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Ocean acidification causes variable trait shifts in a coral species

2

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27

For Review Only

28 Abstract

29 High $p\text{CO}_2$ habitats and their populations provide an unparalleled opportunity to assess how
30 species may survive under future ocean acidification conditions, and help to reveal the traits that
31 confer tolerance. Here we utilize a unique CO_2 vent system to study the effects of exposure to
32 elevated $p\text{CO}_2$ on trait-shifts observed throughout natural populations of *Astroides calycularis*, an
33 azooxanthellate scleractinian coral endemic to the Mediterranean. Unexpected shifts in skeletal
34 and growth patterns were found. Colonies shifted to a skeletal phenotype characterized by
35 encrusting morphology, smaller size, reduced coenosarc tissue, fewer polyps, and less porous and
36 denser skeletons at low pH. Interestingly, while individual polyps calcified more and extended
37 faster at low pH, whole colonies found at low pH site calcified and extended their skeleton at the
38 same rate as did those at ambient pH sites. Transcriptomic data revealed strong genetic
39 differentiation among local populations of this warm water species whose distribution range is
40 currently expanding northward. We found excess differentiation in the CO_2 vent population for
41 genes central to calcification, including genes for calcium management (calmodulin, calcium-
42 binding proteins), pH regulation (V-type proton ATPase), and inorganic carbon regulation
43 (carbonic anhydrase). Combined, our results demonstrate how coral populations can persist in high
44 $p\text{CO}_2$ environments, making this system a powerful candidate for investigating acclimatization and
45 local adaptation of organisms to global environmental change.

46

47

48 1. INTRODUCTION

49 Understanding the effects of environmental variability and extremes on natural populations and
50 ecosystems is a key priority as global environmental change intensifies (Bennett, Duarte, Marba,
51 & Wernberg, 2019; Bozinovic, Calosi, & Spicer, 2011). High local variability in physical and
52 chemical ocean properties can create extreme climatic environments, where marine species persist
53 under suboptimal environmental conditions such as highly variable temperatures, marginal
54 habitats at latitudinal extremes, and acidification at CO₂ vent sites (Camp et al., 2018; Kapsenberg
55 & Cyronak, 2019; Kroeker et al., 2019). Populations living in these unique settings experience
56 high environmental variability and can have broad physiological tolerance to environmental
57 stressors that would prevent survival of conspecifics living in less variable micro-environments
58 (Bozinovic et al., 2011; Thomas et al., 2018). Two important mechanisms for intraspecific
59 variation in tolerance to environmental variability and extremes are adjusting life traits through
60 phenotypic plasticity and local adaptation, and these processes may interact synergistically
61 (Hoffmann & Sgro, 2011; Savolainen, Lascoux, & Merilä, 2013). Phenotypic plasticity (also
62 referred to as acclimatization) is the ability of a genotype to produce different morphological and
63 physiological responses when exposed to different environmental conditions within an organism's
64 lifespan, resulting in a phenotypic shift that is plastic and often reversible (Savolainen et al., 2013;
65 Thomas et al., 2018). Adaptation is the result of natural selection on beneficial genotypes in a
66 population where these changes are heritable and passed on to the next generation (Hoffmann &
67 Sgro, 2011; Savolainen et al., 2013). Natural extreme environments are potential locations for
68 climate-adapted populations where, for example, microhabitats experiencing periodic temperature
69 extremes have shown to generate high-tolerance in some reef-building corals (Palumbi, Barshis,
70 Traylor-Knowles, & Bay, 2014; Thomas et al., 2018). However, there is still much to learn about

71 the underlying mechanisms of acclimatization and adaptation to climate variability and extremes
72 by studying populations in naturally variable environments. Such studies are critical for predicting
73 future biological responses to rapid global environmental change.

74 Insights into species' tolerance to environmental change may be gained by analyzing traits that
75 directly influence an organism's performance (Mouillot, Graham, Villéger, Mason, & Bellwood,
76 2013). Shifts in the occurrence of these traits under variable environmental conditions can reflect
77 patterns of differential survival and growth strategies; for example, different morphological forms
78 (e.g. massive or encrusting), longevity, size, growth rates, physical defenses and dispersal ability
79 (Darling, Alvarez-Filip, Oliver, Mcclanahan, & Côté, 2012; Teixidó et al., 2018). These traits
80 provide relevant information about life strategies that are the result of different evolutionary and
81 ecological processes and influence, both the fitness of individuals and the viability of natural
82 populations (Darling et al., 2012; Mouillot et al., 2013; Teixidó et al., 2018). However, we still
83 know comparatively little about trait-shifts within natural populations and the capacity to adapt to
84 long-term novel environmental conditions.

85 Natural volcanic CO₂ vents cause local acidification of seawater and are used as a proxy to study
86 future ocean acidification (Enochs et al., 2015; Fabricius et al., 2011; Hall-Spencer et al., 2008).
87 Ocean acidification reflects a suite of changes in seawater carbonate chemistry due to the uptake
88 of excess anthropogenic CO₂ by the ocean, resulting in a decline in the surface ocean pH, carbonate
89 ion concentration, and saturation state of calcium carbonate minerals (e.g. aragonite), while
90 increasing the partial pressure of carbon dioxide ($p\text{CO}_2$) and bicarbonate ion concentrations
91 (Doney, Fabry, Feely, & Kleypas, 2009). Low pH levels in natural CO₂ vents represent future
92 climatic conditions where, relative to 1870, surface pH is projected to decline by -0.14 to -0.4 pH
93 units by 2100, under IPCC Representative Concentration Pathways (RCP) RCP 2.6 (low CO₂

94 emissions) and RCP 8.5 (high CO₂ emissions) (Fabricius et al., 2011; Gattuso et al., 2015;
95 Goffredo et al., 2014; Teixidó et al., 2018). Although these pH conditions can provide some insight
96 into future acidification scenarios, they are not perfect proxies. One important assumption to
97 consider is that variability of seawater pH increases with decreasing means at CO₂ vent systems.
98 Although variability in pH/*p*CO₂ will increase with dissolved inorganic carbon due to the
99 thermodynamics of the carbonate system in the future ocean (Takeshita et al., 2015), it is not
100 possible to disentangle the effects of changes in the mean versus variability in this system. Thus,
101 the conditions in the pH zones should be considered as pH regimes, with decreases in mean pH
102 coinciding with increases in variability. Nevertheless, these high *p*CO₂ environments and their
103 populations provide an unparalleled opportunity to assess how species may survive into future pH
104 conditions and to reveal if general traits that confer tolerance can be identified.

105 Corals are key marine organisms that are particularly vulnerable to the impacts of climate change
106 and ocean acidification (Brandl et al., 2019; Gattuso et al., 2015). They create habitats for many
107 species, enhancing biodiversity, playing fundamental ecological roles and sustaining ecosystem
108 processes and services such as fisheries, coastal protection and tourism (Brandl et al., 2019;
109 Gattuso et al., 2015). Ocean acidification may pose a major threat to corals because their growth
110 relies on the precipitation of calcium carbonate (calcification), a process that is expected to
111 decrease as seawater acidity increases (Chan & Connolly, 2013). Studies conducted at CO₂ vent
112 ecosystems on native corals have reported an overall decline in species abundances, decreases in
113 calcification and skeletal density with increasing acidification (Fabricius et al., 2011; Fantazzini
114 et al., 2015; Goffredo et al., 2014).

