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Simulation of observed temperature field below a building in Bologna, Italy

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(Article begins on next page)

1 Simulation of observed

² temperature field below a building

3	Author 1
4	Eng. PhD. Sara Focaccia
5	SERENGEO srl, Bologna, Italy
6	• 0000-0002-1293-4914
7	Author 2
8	Eng. PhD. Alberto Barbaresi
9	Department of Agricultural Sciences, University of Bologna, Bologna, Italy
10	• 0000-0003-1341-984X
11	Author 3
12	Eng. PhD. Francesco Tinti*
13	• Department of Civil, Chemical, Environmental and Materials Engineering, University of
14	Bologna, Bologna, Italy
15	• 0000-0002-6750-9368
16	Paper submitted: 14/12/2017.
17 18	Paper revised: 15/05/2018.
19	*Corresponding Author: Francesco Tinti.
20	Department of Civil, Chemical, Environmental and Materials Engineering, University of Bologna,
21	via Terracini 28, 40100 Bologna, Italy.
22	E-mail: francesco.tinti@unibo.it
23	Phone: +390512090477
24	
25	

26 Abstract

Urban settlements, whether single buildings or apartment blocks, influence near-surface ground temperatures. Heat transfer by buildings to the ground must therefore be considered when designing both vertical probes and energy geostructures in urban areas. However, assessment of ground temperature variability in urban areas is still uncommon for shallow geothermal energy purposes, the standard temperature gradient based on climatic conditions usually being employed during the design phase. Yet precise assessment of the heat transfer between buildings, infrastructures and the underground could improve the planning of geothermal systems.

This work presents a numerical simulation of a finite-element model of heat transfer to the underground due to both a single building and climate conditions with the aim of reproducing the temperature waves at each depth. An isolated building was chosen since it allowed exact quantification of its influence on the ground temperature without external interferences. For this purpose, different boundary and initial conditions were applied to the ground thermal model and results were compared with historical data recorded over several years. The idea proposed in the paper for a single building can be considered as the baseline for further ground temperature assessment of wider urban settlements.

43 Keywords

44 Ground temperature; Shallow geothermal energy; Finite-element modelling.

45 List of notation

- T_g is the varying space-time underground temperature,
- T_{me} is the average annual surface temperature,
- *T*_{building} is the temperature inside the building,
- *T_{cellar}* is the temperature inside the cellar,
- A_e is the annual surface wave amplitude,
- t_{T0} is the time at minimum temperature,
- α is the equivalent thermal diffusivity,
- λ is the equivalent thermal conductivity,
- ρ is the density,

C is the heat capacity,

$\rho \cdot C$ is the volumetric heat capacity.

1. Introduction

The subsurface temperature gradient, or geothermal gradient, depends mainly on endogenous geothermal heat flow (Cermak and Rybach, 1979). However, in shallow layers, its contribution is limited, and subsoil temperature follows climate seasonality, which is damped on account of ground thermal insulation that varies with geological and hydrogeological conditions (Kusuda and Achenbach, 1965). As a result, ground temperatures at very shallow depths change in space - vertically, because of climate wave dampening, and horizontally, due to the variation of geological and hydrogeological conditions - and time, following the ambient temperature wave (Baggs, 1985). It follows that at a certain depth and thickness, depending on local conditions, a ground layer, or so-called neutral zone, will exist where temperature variations in space and time become nil before geothermal heat flow resumes, with the resulting geothermal gradient (Pouloupatis et al., 2011).

Shallow layer ground temperature should also take into account urban ground warming (Ferguson and Woodbury, 2004). The replacement of natural soil and vegetation by artificial surfaces increases temperatures of the surrounding air and subsurface throughout the year on account of indirect solar heating of urban structures, building heat losses and land use change (Bornstein, 1968). At a district or city level, this phenomenon is called Urban Heat Island effect – UHI (Landsberg, 1981), and Subsurface Urban Heat Effect – SUHI, when referred to the thermal effect in the ground (Oke, 1982).

Heat loss from an individual building generates a bulb-shaped volume of subsurface temperatures that are higher compared to a non-urbanized context (Taniguchi et al., 2007) Several experimental studies have demonstrated that heat loss from buildings increases the subsurface temperature by several degrees (generally from 2 to 5 °C) and that this thermal impact is more persistent in the subsurface rather than in the air. Research on the topic has been performed in many part of the world, such as Japan (Huang et al., 2009; Taniguchi et al., 2007), Canada (Ferguson and Woodbury, 2007), Germany (Zhu et al., 2010; Menberg et al., 2013), Ireland (Allen et al., 2003), Turkey (Yalcin and Yetemen, 2009); United Kingdom (Headon et al., 2009); Finland (Arola and Korkka-Niemi, 2014) and Switzerland (Rivera et al.,

86 2017). While calculating and estimating ground temperature distribution is useful in many fields87 of application, it becomes a key design element of:

building basements, to minimize heat losses via the ground (e.g. Claesson and
Hagentoft, 1991; Hagentoft and Claesson, 1991; Rees et al., 2000),

90 - cellars and underground spaces, to ensure appropriate goods storage conditions (e.g.
91 Mazarron et al., 2012; Barbaresi et al., 2014),

92 - ground heat exchangers and energy geostructures, to assess shallow geothermal
93 potential and simulate ground thermal behaviour during heat extraction and/or injection (e.g.
94 Bandos et al., 2009; Kurevija and Vulin, 2010).

95 Dynamic simulators of flow, mass, and heat transfer in the underground are commonly used in 96 the georesources and environmental fields (Pruess and Garcia, 2002; Zhao et al. 2005; Herbert 97 and Chillingworth, 2013; Focaccia et al., 2016). In the geothermal sector, it is common practice 98 to assign a constant ground temperature value to each subsoil layer as an input parameter. In 99 shallow layers, this can be taken as the average ground temperature over the year, which is 910 equivalent to the average ambient temperature (Al-Zyoud et al., 2014; Poulsen et al., 2015). In 911 other case studies, the shallow layer temperature was experimentally measured in the field over 92 a specific time period (Wu et al., 2010; Focaccia, 2013; Barla et al., 2018). In both cases, since 93 shallow ground temperature is time dependent, the values assigned did not effectively represent 93 the ground temperatures over the whole year. Numerical simulations of underground heat flow 94 and mass transfer are not commonly designed to include the thermal effect of aboveground 95 and mass transfer are not commonly designed to include the thermal effect of aboveground 96 temperature variations in these shallow layers could improve the efficiency of exploitable 97 qeothermal energy through the use of energy piles.

The aim of this paper is to provide specific settings for dynamic simulations to include both timerelated temperature variations and the thermal effects of buildings when designing energy geostructures where shallow layer temperatures may play a significant role. The simulation results were validated with a case study of a single building in a rural area that affected the ground temperature around it. The dynamic simulator settings were calibrated accordingly. Located in the countryside around Bologna, Italy, the detached house had already been used in several experimental and analytical investigations and been the object of a multiyear

temperature and humidity measurement campaign. Measurements had been taken inside the building (Barbaresi et al., 2014), in the surrounding undisturbed ground at different depths (Tinti et al., 2014), inside the building cellar (Barbaresi et al., 2015) and beside and below the building at different depths (Tinti et al., 2015). Calculations and modeling were performed to improve cellar air temperature management (Benni et al., 2016) and estimate ground temperature evolution as a result of the building's presence (Tinti et al., 2017a).

122 2. Materials and methods

123 2.1. Choice of the numerical simulator

The numerical simulator *FEFLOW*[®] (Finite Element Flow simulator), commonly used in the shallow geothermal sector (Al-Khoury et al., 2005; Al-Khoury et al., 2006), was chosen for setting optimization. *FEFLOW*[®] is an integrated package including dynamic flow, heat and mass transportation simulation tools. It allows users to create as many layers as necessary and upload node by node database information for each layer (hydraulic properties, thermal properties, initial conditions of temperature and water level).

