of the wg and on the value of $(u_{th} - l_{th})$ adopted to detect the minimum h_1 -value.

404 It is important to remark that the coupling algorithm and the definition of dt_{min} and dt_{max} of Eq. (1)
 405 are valid for non-oblique waves only. For 3D waves, there might be a change in the flow direction
 406 over the structure crest, and therefore the wave celerities should be derived from a least-square
 407 fitting of the different time lags at a larger number of wgs (minimum 3, Lykke Andersen et al.,
 408 2011).

409 Based on c_{sw} and c_{dw} , the coupling procedure associates 2 waves if the time lag between the 2
 410 instants of tdc is included within $[dt_{min}; dt_{max}]$. All the identified events that cannot be coupled are
 411 discarded.

412 In Figure 6-a, the coupled events are marked with the arrows and numbered progressively to
 413 highlight the correspondences between the two plots. The Figure shows that 8 events have been
 414 coupled, while 1 event from wg1 and 3 events from wg2 have been discarded (black crosses
 415 upon the circles) because:

- 416 • they actually do not have a corresponding event at the other wg (see, for example, the event
 417 around 6 s or the wave around 11 s at wg2);
- 418 • despite one event is identified at wg1, its amplitude becomes lower than u_{th} or than $(u_{th} - l_{th})$
 419 at wg2 (see the event at 19 s at wg1 in Figure 6);
- 420 • they are just irregularities of the shape of one larger event: this is the case of the waves with
 421 $tdc \approx 3.5$ s and $tdc \approx 4.5$ s at wg2, which clearly belong to the same wave (nr. 2).

422 The result of the coupling procedure for the selected test (Figures 1 and 6-a) is qualitatively
 423 provided in Figure 7, which displays the wave signals at wg1 and wg2, once all the uncoupled
 424 events are discarded.

425 Once the time-ordered sequences of the coupled events DC_{1s} and DC_{2s} are derived, the time-
 426 ordered sequence of the c -values is obtained as follows:

$$427 \quad c(i) = \frac{diswg}{DC_{2s}(i) - DC_{1s}(i)}, \quad (2)$$

428 where i is the i -th element of DC_{1s} and DC_{2s} (i.e. the i -th coupled event).

429 In 5 cases of extremely rare overtopping, the coupling procedure failed because $nw2$ was
 430 significantly (>40-50%) higher than $nw1$. This happened: i) for permeable structures, when

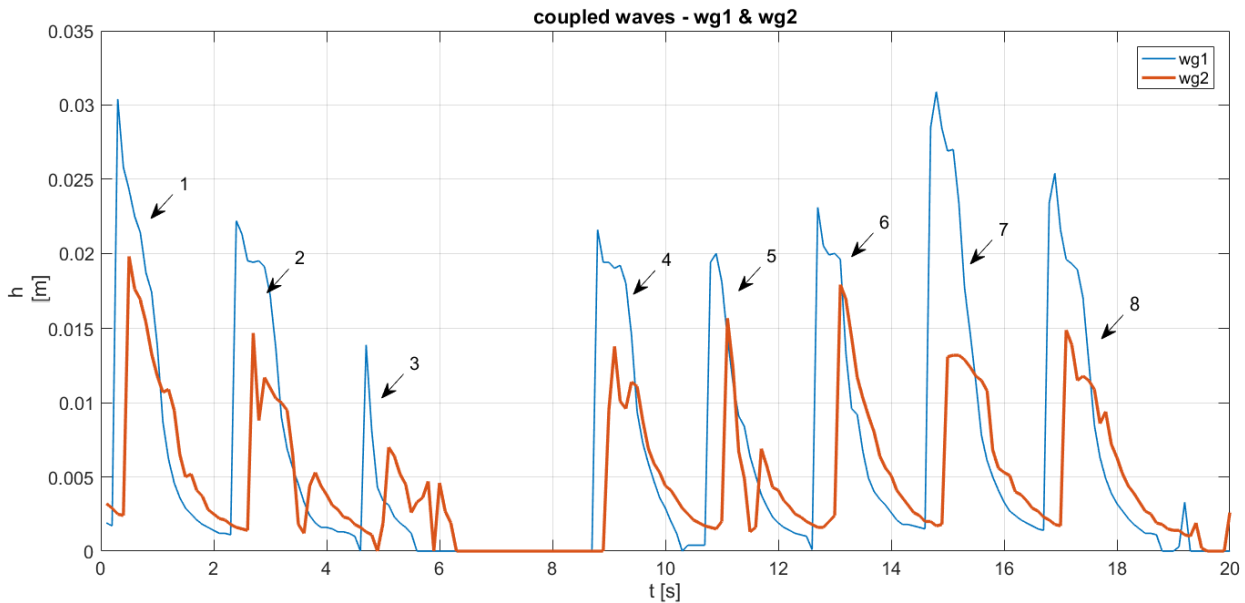
431 contemporarily $R_c > 0$ and $H_s/L_{m-1,0} > 4\%$ (2 out of 33 tests from dataset AAU), and ii) for smooth
 432 dikes when contemporarily $R_c/H_s > 1$ and $H_s/L_{m-1,0} \geq 4\%$ (3 out of 94 tests from dataset UB-num).

433 An algorithm has been implemented in the procedure to achieve the wave coupling also in these
 434 conditions by discarding more bursts at wg2 and consequently reducing nw2. Such algorithm
 435 iteratively increases of 5% the amplitude of $(u_{th} - l_{th})$ by reducing l_{th} and increasing u_{th} with respect
 436 to the first guess until $|nw1 - nw2|/nw1$ is lower than the 30%. The value of l_{th} cannot of course
 437 drop below 0. By applying this algorithm, the wave coupling was achieved for the 5 tests after
 438 one iteration only. The application of this algorithm is optional, but the user should be aware that
 439 if $(nw2 - nw1)/nw1 > 40\text{-}50\%$, the wave coupling might fail. This may happen in the following cases.

- 440 • Relatively rare overtopping, i.e. $q < 10^{-4} \text{ m}^3/(\text{sm})$;
- 441 • smooth structures at $R_c/H_s > 1$;
- 442 • permeable structures at $R_c/H_s > 0.5$ subjected to relatively steep waves, $H_s/L_{m-1,0} > 4\%$.

443 It was instead verified that in the opposite case, i.e. when $nw2 \ll nw1$, the coupling procedure
 444 does not fail, because all the waves that are detected at wg1 but not at wg2 are simply
 445 discarded. Typically, the case $nw2 \ll nw1$ occurs for rubble mound structures and high rates of
 446 volume percolation along the structure crest. The robustness of the procedure has been tested
 447 up to cases of percolation rates of 60-70%, which correspond to $(nw2/nw1) < 40\%\text{-}50\%$.

448



449

450 Figure 7 – Coupled h -signals registered at the off-shore and in-shore edges of the dike crest
 451 (wg1 and wg2, respectively). Data belonging to the same numerical test of Fig. 5.

452

453 As a final remark, the coupling algorithm can be applied to more than two wgs. In this case, the
 454 user is required to provide the h -signals and the $diswg$ values for each pair of wgs to be
 455 processed and the algorithm will compute the time-ordered sequences of coupled events for
 456 each pairs of wgs.

457 3.3. Validation

458 In this Sub-section the procedure is applied to the datasets HT and HS to respectively validate
459 the identification (Sub-section 3.3.1) and the coupling (Sub-section 3.3.2) step.

460

461 3.3.1. Identification of the overtopping volumes

462 Hughes and Thornton (2016) noticed that the identification of overtopping waves by means of
463 standard procedures may lead to inaccurate results and therefore adopted a “human supervised
464 automatic procedure”, which essentially consisted of a manual correction of the outcomes
465 provided by their automated procedure.

466 Here the new semi-automatic procedure is applied to the identification of the overtopping
467 volumes (V) of the same 8 experiments (Table 3) on wave overtopping selected by Hughes and
468 Thornton (2016) and elaborated with the supervised method (dataset HT).

469 The identification of the overtopping volumes V follows essentially the methodology proposed by
470 Hughes and Thornton (2016) and it is structured in the following steps.

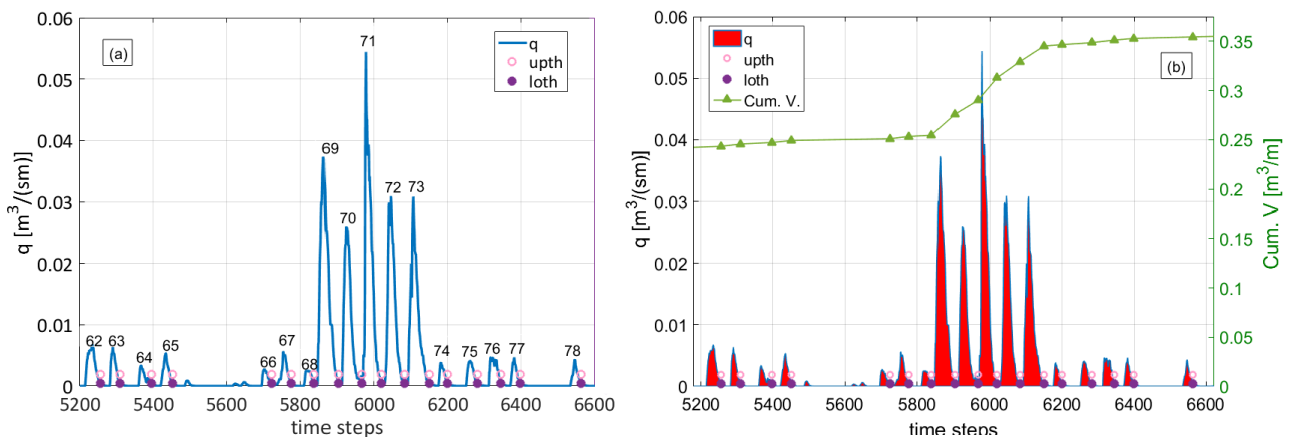
- 471
- 472 • Extraction of the time series of the numerical flow discharge (q) measured at the offshore
473 edge of the structure crest. For each time step, the instantaneous value of q is the result of
474 the product of the average value (along the vertical profile) of the cross-shore directed flow
475 velocity (u , m/s) by the value of the corresponding flow thickness (h , m). In case of an
476 offshore directed flow ($u < 0$), the value of q is set to 0. The example discharge time series
477 used by Hughes and Thornton, 2016 is here reported in Figure 8.
 - 478 • Automatic detection of the individual overtopping volumes V by means of the tdc analysis of
479 the discharge time series (Fig. 8-a). Note that the identification algorithm is here applied to
480 process a time-discharge signal (q) instead of a sea surface level signal (h) as in the
481 description of the algorithm itself (Sub-section 3.1). Therefore, in this case the values of the
482 2 threshold l_{th} and u_{th} are set as a functions of q (instead of H_s and R_c , as described in Sub-
483 section 3.1, where the procedure was applied to process a water level signal), and precisely:
 - 484 ○ the upper threshold u_{th} is set equal to the minimum between $q_{mean}/4$ and 10^{-4} m³/s per
485 m, where q_{mean} is the mean of the values of q ;
 - 486 ○ the lower threshold l_{th} is set equal to $q_{mean}/10$.

