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Forming and preserving aragonite in shear zones: First report of blueschist facies metamorphism in the Jabal Akhdar Dome, Oman Mountains

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Forming and preserving aragonite in shear zones: first report of blueschist facies metamorphism in the Jabal Akhdar Dome, Oman Mountains

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Abstract:

We report the first occurrence of high-pressure metamorphic aragonite in Precambrian carbonates of the Jabal Akhdar Dome in the Oman Mountains. We offer a model for its formation at blueschist facies metamorphic conditions and its subsequent preservation to the surface within the tectonic framework of the late Cretaceous obduction of the Semail Ophiolite. Aragonite formed at $T \sim 350^{\circ}\text{C}$ and $P \geq 0.9$ GPa and is preserved within mylonitic shear zones and in stretched fiber dilational veins where the necessary conditions for its formation and preservation, such as plastic strain accommodation, fluid-enhanced mineralogical reactions, and an anisotropic permeability structure, were preferentially met with respect to the surrounding rock. High-strain structural domains are thus documented to potentially be ideal sites where to look for- and study pro- and retrograde high-pressure metamorphic histories in deeply subducted and exhumed terrains.

1 Forming and preserving aragonite in shear zones: first report of
2 blueschist facies metamorphism in the Jabal Akhdar Dome,
3 Oman Mountains

4
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27 **ABSTRACT**

28 We report the first occurrence of high-pressure metamorphic aragonite in Precambrian
29 carbonates of the Jabal Akhdar Dome in the Oman Mountains. We propose a model for both its
30 formation at blueschist facies conditions and its subsequent preservation to the surface within
31 the tectonic framework of the late Cretaceous obduction of the Semail Ophiolite. Aragonite
32 formed at $T \sim 350^\circ\text{C}$ and $P \geq 0.9\text{GPa}$ and is preserved within mylonitic shear zones and in
33 stretched fiber dilational veins where the necessary conditions for its formation and
34 preservation, such as plastic strain accommodation, fluid-enhanced mineralogical reactions, and
35 an anisotropic permeability structure were preferentially met with respect to the surrounding
36 rock. High-strain structural domains are ideal sites where to look for and study pro- and
37 retrograde high-pressure metamorphic histories in deeply subducted and exhumed terrains.

38

39

40 **1. INTRODUCTION**

41 The identification and characterization of relic high-pressure (HP) assemblages in exhumed
42 rocks are key to constraining the thermobaric conditions and the deformation processes typical
43 of otherwise inaccessible subduction zones (e.g., [Agard et al., 2009](#); [Rubatto et al., 2011](#)). Even
44 though exhumed terrains may be affected by retrograde metamorphism, relic assemblages may
45 locally survive within favourable structural sites that act as sheltering capsules to specific HP
46 phases. Formation and later preservation of such phases therein is favoured or inhibited by the
47 interplay of numerous factors, such as the exhumation rate and geothermal gradient, which may
48 inhibit retrograde mineralogical re-equilibrations, presence and composition of fluids, which
49 can favour element transport and enhance mineralogical reactions during pro- and retrograde
50 histories, amount of strain and type of involved mineralogical phases (e.g., [Goffé and Velde,](#)
51 [1984](#); [Bucher and Frey, 2002](#)). The documentation of index minerals, such as lawsonite and
52 carpholite, has facilitated the study of blueschist facies conditions within metabasite and

53 metapelite (e.g., [Evans and Brown, 1986](#)). Aragonite, on the other hand, is a metastable
54 polymorph of calcite that represents a reliable blueschist facies indicator for “pure” carbonates
55 that otherwise lack other P-sensitive mineral phases (e.g., [Hacker et al., 1992](#); [Stöckhert et al.,](#)
56 [1999](#); [Giuntoli et al., 2020](#)). Aragonite formation and preservation are mostly limited to
57 <10°C/km thermal gradients and rapid exhumation in dry conditions (e.g., [Carlson, 1980](#);
58 [Hacker et al., 1992](#)).

59 We report the first record of HP aragonite in the Precambrian carbonates of the Jabal Akhdar
60 Dome (Oman Mountains), universally acknowledged as a non-subducted portion of the Arabian
61 Plate (e.g., [Breton et al., 2004](#)). We demonstrate that aragonite is therein selectively preserved
62 within narrow mylonitic shear zones that formed due to cyclic brittle-ductile deformation under
63 blueschist facies conditions during the Cretaceous obduction of the Semail Ophiolite. We
64 propose a model wherein the identified mylonitic shear zones acted as sheltering structural
65 capsules, within which aragonite initially formed at HP metamorphic conditions and through
66 which it survived exhumation all the way to the surface.

67

68 **2. GEOLOGICAL SETTING**

69 The Jabal Akhdar Dome is the largest tectonic window in the NW-SE Oman Mountains ([Fig.](#)
70 [1A-B](#)). It is composed of a pre-mid Carboniferous series overlain by a 2.5km-thick Permian to
71 Cretaceous carbonate succession (Autochthonous A and B, respectively). This stack was
72 overthrust by the Hawasina Nappe and the Semail Ophiolite during the Cretaceous (e.g.,
73 [Searle, 2007](#); [Fig. 1B](#)). The autochthonous series is part of the Arabian passive margin, which
74 is reported as having experienced subduction, exhumation and obduction only in its
75 northeasternmost external portion (e.g., [Lippard, 1983](#)). Therefore, deformation in the
76 Precambrian JAD has so far been predominantly related to a complex sequence of contractional
77 and extensional phases from the late Ediacaran to the early Cambrian (e.g., [Callegari et al.,](#)
78 [2020](#); [Scharf et al., 2021](#); [Fig. 1C](#)), with evidence of Cretaceous continental subduction

79 exclusively confined to the Saih Hatat Dome (Fig. 1B), where rocks experienced up to eclogite
80 facies conditions (e.g., Warren et al., 2003; Agard et al., 2010). The JAD in the inner part of the
81 Arabian margin is thus considered to belong to the un-subducted portion of the Arabian passive
82 margin, and to have experienced only anchizonal deformation conditions during the Semail
83 Ophiolite obduction (e.g., Searle, 2007). Exhumation of the subducted Arabian margin through
84 a still ill-defined mechanism started at the end of the Cretaceous (e.g., Hansman et al., 2021),
85 leading to rapid cooling of the whole region (e.g., Grobe et al., 2019).

