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Energy and water balance of a treatment wetland under mediterranean climatic conditions

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A B S T R A C T

During 2013–2014 period, the surface renewal (SR) – energy balance (EB) method combined with the water balance approach was adopted to define the energy and water budget of a treatment wetland (TW) in Eastern Sicily (Italy). The main objective was to evaluate the differences between energy and water balances and understand the underlining mechanisms of environmental control. Results shown that the seasonal variation in evapotranspiration (ET) had similar trend as the seasonal variation of leaf area index (LAI) and net radiation (R_n). During the study period, the daily rate of ET by EB ranged from -0.1 mm d^{-1} to 9.3 mm d^{-1} , with a total ET of about 1837 mm in 17 months. The crop coefficient K_c (e.g. obtained using the FAO-56 mean K_c approach) for the TW ranged from 0.5 to 2.4, with a mean annual value of 1.24. On a 10-day mean basis, ET fluxes from the TW water balance were 15% higher than ET by EB.

Keywords:

Treatment wetland
Evapotranspiration
Surface renewal
Water balance

1. Introduction

Treatment wetlands (TW_s) are designed plants where the natural mechanisms and interactions of vegetation, soils and the microbial populations contribute to reduce surface water or wastewater contaminants (Headley et al., 2012). Their use, as decentralized systems for the treatment of domestic water, is increasingly mainly due of their low operation and maintenance requirements (Cirelli et al., 2007, Barbagallo et al., 2003, 2011) and low risk associated to human contact.

In TW_s, water cycle and the associated energy exchange process have a fundamental role in the operation of these systems (Rejšková et al., 2012). The wetland plants, both in natural wetlands and in those constructed, are adapted to grow in a not limited water availability conditions, with associated high evapotranspiration (ET) rate (Borin et al., 2011).

A detailed knowledge of ET can be crucial for the functioning of TW_s. In fact, especially when the treated water are indispensable resources to be reused for irrigation (Barbagallo et al., 2012, Toscano et al., 2013, Castorina et al., 2016), excessive ET losses may increase salt concentrations in the effluent (Headley et al., 2012).

The hydrological processes occurring in TW_s, such as horizontal flow wetlands, are different from those in the natural wetland systems, due to the infiltration and surface runoff exclusion and the

identification of stream flow with wastewater influent and effluent (Headley et al., 2012), which can be measured with flow metering devices. In these systems, the direct measurement of ET fluxes generally present difficulties (Price, 1994, South et al., 1998, Moro et al., 2004, Noormets et al., 2006).

Several studies have explored how to directly measure or estimate ET fluxes from natural or TW systems (e.g. Idso, 1981, Idso and Anderson, 1988, Abtew and Obeysekera, 1995, Soucha et al., 1996, Wilson and Baldocchi, 2000, Drexler et al., 2004, Toscano et al., 2015), and ET term has become a subject of debate. However, few studies have dealt with the analysis of the relationship between ET, the surface energy fluxes and the related physical processes especially in TW systems (Goulden et al., 2007, Clulow et al., 2012). Different procedures of direct measurement or estimation of ET fluxes were adopted, and, as consequence, significantly different results were obtained, thus contributing to increase the uncertainty about the most appropriate methods for treatment wetland ET determination.

Relevant studies were conducted by Abtew and Melesse (2013) on ET direct measurement by lysimeters, other by Burba et al. (1999) and Peacock and Hess (2004) on the adoption of micrometeorological methods, based on the energy fluxes budget, to measure ET from wetland systems.

Among the micrometeorological methods, the eddy covariance (EC) technique is recognized as one of the most accurate for estimating

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wetland ET (Acreman et al., 2003, Goulden et al., 2007, Siedlecki et al., 2016), even if it involves relevant instrumental costs (Clulow et al., 2012) and specialized personnel for data processing. An important advantage of the EC method is that the obtained water vapour flux comes from a source area within a distance of 200–300 m from the sensor.

In the recent past, the applied research has recognized the Surface Renewal (SR) technique as valid alternative to EC sonic anemometer for sensible heat flux (H) direct measurements (Drexler et al., 2004). The SR is reasonable accurate, cheaper than EC and easily maintained for heterogeneous systems measurements like wetlands. When SR is integrated with direct measurements of net radiation and soil heat flux, it allows to estimate ET as residual of the energy balance equation.

Other studies conducted within treatment wetlands have calculated ET fluxes by the resolution of the water balance using off-site data (Kadlec and Wallace, 2008) or by simple meteorological equations such as the Penman-Monteith model, and other hydrological and crop-growth models (Annandale et al., 2003, van Heerden et al., 2009). The use of these models requires vegetation specific parameters for calibration (Mao et al., 2002).

In this study, the eddy covariance (EC) method was adopted over a period from July 2013 to December 2014, to determine energy fluxes and ET from a treatment wetland system in Eastern Sicily (Italy). The results on ET fluxes were compared to ET estimates from both a hydrologic water balance of the wetland and the FAO-56 Penman-Monteith model. The valuable alternative technique surface renewal (SR) was also adopted at the site to estimate sensible heat fluxes (H_{SR}) from TW areas. The results on ET fluxes were compared to ET estimates from both a hydrologic water balance of the wetland and the FAO-56 Penman-Monteith model.