487 The value of $q=10^{-4}$ m³/s was selected by Hughes and Thornton (2016) as the minimum
488 “arbitrarily small instantaneous discharge value of q ”. In other words, Hughes and Thornton
489 (2016) “arbitrarily” decided to neglect the volumes corresponding to discharges $< 10^{-4}$ m³/s.
490 To identify the closest number of events to the number detected by Hughes and Thornton
491 (2016), the values $q_{mean}/4$ and $q_{mean}/10$ were introduced. By changing the values of l_{th} and
492 u_{th} , a higher number of overtopping volumes can be detected. The procedure is able to catch
493 any overtopping event by simply lowering u_{th} and the distance between u_{th} and l_{th} , no matter
494 how small the wave is.

494 In the example of Figure 8-a, the procedure recognized 17 overtopping waves (marked by
 495 the circles representing the thresholds) and discarded the small waves around 5600 time
 496 steps because their amplitude is lower than U_{th} . The automated procedure of Hughes and
 497 Thornton (2016) failed in the identification of the waves #69, #71 and #72, which were
 498 recovered by a visual examination of the signal.

- 499 • Wave-by-wave integration of q : for each overtopping wave recognized by the tdc procedure,
 500 the V -value is computed through the numerical integration of the overtopping signal. The
 501 integration is performed with the trapezoidal method, using a 0.025 s time-resolution, i.e. the
 502 highest possible resolution imposed by sf . Figure 8-b gives an idea of the degree of
 503 approximation of the integrating method by highlighting the areas swept by the single
 504 discharge waves corresponding to the single volumes (filled-in areas beneath the q signal).
 505 The line of the cumulated volumes is shown with reference to the right scale of Figure 8-b.

506



507

508 Figure 8 – Time evolution of the overtopping discharge (q) measured at the offshore edge of the
 509 dike crest (a) for one of the tests from the dataset HT, and corresponding overtopping volumes
 510 (b). In panel (b), the line of the cumulated volumes is indicated with reference to the right scale.
 511 The labelling of the waves is the same used by Hughes and Thornton (2016).

512

513 The results of the application of the new procedure to the dataset HT are summarized in Table
 514 3 in comparison to the results achieved with the supervised procedure of Hughes and Thornton
 515 (2016). The section of Table 3 dedicated to the results of Hughes and Thornton (2016) includes
 516 the number of V automatically (Auto V_{HT}) and manually (Man V_{HT}) identified, their sum (Total
 517 V_{HT}) and their ratio (Auto V_{HT})/(Total V_{HT}). The section dedicated to the results of the new
 518 procedure includes the automatically identified V (Auto V) and the ratio (Auto V)/(Total V_{HT}).
 519 Based on the total number of overtopping volumes (Total V_{HT} and Total V) recognized by the two
 520 procedures and the total number of incident waves (N_w), Table 3 reports also the resulting values
 521 of the probability of overtopping (Pow) for each test: $Pow = (Auto\ V)/N_w$, $Pow_{HT} = (Total\ V_{HT})/N_w$.

522 In 7 out of 8 cases, the new procedure recognizes more than the 90% of Total V_{HT} and for 5
 523 experiments, it identifies even more volumes than Total V_{HT} (percentages >100%), leading to

524 more cautious estimations of Pow than of Pow_{HT} . In the worst case (test #0199, 78%), the new
525 procedure recognizes anyway 1.6 times the number of the volumes recognized by HT without
526 the manual analysis (347 volumes instead of 211).

527 The new procedure provides a remarkably improved identification of the overtopping volumes,
528 because it gives at least the same results of HT without requiring any manual supervision.

529

530 3.3.2. Wave coupling

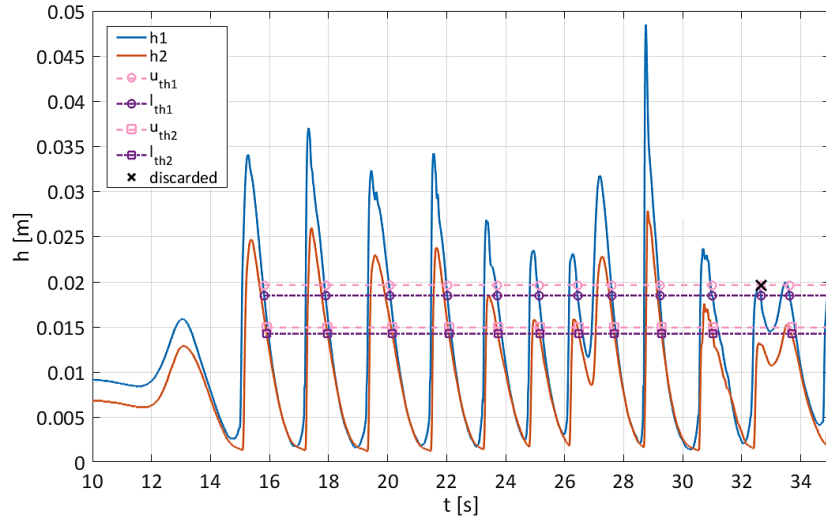
531 The wave celerity c , or the velocity of propagation of a wave front, can be estimated by measuring
532 the time interval occurring between the passage of a wave from one gauge to the following one.
533 In other words, c can be directly derived from the coupling of the wave signals registered at 2
534 consecutive wgs (see Eq. (2)).

535 The accuracy of the coupling step of the new procedure is here assessed by comparing
536 estimated values of c to experimental measures of flow velocity u over the crest of a dike
537 available from 3 tests by Hughes and Shaw (2011) (dataset HS, Table 4). Typically, the flow over
538 the crest of a dike occurs in shallow water and is characterized by broken waves only. In such
539 conditions, u and c should be approximately equal. More precisely, it can be assumed that the
540 maximum u value for breaking waves, u_{max} , is roughly similar to celerity c of the wave crest
541 (Losada et al., 2005). For example, Schüttrumpf and Oumeraci (2005) found the average ratio
542 between c and u_{max} equal to 0.97 with a standard deviation $\sigma\%$ of 29.9%. Since the wave breaking
543 generally occurs when the water depth is approximately equal to the wave height, it can be
544 concluded that in the case of the 3 tests from the dataset HS, the flow over the crest occurs in
545 broken conditions, being the values of $H_g/|R_c|$ range from 2.33 to 8.25 (see Table 4). In such
546 conditions, the approximation $u \approx c$ should be good and therefore the proposed validation for the
547 coupling step meaningful. More details about this topic are discussed in Sub-section 3.4.

548 The new procedure has been applied to the water level signals registered at the 2 wgs (ADVs)
549 placed over the crest of the structure. Based on the resulting time lags occurring between the
550 couples of overtopping waves ($Dc_{2s}-Dc_{1s}$), the outcomes of the procedure have been used to
551 derive the wave celerities (c) of the single waves. The time averaged values of c (c_{mean}) derived
552 for each tests are compared in Table 4 to the corresponding upper 2% values $u_{2\%}$ computed
553 from the time series of u measured by the ADV. For all the tests, c_{mean} is always slightly lower
554 than $u_{2\%}$. This can be explained considering that $u_{2\%} < u_{max}$ is used as estimator of the maximum
555 flow velocity. Anyway, the agreement between $u_{2\%}$ and c_{mean} is remarkable, being the ratio
556 $c_{mean}/u_{2\%}$ included between 1.01 (test R109) and 1.14 (test R18).

557 Figure 9 gives a qualitative idea of the accuracy of the procedure showing the identified and
558 coupled waves during the first few seconds of the water level signals registered at the 2 wgs for
559 the test R14. Note that the wave identified at wg1 around 31 s has been correctly discarded as
560 it is just a noise in the wave signal. The procedure has correctly coupled most of the overtopping
561 waves for each test providing meaningful estimates of the celerities. 2 special cases can be

562 observed in this plot: the wave between 26 and 28 seconds and the wave at 32 s. In both cases
 563 we observe a single double-peaked wave, however in the first case the 2 peaks are recognized
 564 as two waves, while in the second case the first peak is disregarded because of the threshold
 565 values.



566
 567 Figure 9 – Identification and coupling of the overtopping waves performed by the new procedure
 568 for one of the lab tests on levees belonging to the dataset HS (test R14). $R_d/H_s=-0.155$.