115 Here we utilize a unique CO₂ vent system to investigate the effects of exposure to elevate *p*CO₂
116 on trait-shifts on *Astroides calycularis*, an azooxanthellate scleractinian coral endemic to the

117 Mediterranean, that naturally occurs in the acidified environment of a newly discovered CO₂ vent
118 system in Ischia, Italy. This CO₂ vent system locally acidifies the seawater with gas comprising
119 92-95% CO₂ (no sulphur, and no temperature anomaly). *A. calycularis* is a long-lived coral (large
120 colonies may have a life span of several decades), considered a warm-water species with a narrow
121 temperature tolerance confined to 14°C during winter (Bianchi, 2007; Zibrowius, 1995). *A.*
122 *calycularis* has low dispersal capacities, and therefore restricted gene flow (Casado-Amezúa,
123 Goffredo, Templado, & Machordom, 2012). Because *A. calycularis* is a calcifying and long-lived
124 species with low dispersal capacity, and found throughout the CO₂ vents, it is a great model system
125 for investigating variation in local climate phenotypic plasticity and adaptation. Previous research
126 on the effects of ocean acidification on *A. calycularis* has shown contrasting results: a reduction
127 of net calcification rates was found when colonies growing in ambient pH conditions were
128 transplanted to a vent system with pH below 7.7 (Prada et al., 2017), while no change in
129 calcification under acidification was observed in controlled laboratory conditions (Movilla et al.,
130 2016). We compare populations living near the vent to two reference areas outside the influence
131 of CO₂ venting to examine the effects of low pH conditions on *A. calycularis* traits, to characterize
132 the genetic population structure, and to identify differentiation in genes that are central to
133 calcification. Specifically, we addressed the following questions: i) do populations at the CO₂ vent
134 and reference sites exhibit significant trait variation?, ii) do these nearby populations display
135 genetic differentiation?, and iii) does the CO₂ vent population have highly divergent SNP
136 genotypes from calcium- related loci? To answer these questions, we characterized the physical
137 and chemical parameters of the study sites and combined *in situ* population demographics, skeletal
138 characteristics, computed tomography and transcriptomic approaches to assess changes in
139 population abundance, skeletal properties, age, and genomics of differentiation of *A. calycularis*.

140 2. MATERIAL AND METHODS

141 2.1 Experimental design and study sites

142 Here we compare natural populations of the scleractinian coral *A. calycularis* at a volcanic CO₂
143 vent and two nearby reference sites with ambient pH and no vent activity along the coast of Ischia,
144 Italy (Figure 1). The CO₂ vent system is located at a 5 m depth inside a semi-submerged cave of
145 volcanic origin named *Grotta del Mago* (Magician's Cave, 40°42'41.87"N, 13°57'51.06"E,
146 hereafter Vent system) (Figure 1). The cave (total length of 110 m) consists of a main outer
147 chamber (10 m wide x 30 m long), connected to an inner chamber by a long narrow passage
148 (Cinelli et al., 1977). Published studies and personal observations indicated an increase in the CO₂
149 vent activity over the last 50 years in the main chamber, with limited vent activity in the 70's
150 (Cinelli et al., 1977) and 2000's (Dappiano & Gambi, 2004) developing into intense activity from
151 2014 onwards. The abundance of *A. calycularis* in the cave has increased over time, with a low
152 and patchy distribution between 1-2 m depth in the main chamber in the 1970's (Cinelli et al.,
153 1977) to a high and continuous distribution in the 2000's (Dappiano & Gambi, 2004). The present
154 study was performed in the main chamber of the cave. The reference sites with ambient pH were
155 chosen following the criteria: i) *A. calycularis* naturally occurred there, ii) they hosted similar
156 habitats and depths as the CO₂ vent site, and iii) and no venting activity was evident. Two reference
157 sites were selected: Punta Vico (40°45'32.28"N, 13°52'55.38"E, another semi-submerged cave,
158 with a main chamber 10 m wide x 30 m long, 5 m maximum depth, hereafter Ambient 1); and
159 Sant'Angelo (40°41'33.78"N, 13°53'38.88"E, an overhang, also a natural habitat of *A. calycularis*,
160 located on a natural arch, with an opening of 10 m wide x 10 m height, 10 m maximum depth,
161 hereafter Ambient 2). Initial investigations of the natural systems and environmental parameters
162 started in June, 2016. These preliminary environmental data were used to plan subsequent field

163 samplings of the carbonate chemistry associated with the CO₂ vent system and reference sites in
164 September, 2018 and June, 2019.

165 **2.2 The coral**

166 *A. calycularis* (Pallas, 1766) is an azooxanthellate scleractinian colonial coral endemic to the
167 Mediterranean, characterized by the bright orange color of its coenosarc and polyps (Zibrowius,
168 1995). It is considered a long-lived species (*e.g* large colonies may have a life span of several
169 decades) and commonly found in dimly lit, shallow rocky habitats (vertical walls, cave entrances,
170 overhangs, from the intertidal fringe to 50 m depths) (Zibrowius, 1995). It can be highly abundant
171 covering more than 90% of local reefs. It has a limited geographic distribution, with a southwestern
172 distribution in the Mediterranean Sea (Zibrowius, 1995). This coral is considered a warm-water
173 species with a narrow temperature tolerance confined to 14°C during the winter (Bianchi, 2007).
174 Fossil records reveal this species lived in the northwestern Mediterranean during part of the
175 Pleistocene, where climatic fluctuations occurred leading to a reduction of the species (Zibrowius,
176 1995). Interestingly, observed records north of its known distribution range in Italy and Croatia
177 suggest that it is currently expanding northward (Bianchi, 2007). Currently, *A. calycularis* is
178 assessed as vulnerable on the IUCN Red List due to its limited geographic distribution and the
179 historical and current regression caused by human activities in the littoral zone. *A. calycularis*
180 broods its larvae and has relatively low dispersal capacity (Casado-Amezúa et al., 2012).

181 **2.3 Gas and temperature**

182 Gas samples were collected in 200 ml glass bottles and analyzed using gas chromatography
183 (Agilent 7890B combined with a Micro GC analyzer-INFICON, held at a constant temperature of
184 80 °C). The mean composition of the bubbling gas was predominantly CO₂ (92-95%, with
185 undetectable levels of sulfur gas <0.0002 %) and did not elevate the temperature (see Supporting

186 Note 1, Figure S1), subsequently resulting in water acidification. Vent activity was sampled by
187 counting the number of vents in randomly placed 1 m² quadrats (n= 11) in the main chamber, with
188 approximately 5 vents m⁻² (mean \pm SE= 4.9 \pm 2.7 vents m⁻², min= 2 vents m⁻²; maximum = 11
189 vents m⁻²). Temperature was registered every hour by *in situ* temperature data loggers (Hobo
190 TidbiT v2, Onset) in the cave and the reference areas and followed ambient seasonal fluctuations,
191 from 14.7 to 15.2 °C in winter (n=16,754), and from 25.5 to 26.5 °C in summer (n= 19,011) over
192 a 3-year period from 2016 to 2019 at 2 m depth (Figure S1, Table S1).

193 **2.4 pH_T time series, pH_T variability and pH sensor calibration**

194 SeaFET™ Ocean pH sensors (Satlantic) were deployed to quantify variation in pH inside the cave
195 at 2, 3 and 4 m depth. They were deployed in May-June (before summer) and in September (after
196 summer) to assess whether differences in water temperature stratification could influence pH
197 across depths. Dates of deployment were from September 8 to September 24, 2018 and from May
198 30 to June 18, 2019. Two sensors were deployed in the reference areas during the same period
199 (Ambient 1 in September, 2018 and Ambient 2 in June, 2019). Before deployment, the SeaFETs
200 were calibrated with ambient pH water in the aquarium facilities at the Center Villa Dohrn (Ischia,
201 Italy) (for full details of pH sensor calibration, see Supporting Methods). The mean offset between
202 calibration samples and calibrated SeaFET pH was \pm 0.002 units, indicating high quality pH
203 dataset (Figure S2).