130 2.2. Model parameters: geology, hydrogeology and building dimensions.

The case study involved a two-storey building with an underground wine-ageing room in a rural
area in the countryside of Bologna (Italy). The bearing is 32° north-east. The building's key
features are reported in *Table 1*.

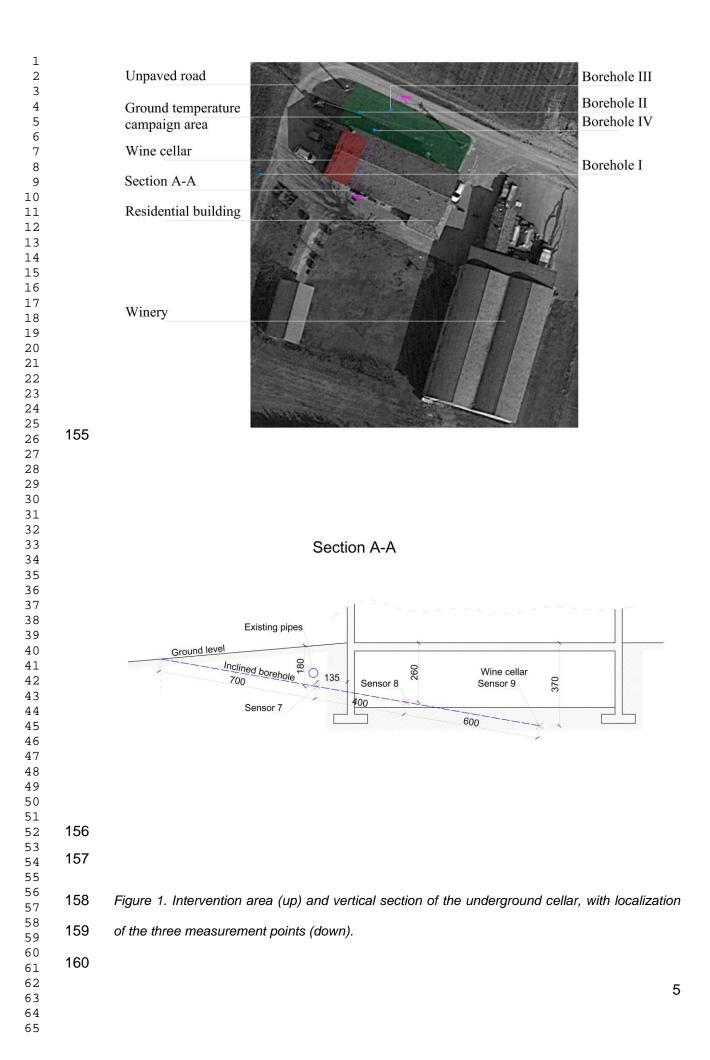
134 Table 1: Key building features.

	Two-storey building	Wine-ageing room
Width	9.8 m	5.6 m
Length	20.5 m (above ground) + 5.6 m (above cellar)	9.8 m
Height	7.3 m (borders), 8.4 m (centre)	2.6 m (underground)
Walls	25.0 cm thick masonry	25.0 cm thick masonry
Floor	30.0 cm hollow concrete slab	20.0 cm concrete slab
Ceiling		30.0 cm hollow concrete slab
Air conditioning	Heating and natural ventilation	Natural ventilation

A model PCE-FWS20 weather station located 100 m from the building collected the main weather data. The two-storey building is heated during the winter season, and there is no mechanical cooling during summer. A set of stand-alone temperature and humidity data loggers, model PCE-HT71, with a 0.1°C resolution, an accuracy of ±0.5°C and a 1 minute registration interval were chosen to collect data. Three were used to measure the internal

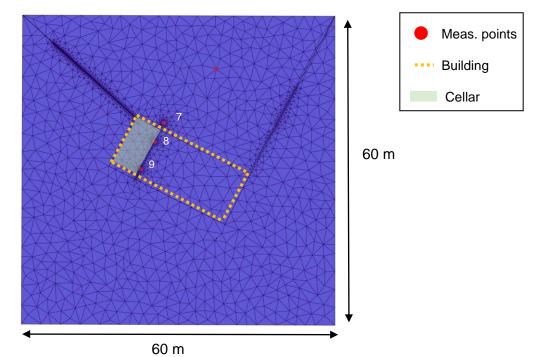
ambient data (temperature and humidity) of the rooms of the two-storey building (Barbaresi et al., 2014). Seventeen data loggers were subsequently placed inside the exclusively naturally
ventilated wine cellar (Barbaresi et al., 2015).

Subsoil temperature distribution in the intervention area had already been thoroughly analysed. Four monitoring boreholes, each with three data loggers, with integrated temperature sensors, placed at different depths, were drilled at designated positions with respect to the building. Boreholes I and II are vertical and are both equipped with three data loggers at depths of 2, 4 and 6 m (Numbered 1 to 6, Tinti et al., 2014). Borehole III, with a total length of 17 m, is inclined (10° from horizontal plane) and has 3 data loggers at depths of 1.8, 2.6 and 3.7 m, 2 of them (the deepest ones) being located below the building and beside the cellar (Numbered 7, 8 and 9, Tinti et al., 2015). Borehole IV is vertical and shallower than the others. It has 3 data loggers at depths of 0.10, 0.65 and 1.20 m (Numbered 10, 11 and 12, Tinti et al., 2015). The whole intervention area and the location of the 3 measurement points on Borehole III with respect to the underground cellar, are shown in Figure 1.



161 The shallow underground of the whole area is prevalently composed of low moist clay down to a 162 depth of 6 m, even lower in some places. Data sources are the geological investigations related 163 to the installation of data loggers. All geological data were collected by the borehole data 164 loggers.

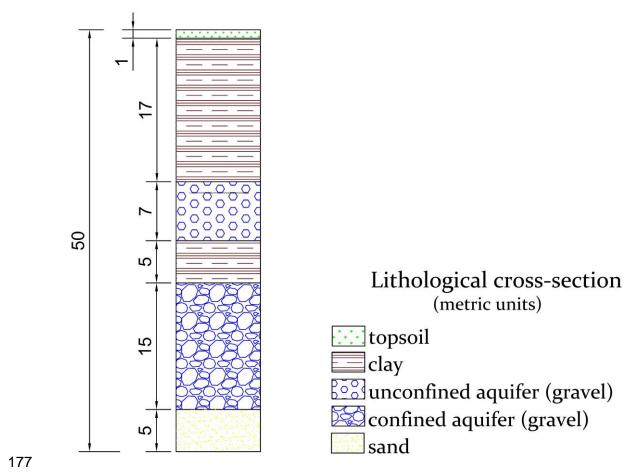
The model domain implemented in *FEFLOW*[®] extends beyond the building itself for a total surface area of 60×60 m, a depth of 50 meters and 0.25 meters of boundary surface area, the latter chosen to guarantee a swathe of undisturbed ground around the boundary of the house. A tetrahedral mesh was used with refinement regions close to three observation points, corresponding to measurement points 7, 8 and 9. The finite elements generated range from $5\cdot10^{-2}$ m² in the zones close to the sensors to 5 m² from the building (*Figure 2*).



, **171**

171 Figure 2. Mesh of the area chosen for the dynamic simulation.

A series of 18 layers was used to detail the geology of the area as well as a first buffer layer. Geological detail up to 50 m was obtained from the Geological Survey of the Emilia Romagna Region (*Figure 3*). The study does not consider groundwater flow in the two aquifers located respectively at 18 m and 30 m below the ground surface since their thermal impact is nil with respect to the intervention area (between 0.5 and 10 m below the ground surface).