86

87 3. METHODS AND RESULTS

88 3.1 ARAGONITE-BEARING MYLONITIC SHEAR ZONES

89 In the study area (Fig. 1C), contractional calcmylonitic shear zones cut across the Precambrian
90 organic-matter-rich carbonate of the Hajir Formation (Fig. 2; Fig. S1). Strain localised in 1)
91 discrete, up to 20cm thick mylonitic to ultramylonitic shear zones (Figs. 2A-B; S1) that
92 discordantly cut across the bedding and 2) thinner shears along bed-bed interfaces within NE-
93 verging folds in response to a flexural slip folding mechanism (Fig. 2). In the former, the
94 mylonitic foliation is penetrative and defined by lenses and laterally continuous films of
95 graphite (Figs. 2C; S1), dips to the SW (Fig. 2B) and is invariably associated with a SW-
96 plunging stretching lineation and thrusting-related top-to-NE S/C-structures and oblique
97 foliation (Fig. 2B). The mylonitic foliation commonly transposes quartz and calcite veins (Fig.
98 2B). Carbonates exhibit a strong shape preferred orientation (SPO, Fig. 2C) with grains
99 elongated parallel to the transport direction and contained within the foliation. They preserve a
100 peculiar “rod-like” shape (Fig. 2C), which is commonly reported as typical for calcite
101 pseudomorphs after aragonite (e.g., Brady et al., 2004; Seaton et al., 2009; Giuntoli et al.,
102 2020), and as ensuing during shearing at the calcite to aragonite phase transformation. Such
103 rod-shaped fabric is not found within the less deformed Hajir Fm outside of the shear zones
104 (Fig. S1).

105 Shear zones related to flexural-slip folding along bed-bed interfaces are extremely localised and
106 thin (<5cm; Fig. 2D-E) and contain sheared quartz and calcite-aragonite veins and fibres (Fig.
107 2E). They exhibit a strong SPO with calcite and aragonite grains aligned parallel to the
108 mylonitic foliation, embedding and transposing lenses and films of graphite (Fig. 2F).
109 All shear zones document broadly coeval brittle and ductile deformation, with the mylonitic
110 fabric cut across by mode-I veins with opening directions subparallel to the regional SW-NE
111 shear direction. These veins contain aragonite and quartz fibres (Fig. 2F; Fig. S2A-B) and are,
112 themselves transposed along the mylonitic foliation, suggesting cyclic brittle-ductile
113 deformation (Figs. 2B-F; S2C-D; S3A-B).

114

115 ***3.2 P-T CONSTRAINTS***

116 Raman Spectroscopy on Carbonaceous Material (RSCM) spectra and high-resolution micro-
117 Raman maps (see Supplementary Material) were acquired to discriminate calcite from
118 aragonite in both mylonites and mode-I veins. Representative RSCM spectra from graphitic
119 films and patches, which outline the mylonitic foliation, constrain synkinematic deformation
120 temperatures in the 336-347°C range (Fig. 3A). Micro-Raman maps and spectra document
121 aragonite (only partially retrogressed to calcite) in both mylonitic shear zones and mode-I veins
122 (Fig. 3B-C), thus confirming inferences from the mylonitic rod-shaped fabric in the shear zones
123 (Fig. 2C and F).

124 Based on the calcite-aragonite stability field (e.g., Johannes and Puhan, 1971), the obtained
125 deformation temperature range requires 0.8-0.9GPa (Fig. 3D) as minimum pressure during
126 shearing and coeval brittle deformation, thus within the range of blueschist facies conditions.
127 Trace element analysis on aragonite fibres and grains in mylonites constrains Sr to <1wt%
128 (Table S1; Fig. S4 and S5), thus excluding lower P values for the calcite-aragonite
129 transformation (e.g., Carlson, 1980).

130

131 **3.3 AGE OF DEFORMATION**

132 In order to temporally place the formation of the studied shear zones in the evolution of
133 northeastern Oman, we dated by U-Pb the described tectonic carbonates. Ages from mode-I
134 carbonate fibres cluster at 96.5 ± 31.6 Ma (2σ ; Fig. S6A-C, E). Dating of the calcmylonitic shear
135 zones defines two different clusters at 97.5 ± 23.3 Ma and 573.3 ± 27.6 Ma, respectively (2σ ; Fig.
136 S6D-F). The large analytical uncertainties notwithstanding, we interpret the new dates as
137 related to the Cretaceous subduction-obduction of the Semail Ophiolite, whereas the
138 Precambrian date as the age of the Hajir Fm protolith, which is ascribed to the early-middle
139 Ediacaran (e.g., Scharf et al., 2021).

141 **4. DISCUSSION AND CONCLUSIONS**

142 This first documentation of blueschist facies conditions from genetically connected shear zones
143 and mode-I veins within the JAD and the interpretation of this data require considering the
144 possible role of non-lithostatic pressure during deformation. Tectonic overpressure may
145 promote HP metamorphism, such that analytically obtained P estimates cannot be converted to
146 depth. Several conditions are discussed as leading to tectonic overpressure, among which the
147 following might affect our results: i) significant viscosity contrast between juxtaposed rock
148 types, ii) heterogenous pressure conditions in folds and iii) fluid release (dehydration,
149 decarbonation) in a closed system (e.g., Mancktelow, 2008; Schmalholz and Podladchikov,
150 2013; Luisier et al., 2019).