569

570 **3.4. Assessment of the validity of the assumptions**

571 This Sub-section provides a detailed analysis of the wave celerities as outputs of the coupling
 572 step of the procedure, with the specific aims of i) further investigating and demonstrating the
 573 validity of the assumption that $u \approx c$ for broken or breaking flow conditions (Sub-section 3.4.1); ii)
 574 assess the reliability of the coupling step also in case of non-broken flow conditions (Sub-section
 575 3.4.2).

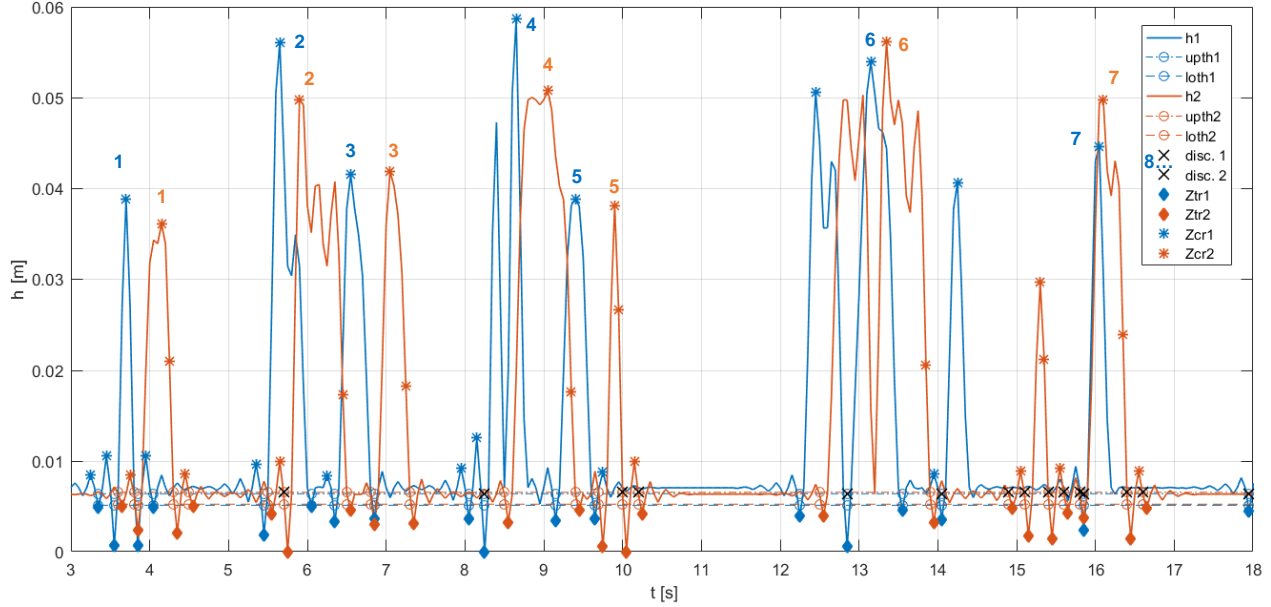
576 **3.4.1. Breaking and broken flow over the crest**

577 The validity of the assumption that for breaking or broken waves $u \approx c$ is checked here by applying
 578 the whole procedure to the UB-exp dataset (Table 5). Since all the tests belonging to UB-exp
 579 were conducted at $R_c \geq 0$, the flow thickness over the dike crest is due to the water that exceeds
 580 the crest level, i.e. at each instant $h \leq (H - R_c)$, where H is the height of the generic incoming wave.
 581 Therefore, the flow over the dike crest certainly occurs in breaking or broken wave conditions
 582 (Losada et al., 2005).

583 To calculate the celerities, the following methodology was adopted. The 2 steps of the procedure
 584 have been consequently applied to the time signals h_1 and h_2 derived from the UVP
 585 measurements at D4 and D6, approximately situated at the off-shore and in-shore edges of the
 586 dike crests during the UB-exp experiments (see Fig. 2 and Tab. 5). The positions D4 and D6

587 represent therefore $wg1$ and $wg2$ for the UB-exp and will be referred in this way hereinafter. An
 588 example of the h_1 and h_2 signals for one of the tests ($R_d/H_s=1$, $H_s=0.05$ m, $s_{m-1,0}=3\%$, $\cot(\alpha_{off})=4$,
 589 $G_c=0.30$ m) is shown in Figure 10. Each couple of waves identified and coupled is numbered in
 590 the Figure, while the waves discarded from the wave coupling are marked with black crosses.

591



592

593 Figure 10 – Identification and coupling of the overtopping waves performed by the new procedure
 594 for one of the test from the dataset UB-exp.

595

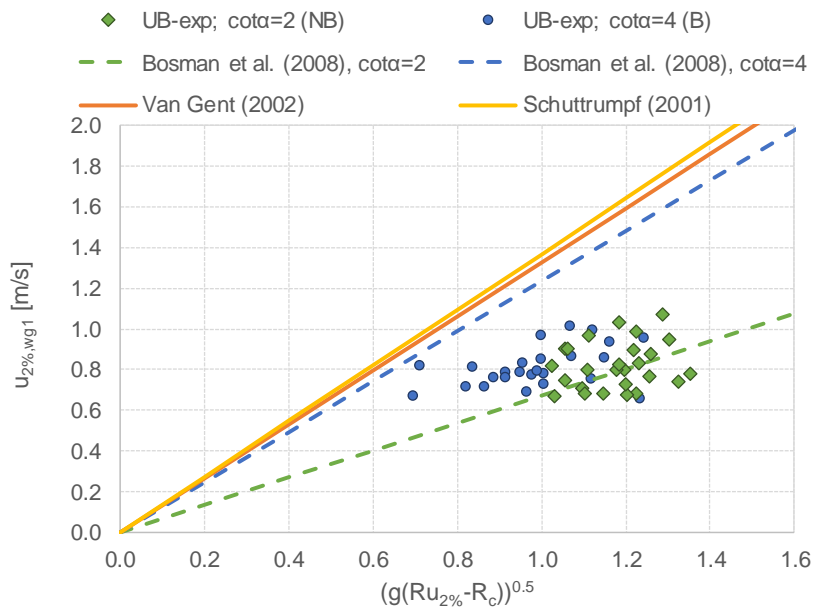
596 For each test, the outputs of the wave coupling has been further processed to derive the time
 597 series of c based on Eq. (2). Finally, the time-average values of c , c_{mean} , have been calculated.

598 As for the flow velocities, the vertical profiles of u measured with the UVPs at $wg1$ have been
 599 elaborated as follows. Firstly, for each instant, the depth-averaged values of u have been derived
 600 from the vertical profiles. The resulting time series of the average u -values have been sorted
 601 and the upper $u_{2\%}$ velocity values – i.e. the values of u exceeded by the 2% of the incoming
 602 waves – have been derived. The maximum velocities, u_{max} , have been also calculated to be
 603 compared to c_{mean} , according to Losada et al. (2005) and Schüttrumpf and Oumeraci (2005).

604 To qualitatively characterize the accuracy of the data collected with the UVPs, the $u_{2\%,wg1}$ values
 605 are displayed in Figure 11 as a function of the quantity $(g(Ru_{2\%}-R_c))^{0.5}$, where $Ru_{2\%}$ is the wave
 606 run-up exceeded by the upper 2% of the incoming waves and is computed following EurOtop
 607 (2018). The quantity $(g(Ru_{2\%}-R_c))^{0.5}$ is used in the formulae by Schüttrumpf (2001), Van Gent
 608 (2002) and Bosman et al. (2008) to estimate $u_{2\%}$ at the dike off-shore edge according to the
 609 following expression:

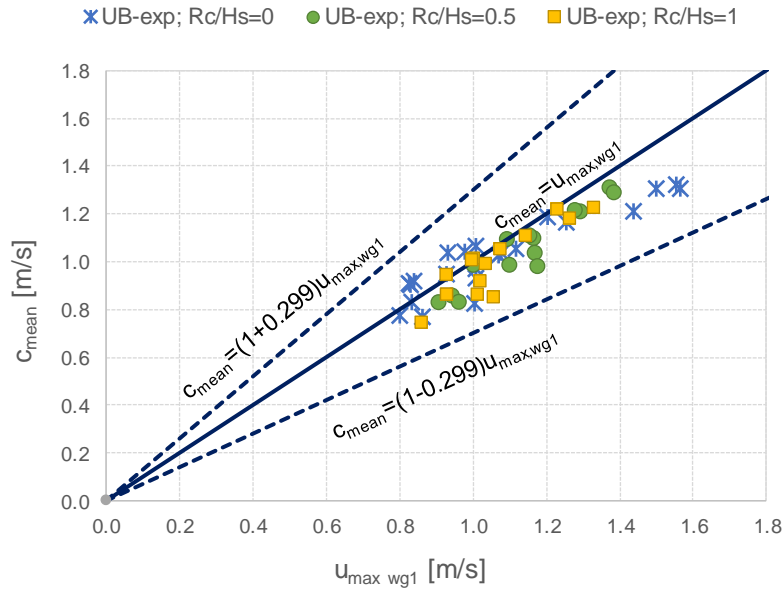
610
$$u_{2\%} = c_u \cdot [g(R_{u,2\%} - R_c)]^{0.5}, \quad R_c \geq 0, \quad (3)$$

611 where $c_u=1.37$ for Schüttrumpf (2001), $c_u=1.33$ for Van Gent (2002) and $c_u=0.30/\sin(\alpha_{off})$ for
 612 Bosman et al. (2008). In the formulation by Bosman et al. (2008), the coefficient c_u is made
 613 varying with α_{off} to account for the effects of the structure slope, and therefore it is supposed to
 614 update the other formulations by Van Gent (2002). It is worthy to remember that Bosman et al.
 615 (2008) fitted the formulation of c_u on values of $\cot(\alpha_{off})=4$ and 6, only. By comparing the $u_{2\%,wg1}$
 616 values to the curves representing the literature formulae in Figure 11, it can be appreciated how
 617 all the data generally follow the same trend with $(g(Ru_{2\%}-R_c))^{0.5}$ indicated by the formulae. The
 618 data at $\cot(\alpha_{off})=2$ (diamonds) seems to be well fitted by the curve by Bosman et al. (2008)
 619 extrapolated for $\cot(\alpha_{off})=2$. The data at $\cot(\alpha_{off})=4$ (circles) are on average slightly higher than
 620 the data at $\cot(\alpha_{off})=4$, though no particular effect of the slope angle is evident. The expressions
 621 by Bosman et al. (2008) for $\cot(\alpha_{off})=4$, by Schüttrumpf (2001) and by Van Gent (2002) represent
 622 upper envelopes to the data.
 623



624
 625 Figure 11 – Comparison among the upper 2% values of u measured at $wg1$ ($u_{2\%,wg1}$) and the
 626 literature formulae as a function of the wave run-up $(g(Ru_{2\%}-R_c))^{0.5}$. Data from UB-exp.