204 **2.5 Carbonate Chemistry and Nutrients**

205 Discrete water samples were collected using Niskin bottles at the vent and reference areas with
206 ambient pH to measure: i) the carbonate system parameters during the pH sensor deployment, and
207 ii) dissolved inorganic nutrients. Salinity was measured using a CTD (CTD Sea Bird Electronics
208 SBE 19 Plus Seacat). Samples for total alkalinity (A_T) were collected using standard operating

209 protocols (for full details, see Supporting Methods). The HCl (0.1 M) titrant solution was
210 calibrated against certified reference materials distributed by A.G. Dickson (CRM, Batches #153,
211 #171, and #177). Precision of the A_T measurements of CRMs was $< 2.0 \mu\text{mol kg}^{-1}$ from nominal
212 values. Means were reported as (mean \pm SD): $A_T = 2562.41 \pm 7.8 \mu\text{mol kg}^{-1}$, $n = 27$ in September
213 2018; and $A_T = 2543.57 \pm 21.78 \mu\text{mol kg}^{-1}$, $n = 21$ in June 2019. A_T and pH_T were used to determine
214 the remaining carbonate system parameters at *in situ* temperature and depth of each sampling
215 period in the R package seacarb v3.2.12 (for constant details, see Supporting Methods). Dissolved
216 inorganic nutrients (nitrite NO_2 , nitrate NO_3 , ammonium NH_4^+ , phosphate PO_4 and silicate
217 $\text{Si}(\text{OH})_4$) were determined using a colorimetric method (Supporting Methods) (Table S2).

218 **2.6 Coral field surveys: cover, population structure and morphology**

219 The *A. calycularis* cover was quantified using 24 photoquadrats (25 x 25 cm) positioned along six
220 transects at four depths: 1, 2, 3, and 4 m in the three study sites (Vent, A1, A2). Percentage cover
221 was analyzed with ImageJ image software (National Institutes of Health,
222 <https://imagej.net/ImageJ>). Size frequency-distribution was calculated at 1 and 3 m depths by
223 counting the number of polyps of each colony and each colony was then pooled into one of five
224 size classes (I: 1-5 polyps; II: 6-10 polyps; III: 11-15 polyps; IV: 16-20 polyps; V: > 20 polyps).
225 These size classes were selected to span the range of colony sizes encountered in the field. We
226 also assessed necrosis as the percentage of the colony exhibiting dead tissue or denuded skeleton,
227 from white-grey to unpigmented or denuded skeleton. Finally, visual assessments were used to
228 classify the colonies into two morphological categories: encrusting (flat growth form) and massive
229 (extensive vertical and lateral growth). Encrusting colonies extended laterally over the surrounding
230 substrate, whereas massive presented a greater vertical accretion which resulted in semi-spherical

231 shapes. This categorical criterion allowed us to obtain two simple morphological variables to
232 capture biologically relevant axes of variation.

233 **2.7 Sample collection for presence of coenosarc, skeletal characteristics and growth**

234 Sixty-six colonies of *A. calycularis* were sampled haphazardly for coenosarc, biometric, growth,
235 and skeletal parameters. Thirty-four colonies were collected at the vent site: 16 colonies were
236 obtained from the vicinity of the CO₂ vents at 3 m depth (vent system deep, Vd) and 18 colonies
237 at 1- 2m depth (vent system shallow, Vs). Thirty-two colonies were collected from areas with
238 ambient pH conditions: 17 colonies in Punta Vico, Ambient 1, 1 – 2 m depth; and 15 colonies in
239 Sant'Angelo, Ambient 2, 1 -2 m depth. The 66 colonies were photographed and the percentage of
240 coenosarc (*i.e.*, the living tissue connecting the polyps) was determined. The % of coenosarc was
241 determined from the edges of the polyp tissue. The % of coenosarc was classified into ten classes
242 at every 10% interval, from 100% to 0%. Loss of coenosarc in *A. calycularis* may occur mainly
243 by two mechanisms: 1) loss of tissue due to necrosis (when colony exhibits dead tissue, from
244 white-grey to unpigmented or denuded skeleton), or 2) the coenosarc is already absent due to
245 physiological and morphological characteristic of the colonies.

246 **2.7.1 Biometric parameters**

247 Coral skeletons were rinsed in a solution of 10% commercial chlorine bleach for 3-4 days to
248 dissolve polyp tissue, then they were dried at room temperature for 3 days. Colony was defined as
249 the whole calcareous skeleton, which included the polyps (corallites) and the coenosteum. The
250 following parameters were measured: colony length (L_c , major axis of the colony) and colony
251 width (W_c , minor axis of the colony); number of polyps in each colony (NP_c), corallite length (L_p ,
252 maximum axis of the oral disc) and corallite width (W_p , minimum axis of the oral disc) (for full
253 details, see Supporting Methods, Figures S3-S4, and Tables S3-S4).

254 **2.7.2 Growth and age estimations**

255 The age of each corallite skeleton was determined by counting the growth bands of 49-70 randomly
256 selected corallites per site, by means of computerized tomography (CT). Growth bands are
257 distinguished by a high-density band in winter and a low-density band in summer in temperate
258 corals (see Supporting Methods). The age of all collected corallites was estimated using the von
259 Bertalanffy's length-age growth function derived from the CT growth bands analysis. Coral
260 growth is described by three parameters: linear extension rate (linear growth), net calcification rate
261 (net mass deposited) and bulk skeletal density (mass per volume unit) (Goffredo et al., 2009). The
262 measurement of all three components is fundamental when assessing the effect of the environment
263 on coral growth, since none of the three parameters is a perfect predictor for the other two and
264 each species can respond differently to environmental conditions. Then, the following three coral
265 growth parameters were calculated for both polyp and colony levels: 1) linear extension rate; 2)
266 net calcification rate and 3) bulk skeletal density (see below for bulk skeletal density
267 measurements) (for full details, see Supporting Methods).

268 **2.7.3 Skeletal parameters**

269 Skeletal parameters of colonies were calculated by applying the buoyant weight technique through
270 the density determination kit of the Ohaus Explorer Pro balance (± 0.0001 g; for further details,
271 see Supporting Methods). This method is based on the Archimedes principle applied to a specimen
272 after full saturation with the same fluid in which it was submerged. The measurements required to
273 calculate the skeletal parameters were: density of the fluid medium (ρ); dry mass (DW), buoyant
274 weight of the skeleton (BW= weight of the skeleton minus weight of the water displaced by it),
275 SW (saturated weight of the coral = weight of the skeleton plus weight of the water enclosed in
276 its). Measurements were repeated three times to get an average for BW and SW. Based on these

277 measurements, the following parameters were calculated: V_{MATRIX} (matrix volume = volume of
278 the skeleton, excluding the volume of its pores); V_{PORES} (pore volume=volume of the pores in the
279 skeleton), V_{TOT} (bulk volume = total volume of the skeleton including its pores). Finally, the
280 skeletal parameters of colonies were calculated: the micro-density (ratio of DW to V_{MATRIX}); the
281 bulk density (ratio DW to V_{TOT}); and the porosity (ratio V_{PORES} to V_{TOT}).

282 **2.8 Sample collection for genetics and transcriptome assembly**

283 *A. calycularis* colonies were haphazardly collected between 1- 2 m depth to compare gene flow
284 among populations at the study sites. Nineteen colonies from the Vent, 14 colonies from Ambient
285 1 and 8 colonies from Ambient 2 were sampled for genetic analysis. RNA was extracted from a
286 single polyp of each colony using a RNeasy kit (Qiagen Inc., Valencia, CA, USA) according to
287 the manufacturer's instructions (for full details, see Supporting Methods). Approximately 1 μ g of
288 RNA was used to construct a cDNA library for each sample using the Illumina TruSeq RNA v2
289 Kit (Illumina, San Diego, CA, USA) (see Supporting Methods).