178 Figure 3: Lithological cross-section for the case study (from Tinti et al., 2017a).

179 Measurement points 7, 8 and 9, expressed in relation to the cellar and building walls, are given

180 in *Table 2.* All of them are located in low moist clay.

181 Table 2. Localization of measurement points 7, 8 and 9.

Sensor number	Depth (m)	Qualitative localization	Orthogonal distance from the building northern wall (m)	Geological layer number (refer to Table 4)
7	1.8	Beside the building and cellar walls	+1.35	6
8	2.6	Under the building, beside the cellar wall	-2.57	7
9	3.7	Under the building, under the cellar	-8.47	9

182 2.3. Model boundaries: ground and building temperature

The basic equation for assessing vertical temperature (T_g) distribution at low depth is a function of the ambient temperature wave and the thermal properties of the ground (Kusuda and Achenbach, 1965). Since all the function variables are regionalized, the target variable is four

dimensional, varying in space (x,y,z) and time (t). Since T_g has a sinusoidal behaviour, the year is usually given as a wave period.

Equation 1 summarizes the well-known distribution of temperatures in the subsoil at low depths189 (Baggs, 1985).

$$T_g(x,y,z,t) = T_{me}(x,y) - A_e(x,y) \cdot \exp\left[-z \cdot \sqrt{\left(\frac{\pi}{365 \cdot \alpha(x,y,z)}\right)}\right] \cdot \cos\left[\frac{2 \cdot \pi}{365 \cdot \left(t - t_{T_0}(x,y) - \frac{z}{2} \cdot \sqrt{\frac{365}{\pi \cdot \alpha(x,y,z)}}\right)\right]$$

193 Where:

1.

 T_g is the space-time varying underground temperature (°C),

T_{me} is the annual surface average temperature (°C),

 A_e is the annual surface wave amplitude (°C),

 t_{T0} is the time at minimum temperature (days),

 α is the equivalent thermal diffusivity (m²/days).

Although other parameters such as the vegetation coefficient, the insulation coefficients, and the geothermal heat flow, can be included in the analytical equation, we decided to omit these in the present work since our aim was to define settings to recreate the bulb-shaped volume of subsurface temperatures below a building in the numerical simulator without the inclusion of corrective coefficients. In addition, since the area of investigated was limited to the shallow layers, disregarding the contribution of geothermal heat flow was in this case an acceptable simplification.

206 Experimental temperature measurement in Boreholes I and II (Tinti et al., 2014) allowed
 207 estimation of ground thermal wave parameters by non-linear regression based on *Equation 1*.
 208 Parameters obtained by non-linear regression are summarized in *Table 3*.

209 Table 3. Entering parameters to estimate ground temperature in the survey area with the 210 analytical method.

Yearly ambient temperature (°C)	Yearly amplitude (°C)	Day of minimum temperature (days)	Ground thermal diffusivity (m²/days)
15.0	15.5	10 th	0.0288

211 January 10th is statistically the day of minimum temperature in the climatic area of Bologna, so

212 in the climatic wave it represents the position of the inflection point.

attributed to parameters. The temperature of the ground below the study area (i.e. from 18 m) was calculated also considering geological and hydrogeological variations, as reported in Tinti et al., 2017a. Since the effect of deep temperature on the surface layers is very limited, no further investigations were conducted during this study below 18 m. The temperatures indicated by Tinti et al., 2017a were used in the dynamic simulator as boundary and initial conditions. The baseline condition for each layer was that layer's average temperature on 1st January (starting date of the simulation) as shown in Table 4, while the Dirichletian boundary conditions at the borders of each layer were represented as the time-varying temperature evolution over the year. As no groundwater flow was considered, a Dirichletian boundary condition of 0 m hydraulic head was set for all the sides and layers of the model.

Since the ground was considered homogeneous in the survey area, no spatial variations were

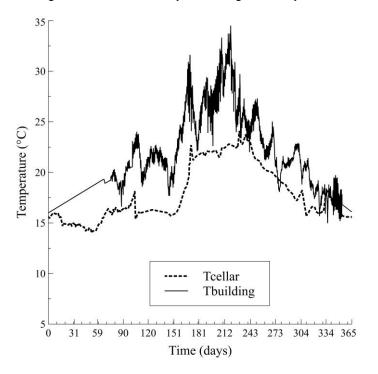
Table 4. Initial temperature conditions for each layer of the model.

Layer n°	Thickness (m)	Depth (m)	Initial temperature (°C)
2	0.25	0.25	4.67
3	0.25	0.50	6.24
4	0.50	1.00	7.80
5	0.50	1.50	10.44
6	0.50	2.00	12.51
7	0.50	2.50	13.47
8	0.50	3.00	14.96
9	1.00	4.00	15.85
10	2.00	6.00	16.45
11	2.00	8.00	16.10
12	2.00	10.00	15.58
13	2.00	12.00	15.52
14	6.00	18.00	15.66
15	7.00	25.00	15.93
16	5.00	30.00	15.37
17	15.00	45.00	15.84
18	5.00	50.00	15.60

The ground profile was defined based on the lithological section (Figure 3). Layers 2 to 4 are all part of the first lithological type (topsoil); layers 5 to 14 are part of the second lithological type (clay). Both were refined to obtain better representation of the superficial layers, which are mostly influenced by the external temperature. In contrast, layers 15 to 18 were defined according to the original thickness of each lithological type.

The basement of the two-storey building was considered as corresponding to the top of Layer 2. Since Layer 1 was considered an ambient temperature buffer, Layer 2 was considered as located at 0 m depth. The time-varying temperature over the year registered inside the building (T_{building}) at this 0 m depth was considered as the Dirichletian boundary. The basement of the

wine ageing room was, on the other hand, placed at the top of Layer 8 (located at a depth of 3 m). The Dirichletian boundary condition at this depth was the time-varying temperature over the year registered in the wine room (T_{cellar}). *Figure 4* shows the reference temperatures inside the building and cellar over one year during the study.



239 Figure 4. Internal temperature of building and cellar

The two temperature evolutions were used as wall temperature conditions in the simulation toexpress the heat effect of the building.

242 2.4. Model parameters: thermal conductivity, heat capacity and effective porosity

Each of the 18 layers was characterized in terms of thermal properties, namely thermal conductivity, thermal capacity, and effective porosity. Respective values are shown in *Table 5*. For the specific case study, the value of thermal diffusivity was calculated through the multiyear measurement campaign up to a depth of 6 m. The value was later additionally verified by a Thermal Response Test (TRT) performed on a geothermal basket installed on site (Ferrari et al., 2016; Tinti et al., 2017b).

Other parameters were obtained on the basis of the relationship between thermal diffusivity andconductivity, as shown in Equation 2.

$$\alpha = \frac{\lambda}{\rho \cdot C}$$

252	2.
253	Where
254	λ is the equivalent thermal conductivity (W/(m·K)),
255	ho is the density (kg/m³),
256	C is the heat capacity (J/(kg·K)),
257	$\rho \cdot C$ is the volumetric heat capacity (J/(m ³ ·K)).
258	Lithology-based bibliographic values were taken for the layers deeper than 18 m below the
259	ground surface (Tinti et al., 2017a). Effective porosity was set at 10% (i.e. low effective porosity,
260	typical of low moist clay) down to the first aquifer (18 m deep), where it was increased up to
261	30%. Effective porosity was used in the simulation to estimate the amount of water in the aquifer
262	and the quantity of air in the dry layers. This, together with the thermal conductivity and
263	volumetric heat capacity provided for each of the lithologies generated the actual thermal
264	properties of each layer.
265	Table 5. Thermal characteristics associated with the layers.