The results of this study are crucial from the point of view of the works aimed at assessing the role of ET in the functioning of the TW systems. Moreover, despite the extensive research on the topic of TW, there are still few studies on the complex characteristics of evapotranspiration in Italy's full-scale wetland systems.

2. Materials and methods

2.1. The study site

The study site is the largest TW of Sicily (South Italy), used as tertiary wastewater treatment of 5000 Persons Equivalent (PE) (Fig. 1).

It is located in the small village of S. Michele di Ganzaria in Eastern Sicily ($37^{\circ}17'0''N$, $14^{\circ}26'0''E$). The TW areas ($12,700\text{ m}^2$) is part of the wastewater reuse project of the village, consisting of four horizontal sub-surface (H-SSF) reed beds (H-SSF1, H-SSF2, H-SSF3 and H-SSF4), followed by three batch wastewater storage reservoirs (S1, S2 and S3)

working in parallel (Fig. 2) (Castorina et al., 2016).

The sewerage system collects wastewater from almost every household in a conventional wastewater treatment plant (WWTP), north-west of San Michele di Ganzaria, where urban wastewater receive a pre-treatment step followed by two parallel water lines (Imhoff tank, trickling filter and a secondary sedimentation tank). For tertiary treatment, the effluent is sent in four H-SSF beds (about 2 L s^{-1} per bed), of which H-SSF1 and H-SSF2 are in operation respectively since 2001 and 2006, while H-SSF3 and H-SSF4 since summer 2013.

The research focused on the study of the H-SSF3 TW bed, that has a surface area of about 2300 m^2 ($36 \times 64\text{ m}$) and a theoretical flow rate of about $173\text{ m}^3\text{ d}^{-1}$. The TW depth (e.g. 0.6 m) is filled with 8–10 mm of volcanic gravel. The TW bottom has a slope of 1% and the average water depth is 0.5 m. Both the excavated beds and banks were lined with a 4 mm thick bentonite sheet to prevent groundwater contamination. Earth banks, with a slope of 3:1, were covered with jute netting to facilitate vegetation and prevent soil erosion. In September 2012, the H-SSF3 was planted with *Phragmites australis* (Cav.) Trin. ex Steud at a density of one rhizome m^{-2} ; the complete plant cover was reached a year later. The vegetation had average maximum height of 2.0 m, with homogeneous canopy cover and no areas of fully exposed substrate and plant roots were in water saturation conditions. The leaf area index (LAI) ranged between $0.4\text{ m}^2\text{ m}^{-2}$ in winter and $5.0\text{ m}^2\text{ m}^{-2}$ in summer.

In the TW, the influent is distributed at the bed-head through a perforated 200 mm PVC pipe above the substrate and oriented normal to the flow direction to allow homogenous wastewater distribution. Wastewater is intercepted downstream by a cross perforated pipe located in the final section at the bottom of the bed. Electromagnetic flow meters (ISOIL mod. MS2500), installed at the inlet and at the outlet pipes, measure the flow rate (L s^{-1}) in continuous. In the bed, nine piezometers (i.e. 0.2 m of diameter) were arranged in 3 rows; they consist of open perforated plastic tubes inserted into the substrate to the bottom of the bed, used to measure the water height and to collect wastewater samples.

The wetland area has been instrumented with an automatic weather station (Campbell Scientific, Logan, UT) to measure air temperature, wind speed and direction, rainfall, global radiation and relative humidity. These data were used to calculate the Penman-Monteith equation (Allen et al., 1998) for reference ET (ET_0) estimation. The FAO-56 Penman-Monteith model provides estimates of hourly ET from a hypothetical grass reference surface (ET_0) (Eq. (1)). It is well-known to estimate the crop coefficient (K_c) as ratio between crop evapotranspiration during no water-stress conditions (ET_c , in turn ET from EC or ET from TW water balance) and ET_0 .

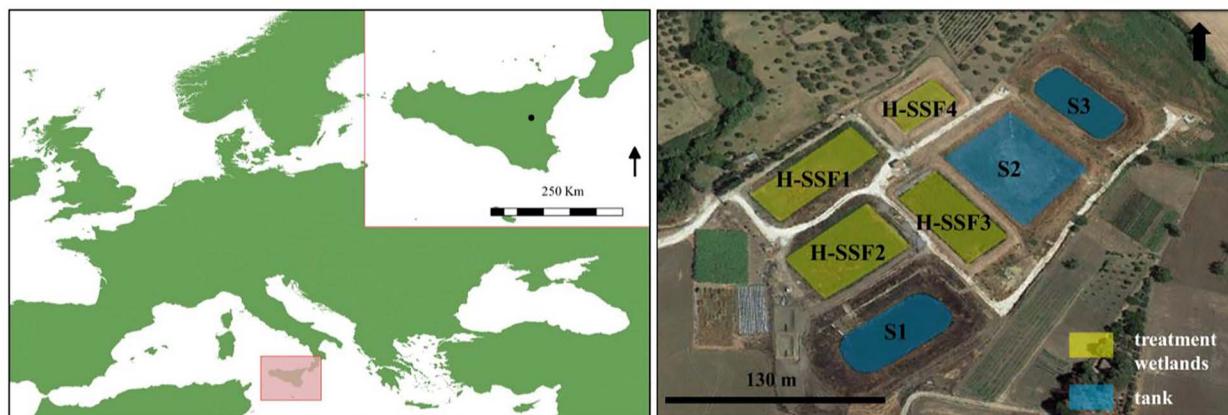


Fig. 1. Geographical location of the treatment wetland areas in Eastern Sicily.