290 Libraries were sent to the Genomics Core Facility of the Health Sciences Cores at the University
291 of Utah (Salt Lake City, Utah, USA) and samples were quantified using a Bioanalyzer (Agilent,
292 Santa Clara, CA, USA) (see Supporting Methods). We assembled the first *de novo* transcriptome
293 of *A. calycularis* with samples collected from Ischia. The following programs and scripts were run
294 on Stanford University's Sherlock cluster and all scripts used in this pipeline
295 (https://github.com/bethsheets/Astroides_transcriptomics) along with a general guide are available
296 on GitHub (doi:10.5281/zenodo.2580291). Four population-specific *de novo* assemblies were
297 generated using three individuals per population for each population in the program Trinity-2.8.4
298 (see Supporting Methods). Prior trials mapping to available corallimorpharian genomes produced
299 incomplete assemblies. Therefore, assembled contigs were validated to be included in the

300 assembly by filtering for only metazoan matches found in the combined UniProt's Swiss-Prot and
301 TrEMBL databases using BLASTX in the BLAST+ toolkit (see Supporting Methods). Matches
302 were considered significant at values of $\leq 1 \times 10^{-3}$ and the top hit for each contig was kept for
303 assembly filtering and annotation. Transcriptome completeness was assessed using BUSCO v3
304 (see Supporting Methods) against the metazoan (obd9) set. BUSCO analyses revealed that the final
305 combined transcriptome was 97% complete (949 complete BUSCOs out of 978 searched: 366
306 complete single-copy, 583 complete and duplicated; 12 fragmented, and 17 missing). For the 41
307 individuals used for population analyses, the average overall mapping rates for each population
308 were as follows: Vent-Grotta Mago 79% (range 71.91 - 85.28), Ambient 1- Punta Vico 80% (range
309 65 – 84.74), Ambient 2- Sant' Angelo- 80.26% (range 71.91 - 85.28). After filtering, we detected
310 46,784 biallelic SNPs among the vent and two ambient populations.

311 **2.8.1 Mapping and SNP detection**

312 For all 41 samples, raw paired- and single-end sequence files were mapped to the *de novo* assembly
313 using HISAT2 (see Supporting Methods) with the very-sensitive setting. Duplicate reads due to
314 PCR were removed with Picard tools (<http://broadinstitute.github.io/picard/>) MarkDuplicates
315 using the lenient validation stringency. Overall mapping rates were compared among populations
316 to assess whether certain populations were preferentially mapping to the *de novo* assembly.
317 Transcriptome-derived single nucleotide polymorphisms (SNPs) were called on all individuals
318 using SAMtools mpileup and BCFtools (for filter settings, see Supporting Methods).

319 **2.8.2 Identifying SNP candidates for environmental selection in high CO₂ conditions and** 320 **enrichment analysis**

321 We identified potential outlier SNPs related to the CO₂ vent location. We calculated pairwise Fst
322 estimates (Vent - Ambient 1, Vent - Ambient 2, Ambient 1 - Ambient 2) per locus using the basic

323 stats function with HIERFSTAT package in R (see Statistical analyses). We used these
324 estimates to compare the genetic distance for each SNP between the three populations
325 $[F_{ST}(\text{Vent-A1}) + F_{ST}(\text{Vent-A2})]/F_{ST}(\text{A1-A2})]$. To identify potential outlier SNPs related to
326 the CO₂ vent location, we compared the genetic distance for each SNP for the population
327 comparisons including Grotta Mago to the genetic distance between the ambient
328 populations A1 and A2 $[F_{ST}(\text{Vent-A1}) + F_{ST}(\text{Vent-A2})]/F_{ST}(\text{A1-A2})]$. SNPs with values over
329 2 showed an excess of genetic differentiation in the CO₂ Vent compared to the other
330 ambient pH populations. Using the transcriptome assembly annotations, we searched for
331 enrichment patterns in the contigs holding these candidate SNPs by using their UniProt
332 identifiers (<https://www.uniprot.org/>) in a Gene Ontology (GO) search
333 (<http://geneontology.org/>).

334 **2.9 Statistical analyses and data visualization**

335 *Environmental data analyses:* Temperature, pH_T, SeaFET sensor calibration, carbonate chemistry
336 and figures were performed using the R packages: seacarb v3.2.12 and ggplot2 v3.1.1 (see
337 Supporting Methods for R package references). Carbonate system parameter figures in the vent
338 system were created with Ocean Data View software (version 5.1.2, <http://odv.awi.de>). *Biological*
339 *surveys:* A linear mixed model was used to test for differences in % cover (logit transformation)
340 as a function of site (fixed factor, 3 levels), depth (fixed factor, 4 levels) and quadrat as a random

341 effect. Chi-square contingency tables were used to compare the size–frequency distributions
342 among sites, as well as the frequency of encrusting and massive forms. Kolmogorov-Smirnov two-
343 sample tests were used to determine whether there were significant differences in necrosis between
344 the CO₂ vent and ambient pH sites. These analyses were computed using lme4 v1.1.21 package
345 implemented in R. *Skeletal characteristics and growth*: Relationships between biometrical and
346 skeletal parameters were calculated using the power function model. Pearson’s correlation
347 coefficients were calculated for the relationships among biometric and skeletal parameters at both
348 colony and polyp levels. Spearman’s rank correlation coefficient was used to calculate the
349 significance of the correlations between colony biometric and skeletal parameters and pH_T.
350 ANOVA was used to test % cenosarc (with arcsin transformation), mean mass, polyp number,
351 bulk density, linear extension rate, calcification rate, and porosity of the colonies and mean length
352 of corallites among sites. We used the non-parametric Kruskal-Wallis test for differences in means
353 for data that did not meet the assumption for normality and equal variance. Kruskal-Wallis tests
354 were applied to mean area, length, width, micro-density of colonies, corallite mean age, polyp
355 linear extension rate, net calcification rate, and length of central and all corallites. These analyses
356 were computed using IBM SPSS Statistics 12.0 (IBM Corporation). The Von Bertalanffy growth
357 model and confidence intervals (CI) were estimated through a regression analysis by least squares
358 procedure using raw data of corallite length and age (measured by computerized tomography) (see
359 Supporting Methods). These analyses were carried out in the software MATLAB R2012a
360 (MathWorks, Natick, USA). *Population genetics*: Population genetic analyses of SNPs, Weir and
361 Cockerham’s pairwise F_{ST} estimates among populations, and the heatmap of divergent SNP
362 genotypes were conducted in the R-based program HIERFSTAT v4.22 and ComplexHeatmap
363 v2.4.2, respectively.

364 3. RESULTS

365 3.1 Carbonate Chemistry Associated with the CO₂ vent system in the *Grotta del Mago*' cave

366 The CO₂ vent system occurs at a 5 m depth inside the main chamber of the cave *Grotta del Mago*
367 with approximately 5 vents m⁻² (Figure 1) and do not elevate temperature (Table1, Figure S1,
368 Table S1). The carbonate chemistry derived from discrete water samples and *in situ* monitoring of
369 seawater pH_T (pH on the total scale) delineated a pH_T gradient from 4 m to 2 m depth caused by
370 the distance from the bubbling of CO₂ gas from the seafloor (92-95% CO₂, with undetectable levels
371 of sulfur gas < 0.0002 %, see Supporting Note 1) (Figure 1, Table 1). pH sensors revealed
372 reductions in mean pH_T at each depth associated with increased temporal variability in pH_T (Figure
373 1, Table 1, Figure S5, Table S5). Mean pH_T were: 7.65 to 7.88 at 2 m; 7.62 to 7.74 at 3 m, and
374 7.60 to 7.60 at 4 m, respectively (see Table 1 and Table S5 for detailed pH statistics). At 2 m depth,
375 14% and 56% of the pH_T measurements were below 7.8 (projected average global sea surface pH
376 value for the year 2100 with the high emission scenarios RCP8.5) (Gattuso et al., 2015) in June
377 and September, respectively (Table S5). The percentage rose to 34% and 61% at 3 m depth, and
378 55% and 66% at 4 m depth, in June and September, respectively (Table S5). This pattern of depth-
379 dependent low pH_T was also manifested as extreme pH events (defined as the pH value of 0.4 units
380 less than the mean pH for each depth) that increased in number and duration with depth (Table
381 S6). Mean pH_T and variability were influenced by temperature stratification in June and September
382 (Figure 1, Figures S5-S6). This is because during periods of stratification, and hence reduced
383 vertical mixing (Turner, 1973), the input of CO₂ is likely to be primarily confined to the lower part
384 of the water column, leading to lower pH values near bed than when the water column is well
385 mixed. The mean temperature difference between 2 m and 4 m in June was 0.25 °C, whereas the
386 mean temperature difference was only 0.02 °C in September (Figure S6). In September, reductions