Layer n°	Thickness (m)	Depth (m)	Thermal conductivity (W/(m·K))	Volumetric heat capacity (MJ/(m³·K))	Effective porosity (-)
2	0.25	0.25	0.50	1.0	0.1
3	0.25	0.50	0.65	1.5	0.1
4	0.50	1.00	0.65	1.5	0.1
5	0.50	1.50	0.65	1.5	0.1
6	0.50	2.00	0.65	1.5	0.1
7	0.50	2.50	0.65	1.5	0.1
8	0.50	3.00	0.65	1.5	0.1
9	1.00	4.00	0.65	1.5	0.1
10	2.00	6.00	0.65	1.5	0.1
11	2.00	8.00	0.65	1.5	0.1
12	2.00	10.00	0.65	1.5	0.1
13	2.00	12.00	0.65	1.5	0.1
14	6.00	18.00	0.65	1.5	0.1
15	7.00	25.00	1.80	2.4	0.3
16	5.00	30.00	1.00	1.6	0.3
17	15.00	45.00	1.80	2.4	0.3
18	5.00	50.00	2.40	2.4	0.3

265 Table 5. Thermal characteristics associated with the layers.

A thermal conductivity of 0.35 W / (m·K) was assigned to the building walls and floors. This represents the characteristic thermal conductivity of a hollow concrete slab with a high degree of

insulation.

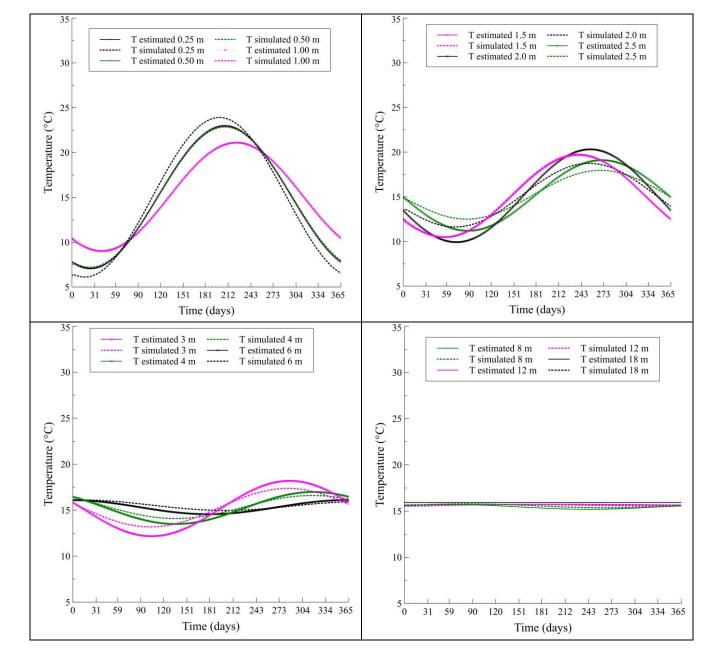
3. Results

The above model, implemented in FEFLOW®, was run for 1825 days under the boundary and initial conditions described. The initial time step was set at 0.5 days and subsequently defined automatically by the convergence of the equations. Local temperatures at the measurements points were retrieved from the simulation each day and compared with the on-site measurements and estimated data based on non-linear regression on Equation 1. The undisturbed ground temperature of each layer was recorded along a vertical profile at observation points positioned one on top of the other at 15 m from the building (where the heat island effect is negligible). Since the new dynamic boundary conditions settle after 2 years of simulation, comparison between simulated and estimated natural ground temperatures can be made taking either the 3rd, 4th or 5th year of simulation (*Equation 1*). In this study we carried out the comparison on the basis of simulated temperatures of the 5th year.

281 The match between simulated and estimated values of space-time varying ground temperature 282 was obtained in all parts of the domain (including below the building and beside the cellar) using 283 the following settings:

- Tetrahedral mesh refined around sensor positions, where observation points were defined;
- Initial temperatures of the different layers set equal to the average 1st January temperature;
- Dirichletian boundary conditions for the temperature of each layer equal to the time-varying
 evolution of each layer temperature over the year;
 - Dirichletian boundary condition for building wall temperature equal to the time-varying
 temperature registered over the year inside the building;
- Dirichletian boundary condition for cellar wall temperature equal to the time-varying
 temperature registered over the year inside the cellar;
- Dirichletian boundary condition of 0 m hydraulic head set throughout the model.
- 293 3.1. Reconstruction of undisturbed underground temperature

The comparison between simulated and estimated values of ground temperature evolution is presented in *Figure 5.* Ground temperatures at the different depths were estimated using nonlinear regression based on *Equation 1,* and validated at the different depths (2, 4 and 6 m) against those measured (Tinti et al., 2014).



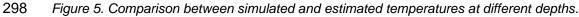


Table **6** reports the average simulated wave error with respect to that estimated at each depth.

301 The values estimated with non-linear regression are considered as reference values. Error

tends to decrease with depth because the temperature becomes more homogeneous.

 304 Table 6. Maximum, minimum and average error of the simulated ground temperature at different

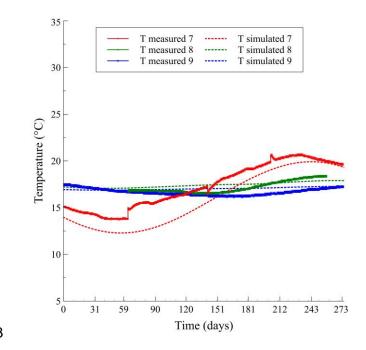
305 depths compared to the estimated values by non-linear regression of Equation 1 on

306 measurement points (Tinti et al., 2014).

Depth	Maximum Relative Error	Minimum Relative Error	Average Relative Erro
(m)	(%)	(%)	(%)
0.00	6.1%	0.0%	2.2%
0.25	5.5%	0.0%	2.1%
0.50	5.1%	0.0%	2.3%
1.00	4.7%	0.0%	2.3%
1.50	17.8%	0.1%	9.3%
2.00	13.4%	0.0%	7.3%
2.50	10.3%	0.0%	5.8%
3.00	6.3%	0.0%	3.6%
4.00	4.4%	0.0%	2.5%
6.00	1.9%	0.0%	0.9%
8.00	0.7%	0.0%	0.3%
10.00	0.4%	0.0%	0.2%
12.00	1.5%	1.4%	1.5%
18.00	2.1%	1.9%	2.0%

3.2. Reconstruction of sensor recorded temperatures.

The temperatures measured at points 7, 8 and 9 were compared with the products of the simulation. Since the three investigation points come within the impact area affected by building heat loss, the values returned at these points do not follow the natural ground thermal state. *Figure 6* shows the comparison.



314 Figure 6. Comparison between simulated and measured temperature values beside and below

the building at measurement points 7 (1.8 m), 8 (2.6 m) and 9 (3.7 m).

316 The comparison led to the average error presented in *Table 7*.

317 Table 7. Maximum, minimum and average error of the simulated sensor-based temperature

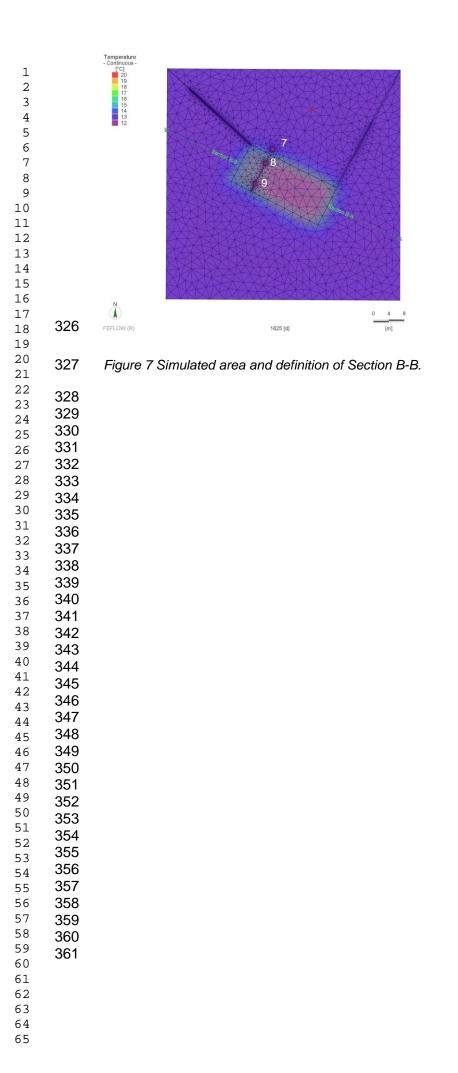
Sensor	Depth	Maximum Relative Error	Minimum Relative Error	Average Relative Error
7	1.8 m	19.4%	0.8%	9.0%
8	2.6 m	6.1%	0.1%	2.4%
9	3.7 m	5.3%	0.1%	2.8%

318 compared to the actual values measured.