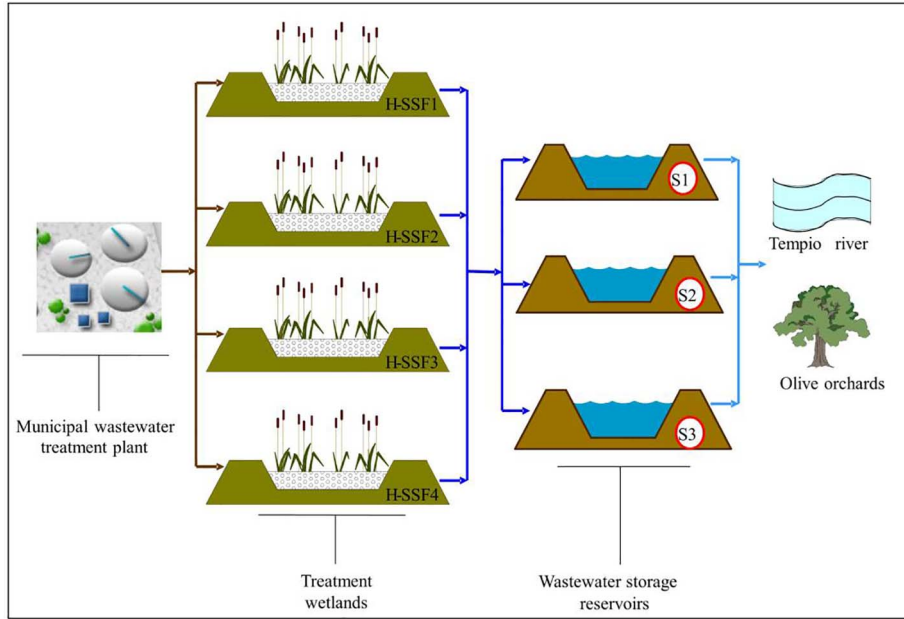


Fig. 2. Layout of the treatment wetland areas in Eastern Sicily.

$$ET_0 = \frac{0.408\Delta(R_n - G) + \gamma \frac{900}{T+273} u_2 (e_s - e_a)}{\Delta + \gamma(1 + 0.24u_2)} \quad (1)$$

In Eq. (1), Δ is the slope of the saturation vapour pressure at the air temperature (kPa °C⁻¹), γ is the psychrometric constant (kPa °C⁻¹), T is the daily (or hourly) mean temperature (°C), u_2 is the mean wind speed in m s⁻¹, and $e_s - e_a$ is the vapour pressure deficit (kPa).

2.2. Micrometeorological measurements

In July 2013, a micrometeorological energy balance tower was installed at the H-SSF3 TW bed. It was equipped with eddy covariance

(EC) sensors (Kaimal and Finnigan, 1994a,b), mounted at 4 m of height (about twice the plant maximum height) (Fig. 3).

At the site, net radiation (R_n , W m⁻²) was measured with a NR Lite net radiometer (Kipp&Zonen, Campbell Scientific Ltd) at a height of 6 m. Heat flux density (G , W m⁻²) was estimated as mean value of three heat flux plates measurements (HFP01, Campbell Scientific Ltd). The plates were inserted 0.05 m below the saturated substrate surface.

The air temperature and the three wind speed components were measured using a fine wire thermocouples (CS FW3, 0.0762 mm diameter) located at the roughness sub-layer, about 0.5 m above the canopy top, and a sonic anemometer (CSAT, Campbell Sci., at 8 m) at 4 m



Fig. 3. Eddy covariance footprint area of the TW system of S. Michele di Ganzaria.

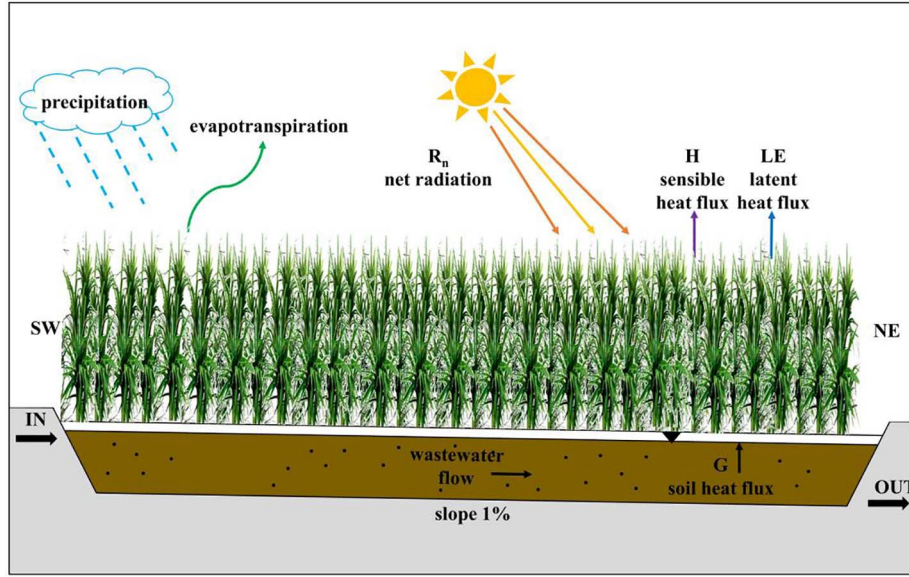


Fig. 4. Energy and water balance at the TW system.

of height, respectively. An infrared gas analyzer (LI-7500, LI-COR) located at 4 m was used for H_2O and CO_2 acquisitions. The EC data were recorded at 20 Hz of frequency. Fig. 3 shows the footprint area at the study site.