627
 628 As long as the approximation $U_{max} \approx C_{mean}$ holds, it is expected that the c_{mean} values are in good
 629 agreement with the $U_{max,wg1}$ values. The comparison among $U_{max,wg1}$ and c_{mean} is proposed in
 630 Figure 12. The chart includes the bisector (continuous) line representing the optimal condition
 631 $U_{max,wg1}=C_{mean}$ and the range of $\pm 29.9\%$ around it (dashed lines), where 29.9% is the standard
 632 deviation between u and c found by Schüttrumpf and Oumeraci (2005). In the Figure, all the data
 633 are straightly included within the $\pm 29.9\%$ bands, denoting a higher level of agreement between
 634 celerities and velocities. Specifically, it is found that on average $c_{mean}/U_{max,wg1} = 0.95$ with a
 635 standard deviation $\sigma\% = 7.4\%$. These findings support and prove the validity of the assumptions
 636 made to verify the accuracy of the coupling step for flow in broken or breaking wave conditions
 637 (Sub-section 3.3.2).



638

639 Figure 12 – Comparison among mean wave celerities (c_{mean}) estimated with the new procedure
 640 applied at wg1 and wg2 and the max values of u measured at wg1 ($u_{max,wg1}$). Data from UB-exp.

641

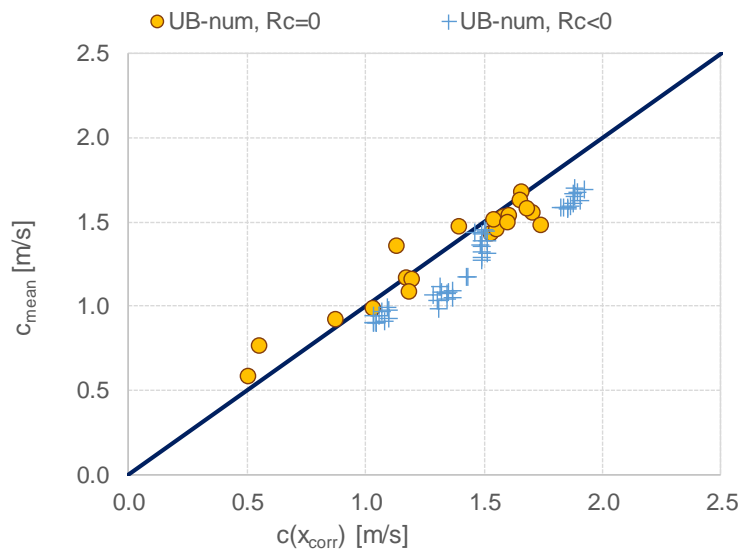
642 3.4.2. Non-breaking flow over the crest

643 In case of structures at negative crest freeboard, the assumption that the flow is characterized
 644 by broken waves only holds no more. The overtopping process includes a storm surge overflow
 645 component. In such conditions, the approximation $c \approx u$ is no more realistic and cannot be used
 646 to verify the consistency of the celerity values computed by the new procedure. With the
 647 exception of human visualization of the coupled waves, no reference method is available for a
 648 rigorous verification of the coupling step of the procedure in case of $R_c \leq 0$.

649 An indication about the reliability of the coupling step of the new procedure at any crest level,
 650 and specifically in case of $R_c \leq 0$, can be obtained by cross-correlating the 2 h -signals at the 2
 651 consecutive wgs. The average time lags between the 2 h -signals resulting from the cross-
 652 correlation, $lag(x_{corr})$, can be compared to the mean values of the array $(Dc_{1s}-Dc_{2s})$ calculated
 653 with the new procedure. The mean wave celerities $c(x_{corr}) = lag(x_{corr}) / diswg$ and $c_{mean} = mean(Dc_{1s}-$
 654 $Dc_{2s}) / diswg$ can be also compared. It is expected that the statistical distribution of the c -values
 655 is more uniform for submerged than for zero-freeboard structures for the presence of the
 656 constant over-flow component over the structure crest. Therefore, the average $c(x_{corr})$ resulting
 657 from the cross-correlation are expected to be more representative of the actual distribution of
 658 the c -values, i.e. that $c(x_{corr}) \approx c_{mean}$ when $R_c < 0$.

659 The cross-correlation function has been applied to the available data at $R_c/H_s \leq 0$ from the dataset
 660 UB-num. In this case, wg1 and wg2 are the numerical gauges placed in proximity of the off-
 661 shore and in-shore edges of the dike crest (see Fig. 1), respectively providing the time signals
 662 h_1 and h_2 . The same methodology described in Sub-section 3.4.1 has been followed to derive
 663 the c_{mean} values from the outcomes of the procedure. The comparison among the values of

664 $c(x_{corr})$ and c_{mean} is qualitatively provided in Figure 13, by grouping the data at $R_c=0$ and $R_c<0$.
 665 This Figure suggests that the results of the procedure and of the cross-correlation are very
 666 similar and the scatter is limited ($R^2=0.96$ for both the groups of data). The data at $R_c=0$ are
 667 more symmetrically distributed around the bisector line ($c_{mean}/c(x_{corr})=1.00$, with $\sigma\%=11\%$), while
 668 the values of c_{mean} tend to be slightly lower than $c(x_{corr})$ when $R_c<0$ ($c_{mean}/c(x_{corr})=0.88$, with
 669 $\sigma\%=5\%$). Actually, the greater the water layer over the structure crest h , the higher the c -values.
 670 In case of $R_c<0$, it is likely that the real c values exceed the upper limit of c (≈ 2.7 m/s) that can
 671 be estimated with the new procedure due to $sf = 20$ Hz. In conclusion, this analysis confirms
 672 that the results of the procedure are meaningful and physically coherent, and the main limit is
 673 represented by sf .
 674



675
 676 Figure 13 –Values of the wave celerities ($c(x_{corr})$) derived from cross-correlation of the 2 h -signals
 677 at wg1 and wg2 compared to the mean wave celerities (c_{mean}) estimated with the new procedure.
 678 Data from UB-num at $R_c/H_s \leq 0$.

679

680 4. Wave overtopping at smooth dikes

681 In this Section the principal results and applications of the procedure are presented: i) the
 682 estimation of the probability of overtopping (Sub-section 4.1); ii) the calculation of the average
 683 overtopping discharge (Sub-section 4.2); iii) the identification of the extreme overtopping
 684 volumes (Sub-section 4.3).

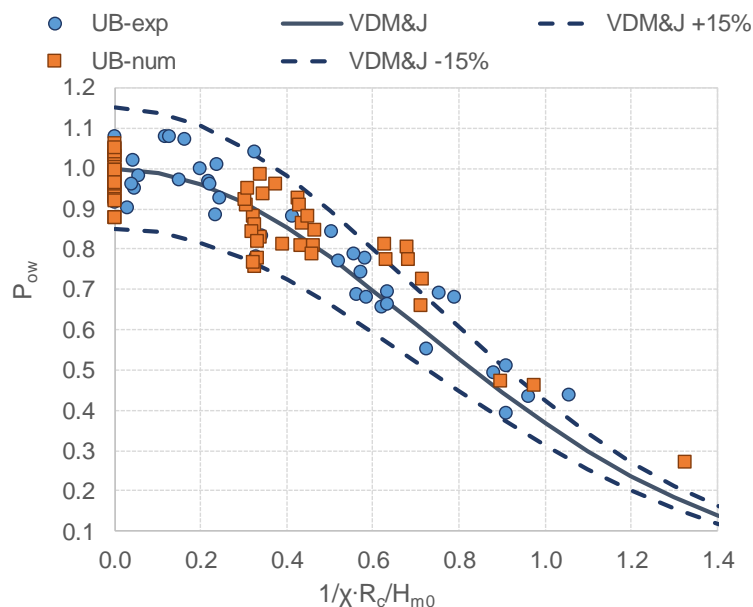
685

686 **4.1. Probability of overtopping**

687 One direct application of the procedure is the calculation of the probability of overtopping of a
 688 structure P_{ow} , which can be easily determined by dividing the number of the overtopping waves
 689 (or volumes) identified by the first step to the total number of incident waves (see Sub-section
 690 3.3.1). According to Van der Meer and Janssen (1994), P_{ow} at dikes can be estimated from the
 691 following expression:

692
$$P_{ow} = \exp\left(-\left(\frac{1}{\chi} \cdot \frac{R_c}{H_{m0}}\right)^2\right), \text{ with } \chi \approx 0.51 \cdot \frac{R_{u,2\%}}{H_{m0}} \text{ and } R_c > 0. \quad (4)$$