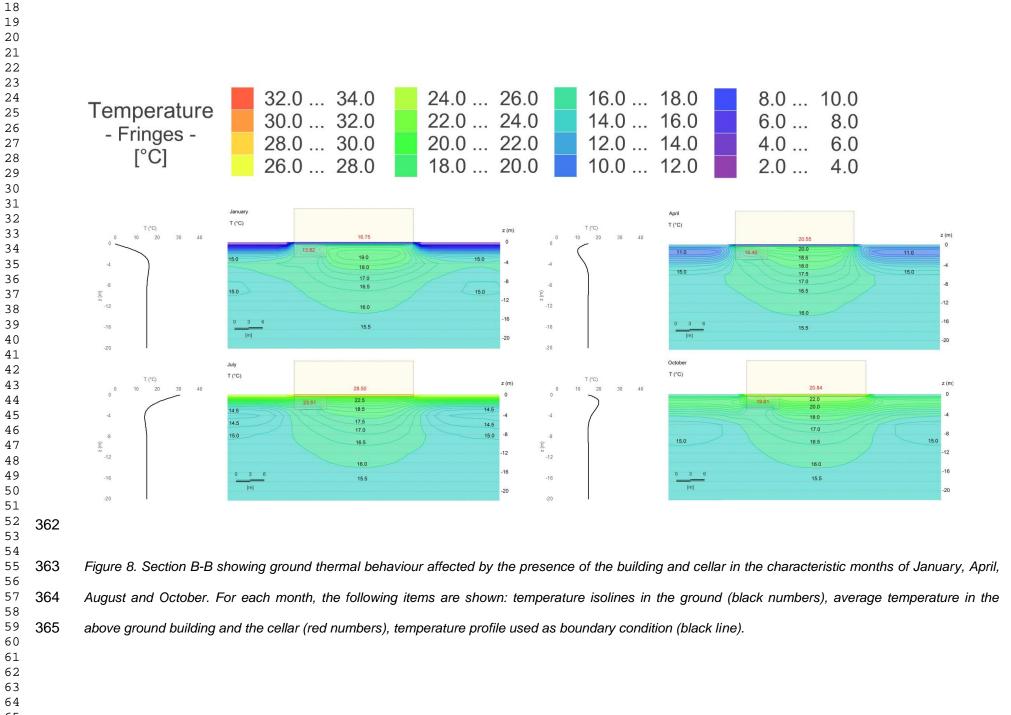
The low average error shows that the settings were correctly chosen to include the building'sthermal contribution in the dynamic simulator.

321 3.3. Simulation of building thermal effect on the ground

Having calibrated the model by comparing simulated and measured data, it was then possible to
simulate the entire ground thermal behaviour as affected by the building over time. The results
for four characteristic periods of January, April, July and October are given in section *B-B* of *Figure 7.* Ground temperature evolution is shown in *Figure 8.*



1	6



Simulation of ground temperature evolution under the building follows the bulb-shaped pattern described consistently in the literature and presented in the introduction to this paper. Temperature variation with depth is notably vertically dampened compared with the undisturbed ground, with no relevant changes occurring over time. This is because the building acts as an insulator against the external weather. Below the building, at depths from 2 to 8 m, temperature remains constant throughout the year at around 17-18°C. In contrast, temperatures around the building down to 4 m depth, in winter, vary from 2 to 15 °C, while, in summer, vary from 28 to 14.5°C. From 4 down to 8 m, there is a transitional zone, with very low temperature variations along the year, while below 8 m temperature remains constant in time.

375 The numerical simulation also evidences the thermal contribution of the indoor air temperature376 of the cellar, which slightly affects the temperatures of its surroundings.

Being validated by the measurements taken, the simulation values may be used in the design ofenergy geostructures and ground heat exchangers when these are located below buildings.

Results can be used both directly in *FEFLOW*[®] as initial conditions for further simulations of
geothermal energy exploitation or extrapolated for the dedicated design of energy geostructures
and ground heat exchanger fields.

Since the heat exchange rates of ground heat exchangers in winter and summer are proportional to initial ground temperature, it follows that exchanger systems for standard heating and cooling loads will be more profitable below a building than underneath an uncovered area. So correct simulation in all grid nodes improves the quality of design. Considering the correct thermal impact of buildings on the ground should lead to more precise forecasts of geothermal energy exploitation rates for heating and cooling needs.

388 4. Conclusions

The study defined optimal settings to reconstruct the time-space varying temperature field deriving from the presence of a building in a dynamic simulator. The settings were calibrated performing numerical simulation and comparing the results with ground temperature values measured on site in a rural area beside and below a detached building. As far as the authors are aware, this is one of the few studies of its kind. It could pave the way for better thermal behaviour simulation, which would lead to improved energy geostructures and ground heat

 exchanger design. In particular, the results of this paper are especially pertinent for ground heat exchanger design in urban areas, where the presence of buildings can considerably affect thermal system performance. This is especially the case of energy piles with absorption pipes installed below buildings, where standard climate-dependent ground thermal conditions do not apply. Moreover, energy pile length is usually limited to a few meters below any building, with the result that thermal losses from buildings will have a large impact on energy performance. The study further underlines that the exploitable energy used by shallow ground heat exchangers located in urban areas is only in part 'natural' (renewable geothermal and solar energy) since heat recovery of energy losses from buildings and districts (subsurface urban heat island effect) make up a sizeable component of the energy available to these systems.

Further studies will be performed to quantify this contribution in the short and long term, both for energy geostructures located below single buildings and for ground heat exchanger fields located in an urban area, with estimation and simulation of the effect of multiple buildings on ground heat exchanger fields.

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 - 522 Figure captions
 - 523 Figure 1. Intervention area (up) and vertical section of the underground cellar, with localization
 - 524 of the three measurement points (down).
 - 525 Figure 2. Mesh of the area chosen for the dynamic simulation.
 - 526 Figure 3: Lithological cross-section for the case study (from Tinti et al., 2017a)
 - 527 Figure 4. Internal temperature of building and cellar
 - 528 Figure 5. Comparison between simulated and estimated temperatures at different depths
 - 529 Figure 6. Comparison between simulated and measured temperature values beside and below
 - 530 the building at measurement points 7 (1.8 m), 8 (2.6 m) and 9 (3.7 m).
- 531 Figure 7 Simulated area and definition of Section B-B.
- 532 Figure 8. Section B-B showing ground thermal behaviour affected by the presence of the
- 533 building and cellar in the characteristic months of January, April, August and October. For each
- 534 month, the following items are shown: temperature isolines in the ground (black numbers),
- 5 535 average temperature in the above ground building and the cellar (red numbers), temperature
- 536 profile used as boundary condition (black line).
- 537 Table captions
- 538 Table 1: Key building features.
- 539 Table 2. Localization of measurement points 7, 8 and 9.
- 540 Table 3. Entering parameters to estimate ground temperature in the survey area with the 541 analytical method.