The standard EUROFLUX rules (Aubinet et al., 2000) were used for EC measurements and data processing. A data quality check was applied during the post processing together with some routines to remove the common errors: running means for detrending, three angle coordinate rotations and despiking. Stationarity and surface energy closure were also checked (Kaimal and Finnigan, 1994a,b).

To avoid underestimation of evapotranspiration, data gap on EC measurements were filled before determining the daily values. Gaps of several hours were filled using the average daily pattern of values calculated over a period of 14 days (Foken, 2008). On the other hand, in assessing the daily pattern of evapotranspiration, series with unfilled measurement gaps were used. This approach allowed to determine the statistic variability only on the basis of the EC data.

The eddy covariance sensible heat flux H ($W m^{-2}$) was calculated as:

$$H_{EC} = \rho \cdot c_p \cdot \sigma_{wT} \quad (2)$$

with, ρ ($g m^{-3}$): air density, c_p ($J g^{-1} K^{-1}$): air specific heat capacity at constant pressure and σ_{wT} ($m s^{-1} K$): covariance between the vertical wind speed and air temperature.

The flux of water vapour content, i.e. the latent heat flux (LE, $W m^{-2}$), was expressed as:

$$LE_{EC} = \lambda \cdot \sigma_{wq} \quad (3)$$

with, λ ($J g^{-1}$): latent heat of vaporization and σ_{wq} ($m s^{-1} K$): covariance between the vertical wind speed and water vapour concentration.

30-min latent heat fluxes, acquired in $W m^{-2}$, aggregated to a daily scale were transformed to ET rate ($mm d^{-1}$) (Mauder and Foken, 2004; Mauder et al., 2007).

In the study, the surface renewal (SR) technique (Paw and Brunet, 1991; Snyder et al., 1996) was adopted to estimate the sensible heat flux density (Eq. (4)); SR was considered as alternative technique to the EC sonic anemometer. The technique required high frequency temperature data (4 Hz), collected using one $76.2 \mu m$ thermocouple mounted 0.5 m above the canopy top. The traces of temperature have ramp like characteristics, which were analysed to compute the mean ramp amplitude (a) and the inverse ramp frequency ($d + s$) using structure functions (Van Atta, 1977).

$$H_{SR} = \alpha \rho C_p \left(\frac{a}{d + s} \right) z \quad (4)$$

The factor α accounts for the unequal heating below the sensors (Paw et al., 1995; Snyder et al., 1996; Spano et al., 1997); ρ : air density ($g m^{-3}$) and C_p : specific heat of air at constant pressure ($J g^{-1} K^{-1}$). The ratio $\frac{a}{d+s}$ is the mean change in temperature ($^{\circ}C$ or K) with time (seconds) during the sampling interval (i.e. 30 min in these experiments).

In the study, sensible heat flux data from SR technique (H_{SR}) were calibrated with independent measurements of EC by the 3-D sonic anemometer. The calibration data allowed to obtain the α value of Eq. (3) by simple linear regression forced through the origin.

After calculating H_{SR} , LE was the residual term of the energy balance equation (Eq. (5)) on a half-hourly basis:

$$LE = R_n - G - H \quad (5)$$

Daily ET_{EC} were compared with ET_{SR} and the FAO-56 Penman-Monteith model.

2.3. The water balance equation

At the TW, a daily water balance was computed as follows:

$$P - ET + IN - OUT + \Delta S = 0 \quad (6)$$

with P: precipitation (mm), ET: evapotranspiration (mm), IN and OUT: inflow (mm) and outflow (mm), which represent the wastewater inflow and outflow controlled by pumps (Fig. 4) and ΔS : change in water storage (mm) in the saturated substrate. The water storage ΔS was neglected, since piezometers levels were similar during the experimental period. The simplified Eq. (6) was then solved to determine daily ET rates.

3. Results

3.1. Results on micrometeorological measurements and energy balance

A seasonal variation in air temperature characterized the study period (Fig. 5), with the warmest monthly mean temperature of $23.6^{\circ}C$ in July and the coldest value of $9.0^{\circ}C$ in December. Monthly precipitation showed a large variability, with most rainfall during December-March as typical of semi-arid climatic conditions. In 2014, the annual precipitation was 470 mm, and the corresponding annual mean