693 Figure 14 shows the distribution of the values of P_{ow} obtained from the application of the new
 694 procedure to the tests at $R_c \geq 0$ of the datasets UB-exp and UB-num in comparison to the curve
 695 representing Eq. (4). For both UB-exp and UB-num, the P_{ow} values are calculated based on the
 696 results of the identification step of the procedure applied at wg1 (see Figures 1 and 2). In Figure
 697 14, the data are plotted as function of the quantity $\frac{1}{\chi} \frac{R_c}{H_{m0}}$ and are grouped by dataset. The Figure
 698 indicates that all the data follow the trend of the formula with $\frac{1}{\chi} \frac{R_c}{H_{m0}}$ and most of them are included
 699 in a range of $\pm 15\%$ (dashed lines) around Eq. (4) (continuous line). The 7 outliers fall anyway
 700 within $+20\%$ and are all positioned over the curve. This slight bias suggests that the procedure
 701 tends to give cautious estimations of P_{ow} for $\frac{1}{\chi} \frac{R_c}{H_{m0}} > 0.6$. In a few cases at zero or modest
 702 freeboard, there is a small overestimation of $P_{ow}=1$: this happens when the procedure does not
 703 discard all the bursts in the wave signal generated by the wave breaking and identifies a number
 704 of waves slightly larger (less than 10%) than the number of the incident waves.
 705



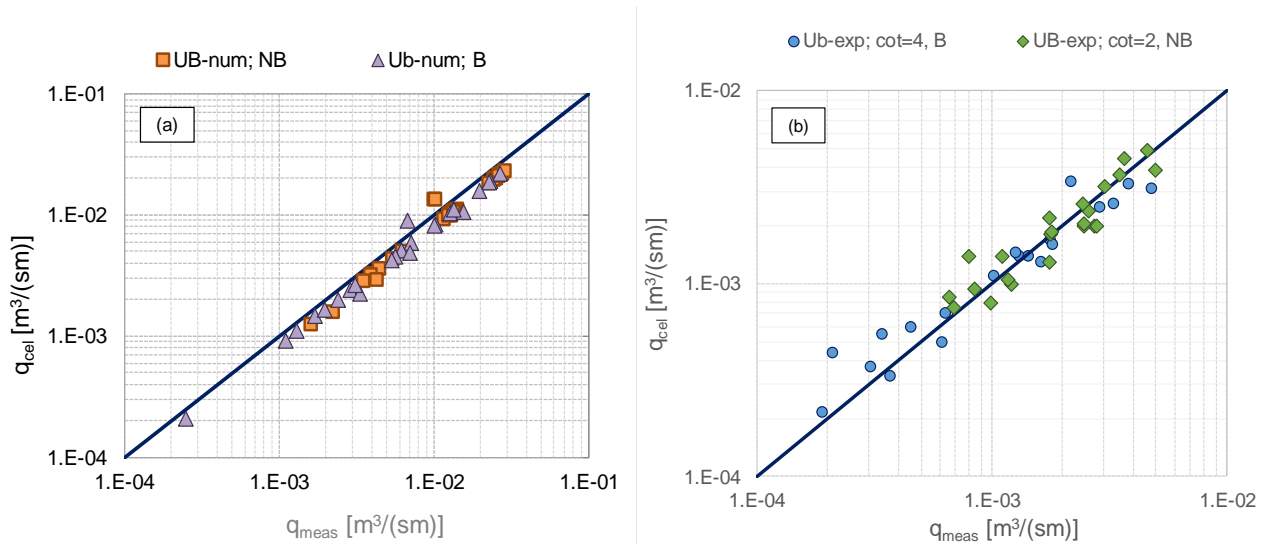
706
 707 Figure 14 – Probability of overtopping (P_{ow}) calculated by applying the identification step of the
 708 procedure to the datasets UB-exp and UB-num compared with the predictions by the formula by
 709 Van der Meer and Janssen, 1994 (VDMandJ, continuous line). The range of $\pm 15\%$ around the
 710 formula is shown (VDMandJ $\pm 15\%$, dotted lines). Data at $R_c/H_s \geq 0$.

711 4.2. Overtopping discharges

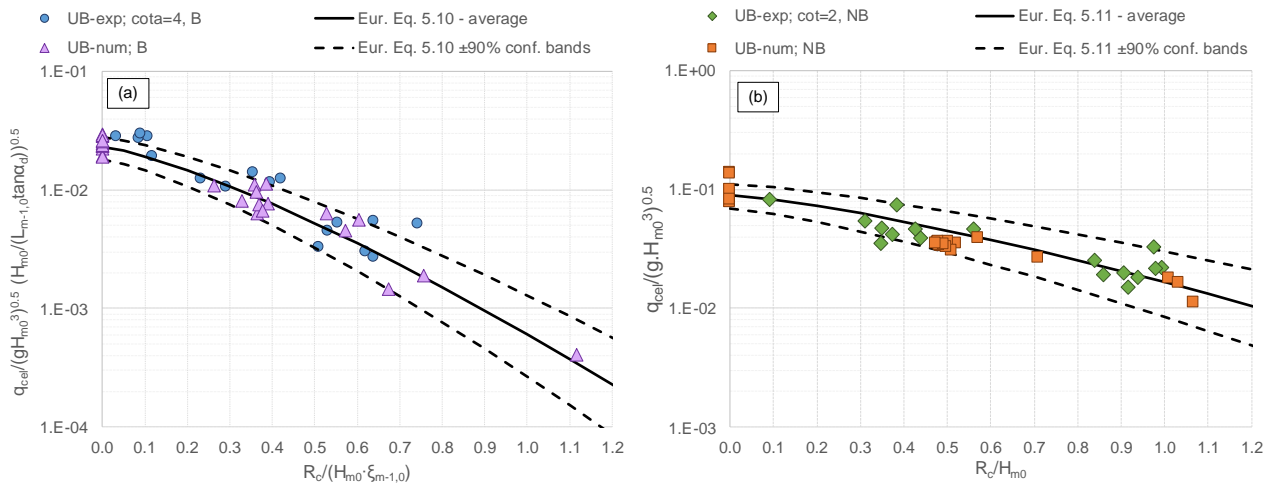
712 The wave celerities c can be also integrated with the water levels h and used to estimate the
713 instantaneous and average wave overtopping discharge q . As further application of the
714 procedure, this Sub-section presents the average values of q derived from the c values of the
715 datasets UB-num and UB-exp, q_{cel} . Since the assumption that $c \approx u$ holds only for flow in breaking
716 wave conditions (see Sub-section 3.4), the extraction of q_{cel} from c_{mean} was done for the tests at
717 $R_c/H_s \geq 0$ only. Starting from the time series of c and h_1 (water level signal at wg1 at the dike off-
718 shore edge), the time series of q were firstly obtained by multiplying the instantaneous values of
719 c by h_1 . Then, the q_{cel} values were calculated for each test as the time-averages of the time series
720 of q .

721 In Figure 15, the values of q_{cel} derived for UB-num and UB-exp (panels a and b, respectively)
722 are directly compared to the average values of q available from the numerical and experimental
723 measurements, named q_{meas} . In case of UB-num, the data of q_{meas} correspond to the time-
724 averaged values of q derived from the integration of the numerical flow velocities with the flow
725 depths, while in case of UB-exp q_{meas} are the average q values measured from the overtopping
726 tank. In both the charts of Figure 15, the data are grouped by breaking (B) and non-breaking
727 (NB) wave conditions. Such distinction follows EurOtop (2018) and it is based on the values of
728 the Iribarren-Battjes breaker parameter $\xi_{m-1,0} = \tan(\alpha_{off}) / (s_{m-1,0})^{0.5} \leq 2$ (B) or > 2 (NB). It is worthy to
729 stress that the distinction between B and NB based $\xi_{m-1,0}$ refers to the occurrence of the wave
730 breaking along the dike off-shore slope, i.e. before the dike crest. The overarching superposition
731 for the analysis of q_{cel} is that the flow is already broken over the dike crest independently of $\xi_{m-1,0}$.

732 In Figure 16-a,b, the q_{cel} values are made dimensionless through $(gH_{m0}^3)^{0.5}$ and compared to the
733 EurOtop (2018) equations 5.10 and 5.11 for the prediction of the average q at dikes under B and
734 NB waves, respectively. The values of H_{m0} have been calculated from the spectral analysis of
735 the incident wave signals both in the numerical and in the laboratory channel. In these charts,
736 the data are grouped by dataset (UB-num and UB-exp) and are displayed as functions of the
737 relative crest freeboard. The curves representing the formulae are shown as continuous lines
738 while the interval of $\pm 5\%$ around the curves themselves is represented through dashed lines,
739 following EurOtop (2018).