387 in seawater pH_T were driven by increased dissolved inorganic carbon concentrations (C_T) and
388 higher pCO_2 concentrations at relatively constant total alkalinities (A_T) and temperatures across
389 depths (Table 1, Figure S5). Mean pCO_2 ranged from $2905 \pm 1664 \mu\text{atm}$ at 2 m, to 3146 ± 1928
390 μatm at 3 m, to $3192 \pm 1806 \mu\text{atm}$ at 4m depth, and aragonite saturation state (Ω_a) ranged from
391 1.10 ± 0.4 at 2m, to 1.05 ± 0.4 at 3 m, and to 1.02 ± 0.4 at 4m depth (Table 1). In contrast, in June,
392 the vent site was characterized by an increase in temperature along the water column (from ~ 18.5
393 $^{\circ}\text{C}$ to $\sim 25^{\circ}\text{C}$), which created greater difference across the three depths in terms of pH_T and
394 associated carbonate chemistry parameters, particularly for the pCO_2 concentrations (from $1531 \pm$
395 $627 \mu\text{atm}$ at 2 m, to $2082 \pm 1502 \mu\text{atm}$ at 3 m, and $2812 \pm 2310 \mu\text{atm}$ at 4 m depth) and Ω_a (from
396 1.44 ± 0.27 at 2 m, to 1.23 ± 0.35 at 3 m, to 1.05 ± 0.42 at 4 m depth) (Table 1, Figure S5). At the
397 two ambient pH sites, the mean pH_T ranged from 7.97 to 8.05 units, pCO_2 from 322 ± 34 to $586 \pm$
398 $76 \mu\text{atm}$, and Ω_a from 3.54 ± 0.23 to 3.86 ± 0.23 (Table 1).

399 **3.2 Cover, population structure, and morphology of *A. calycularis***

400 The CO_2 vent population was characterized by small colonies (90% colonies had up to 10 polyps)
401 and no large colonies of more than 20 polyps (class V) with massive morphology were found
402 (Figure 2, Figure S7). Larger colonies (*i.e.* > 16 polyps, size classes IV and V) were only found in
403 the two reference areas (percentage of larger size classes ranged from 13% in A1 to 16 % in A2),
404 which differed significantly among the CO_2 vent site ($\chi^2 = 91.9$, $df = 8$, $p < 0.0001$). Additionally,
405 necrosis was significantly higher at the CO_2 vent site ($13 \pm 4\%$) than the reference areas (both
406 $< 0.5\%$, $D = 0.56$, $p < 0.0001$). *A. calycularis*' cover at the CO_2 vent site decreased from 50% at 1 m
407 depth, to 30% at 2m, 9% at 3 m, and 1% at 4m depth, as seawater pH_T also declined (Figure S7).
408 This decline in coral cover was also observed in Ambient 1 (also a cave, from 69% at 1 m to 14%

409 at 3 m) but not in Ambient 2 (overhang, from 72% at 1 m to 62% at 3 m; $F= 14.1$, $df=11,55$;
410 $p<0.001$) (Figure S7).

411 **3.3 Coenosarc, skeletal parameters and growth**

412 Percentage of coenosarc (the living tissue connecting the polyps) significantly decreased between
413 ambient pH sites (87 and 70%) and the CO_2 vent (28 and 14%) (Figure 2, $p<0.0001$). Mean colony
414 area decreased by $\sim 80 - 71\%$ and mean polyp number by $\sim 27 - 18\%$ at the Vent deep compared
415 to ambient pH sites ($p < 0.001$) (Figure 3A). The skeletal parameters that characterize the
416 architecture of colonies showed different patterns in relation to pH (Figure 3A-B, Figures S8-S9).
417 Bulk density (ratio dry mass to bulk volume) and micro-density (ratio of dry mass to matrix
418 volume) increased at low pH, while porosity (ratio pore volume to bulk volume; see Methods)
419 decreased at low pH ($p < 0.001$; Figure 3A-B). Colonies at the CO_2 vent deep presented higher
420 bulk density ($\sim + 27 \%$) and micro-density ($\sim + 7 \%$) and lower porosity ($\sim -28 \%$) compared to
421 colonies from the ambient pH sites ($p < 0.001$) (Figure 3A-B, Figure S9).

422 Growth parameters of *A. calycularis* differed significantly between the CO_2 vent site and ambient
423 pH sites (Figure 3B, Figure 4, Figure S10, Table S7). Mean polyp growth rate decreased
424 exponentially with age at all sites (Figure S10). Young individuals (1–3 years old) grew relatively
425 rapidly ($> 2 \text{ mm yr}^{-1}$), but, as they aged, their skeletal growth rate decreased ($< 1.3 \text{ mm yr}^{-1}$ at 8–
426 10 years old) (Figure 4, Tables S7-S8). A trend toward higher linear extension and net calcification
427 rate was observed at low pH at the polyp level (Figure 3B, Table S7). Polyp net calcification rate
428 ranged from $3.95 \text{ mg mm}^{-2} \text{ yr}^{-1}$ at Vent deep, to $3.04 \text{ mg mm}^{-2} \text{ yr}^{-1}$ at Vent shallow, to 2.39 mg
429 $\text{mm}^{-2}\text{yr}^{-1}$ and $2.06 \text{ mg mm}^{-2} \text{ yr}^{-1}$ at ambient pH sites (Table S7). This indicates that net calcification
430 rates increased approximately by 48%-93% from the more acidified (Vent deep) to the less
431 acidified (Vent shallow) to the non-acidified (ambient pH) locations at polyp level. Linear

432 extension and net calcification rates at colony level were homogenous in ambient pH and acidified
433 conditions (Figure 3B, Table S7).

434 **3.4 Transcriptome assembly and population genomics**

435 The *A. calycularis* transcriptome composed of 12 colonies contained 113,351 contigs with an N50
436 of 2,285 (range 501 – 38,179). Based on 46,784 SNPs in 41 individuals, PCA analysis revealed
437 strong clustering by population (Figure 5). The vent population in Grotta Mago was most distinct
438 along PCA axis 1, but overlapped with Ambient1 (PV) along axis 2. Pairwise F_{ST} measurements
439 also support strong population structure within each of the three locations: CO₂ vent (Grotta Mago)
440 – Ambient 1 (Punta Vico) = 0.034, CO₂ vent (Grotta Mago) – Ambient 2 (Sant'Angelo) = 0.026,
441 Ambient 1 (Punta Vico) – Ambient 2 (Sant'Angelo) = 0.024. SNPs with values over 2 show an
442 excess of genetic differentiation in GM compared to the other populations. There were 334 loci
443 with an excess F_{ST} ratio of 10 or more, out of 2246 loci in our data set with known molecular
444 function. An analysis of the 402 unique molecular function Gene Ontology (GO) terms associated
445 with these loci showed there to be three significant enrichment classes (Table S9): calcium ion
446 binding (12 loci, $p_{adj} = 7 \times 10^{-23}$), catalytic activity (4 loci, $p_{adj} = 2 \times 10^{-9}$), and
447 oxidoreductase action (12 loci, $p_{adj} = 0.05$).