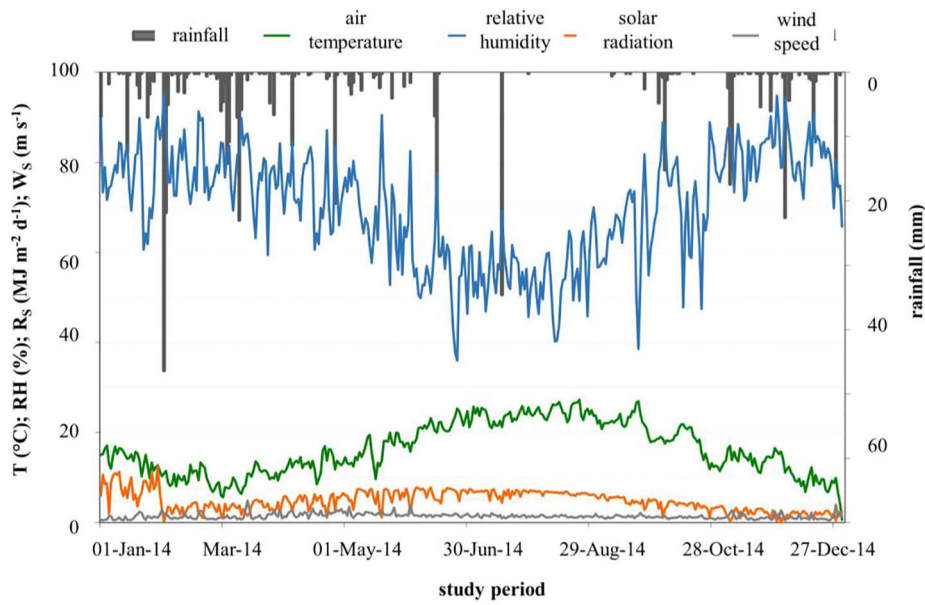


Fig. 5. Climatic variability during the study period in 2014: mean daily air temperature, $T(^{\circ}\text{C})$, mean daily relative humidity, $\text{RH}(\%)$, daily solar radiation, $R_s(\text{MJ m}^{-2} \text{d}^{-1})$, mean daily wind speed, $W_s(\text{m s}^{-1})$.

value of air temperature was 16°C .

The sum of latent (LE) and sensible (H) heat measured by the eddy covariance (EC) system located within the TW system was linearly correlated with the available energy, R_n-G at the 30-min time scale (Fig. 6a and b).

The energy balance closure was 0.89 and 0.90 in 2013 and 2014, respectively. These high values suggested reliable energy balance measurements (Consoli and Papa 2013) to examine the energy partitions within the studied treatment wetland.

The mean hourly net radiation (R_n) flux was of $120(\pm 212) \text{ W m}^{-2}$. R_n during the summer period (July-September) of the two years of trial had an hourly mean of $166(\pm 259) \text{ W m}^{-2}$. The latent heat flux from eddy covariance (LE_{EC}) was correlated with R_n , with a determination coefficient of 0.78; LE_{EC} had an hourly mean of $110(\pm 286) \text{ W m}^{-2}$, with a mean hourly value during summer of $213(\pm 193) \text{ W m}^{-2}$. The hourly mean sensible heat flux from eddy covariance (H_{EC}) was of $20(\pm 123) \text{ W m}^{-2}$, with a summer mean value of $30(\pm 76) \text{ W m}^{-2}$. Hourly average soil heat flux (G) was of $-4.3(\pm 9.7) \text{ W m}^{-2}$, with few positives during summer. The maximum value of 30-min G at a 5-cm depth was of 33.5 W m^{-2} . The minimum 30-min G was of -57 W m^{-2} during early morning hours. Overall G was a minor component of the overall energy balance of the TW system.

Energy partitioning showed a clear seasonal variability (Fig. 7).

The LE_{EC} dominated the energy balance, reaching the maximum of

$22.8 \text{ MJ m}^{-2} \text{d}^{-1}$ at the end of July 2014. During the dormant period (October-February) and when rainfall increased the soil water content, LE was similar to H.

The mean daily ET_{EC} from eddy covariance at the TW during July 2013-December 2014 was $3.6(\pm 1.8) \text{ mm}$; in summer the maximum daily ET_{EC} had a peak of 9.3 mm and a mean daily value of 6.3 mm (Fig. 8).

ET_{EC} variability was correlated with ET_0 and leaf area index (LAI) variation. The total ET_{EC} in late July 2013-December 2014, was of 1836.6 mm . The mean ET herein obtained is very similar to the results of other related researches with similar TW plantations (Fermor et al., 2001, Zhou et al., 2009, Headley et al., 2012, Irmak et al., 2013).

The hourly mean sensible heat flux from surface renewal (H_{SR}) was of $27.2(\pm 54.4) \text{ W m}^{-2}$, with a summer mean value of $35(\pm 63) \text{ W m}^{-2}$. ET_{SR} , obtained as residual of Eq. (4), had a mean of $2.7(\pm 1.8) \text{ mm d}^{-1}$ (Fig. 8). The SR method was found reliable, easy to operate in the field and sustainable for long-term use over treatment wetland systems. The SR technique requires to be calibrated with EC system (Castellví et al., 2012). Fig. 8 shows the comparison between ET_{EC} by eddy covariance system, ET_{SR} estimated as the residual term of the energy balance equation when H is surface renewal-based and ET_0 which was calculated with the FAO Penman-Monteith method. The values of ET_{SR} and ET_0 were on average about 89% of ET_{EC} .