740
 741 Figure 15 – Comparison among values of q (q_{meas}) measured from the overtopping tank (UB-
 742 exp) or derived by the integration of the numerical flow velocities with the flow depths (UB-num)
 743 and corresponding values (q_{cel}) calculated from the wave celerities obtained with the new
 744 procedure. Data at $R_c/H_s \geq 0$ grouped by breaking (“B”) and non-breaking wave conditions (“NB”).
 745



746
 747 Figure 16 – Dimensionless values of q calculated from the wave celerities (q_{cel}) obtained with
 748 the new procedure compared to the curves (continuous lines) representing the EurOtop (2018)
 749 formulae for the prediction of q . Data at $R_c/H_s \geq 0$ grouped by breaking (“B”) and non-breaking
 750 wave conditions (“NB”). Datasets UB-exp and UB-num in panel a and b, respectively.
 751

752 Figure 15 gives a qualitative idea of the level of agreement between q_{cel} and q_{meas} , while Figure
 753 16 is meant to illustrate the trend of the q_{cel} values with the physical parameters and the
 754 predicting formulae. The quantitative indexes assessing the agreement between q_{cel} and q_{meas}
 755 and between q_{cel} and the predictions of q obtained with the EurOtop (2018) equations (q_{Eur})
 756 are collected in Table 7. For each dataset and each group of data (B and NB), Table 7 provides the
 757 values of the average ratios q_{cel}/q_{meas} and q_{cel}/q_{Eur} , the corresponding standard deviations $\sigma\%$
 758 and the coefficients of determinations R^2 . Overall, the agreement of the q_{cel} values with q_{meas} is

759 remarkable, considering that R^2 ranges between 0.92 and 0.98 and that $q_{cel}/q_{meas}=0.80-1.09$.
760 Both Figure 15-a and Table 7 indicate that q_{cel} tend to slightly but systematically underestimate
761 the q_{meas} values of UB-num ($q_{cel}/q_{meas}=0.80$ and 0.81 for B and NB waves, respectively). The
762 cause of this underestimation is again the upper limit of c at ≈ 2.7 m/s imposed by the combination
763 of sf and $diswg$. However, the scatter between q_{cel} and q_{meas} associated to UB-num is extremely
764 limited ($R^2=0.97-0.98$, $\sigma\%=9.9\%-11\%$) and the agreement with the trend suggested by the
765 EurOtop formulae (Fig. 16-a,b) is good, especially in case of B waves (Fig. 15-a, $R^2=0.96$). As
766 for UB-exp, Fig. 15-b shows that most of the data are symmetrically distributed in proximity of
767 the bisector line. The slightly higher standard deviations ($\sigma\%=32\%-34\%$) with respect to UB-num
768 can be explained with the greater noise associated to the experimental signals (compare Fig. 10
769 to Fig. 7) and to the higher level of uncertainty associated to the lab measurements (see Sub-
770 section 2.4).

771

772 Table 7. Comparison between q_{cel} and q_{meas} and between q_{cel} and the predictions by the EurOtop
773 (2018) formulae (q_{Eur}). Datasets UB-exp and UB-num.

Dataset	q_{cel} VS q_{meas}			q_{cel} vs EurOtop (2018)		
	mean(q_{cel}/q_{meas})	$\sigma\%(q_{cel}/q_{meas})$	$R^2(q_{cel}, q_{meas})$	mean(q_{cel}/q_{Eur})	$\sigma\%(q_{cel}/q_{Eur})$	$R^2(q_{cel}, q_{Eur})$
UB-num, B	0.80	11%	0.98	0.98	8.5%	0.96
UB-num, NB	0.81	9.9%	0.97	0.92	32%	0.86
UB-exp, B	1.09	34%	0.92	1.09	35%	0.89
UB-exp, NB	1.06	32%	0.94	1.01	24%	0.95

774

775 4.3. Overtopping volumes

776 The wave overtopping volumes V are commonly treated as stochastic variables and associated
777 with a probability distribution. According to the literature (EurOtop, 2018), the probability
778 distribution of the exceedance of V can be approximated by the Weibull distribution:

$$779 P(V_i \geq \bar{V}) = \exp\left(-\left(\frac{V_i}{a}\right)^b\right), \quad (5)$$

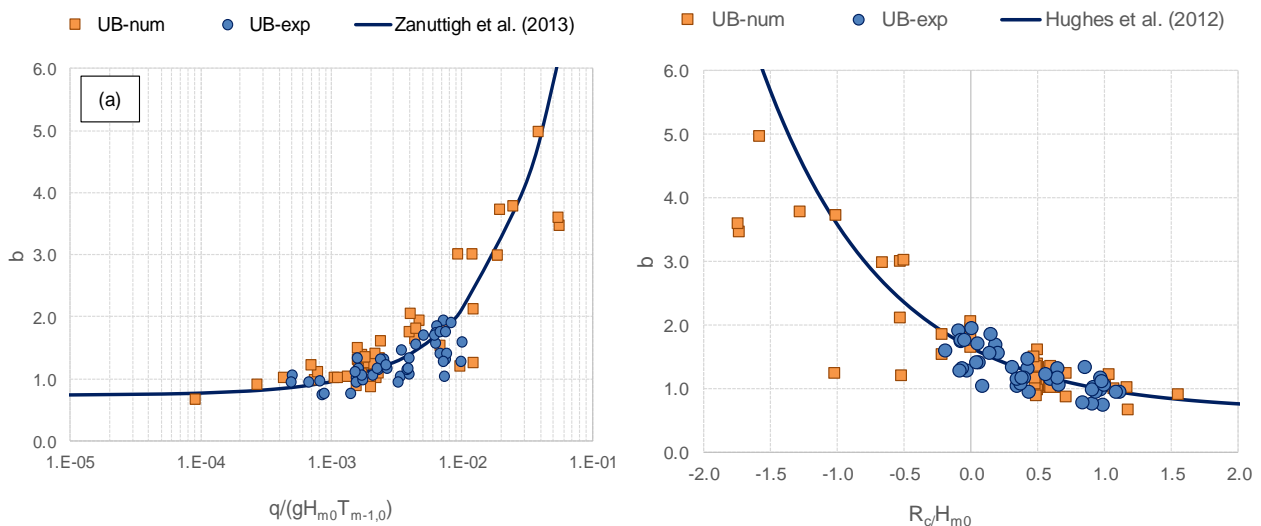
780 where $P(V_i \geq \bar{V})$ (or simply P , hereinafter) is the probability that the i -th individual volume V_i is
781 greater than a specified volume \bar{V} . The parameters a and b are the Weibull's scale and shape
782 factors, respectively. a generally corresponds to the mean of the distribution of the V -values (Van
783 der Meer and Janssen, 1994), while b is the slope of the straight line fitting the Weibull's
784 distribution of the "extreme" V -values in a double-log chart $\ln(V/a)$ vs $\ln(-\ln(P))$. Several methods
785 are proposed in the literature to identify the "extreme" V -values (upper 50%, upper 20%, upper
786 10%, human selection of the upper tail on the distribution), see for instance Pan et al. (2016)
787 and Molines et al. (2019). In this work, the automatic selection of the upper 20% V -values was
788 preferred as best compromise, see the analysis performed by Formentin and Zanuttigh (2018,

789 b). However, the choice of the percentage of the data to be used seems not to significantly affect
 790 the estimation of the maximum V -values as already demonstrated by Pan et al. (2016).

791 The methodology described in Sub-section 3.3.1 for the dataset HT has been also applied to the
 792 datasets UB-num and UB-exp to derive the individual overtopping volumes and the relative
 793 exceedance probability distribution P . Therefore, the values of the Weibull's shape factor b have
 794 been automatically extracted for each test. The results are reported in Figure 17 in comparison
 795 to the 2 relationships for the prediction of b at smooth structures proposed by: i) Zanuttigh et al.
 796 (2013), where b is a function of the dimensionless average q associated to the single test,
 797 $q/(gH_s T_{m-1})$; ii) Hughes et al. (2012), where b is a function of R_o/H_s .

798 Figure 17 shows that on average the b -values follow the trends of both the formulae. The
 799 agreement among values of b and the fitting by Zanuttigh et al. (2013) is characterized by values
 800 of $\sigma_{\%}=34\%$ and 41% and values of $R^2=0.74$ and 0.79 for the UB-exp and UB-num datasets,
 801 respectively. The agreement with the formula proposed by Hughes et al. (2012) gives instead
 802 $\sigma_{\%}=28\%$ and 60% and $R^2=0.78$ and 0.71 for UB-exp and UB-num, respectively. Overall, the data
 803 from UB-num seem to be better represented by Zanuttigh et al. (2013) while the data from UB-
 804 exp are in better agreement with Hughes et al. (2012). In Fig. 17-a, the greatest scatter is
 805 concentrated around $q_{mean}/(gH_s T_{m-1}) \approx 6 \cdot 10^{-2} \text{ m}^3/(\text{s} \cdot \text{m})$, while in Fig. 17-b a few b -values belonging
 806 to UB-num are overestimated by Hughes et al. (2012). Overall, both the charts show no
 807 heteroscedasticity towards neither q_{mean} nor R_o/H_s . These results allow the conclusion that the
 808 identification step can be successfully adopted for the identification of V and the determination
 809 of the shape factor b .

810



811
 812 Figure 17 – Values of b derived with the new procedure for UB-num and UB-exp in comparison
 813 with the relationships for smooth structures by Zanuttigh et al., 2013 (panel a) and by Hughes et
 814 al., 2012 (panel b).

815 **5. Application of the new procedure to permeable structures**

816 This Section suggests 2 potential and original applications of the new procedure to the tests on
817 permeable structures by Kramer et al. (2005) (dataset AAU, Sub-section 2.5, Fig. 3 and Table
818 6). The applications allow a detailed description of i) the evolution of the wave shapes along the
819 structure crest (Sub-section 5.1) and ii) the estimate of the percolation rate (Sub-section 5.2).

820 To this purpose, the procedure was applied to the analysis of the h -signals registered at the 2
821 wgs, wg1 and wg2, placed over the barrier crest at a distance of 0.15 or 0.40 m (according to
822 the width of the crest itself, see Table 6). From the time-ordered sequences of the coupled
823 overtopping waves, we derived the values of c , V and q_{cel} from the integration of c by the
824 corresponding h .

825 **5.1. Evolution of the wave shape over the structure crest**

826 The reconstruction of the shape of the overtopping waves at the crest in-shore edge is of practical
827 use for the parametrization of the wave asymmetries (a.o., Peng et al., 2009) and the set-up of
828 physical and conceptual models of wave overtopping propagation, e.g., the Wave Overtopping
829 Simulator (van der Meer et al., 2006) and of wave transmission (Zanuttigh and Martinelli, 2008).
830 Moreover, the study of the evolution of the wave shape may result in important information for
831 the characterization of the flow thickness and velocity over the crest and the landward slope
832 (EurOtop, 2018).