151 We tend to exclude a significant contribution of tectonic overpressure during the evolution of
152 the studied shear zones in the JAD based on the following:

- 153 - Mode-I fracturing causes instantaneous pressure drops in dilatant fractures to conditions
154 corresponding to the local lithostatic load. Coeval fracture infilling is, therefore, at
155 equilibrium with the lithostatic pressure and aragonite in veins thus reflects the
156 overburden.

- 157 - Mylonitic shearing occurred in the absence of a significant viscosity contrast between
158 shear zone and the undeformed host rock, as evidenced by their similar composition and
159 by aragonite occurring in both shear zones and mode-I veins.
- 160 - Shear zones locally formed by flexural slip along bed-bed interfaces on fold limbs,
161 which are not prone to develop significant overpressure (e.g., [Mancktelow, 2008](#)).
- 162 - Shearing occurred at too low a temperature to release CO₂-rich fluids from
163 decarbonation (e.g., [Samtani et al., 2002](#)).

164 Theoretically, aragonite might form even at low-pressure conditions in highly-strained
165 domains, where shear strain would increase the internal energy of calcite crystals by densifying
166 defects and dislocations in the lattice ([Newton et al., 1969](#)). This process, however, has been
167 described for late calcite veins cutting across mylonitic fabrics and is thought to be rather
168 unlikely in natural shear zones ([Gillet and Goffé, 1988](#)).

169 We therefore consider our minimum P estimate of 0.8-0.9GPa as indicative of at least 25-30km
170 depth, which attests to reaching blueschist facies conditions in the dynamic context of the
171 Oman Cretaceous subduction. This is consistent with the new U-Pb results, which, despite the
172 large analytical errors (24-30%; [Fig. S6E-F](#)), fit available age constraints for the known HP and
173 subduction-related metamorphism in northeastern Oman ([Fig. S6G](#)). We also exclude an
174 Ediacaran sedimentary origin for aragonite because we deem highly unlikely its preservation
175 over such a long time span due to the rapid kinetics of the aragonite/calcite transformation (e.g.,
176 [Theye and Seidel, 1993; Hacker et al., 2005](#)).

177 Cyclic brittle-ductile deformation under HP conditions in the JAD finds analogies with
178 descriptions of similar deformation styles under blueschist facies conditions elsewhere in
179 orogenic belts (e.g., [Molli et al., 2017; Giuntoli and Viola, 2022](#)) and allows us to present a
180 conceptual model for the formation and preservation of aragonite within mylonitic shear zones
181 ([Fig. 4](#)). Formation of aragonite is known to be both favoured and limited by the i) amount of
182 strain, ii) dynamic evolution of the permeability tensor through time and fluid-rock interaction,

183 iii) temperature changes, iv) overall composition of the deforming system and v) exhumation
184 modes (Fig. 4A). The interplay of these factors within the studied shear zones would have
185 made it possible for aragonite to first form and then be preserved. We suggest that the
186 mylonitic fabric triggered transient chemical/physical processes that were, instead, unable to
187 operate in the less deformed Hajir Fm outside of the shear zones (as shown by the lack of
188 aragonite therein; Fig. S1).

189 Shear zones cyclically accommodated broadly coeval ductile and brittle deformation,
190 both taking place under blueschist facies conditions, as demonstrated by aragonite occurring in
191 both coeval mylonites and mode-I veins. Progressive strain localisation within the initially
192 undeformed Hajir Fm under increasing pressure and temperature (Fig. 4B) caused calcite
193 recrystallization and grain size reduction (Fig. 4C). Fluids enhanced the calcite-aragonite
194 transformation at favourable P-T conditions that were met as the Hajir Fm reached
195 progressively greater depth during subduction (Carlson, 1980). We see this process as being
196 favoured by enhanced permeability along the mylonitic foliation due to the progressive
197 alignment of ordered graphite films along the mylonitic foliation planes (Fig. 4C; Upton and
198 Craw, 2008). Fluid ingress and flow within shear zones was facilitated by broadly coeval
199 fracturing and viscous deformation cyclically alternating under HP conditions (e.g., Molli et al.,
200 2017) and by the establishment of a large foliation-parallel permeability, with shear zones thus
201 acting as conduits (Fig. 4D).

202 Close to or at P-T peak conditions, rod-shaped aragonite grains formed (Fig. 4C-D)
203 growing an interwoven network with graphite crystals to develop laterally continuous and
204 extensive layers (Fig. 4E). This would have sealed the shear zones, eventually leading to lower
205 permeability conditions compared to the less deformed hosting Hajir Fm. Low-permeability
206 and sealed shear zones characterised the exhumation retrograde path of this part of the
207 subducted slab (Fig. 4E). Finally, the Hajir Fm re-entered the stability field of calcite at
208 $T \leq 200^\circ\text{C}$ (e.g., Hacker et al., 1992; Fig. 4A and F), with shear zones remaining relatively

209 sealed and dry, thus fulfilling a key condition for aragonite preservation (e.g., [Gillet and Goffé,](#)
210 [1988](#)). When coupled with a low geothermal gradient ($<10^{\circ}\text{C}/\text{km}$) typical of cold subduction
211 zones and fast exhumation ($\sim 1\text{-}3\text{mm/a}$), as indeed reported for NE Oman (e.g., [Grobe et al.,](#)
212 [2019](#)), our proposed structural evolution explains the seldomly observed crystallographic and
213 morphologic preservation of aragonite in the stability field of calcite.

214 In conclusion, we document for the first time the occurrence of metamorphic aragonite
215 in the JAD, attesting to blueschist facies conditions during subduction to at least 25-30 km
216 depth. Its preservation calls for a model based on structural capsules where HP phases
217 selectively formed and were later preserved by the concomitant effects of regional (rapid
218 exhumation) and local factors (low permeability horizons). Our results provide new insights
219 into both the regional evolution of the JAD and the processes allowing for the preservation of
220 HP phases in deeply subducted and exhumed carbonate(meta)sedimentary successions.