The crop coefficient (K_c), calculated as ratio of daily ET_{EC} to ET_0

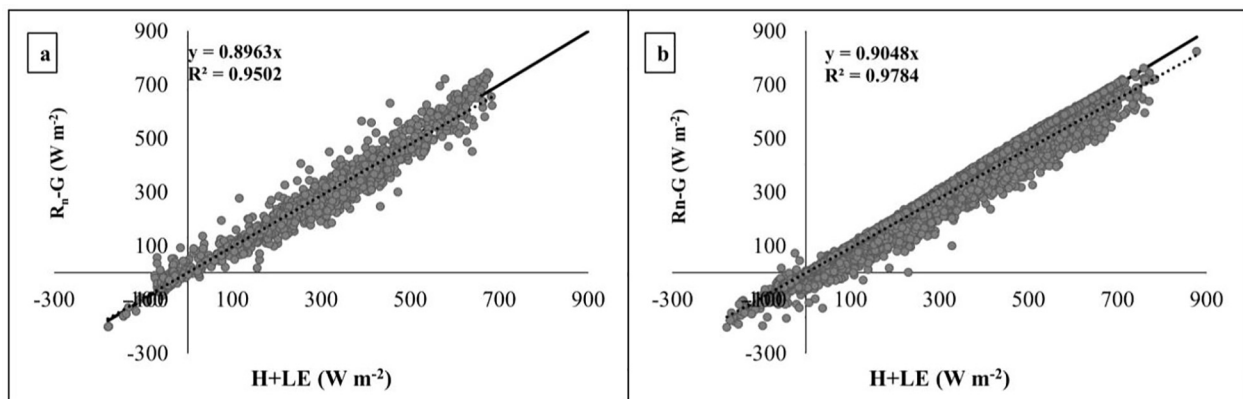


Fig. 6. Energy balance closure during 2013 (a) and 2014 (b) at the treatment wetland system.

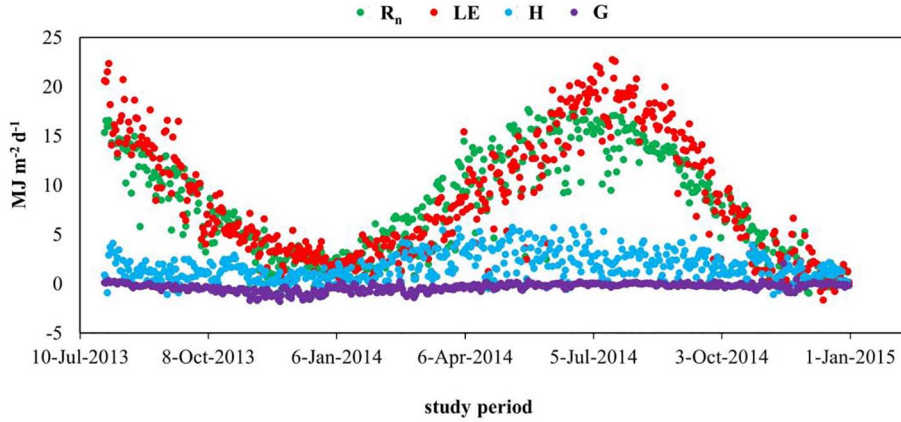


Fig. 7. Daily values of net radiation (R_n), soil heat flux (G), and eddy covariance sensible (H) and latent heat (LE).

showed a daily variability (Fig. 9).

K_c values exhibited a gradual increase from March onward, reached a peak in late July, and then went down again. The values of K_c were within the range 0.1–2.9, with a mean of $1.2 (\pm 0.4)$. The value of K_c for the *Phragmites australis* were within the range of crop coefficient values from other wetland studies (Drexler et al., 2008, Zhou et al., 2009).

3.2. Results on water balance

Decadal water balance for the TW system is presented in Fig. 10, with ET from Eq. (5), and the other variables directly measured.

The developing of the water balance was feasible and accurate in the period February–August 2014, but was problematic for some months of the hydrological year (data not showed) with negative values for the OUT term. The water budget indicated that, in the available period in 2014, ET had a mean of $6.6 (\pm 2.5)$ mm, IN and OUT were in their average of $60.9 (\pm 29.8)$ mm and $56.3 (\pm 28.9)$ mm, respectively. ET from water balance exceeded precipitation of about 80% and accounted for 3–25% of the IN received by the TW system (Fig. 11). Compared to ET coming from the water balance, the mean ET_{EC} measured with the eddy covariance (EC) system was about 15% lower (5.6 ± 2.4 mm). During the monitored period, the wetland lost a total of 1379.4 mm of water through evapotranspiration. The total ET_{EC} measured during the TW water balance period was of 1185.3 mm.

The mean 10-days K_c values relative to eddy covariance ET and water balance ET-based are reported in Fig. 12, with the 1:1 line.

K_c values had a mean of $1.5 (\pm 0.30)$ when ET was directly measured with EC, while the mean K_c from water balance ET-based was $1.9 (\pm 0.30)$, with peaks around 5.0 during July 2014 (data not showed). The values of K_c resulted from the two approaches were

significantly correlated, thus suggesting the reliability of micro-meteorological methods based on energy balance to estimate ET fluxes from TW systems.

4. Discussion

In this study the energy balance approach, based on the eddy covariance (EC), was adopted to determine ET fluxes in a TW used as tertiary wastewater treatment system for a small village in Easter Sicily, Italy. The surface renewal (SR) technique was also tested as alternative technique for estimating sensible heat fluxes (H_{SR}). It only requires calibration using an EC sonic anemometer. But, the re-calibration is necessary if there are significant changes in the vegetation canopy (Paw et al., 2005). SR is a low cost system, which needs simple and basic maintenance operations.