- 542 Table 4. Initial temperature conditions for each layer of the model.
 - 543 Table 5. Thermal characteristics associated with the layers.

545 Table 6. Maximum, minimum and average error of the simulated ground temperature at different

- 546 depths compared to the estimated values by non-linear regression of Equation 1 on 547 measurement points (Tinti et al., 2014).
- 548 Table 7. Maximum, minimum and average error of the simulated sensor-based temperature
- 549 compared to the actual values measured.

Tables

Table 1: Key building features.

	Two-storey building	Wine-ageing room
Width	9.8 m	5.6 m
Length	20.5 m (above ground) + 5.6 m (above cellar)	9.8 m
Height	7.3 m (borders), 8.4 m (centre)	2.6 m (underground)
Walls	25.0 cm thick masonry	25.0 cm thick masonry
Floor	30.0 cm hollow concrete slab	20.0 cm concrete slab
Ceiling		30.0 cm hollow concrete slab
Air conditioning	Heating and natural ventilation	Natural ventilation

Table 2. Localization of measurement points 7, 8 and 9.

Sensor number	Depth (m)	Qualitative localization	Orthogonal distance from the building northern wall (m)	Geological layer number (refer to Table 4)
7	1.8	Beside the building and cellar walls	+1.35	6
8	2.6	Under the building, beside the cellar wall	-2.57	7
9	3.7	Under the building, under the cellar	-8.47	9

Table 3. Entering parameters to estimate ground temperature in the survey area with the analytical method.

Yearly ambient temperature (°C)	Yearly amplitude (°C)	Day of minimum temperature (days)	Ground thermal diffusivity (m²/days)
15.0	15.5	10 th	0.0288

Table 4. Initial temperature conditions for each layer of the model.

Layer n°	Thickness (m)	Depth (m)	Initial temperature (°C)
2	0.25	0.25	4.67
3	0.25	0.50	6.24
4	0.50	1.00	7.80
5	0.50	1.50	10.44
6	0.50	2.00	12.51
7	0.50	2.50	13.47
8	0.50	3.00	14.96
9	1.00	4.00	15.85
10	2.00	6.00	16.45
11	2.00	8.00	16.10

12	2.00	10.00	15.58
13	2.00	12.00	15.52
14	6.00	18.00	15.66
15	7.00	25.00	15.93
16	5.00	30.00	15.37
17	15.00	45.00	15.84
18	5.00	50.00	15.60

Table 5. Thermal characteristics associated with the layers.

Layer n°	Thickness (m)	Depth (m)	Thermal conductivity (W/(m·K))	Volumetric heat capacity (MJ/(m³·K))	Effective porosity
					(-)
2	0.25	0.25	0.50	1.0	0.1
3	0.25	0.50	0.65	1.5	0.1
4	0.50	1.00	0.65	1.5	0.1
5	0.50	1.50	0.65	1.5	0.1
6	0.50	2.00	0.65	1.5	0.1
7	0.50	2.50	0.65	1.5	0.1
8	0.50	3.00	0.65	1.5	0.1
9	1.00	4.00	0.65	1.5	0.1
10	2.00	6.00	0.65	1.5	0.1
11	2.00	8.00	0.65	1.5	0.1
12	2.00	10.00	0.65	1.5	0.1
13	2.00	12.00	0.65	1.5	0.1
14	6.00	18.00	0.65	1.5	0.1
15	7.00	25.00	1.80	2.4	0.3
16	5.00	30.00	1.00	1.6	0.3
17	15.00	45.00	1.80	2.4	0.3
18	5.00	50.00	2.40	2.4	0.3

Table 6. Maximum, minimum and average error of the simulated ground temperature at different depthscompared to the estimated values by non-linear regression of Equation 1 on measurement points (Tinti et al.,2014).

Depth (m)	Maximum Relative Error (%)	Minimum Relative Error (%)	Average Relative Error (%)
0.00	6.1%	0.0%	2.2%
0.25	5.5%	0.0%	2.1%
0.50	5.1%	0.0%	2.3%
1.00	4.7%	0.0%	2.3%
1.50	17.8%	0.1%	9.3%
2.00	13.4%	0.0%	7.3%
2.50	10.3%	0.0%	5.8%
3.00	6.3%	0.0%	3.6%
4.00	4.4%	0.0%	2.5%
6.00	1.9%	0.0%	0.9%
8.00	0.7%	0.0%	0.3%
10.00	0.4%	0.0%	0.2%
12.00	1.5%	1.4%	1.5%
18.00	2.1%	1.9%	2.0%

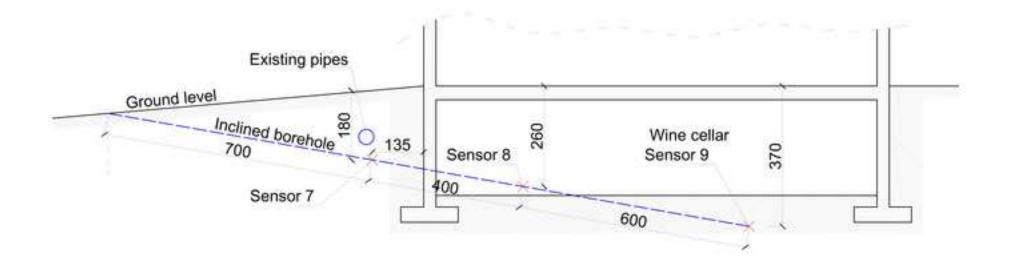
Table 7. Maximum, minimum and average error of the simulated sensor-based temperature compared to the actual values measured.

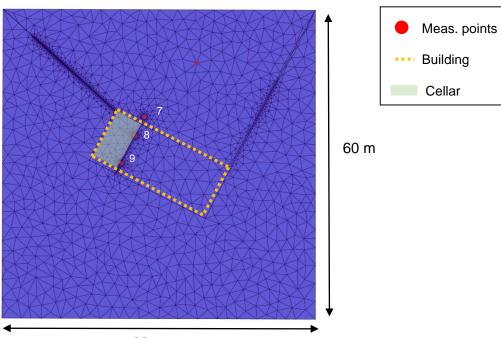
Sensor	Depth	Maximum Relative Error	Minimum Relative Error	Average Relative Error
7	1.8 m	19.4%	0.8%	9.0%
8	2.6 m	6.1%	0.1%	2.4%
9	3.7 m	5.3%	0.1%	2.8%



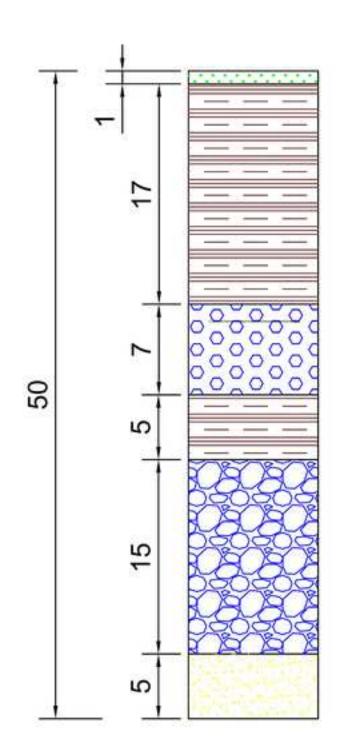
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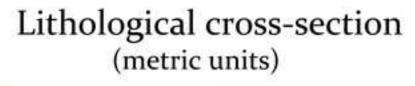
Section A-A