833 The example of Figure 18 shows a 5-second-time-evolution of the water levels registered at wg1
834 and wg2 during one of the tests from the dataset AAU at $R_c > 0$. In this Figure, 4 overtopping
835 events are visible and their correspondence between the 2 wgs is highlighted by the circles,
836 representing the tdc instants filtered and coupled by the procedure.

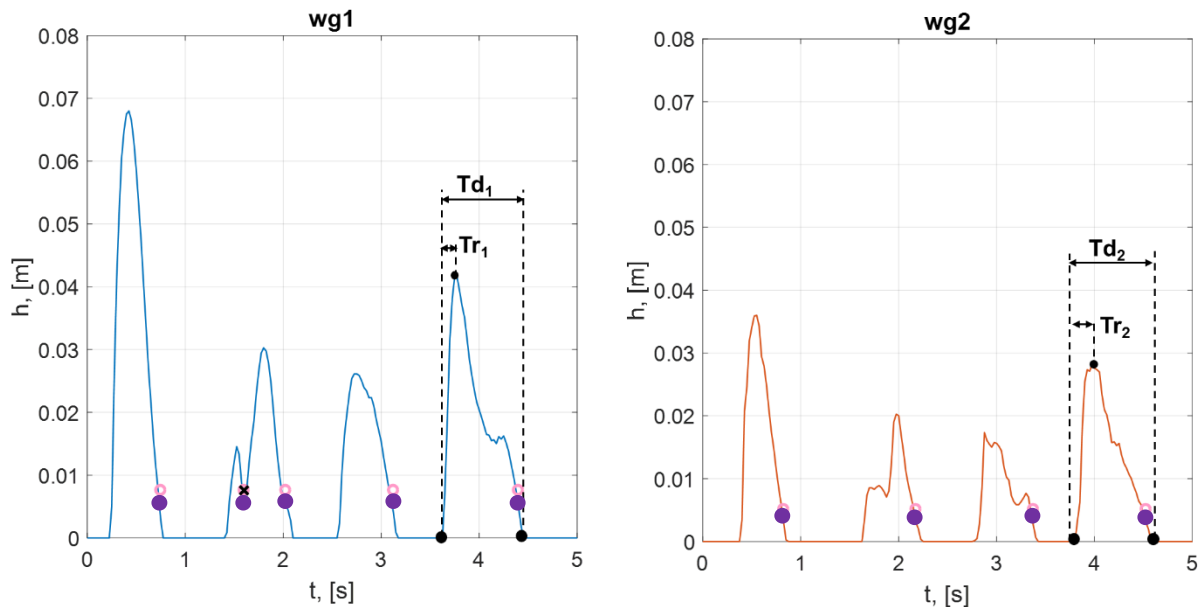
837 The single overtopping events can be schematized as triangular volumes (Zanuttigh and
838 Martinelli, 2008) characterized by a duration time (Td) and a rise time (Tr), which is the time
839 occurring between the beginning of the event and the instant of the passage of the crest (see
840 Fig. 18). The shape of the overtopping events evolves during the passage from wg1 to wg2
841 reducing the wave amplitudes heights ($H = Z_{cr} - Z_{tr}$) and slightly increasing the durations.

842 The time series of the parameters Tr , Td or H can be used to statistically characterize the
843 evolution of the flow over the crest. As an example, Figure 19 illustrates the pdf distribution of
844 the ratio Tr/Td . The values of Tr/Td of each overtopping wave of each test (overall, 21 tests for
845 a total of 6251 overtopping events) are grouped into two histograms in Figure 19, referring to
846 wg1 and wg2. The frequency of the values of $Tr/Td > 0.5$ is negligible and was cut off from the
847 diagrams.

848 Figure 18 indicates that the events are highly asymmetric, being the median, the mean and the
849 mode values of the distributions of Tr/Td between 0.25 and 0.3, i.e. significantly lower than 0.5.
850 The comparison of the 2 diagrams suggests that Tr/Td varies with the evolution of the waves
851 over the crest: the distribution at wg2 tends to be wider (indeed the standard deviation is 0.13

852 for wg1 and 0.17 for wg2) and its mean and median values are slightly higher (respectively
 853 0.30 and 0.29) than at wg 1 (0.27 and 0.26).

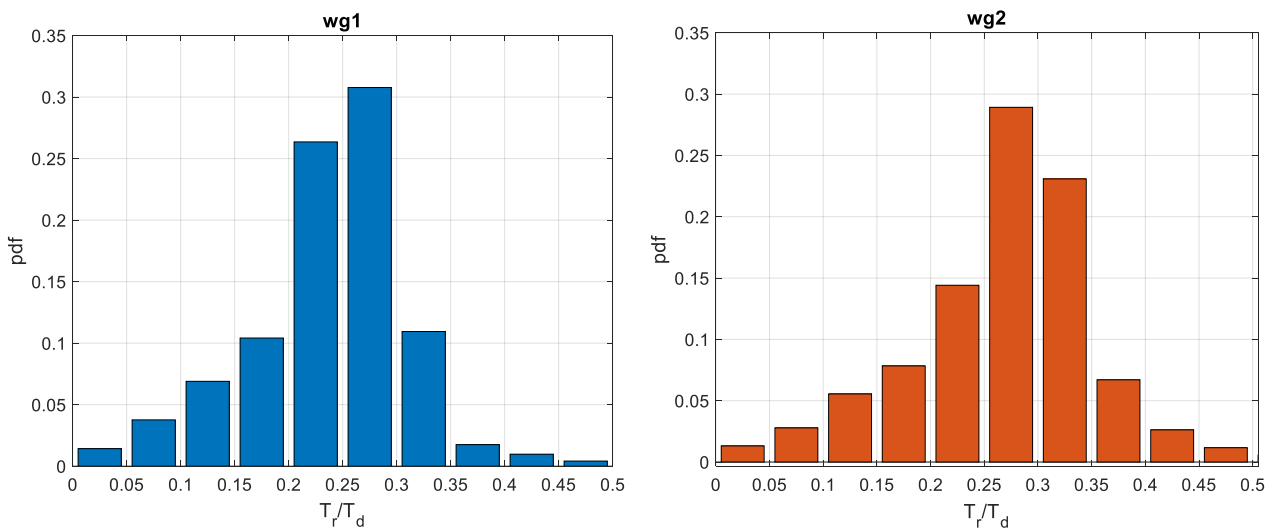
854 These results are similar to the findings of the previous study by Zanuttigh and Martinelli (2008).
 855 The accuracy achieved with the new coupling procedure, and a wider availability of data, may
 856 allow drawing some more general conclusions, leading to a parametrization of the evolution of
 857 the shape of the overtopping volumes.



858

859 Figure 18 – Time evolution of the h -signal measured at wg1 and wg2 (to the left and to the right,
 860 respectively) for a test at $R_c > 0$ belonging to the AAU dataset. The durations of the single
 861 overtopping events are marked by the circles. The duration time and the rise time (Td and Tr ,
 862 respectively) of the 4th event are highlighted as example.

863



864

865 Figure 19 – Histograms of the distributions of the values of the ratio among the rise and the
 866 duration times (Tr/Td) computed at wg1 (a) and at wg2 (b) for each wave identified by the new
 867 procedure for the dataset AAU.

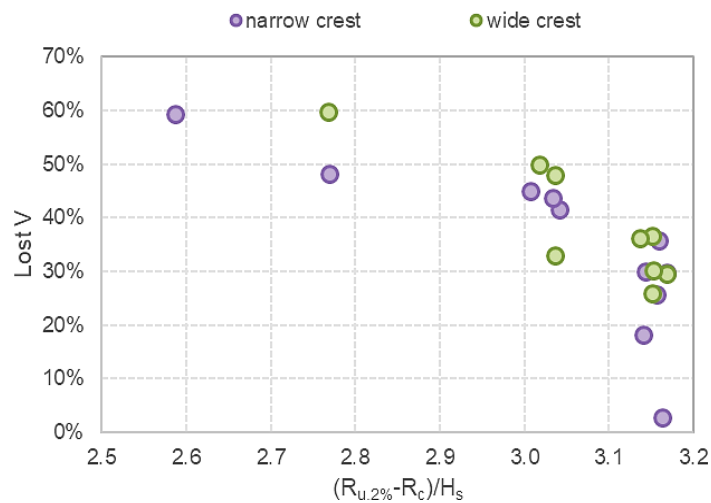
868 **5.2. Estimate of the percolation over the structure crest**

869 The coupled wave signals resulting from the procedure can be numerically integrated by means
 870 of a simple algorithm to derive the individual or total V . In case of permeable structures, part of
 871 the water volume is lost for percolation into the crest during the passage of the waves from wg1
 872 to wg2. The percolation rate is an indicator of the dissipation along the crest and could be used
 873 for the design of the crest width G_c . Furthermore, it may represent a key element for the
 874 assessment of structural failure scenarios induced by the water infiltration.

875 Figure 20 reports the percentages of the volumes V lost for percolation over the crest of a
 876 permeable structure for the tests composing the dataset AAU. In the Figure, the percentages are
 877 grouped by values of G_c and are plotted as functions of the dimensionless quantity $(R_{u,2\%}-R_c)/H_s$.
 878 The diagram suggests that, on average, the percentages of the lost V tend to decrease with
 879 increasing $(R_{u,2\%}-R_c)/H_s$ and with decreasing G_c . Overall, $R_{u,2\%}$ appears to be dominant with
 880 respect to G_c . The increase of lost V associated to the wide crest configuration is indeed modest
 881 and not systematic. This limited effect of G_c on the percolation rate is explained with the
 882 overtopping dynamics at $R_c \geq 0$ and the observed trend of lost V with $R_{u,2\%}$. The higher $R_{u,2\%}$, the
 883 farther the wave impinges on the structure crest from the off-shore edge, resulting in a shortening
 884 of the percolation area.