The energy balance of the TW system revealed a dominance of LE, with H_{EC} rarely exceeding.

Despite the footprint area of the EC tower was a complex surface, with areas of open water (i.e. storage reservoirs), soil and vegetation, the quite high energy balance closure has confirmed the good reliability of the technique to identify the energy partitions within the study system.

Results of ET energy balance were compared for a short period with ET values obtained from the resolution of the TW simplified water balance. The FAO-56 approach was applied for reference ET and K_c determination. Total ET from EC in 2013–14 period of monitoring was about 1837 mm for the *Phragmites australis*. Data showed that daily ET varied in response to the LAI variation and, consequently, to the phenology of the species. ET fluxes increased rapidly once the leaves expanded and reached peak rates in July. ET at the TW system was generally higher than reference ET_0 (total of about 1450 mm). The

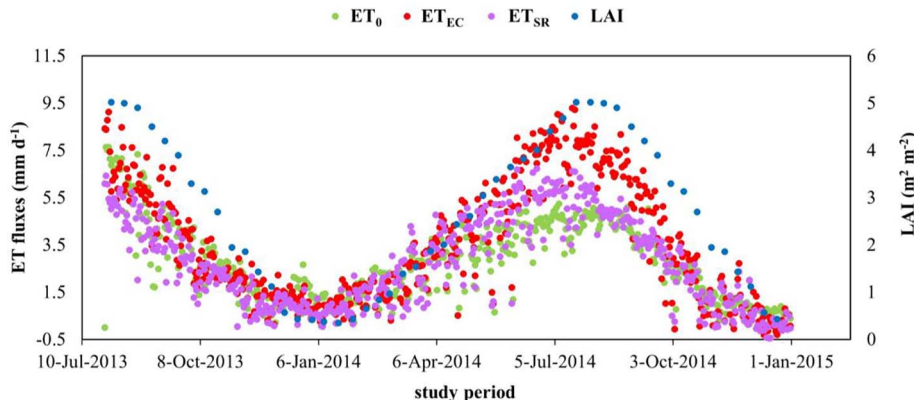


Fig. 8. Daily values ET_{EC} (mm), ET_{SR} (mm), ET_0 (mm) with LAI ($m^2 m^{-2}$) evolution during the study period.

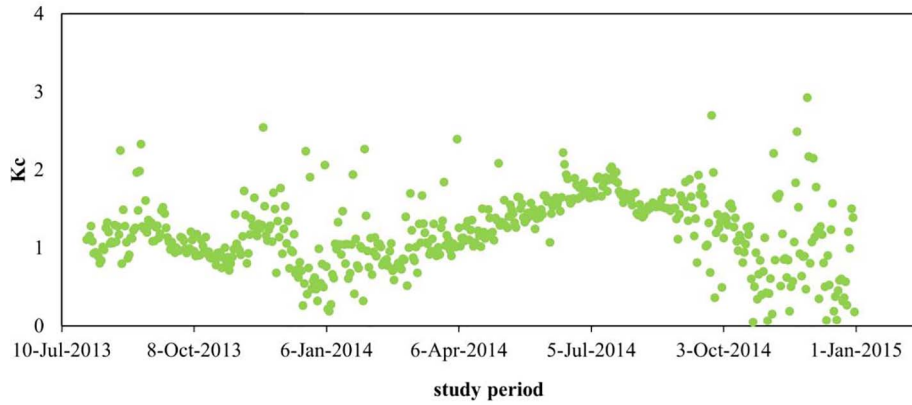


Fig. 9. Evolution of daily K_c during the study period.

greatest difference between ET_{EC} and ET_0 occurred in the spring-summer period of 2014, mainly as a result of the rapid vegetative growth of the species. On annual basis, K_c obtained as ratio between ET_{EC} and ET_0 -FAO56 based, was 1.24 at the study TW. On a daily scale, K_c was highly variable, within a range consistent with outcomes of similar studies.

The results of this study are of a certain interest because, although improvements on measurements techniques, and the dominant role of ET in TW system, there are few studies in Italy with actual measurements of ET in treatment wetland systems. At the same experimental study site, TW ET has been estimated by Borin et al. (2011) (daily maximum of 37.8 mm for reed) and Milani and Toscano (2013) at pilot spatial scale using water balance and climatic methods (K_c up to 8.5 in August 2009).

In the particular case of the TW system of S. Michele di Ganzaria (Eastern Sicily), which is part of the municipal sewage treatment plant, the study of the ET fluxes is fundamental for the analysis of the system functioning and for understanding its role in the environmental control processes (e.g. water vapour, gases and energy exchanges).

Except for the pilot-scale applications, the results obtained in our study, by using both the water and energy budget approaches, are similar to those obtained at international level on comparable or wider scales of applications.

In particular, Zhou and Zhou (2009) adopted the eddy covariance technique to estimate the magnitude and the dynamics of ET (average of 1.9 mm d^{-1} during the growing season) over a reed marsh in Northeast China. Headley et al. (2012) adopted the water balance method for estimating ET in a pilot-scale treatment constructed wetlands in Australia (average $ET = 7.0 \text{ mm d}^{-1}$). Clulow et al. (2012)

applied the micrometeorological SR method for estimating the ET fluxes (average of 3.2 mm d^{-1} in summer) from a coastal wetland in South Africa. Irmak et al. (2013) applied a two-step approach based on the Bowen ration energy balance system to measure ET for a mixed wetland system in Nebraska-USA (average seasonal ET of 3.8 mm d^{-1}). Anda et al. (2014) used a climatic approach to estimate ET (on annual scale average ET was between 2.1 and 7.7 mm d^{-1}) of common reed in Hungary.