448 Calcium ion loci include calmodulin, calcineurin, calnexin, calcium-binding proteins in the
449 sarcoplasmic reticulum, and a set of two poorly characterized proteins with calcium binding motifs
450 (Contigs 3436 and 7780). These proteins are EF-hand calcium-binding protein and C-type lectin
451 calcium-binding in a hmmer database search (<https://www.ebi.ac.uk/Tools/hmmer/>). Because

452 calcium ion control is particularly central to calcification in scleractinians, we examined the 77
453 SNPs from the thirteen calcium-related loci for patterns across populations (Figure 6). As
454 expected, the Grotta Mago population had highly divergent SNP genotypes at these loci (average
455 minor allele frequency difference of 0.24), but these genes also showed a strong degree of
456 linkage among SNP genotypes within a single gene often across 1000s of base pairs (Figure 6).
457 Such multi-SNP haplotypes are rare in our data set yet occur in 8 of 10 high excess, calcium ion
458 loci with multiple SNPs.

459 Given the strong differences in calcium management suggested by excess differentiation
460 of calcium ion genes in Grotta Mago, we queried our transcriptome SNP data base for
461 other genes potentially involved in calcification. In corals, calcification occurs in the
462 calciblastic space through a combination of high calcium concentration and high pH (reviewed
463 in Drake et al., 2020). High pH is achieved through proton transport by specific calcium/proton
464 pumps, including the plasma membrane calcium ATPase (PMCA). There were no PMCA
465 polymorphisms in our data set, but the V-type ATPase proton pump (contig DN1551, SNPs 1701-
466 1805) showed six of seven SNPs with strong differentiation in Grotta Mago (excess 0.03 -28.8,
467 average 11.4, Figure 6). Minor allele frequencies differed by about a factor of 2 in Grotta Mago
468 compared to the other locations, and SNPs show strong linkage.

469 In addition to genes involved in calcium and pH regulation, we explored possible adaptation in
470 genes managing intracellular carbonate chemistry. Coral calcification depends on the delivery of
471 CO₂ to the calciblastic layer where it is converted to carbonate ions (Cohen & McConnaughey,

472 2003). In *A. calycularis*, a plethora of carbonate related genes showed high differentiation and high
473 linkage in Grotta Mago, including carboxylases and transporters, but the most interesting is a
474 carbonic anhydrase (Contig 15508, 5 of 12 SNPs with excess F_{st} above 10, average=10.3).
475 Carbonic anhydrases convert highly diffusive CO_2 into charged carbonate ions, localizing them in
476 the calicoblastic layer and regulating calcification (Chen, Gagnon, & Adkins, 2018), though they
477 also play a role in pH regulation of non-calcifying coral cells (Bertucci et al., 2013).

478 4. DISCUSSION

479 This study contributes to increasing our understanding of how coral populations persist under
480 naturally high pCO_2 environments – and therefore how they might cope under future ocean
481 acidification scenarios – and links trait-shifts with local variation in environmental parameters found
482 in this new CO_2 vent system. Our results expand upon previous research on populations of corals
483 exposed to naturally elevated pCO_2 (Crook, Cohen, Rebolledo-Vieyra, Hernandez, & Paytan, 2013;
484 Enochs et al., 2015; Fabricius et al., 2011; Fantazzini et al., 2015), demonstrating unexpected shifts
485 in patterns of skeleton and growth of the azooxanthellate coral *A. calycularis*. Specifically,
486 colonies shift to a skeletal phenotype characterized by encrusting morphology, smaller size,
487 reduced coenosarc tissue, fewer polyps, and less porous and denser skeletons at low pH. However,
488 while individual polyps calcified more (greater net calcification rates), calcification rate of whole
489 colonies were similar across sites. The resulting colony skeletons showed equal linear extension
490 at low and ambient pH conditions, while their polyp skeletons extended faster in acidified
491 conditions (Figure 7). Transcriptomic data revealed strong genetic differentiation among local
492 populations of this low-dispersal species. We found excess differentiation in the Grotta Mago
493 population for genes central to calcification, including genes for calcium management

494 (calmodulin, calcium-binding proteins), pH regulation (V-type proton ATPase), and carbonate
495 localization (carbonic anhydrase).

496

497 *Environmental variability in the CO₂ vent system*

498 The vent system exhibits high temporal variability in seawater pH due to varying CO₂ venting
499 intensity from the seabed, mixing due to variations in stratification, and fundamental
500 thermodynamics processes fundamental to the carbonate system (Takeshita et al., 2015). The
501 carbonate chemistry and *in situ* monitoring of seawater pH delineated a pH gradient (from 4 to 2
502 m depth) caused by the distance from the venting. This acidification gradient is important for the
503 colonies exposed to more (deep) or less (shallow) acidified conditions, as reflected by the
504 biological response of *Astroides*. The conditions in these zones are comparable with IPCC
505 projections for near future acidification scenarios (RCP2.6 and RCP8.5), which project a decrease
506 in surface pH from -0.14 to -0.4 pH units by 2100 relative to 1870 (Gattuso et al., 2015). pH and
507 its variability found in this study are comparable with the range of natural variation observed in
508 other CO₂ vent systems, with fluctuations of 0.6, 0.7 and 0.5 pH units in coral reefs (Agostini et
509 al., 2018), temperate reefs in Panarea (Prada et al., 2017) and in other CO₂ vents in Ischia (Teixidó
510 et al., 2018), respectively. Likewise, Hofmann et al. (2011) reported pH fluctuations between 0.024
511 to 1.430 pH units, in which pH measurements were taken from different locations, ranging from
512 polar to tropical, and open-ocean to coastal areas. Interestingly, as mean pH was reduced, its
513 variability and the percentage of pH_T measurements registered below 7.8 units increased, when
514 seawater was uniformly warm. In contrast, in June when the water column was stratified, pH and
515 its variability near the bottom was similar to what was observed in September whereas, farther

516 from the bottom, pH was higher and less variable. These results indicate that seawater stratification
517 may play a key role in controlling the temporal and depth patterns of pH/pCO₂.

518

519 *Shifts in coral skeleton, growth, and coenosarc*

520 Each scleractinian coral species may adopt different growth strategies in response to ocean
521 acidification. For example, investing calcification resources in bulk skeletal density by sacrificing
522 the rate of linear extension has been observed in *Orbicella annularis* (Carricart-Gavinet, 2007). In
523 contrast, investing calcification resources in linear extension rate at the expense of bulk density
524 has been reported for some *Porites* (Lough & Barnes, 2000) and Dendrophyllidae species
525 (Goffredo et al., 2009). Both strategies may imply different ecological trade-offs for the coral:
526 investing in a denser skeleton results in greater resistance to mechanical stress, while increasing
527 linear extension rate may be advantageous for space competition and earlier sexual maturity
528 (Goffredo et al., 2009). Unexpectedly, *A. calycularis* revealed unusual patterns in the calcification
529 response to ocean acidification, such as higher bulk skeletal density and lower porosity while
530 maintaining colony linear extension rates and net calcification. This response is different to what
531 was previously shown in solitary corals (e.g. *Balanophyllia europaea*) growing in natural low pH
532 conditions, where a decrease of net calcification resulted from preserving linear extension (to reach
533 the polyp size of sexual maturity) at the expenses of lower bulk density (e.g. increased in skeletal
534 porosity resulting in more fragile skeletons) (Fantazzini et al., 2015). Tambutté et al.(2015) found
535 the same response of decreasing calcification and bulk skeletal density while linear extension of
536 skeleton remained unchanged in the tropical coral *Stylophora pistillata* subjected to low pH
537 conditions in laboratory. Mollica et al. (2018) modelled the skeletal growth to changes in seawater
538 chemistry and predicted declines in *Porites* skeletal density but no linear extension across global