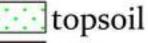




60 m



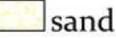


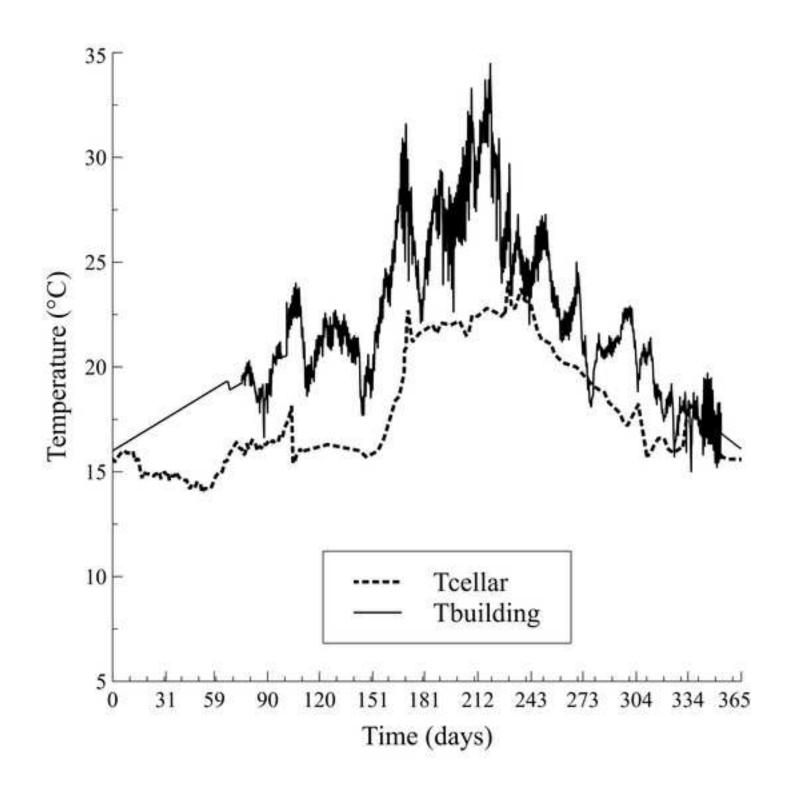


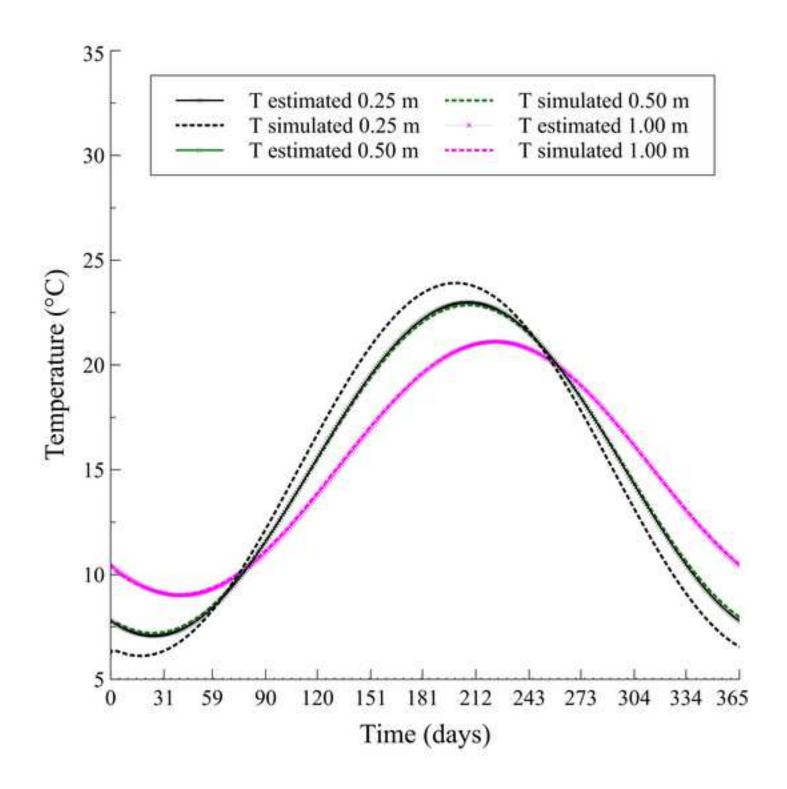


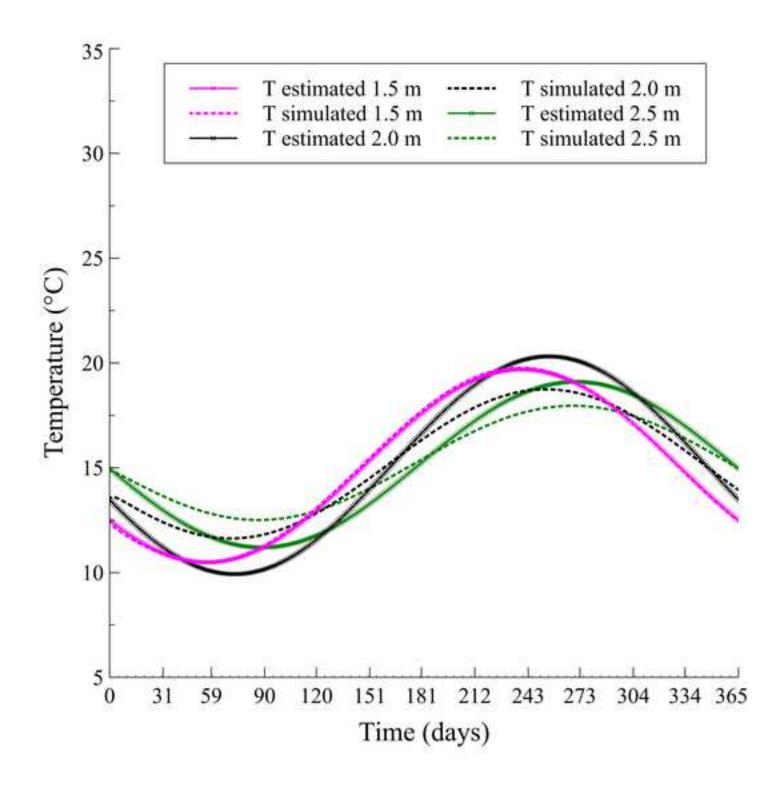
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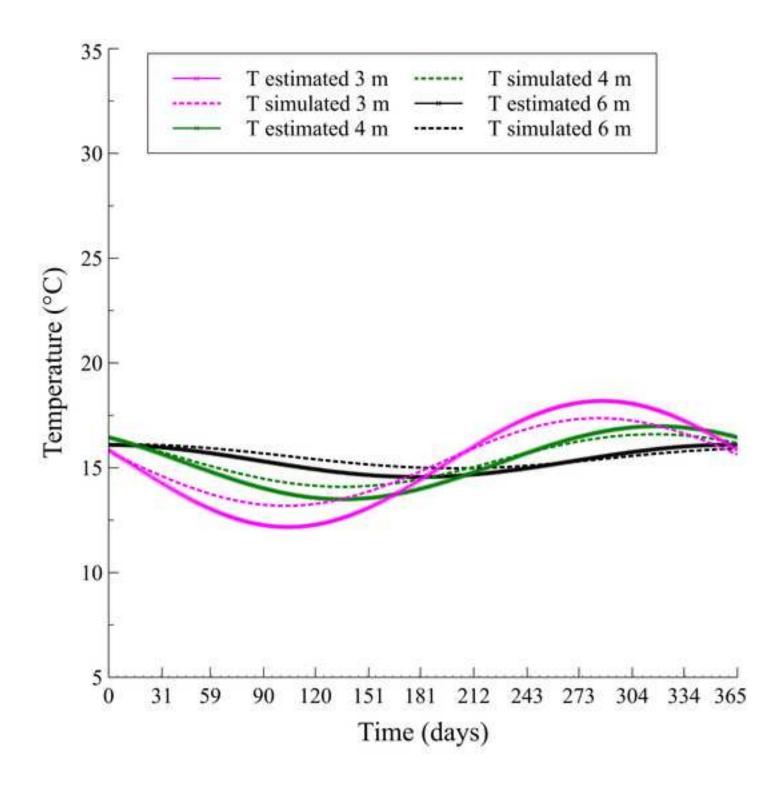
confined aquifer (gravel)