Our results are of interest also in relation to the limits of the application of the water balance technique for real-scale systems. In fact, developing a TW water balance using measured P, IN and OUT and estimated ET did not result accurate in the whole monitoring period, due to management problems of wastewater IN coming from the secondary conventional treatment plant (i.e. tricking filter) serving the area of study in Sicily.

Therefore, the energy balance monitoring approach can be considered as reliable and capable to reasonable well capture the mass and energy partitioning within the heterogeneous TW system.

5. Conclusions

The main result coming from this research is the feasibility of the micrometeorological methods (based on EC or alternatively on SR for sensible heat fluxes) to be applied in such heterogeneous soil-plant-atmosphere environment.

The following main outcomes can be summarized:

- evapotranspiration in the TW system is highly correlated with vegetation growth;

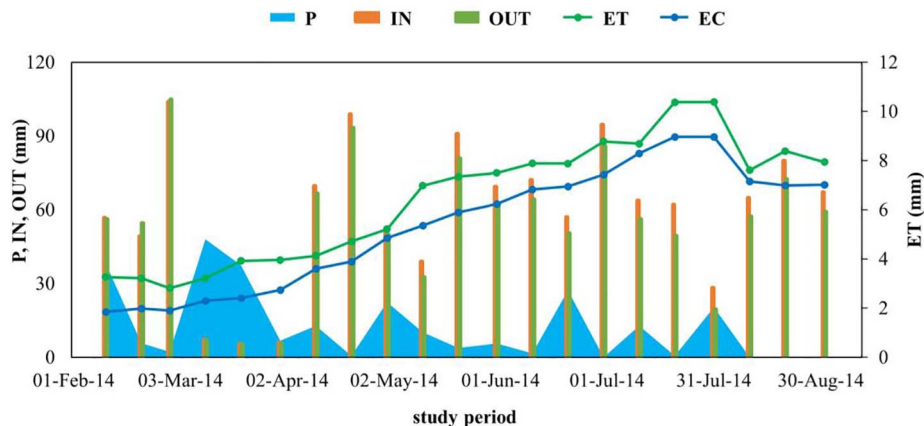


Fig. 10. Decadal mean values of precipitation (P, mm), IN (mm), OUT (mm), evapotranspiration as residual of water balance (ET, mm) and ET from eddy covariance (EC, mm) during February-August 2014.

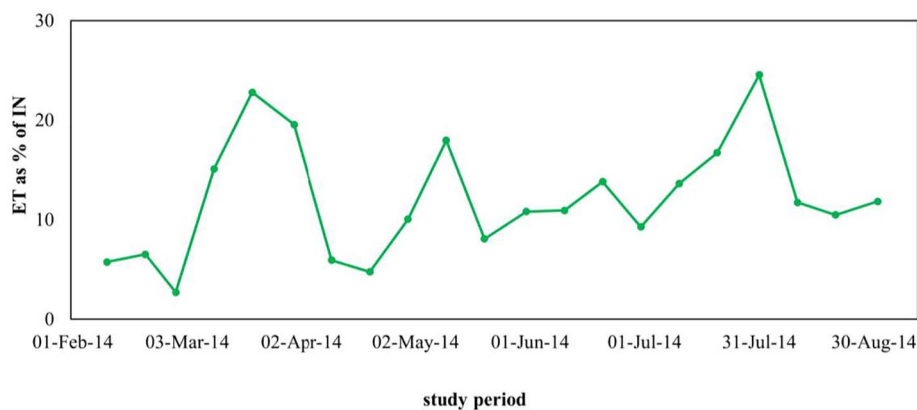


Fig. 11. Mean 10-days mean percentage of influent wastewater IN (mm) that was lost via ET from the TW system.

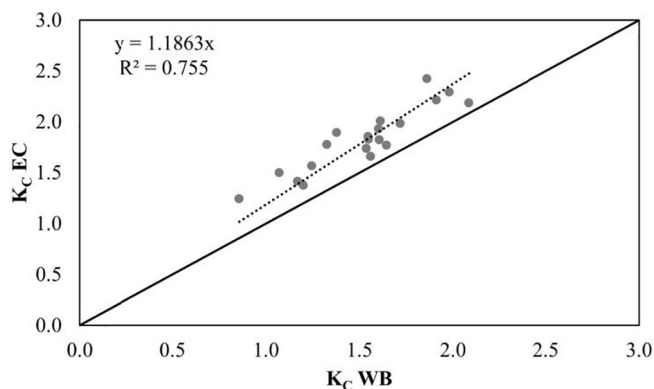


Fig. 12. 10-day average K_c from water balance (WB) and eddy covariance (EC) approaches.

- average loss of available energy for ET is of about 65% during the study periods;
- surface renewal is a valid alternative to EC for estimating sensible heat fluxes;
- evapotranspiration from TW water balance is overestimated by about 15% compared to EC.

Finally, further studies will imply the analysis of the most suitable aquatic vegetation that would minimize ET losses from TW systems used for agricultural reuse purposes.

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