539 reefs, reflecting the large variability in the response of coral calcification to ocean acidification.
540 The authors suggested that under low-pH conditions, the increase in linear extension reflects the
541 elongation of calcium carbonate crystals at the site of calcification, while the increase in skeletal
542 density reflects the lateral thickening of calcified elements of coral skeleton (Mollica et al., 2018).
543 The unusual response of *A. calycularis* to acidification may reflect an overall maintenance of
544 energetic resources allocated to calcification at the level of the colonies, which extended at the
545 same rate but were composed by fewer polyps, thus partitioning the available energy for
546 calcification among fewer polyps (Swain et al., 2018). We can therefore reasonably assume that
547 nutrients levels, and potentially the zooplankton, did not differ among study sites (Supporting
548 Material Table S2). As a result, a single polyp would have more energetic resources available for
549 calcification than in ambient pH conditions, as indicated by their greater skeletal growth
550 parameters. This particular response may be possible in *A. calycularis* due to its colonial nature.
551 We hypothesize that in order to maintain the calcification rate at the colony level, colonies tend to
552 decrease their number of polyps, but in turn, their few polyps extend their skeleton faster to reach
553 the size of sexual maturity. No asexual division (fragmentation) has been observed over 5 years in
554 *A. calycularis* in the three study sites, suggesting that this strategy of reproduction is not common.
555 *A. calycularis* exhibited at the vent site a morphological shift to encrusting and smaller colonies,
556 with less porous, and potentially more robust corallite skeletal architecture.
557 While individuals from the vents were composed of corallites with higher skeletal density, this
558 was less evident at the colony level. While skeletal integrity was not strictly quantified,
559 observations at the vent site showed that the colonies were more fragile and lost their integrity
560 more readily, suggesting that the section of the skeleton located between the polyps (coenosteum)
561 was either less calcified and/or more prone to dissolution. This could be the result of thinner or

562 absent coenosarc (the tissue covering the coenosteum) found in the colonies at the CO₂ vent.
563 Contrasting responses between skeletal parts that either are or are not protected by living tissues
564 has already been reported for corals (Hennige et al., 2015; Ries, 2011; Rodolfo-Metalpa et al.,
565 2011). This loss of coenosarc could indicate the beginning of a further shift from colonial forms
566 towards solitary polyps to ensure survivorship, as has occurred throughout the history of
567 scleractinian coral evolution and in laboratory conditions (Fine & Tchernov, 2007; Kvitt et al.,
568 2015). Interestingly, it was recently shown that following a heatwave and a bleaching event, the
569 Mediterranean coral *Cladocora caespitosa* suffered from apparent mortality but its tissues actually
570 retracted inside the individual corallites before rejuvenescence occurred a few years later (Kersting
571 & Linares, 2019). Here, we hypothesize that a similar, but less extreme phenomenon occurred with
572 the corals reducing their coenosarc in response to low pH. A reduction in tissue thickness can have
573 implications for calcification because the precipitation of calcium carbonate occurs in the
574 calcifying fluid that is a medium semi-isolated from the external seawater by the overlying coral
575 tissues. To promote calcification, corals have the ability to upregulate pH and C_T (dissolved
576 inorganic carbon) concentrations in the calcifying fluid (Comeau, Cornwall, & McCulloch, 2017),
577 a capacity that is reduced under ocean acidification (McCulloch, Falter, Trotter, & Montagna,
578 2012). Here reduced calcification between the polyps was likely due to a reduction of coral ability
579 to maintain optimal carbonate chemistry conditions in the calcifying fluid between the polyps. This
580 could have been the result of natural spatial heterogeneity sensitivity of colonies to acidification
581 (Holcomb et al., 2014). In addition to reduced calcification, thinning or disappearing of the tissues
582 likely led to local dissolution because exposed skeletons are more prone to dissolution (Ries,
583 2011). As a result of reduced calcification and dissolution of the coenosteum, colonies at the vent
584 site were weaker and smaller despite heavily calcified corallites.

585

586 *Genomics of differentiation of corals at the CO₂ vent system*

587 Our results show strong genetic differentiation of all three *A. calycularis* populations (F_{ST}
588 averaging 0.024 – 0.034), with over 5000 SNPs showing $F_{ST}>0.10$. A previous genetic study of *A.*
589 *calycularis* using microsatellites also found significant genetic differentiation at km-scale
590 distances, likely a reflection of this species limited dispersal capabilities (Casado-Amezúa et al.,
591 2012). These data showed strong linkage among SNPs across 1000s of bp within genes and strong
592 linkage across different genes. These patterns could be generated by selective sweeps acting on a
593 small number of founder colonies, but because linkage among genes and SNPs occurs among
594 ambient pH (Ambient 1 and Ambient 2) individuals as well as Vent corals. These linkages were
595 probably not due to recent selection at the CO₂ vent site alone but reflect the underlying
596 architecture of adaptation. *A. calycularis* is a warm-water coral whose distribution range is
597 currently expanding northward (Bianchi, 2007), where Ischia belongs to the north-east margin of
598 the confirmed distribution. As a result, it is also likely that selection is acting at the Ambient 1 and
599 2 sites compared to more southerly populations. Disentangling the ways in which selection at high
600 CO₂ locations combines with selection at higher temperatures may be particularly important in
601 future ocean conditions. The matrix of genetically distinct populations of *A. calycularis*
602 experiencing a variety of selection regimes for heat and CO₂ may be a powerful setting for this
603 future work.

604 The most highly differentiated genes in the vent site population, Grotta Mago, are
605 annotated for calcium regulation, proton pumping, and inorganic carbon regulation. It is
606 possible that they are differentiated in Grotta Mago for reasons other than selection on

607 calcification in the presence of high CO₂. However, the strong shift in alleles at these loci
608 and the linkage among those differentiated SNPs provides a strong set of hypotheses for
609 the way selection might act to favor coral calcification at low pH.

610 Two loci of calmodulin were highly differentiated in Grotta Mago, with linked SNPs in our
611 data set. Calcium transporter genes are thought to be important in delivering calcium to
612 the calciblastic space (Allemand, Tambutté, Zoccola, & Tambutté, 2011). Though the
613 precise mechanism is not known, calmodulin along with calbindin and calreticulin are
614 hypothesized to play a role in managing calcium levels and can be sensitive to pH
615 (Allemand et al., 2011). For example, Kaniewska et al. (2012) showed 8-fold down
616 regulation of calmodulin in a CO₂-treated reef building coral. An increase in calcium at the
617 site of calcification has been shown as a mean to alleviate the negative effect of low pH in some
618 corals (Decarlo, Comeau, Cornwall, & McCulloch, 2018). Our data also showed excess
619 differentiation in genes that depend on calcium concentrations for their function, such as
620 calineurin, calnexin and the sarcoplasmic reticulum histidine-rich calcium-binding protein,
621 which are thought to play a role in calcium homeostasis.

622 Intracellular and vacuolar H⁺ concentrations are central to coral calcification by controlling
623 pH of the calcifying fluid and the calciblastic and gastrodermal cells. The V-type ATPase
624 proton pump is localized in the symbiosomes of corals that contain intracellular symbionts
625 (Tresguerres, 2016), but is also highly expressed in non-symbiotic gastroderm of
626 symbiont-free tips of quickly calcifying corals, suggesting a role in calcification in the
627 absence of symbionts (Perez, 2015). In particular, the protein VHA (V-type proton
628 ATPase) is differentially regulated in corals exposed to low pH, being downregulated
629 more than four fold in experiments on the reef building coral *Acropora millepora*
630 (Kaniewska et al., 2012). The population of *A. calicularis* in Grotta Mago showed strong
631 differentiation at 6 SNPs across a 100 bp region of this gene.

632 CO₂ diffuses very quickly through cells, and is hard to localize in cell regions that need it
633 for photosynthesis or calcification. Carbonic anhydrase catalyzes the reaction to convert
634 CO₂ to carbonate ions that diffuse much less quickly. As a consequence, carbonic
635 anhydrase is central to calcification in many marine species. In fact, Zebral et al. (2019)
636 suggest use of it as a biomarker to monitor effects of acidification. In corals, carbonic
637 anhydrase is thought to favor carbonate ion concentration in the calciblastic space and

638 in symbiosomes through conversion of CO₂ (Chen et al., 2018; Zoccola et al., 2016).
639 However, low-pH experiments on a deep water coral (*Lophelia pertusa*) did not find strong
640 shifts in carbonic anhydrase activity, nor did an examination of carbonic anhydrase in
641 polychaete worms from the Ischia CO₂ vents (Del Pasqua, Gambi, Caricato, Lionetto, &
642 Giangrande, 2019). Experiments on reef building corals have shown mixed results
643 (Kurihara, Takahashi, Reyes-Bermudez, & Hidaka, 2018). We found strong differentiation
644 of one carbonic anhydrase locus in Grotta Mago. An average shift in minor allele
645 frequency from 0.14 to 0.50 in six linked SNPs may signal differential activity of this gene
646 in some functionally important way. These SNPs appear to be downstream of the
647 carbonic anhydrase coding region and if they play a role it may be as allele-specific
648 regulators of expression. These allele differences may be a good tool to understand the
649 role of carbonic anhydrase in reaction to high CO₂.