221

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227

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334

335 CAPTIONS TO FIGURES

336 Fig. 1. A) Geographical setting and B) Geological setting of the southern Oman Mountains
337 with the Jabal Akhdar (JAD) and Saih Hatat (SHD) Domes; from [Searle, 2007](#). C) Geological
338 map of the study area and local stratigraphy (from [Callegari et al., 2020](#)).

339

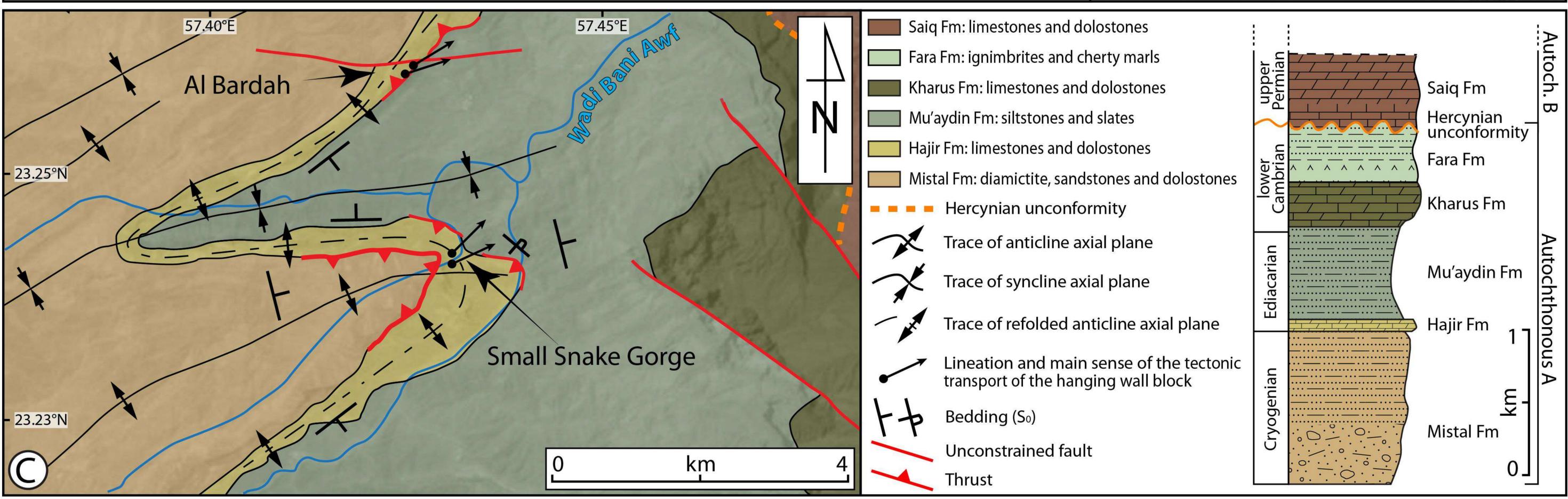
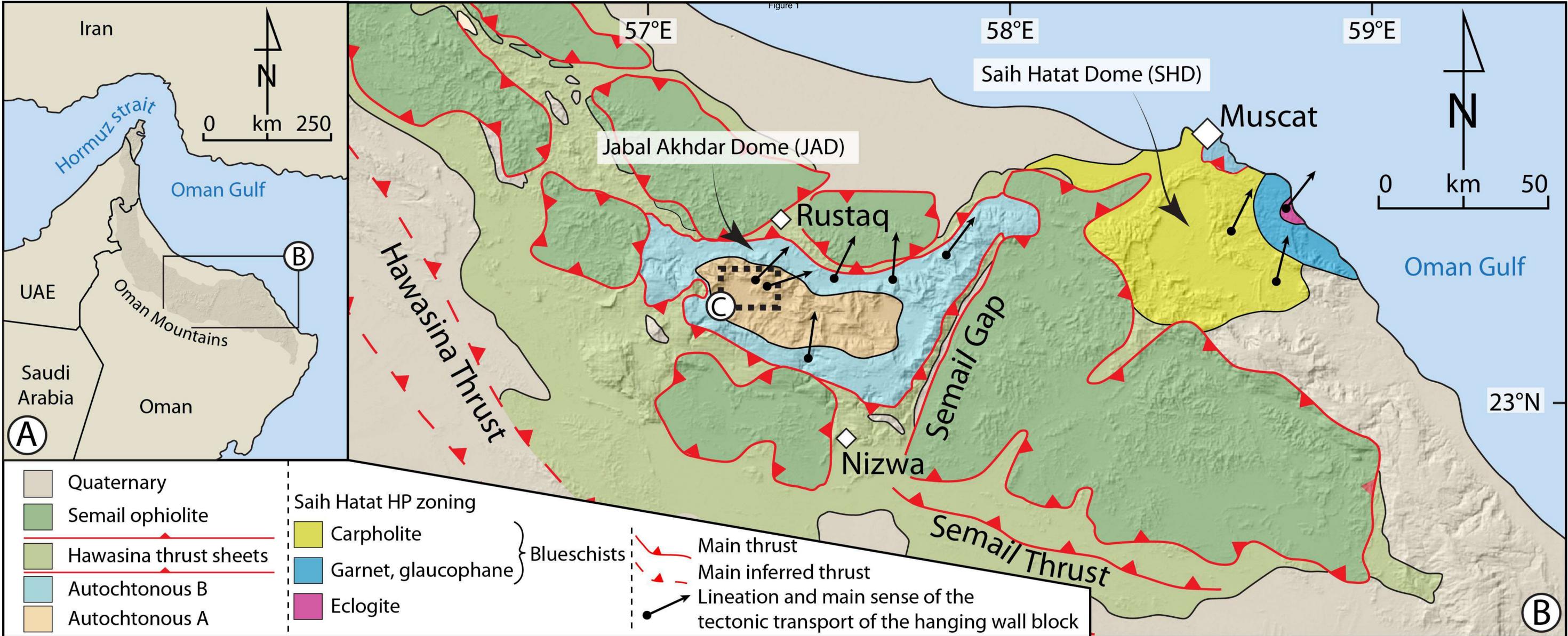
340 Fig. 2. A) Calcmylonitic shear zone cutting a NE-verging fold. B) Detail of mylonitic shear
341 zone in (A). C) Ultramylonitic shear zone in (B), with rod-shaped aragonite crystals interleaved
342 with semi-continuous graphitic films defining the foliation. D) NE-verging fold with flexural-
343 slip related shear zones along bed-bed interfaces. Men for scale. E) Detail of shear zone in (D)
344 deforming fibrous quartz and aragonite. F) Aragonite-calcite mylonitic foliation cut across by
345 mode-I stretched fiber quartz and aragonite veins.