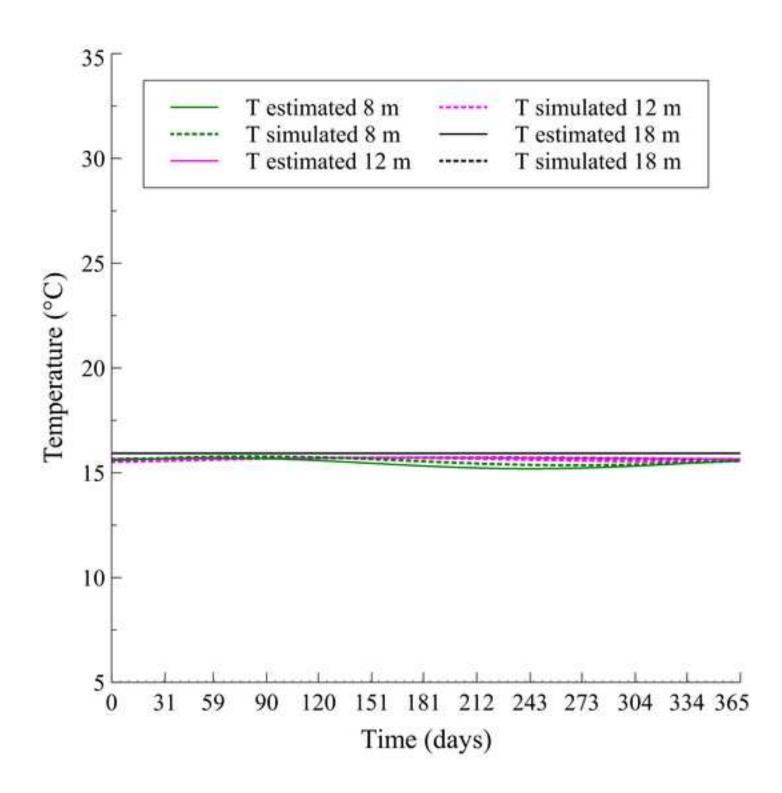


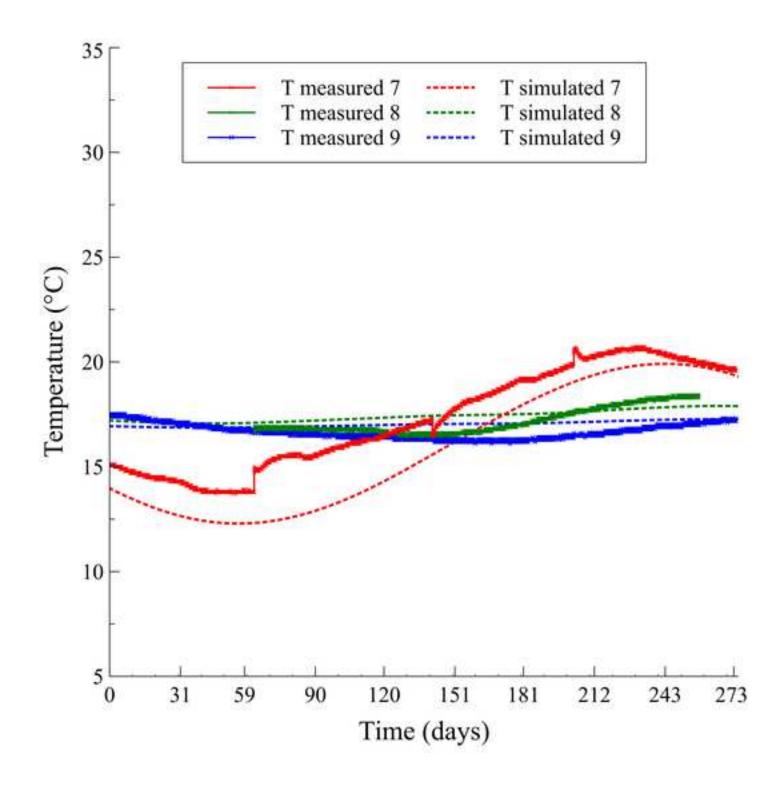


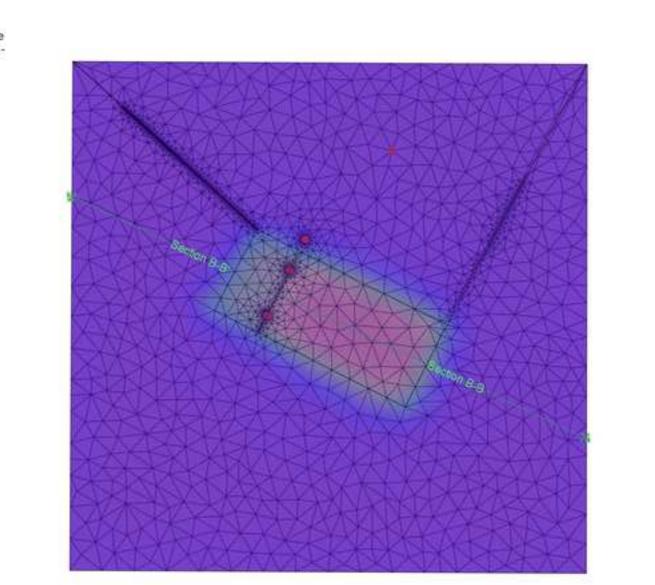












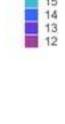
1825 [d]

[m]



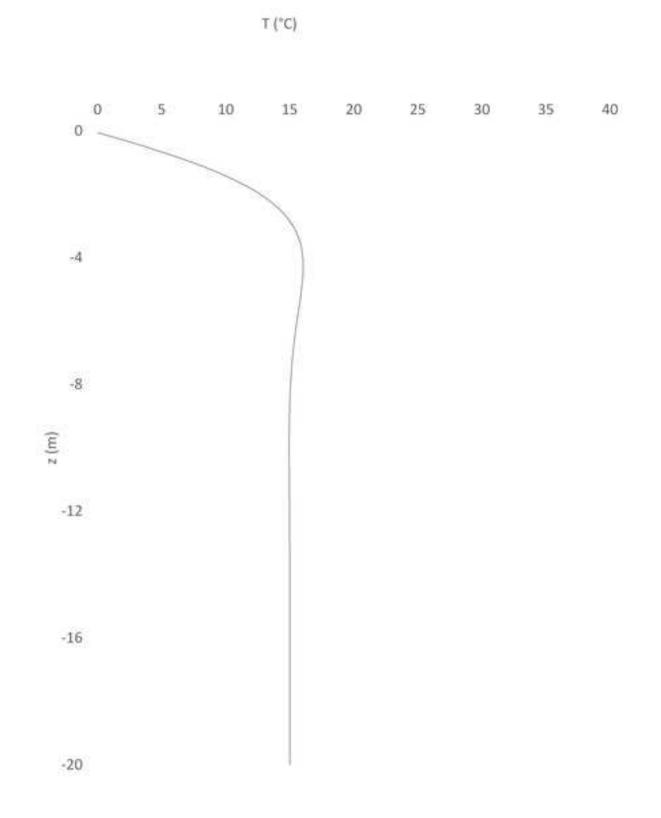


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Temperature - Fringes - [°C]	32.0 34.0	24.0 26.0	16.0 18.0	8.0 10.0
	30.0 32.0	22.0 24.0	14.0 16.0	6.0 8.0
	28.0 30.0	20.0 22.0	12.0 14.0	4.0 6.0
	26.0 28.0	18.0 20.0	10.0 12.0	2.0 4.0



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