885 Figure 20 shows that percolation rate is never higher than the 60% (for both the crest
 886 configurations) and does not drop below the 25% in the case of the wide-crest configuration. The
 887 single test providing a lost V lower than the 5% is associated to the smallest and shortest wave
 888 ($H_s=0.027$ m and $T_p=0.74$ s, check Table 6) which determines in turn the lowest $R_{u,2\%}$ (0.088 m)
 889 and one of the lowest values of Pow (7.5%). The narrow-crest configuration produces, on
 890 average, smaller losses of water with respect to the wide-crest configuration for the same value
 891 of $(R_{u,2\%}-R_c)/H_s$. These (few) data indicate that, up to $(R_{u,2\%}-R_c)/H_s \approx 3$, the percolation rate might be
 892 increased roughly 10% by increasing the crest width on the order of 50-60%.

893



894
 895 Figure 20 – Percentages of the volumes lost for percolation between wg1 and wg2 (% $Lost V$)
 896 as function of the dimensionless wave run-up $(R_{u,2\%}-R_c)/H_s$. Data grouped by values of $G_c=0.2$
 897 or 0.6 m, i.e. narrow and wide configuration, respectively. AAU dataset.

898 6. Conclusions

899 This paper presented a new procedure for the identification and the coupling of the individual
900 wave overtopping events. This procedure has been tested on 5 datasets of new and existing
901 data from the literature, for a total of 192 tests on wave overtopping at smooth dikes (Hughes
902 and Shaw, 2011; Hughes and Thornton, 2016; Formentin and Zanuttigh, 2018 a; Zanuttigh and
903 Formentin, 2018) and permeable structures (Kramer et al., 2005), under perpendicular waves in
904 2D and 3D conditions, with freeboards R_o/H_s in the range [-1,5; 1.5].

905 The procedure consists of 2 algorithms, organized into 2 sequential steps, and respectively
906 dedicated to the wave identification and the wave coupling.

907 The first algorithm takes as input the time signals of the water level recorded at a gauge and
908 identifies the single overtopping waves based on a threshold-down-crossing analysis of the
909 signal. Through the definition of a second threshold value, this algorithm offers the possibility of
910 filtering the wave signal, distinguishing between real overtopping waves and signal noise.
911 Differently from the existing identification algorithms available from the literature, the peculiar
912 introduction of the second threshold allows the user to set his/her own customized level of
913 accuracy in the detection of the waves. The identification step was validated against the data by
914 Hughes and Thornton (2016) resulting in the same accuracy achieved by the authors through
915 the visual examination of the discharge time series. It was also proved that the identification
916 algorithm can be successfully applied to calculate the probability of overtopping of a structure
917 and the reconstruction of the distribution of the extreme overtopping volumes. The results of both
918 the applications are in good agreement with the predictions from the literature formulae, being
919 the relative differences in the range $\pm 15\%$.

920 The main novelty of the procedure is represented by the second step algorithm dedicated to the
921 wave coupling. This algorithm elaborates the time-ordered-sequences of the overtopping waves
922 identified by the first step at 2 gauges placed consecutively in the direction of the overtopping
923 flow, and matches each overtopping event propagating between them, driven by the calculation
924 of the minimum and maximum time lags necessary for the propagation. The outputs of the
925 algorithm are the time-ordered-sequences of the coupled overtopping events and the celerity (c)
926 of propagation of each coupled wave. The distance between the gauges and the sampling
927 frequency of the wave signals do affect the accuracy of the wave coupling. In case of flow over
928 the dike crest in breaking or broken wave conditions, the c -values obtained with the coupling
929 step resulted very similar to the lab measurements of the flow velocities (u) at the off-shore edge
930 of the levee or dike crest (on average $c/u \approx 0.95$ with a standard deviation $\sigma_{\%} = 7.4\%$), in agreement
931 with Schüttrumpf and Oumeraci (2005) and Lykke Andersen et al. (2011). For flow in non-
932 breaking wave conditions, the values of the time lags necessary to the waves to propagate from
933 one wave gauge to the consecutive one computed with the coupling step were consistent with
934 the time lags obtained by cross-correlating the surface elevation signals ($R^2 = 0.88-0.96$, $\sigma_{\%} = 5-$
935 11%).

936 The coupling algorithm allows for a number of original applications such as:

- 937 • The estimation of the overtopping discharges from the integration of the wave celerities. The
 938 discharges obtained from the application of the procedure to experimental and numerical
 939 data are in good agreement with both the corresponding traditional measurements ($R^2=0.92-$
 940 0.98 , $\sigma_{\%}=9.9-34\%$) and the literature formulae ($R^2=0.92-0.97$, $\sigma_{\%}=8.5-35\%$).
- 941 • The evolution of the overtopping characteristics along the structure crest. In this contribution
 942 the coupling procedure was specifically applied to investigate the evolution of the wave
 943 shape along the crest of rubble mound breakwaters, resulting into a slight decrease of both
 944 the wave steepness and the wave asymmetry.
- 945 • The estimation of the water percolation rate over the crest of permeable structures. Based
 946 on the results of the procedure, it was found that the water percolation rate is mainly
 947 determined by the wave run-up and varies between the 25% and the 60%.

948 The new procedure is suitable to process 2D wave signals propagating perpendicularly towards
 949 the structures. To deal with oblique waves, it would be necessary to calculate the time lags of
 950 wave propagation from more than 2 wgs, and update the coupling algorithm accordingly (Lykke
 951 Andersen et al., 2011).

952 The limits to the accuracy of the procedure are essentially determined by the definition of the
 953 threshold values of the identification step and by the combination of the sample frequency of the
 954 input signals and the distance between the gauges for the coupling step. For a proper wave
 955 coupling, it is thus recommended to keep the distance between the 2 consecutive wave gauges
 956 ($diswg$) at least equal to 2 or 3 times the time resolution of the instruments themselves ($1/sf$, with
 957 sf =sample frequency), i.e. $diswg \geq 2/sf-3/sf$.

958

959 List of notations

a	Weibull's scale factor
b	Weibull's shape factor
c	Wave (or flow) celerity
$c_{2\%}$	Value of c exceeded by the 2% of the incident waves
c_{dw}	Wave celerity in deep water
c_{sw}	Wave celerity in shallow water
c_{mean}	Mean of the distribution of the c -values obtained with the coupling step
$c(x_{corr})$	Average value of the wave celerity estimated with the cross-correlation of the h -signals at wg1 and wg2
$diswg$	Distance between two wgs
dt_{max}	Maximum time lag that may occur between the instant of the tdc of a wave at wg1 and at wg2
dt_{min}	Minimum time lag that may occur between the instant of the tdc of a wave at wg1 and at wg2

Dc	Instants of zero-down-crossing identified by the new procedure presented in Sub-section 2.2
Dcs	Instants of zero-down-crossing identified and filtered by the new procedure presented in Sub-section 2.2
g	Acceleration due to gravity
G_c	Structure crest width
h	Water depth or thickness or level
H	Wave height (=Zcr-Ztr, as identified by the tdc procedure)
H_{m0}	Significant wave height at the toe of the structure
H_s	Simplified notation of H_{si} (significant incident wave height)
HS	Acronym for “Hughes and Shaw, 2011”
HT	Acronym for “Hughes and Thornton, 2016”
lcr	Instants occurrence of the wave crests identified by the new procedure
ltr	Instants occurrence of the wave troughs identified by the new procedure
$lag(x_{corr})$	Time lag estimated with the cross-correlation of the h -signals at wg1 and wg2
l_{th}	Lower threshold for the identification of the overtopping waves
nw	Acronym of “number of waves”
L_p	Peak wave length from spectral analysis
q	Instantaneous specific wave overtopping discharge
q_{cel}	Average wave overtopping discharge derived from the wave celerities
q_{mean}	Average specific wave overtopping discharge (generic)
q_{meas}	Average specific wave overtopping discharge derived from the experiments or the numerical tests
P	Probability of occurrence of the overtopping volumes
Pow	Probability of overtopping (number of overtopping waves/nw)
R^2	Coefficient of determination
R_c	Structure freeboard (negative if the structure is submerged) with the respect to the still water level
$R_{u,2\%}$	Wave run-up
$s_{m-1,0}$	Wave steepness calculated based on the spectral wave period
sf	Sample frequency
$T_{m-1,0}$	Spectral wave period
T_p	Peak wave period
Td	Duration time of a single wave event
tdc	Acronym of “threshold-down-crossing”
Tr	Rise time of a single wave event
u	Flow velocity in the cross-shore direction
$u_{2\%}$	Value of the u exceeded by the 2% of the incoming waves
u_{max}	Maximum value of u
u_{th}	Upper threshold for the identification of the overtopping waves

V	Overtopping volumes
wg(s)	Acronym of “wave gauge(s)”
wg1	Wave gauge 1 (wave gauge at the off-shore edge of the structure crest)
wg2	Wave gauge 2 (wave gauge at the in-shore edge of the structure crest)
Z_{cr}	Wave crest heights identified by the new procedure
Z_{tr}	Wave trough heights identified by the new procedure
α_{in}	In-shore slope of a structure
α_{off}	Off-shore slope of a structure
$\xi_{m-1,0}$	Iribarren-Battjes breaker parameter
$\sigma_{\%}$	Standard deviation

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968

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