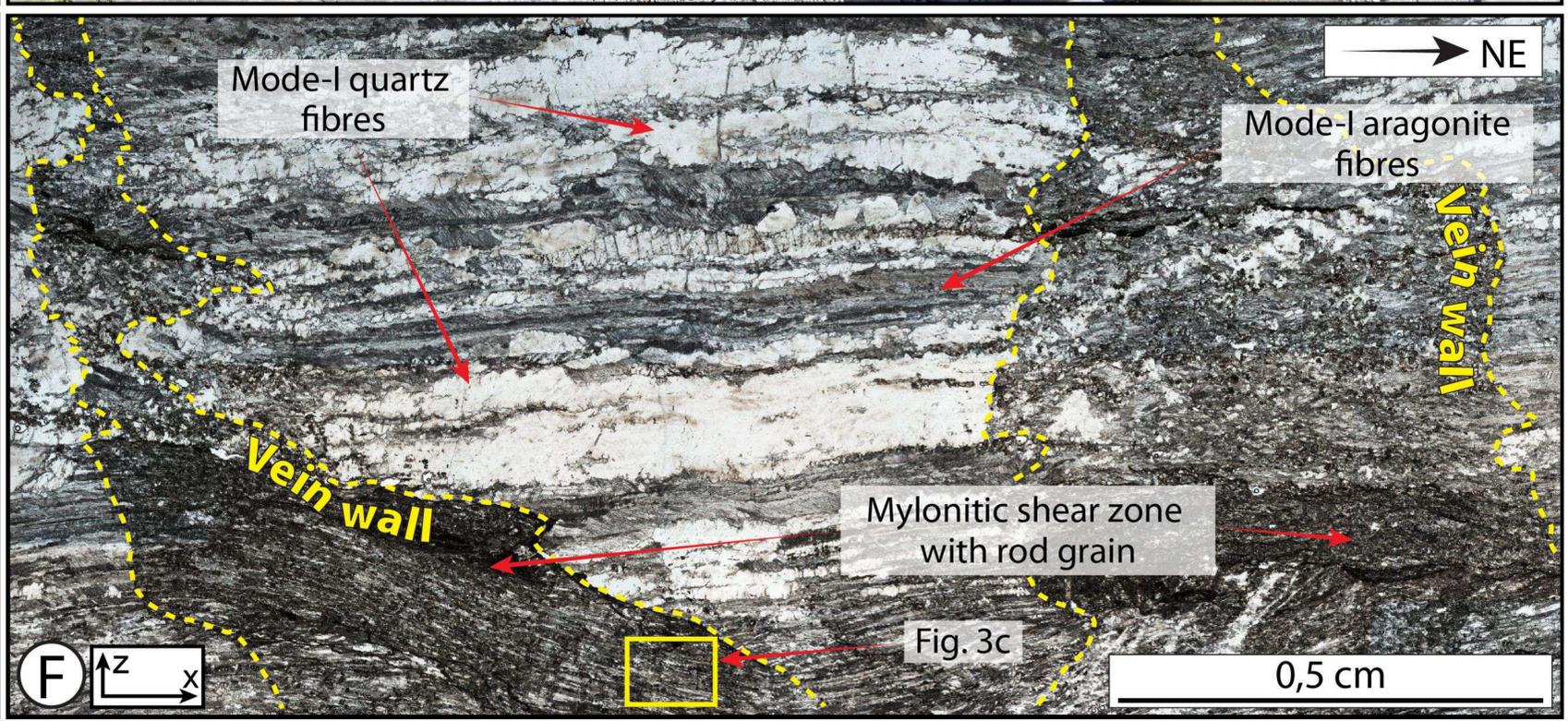
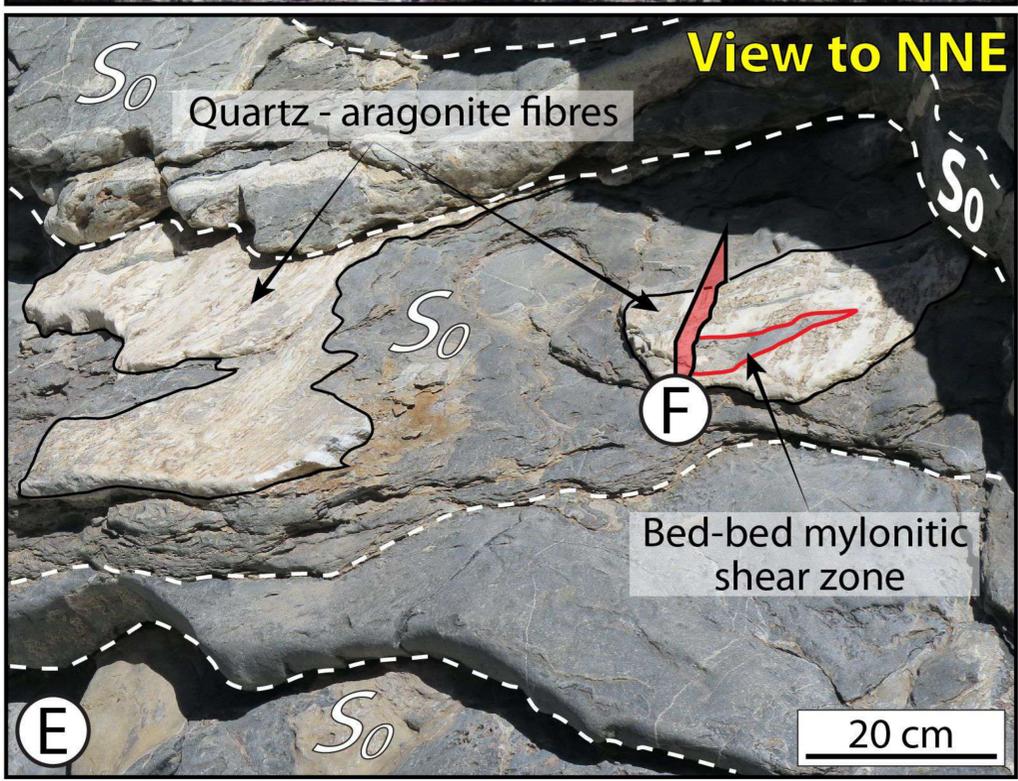
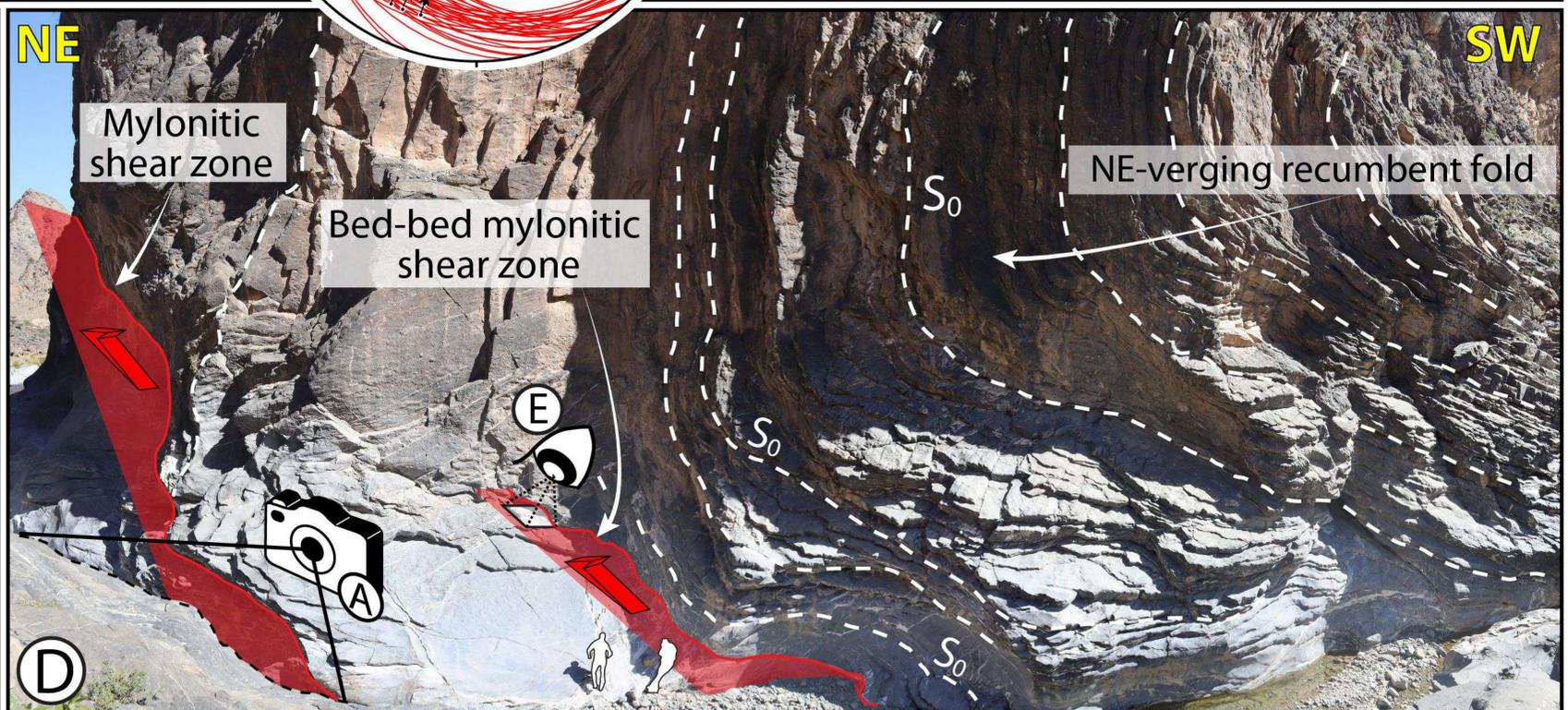
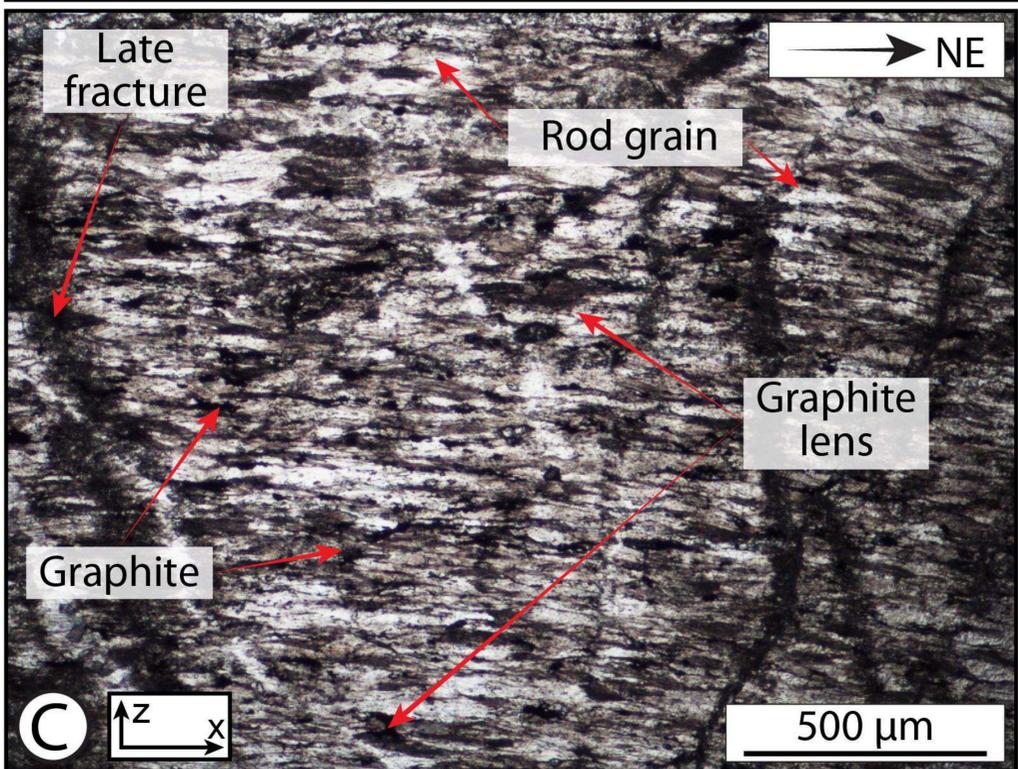
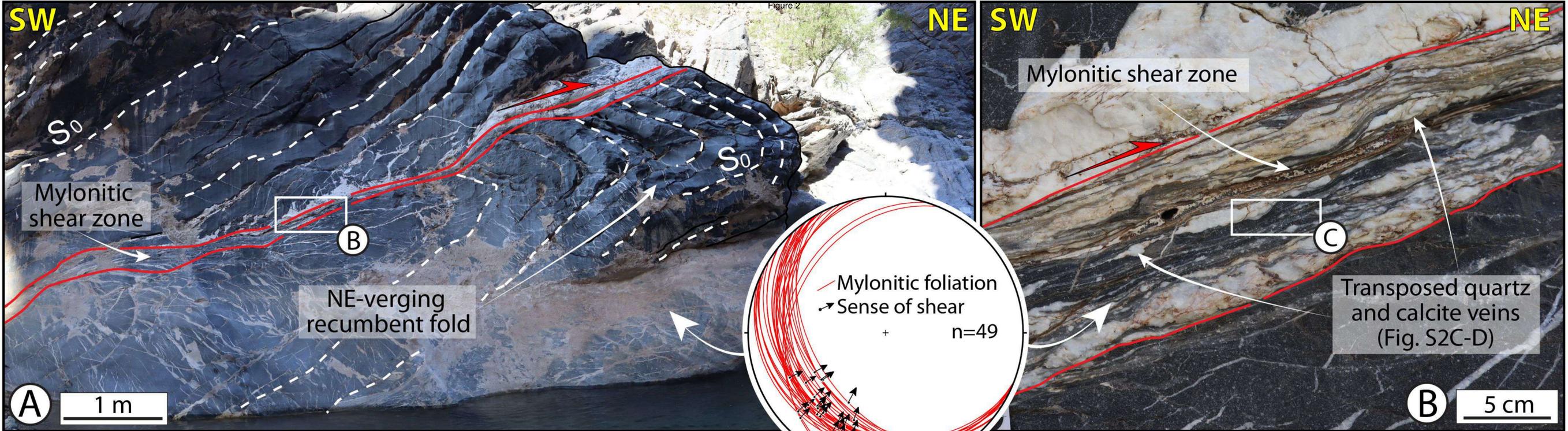
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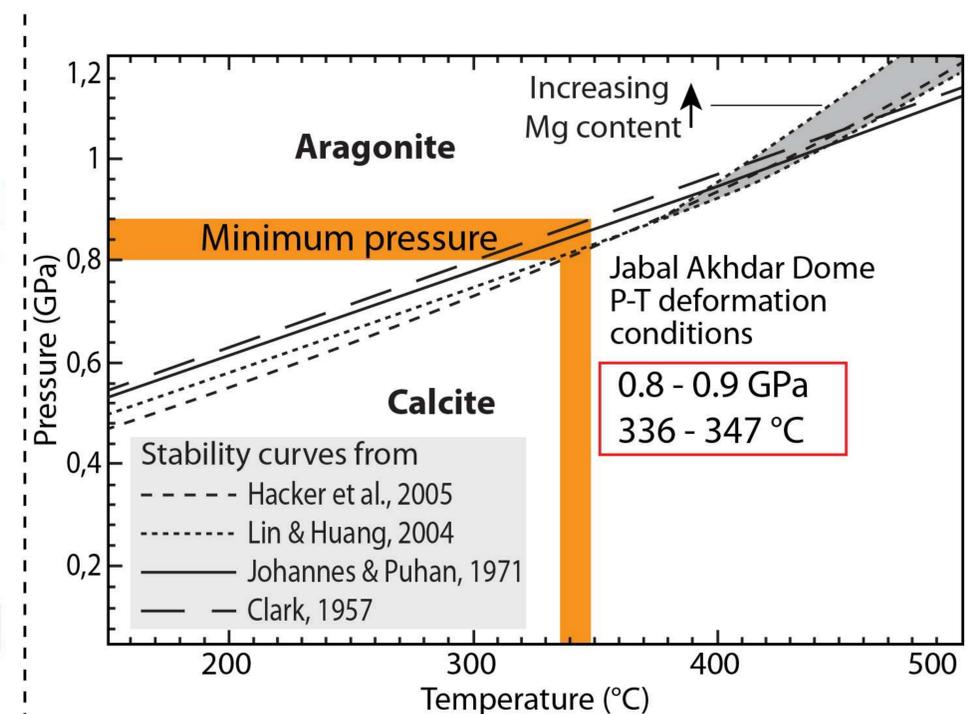
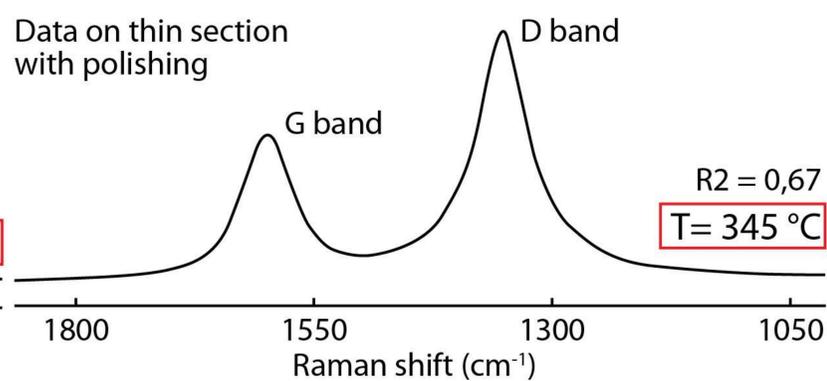
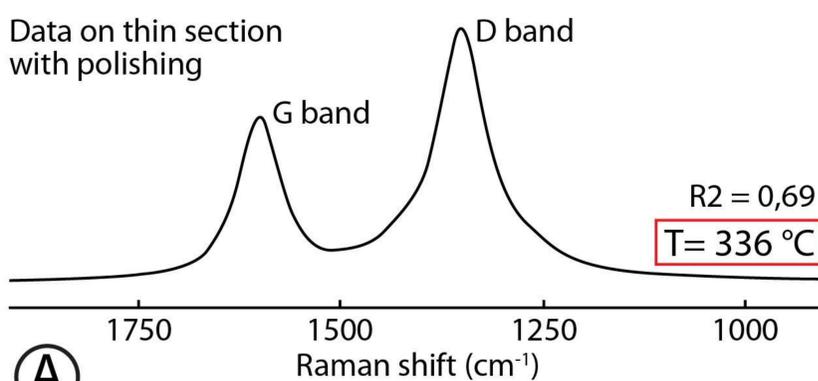
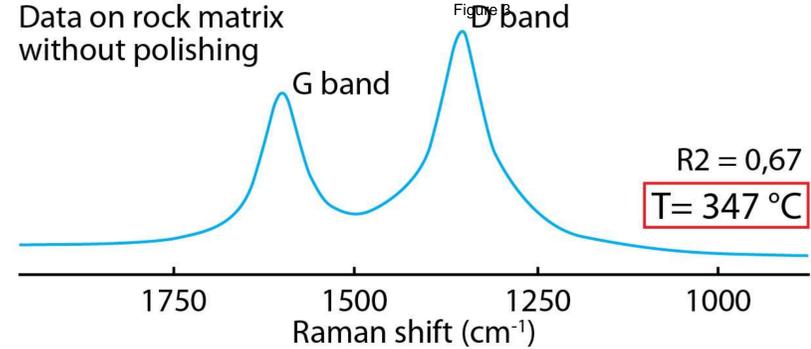
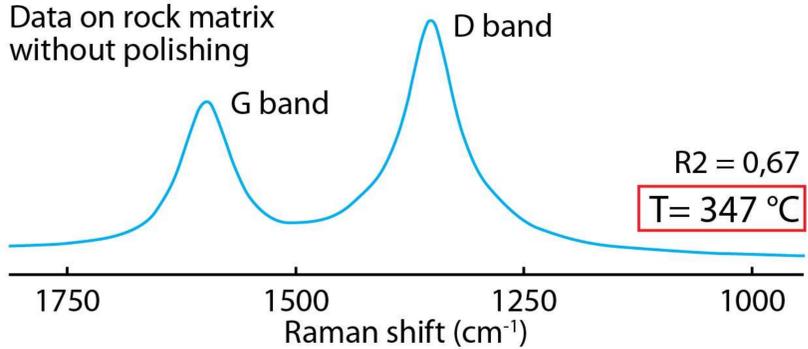
347 Fig. 3. A) Raman spectra from carbonaceous material (RSCM) on rock matrix with (black) and
348 without (blue) polishing. B-C) High-resolution micro-Raman spectroscopy spectra, related
349 microphotograph and phase map. Bright yellow aragonite in blue ellipses. D) Calcite-aragonite
350 stability field. Different experimental stability curves are shown (from Clark Jr., 1957;
351 Johannes and Puhan, 1971; Lin and Huang, 2004; Hacker et al., 2005). The estimated P-T
352 conditions for studied JAD carbonates are indicated.

353

354 Fig. 4. Model for the formation/preservation of aragonite. A) P-T retrograde path. B)
355 Undeformed Hajir Fm carbonate. C) Calcite dynamic recrystallization forming new grains. D)
356 Shear zone structuring at P-T peak conditions and formation of rod-shaped aragonite grains.
357 with Quartz and aragonite fibre mode-I veins locally transposed by mylonitic fabrics constrain
358 brittle-ductile cyclicity under HP conditions. E) Shear zone sealing by high-ordered graphite
359 coalescence, forming thin, permeable and laterally continuous layers. F) Present configuration
360 of the JAD mylonitic shear zones, acting as sheltering structural capsules preserving aragonite.







A

D