650 In conclusion, our study demonstrates that the natural population of the azooxanthellate coral *A.*
651 *calycularis* living near the CO₂ vent system shows variable responses in terms of skeleton and
652 growth patterns that result in a shift in phenotypic and ecological traits. We have shown that these
653 variable responses at the polyp and colony level allow this coral to cope with low and variable pH
654 in the long term by re-allocating energy investments between individual and colony growth as well

655 as mineralogical characteristics. Transcriptomic data revealed strong genetic differentiation among
656 local populations with several candidate loci and linkage blocks under selection. In addition, we
657 found excess differentiation in the CO₂ vent population for genes central to calcification, including
658 genes for calcium management (calmodulin, calcium-binding proteins), pH regulation (V-type
659 proton ATPase), and inorganic carbon regulation (carbonic anhydrase). These patterns highlight
660 both the CO₂ vents as well as the fringes of this species' expansion as potential drivers of
661 adaptation.

662

For Review Only

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868 DATA AVAILABILITY STATEMENT

869 RNASeq FASTQ files for all 41 samples sequenced in this study were deposited in the NCBI Short
870 Read Archive (SRA) under BioProject PRJNA643775 (Accession numbers SRR12135922 -
871 SRR12135962), <https://dataview.ncbi.nlm.nih.gov/object/PRJNA643775?reviewer=5j14na61906tr4dne0bvg9kbhq>.
872 The de novo transcriptome assembly generated in this study and used for mapping the samples has
873 been deposited in the NCBI Transcriptome Shotgun Assembly (TSA) database at
874 DDBJ/ENA/GenBank under accession number GIRZ01000000
875 (<https://www.ncbi.nlm.nih.gov/nucleotide/GIRZ00000000>). The version described in this paper
876 is the first version, GIRZ01000000. The bioinformatics scripts used for assembly, mapping, and
877 SNP-calling are available on Github at DOI: 10.5281/zenodo.3934433.

878

879 AUTHOR CONTRIBUTIONS

880 N.T., E.C., S.P., E.S., S.G., M.C.G designed the study. N.T., E.C., S.A., S.C., F.M., A.M., M.M.,
881 S.P., E.S., L.U., C.d.V, M.C.G were involved with fieldwork. N.T., S.A., J.P.G., A.M., S.G.M.,
882 L.U., C.d.V. analyzed the environmental data; E.C., C.C., P.F., S.G. analyzed the skeletal data; E.
883 S., S. P., and N.T analyzed the genomic data. N.T. drafted the initial manuscript and all authors
884 contributed discussion, writing and interpretation.

885 COMPETING INTERESTS

886 The authors declare that they have no competing interests.

887 ETHICAL STATEMENT

888 All work undertaken in this study complied with current laws of Italy and United States of America
889 for collecting and importing/exporting coral specimens. Cites permits IT/EX/2018/MCE/00170,
890 IT/EX/2017/MCE/00214, IT/EX/2017/MCE/00325.

892 **FIGURES AND TABLES**

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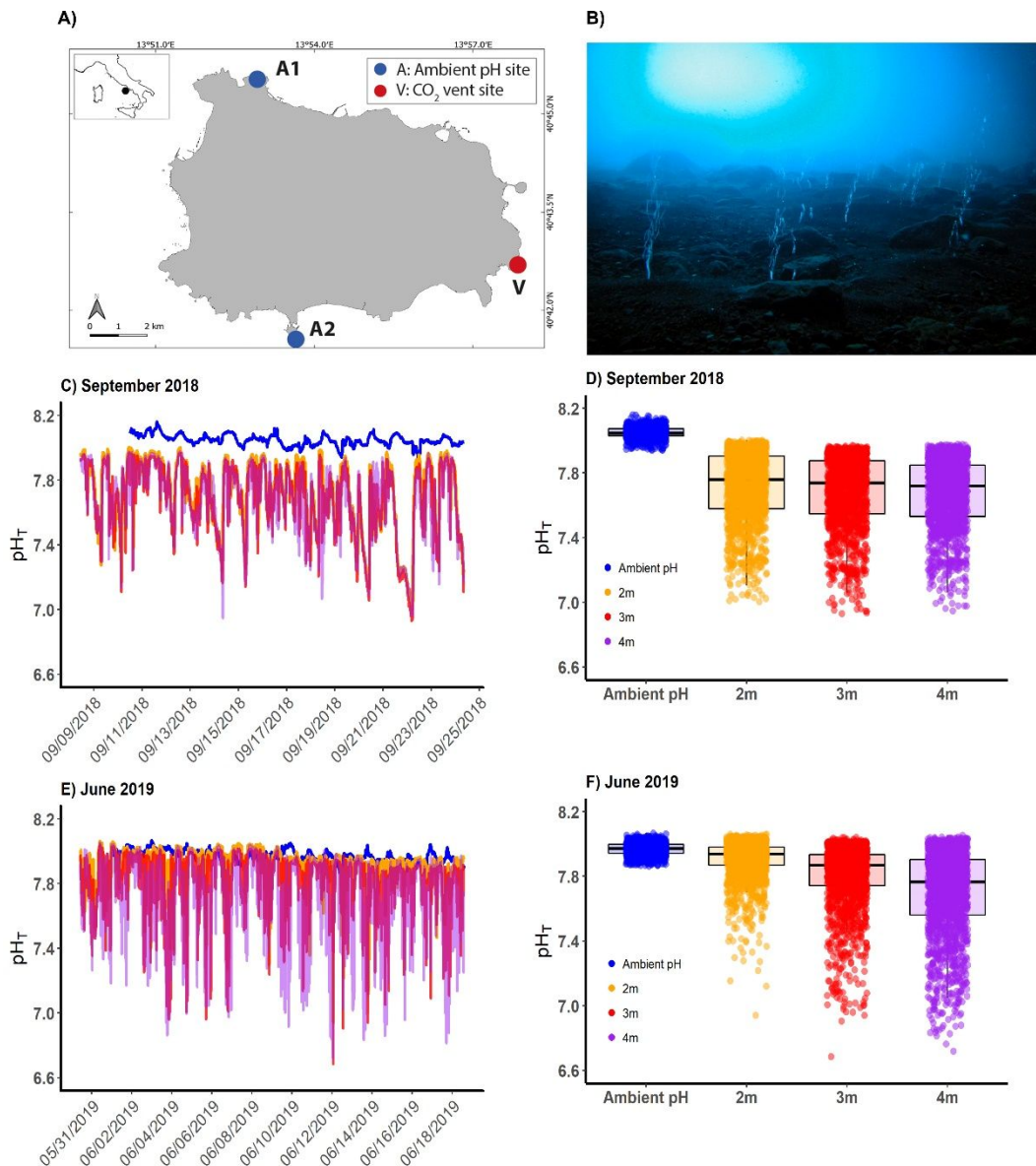
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915 **FIGURE 1. New CO₂ vent system and pH time series and variability.** A) Map showing the

916 location of the study sites along the coast of Ischia Island, Italy. V refers to the CO₂ vent system

917 named *Grotta del Mago*: A1 (Ambient 1, Punta Vico) and A2 (Ambient 2, Sant'Angelo) are off-

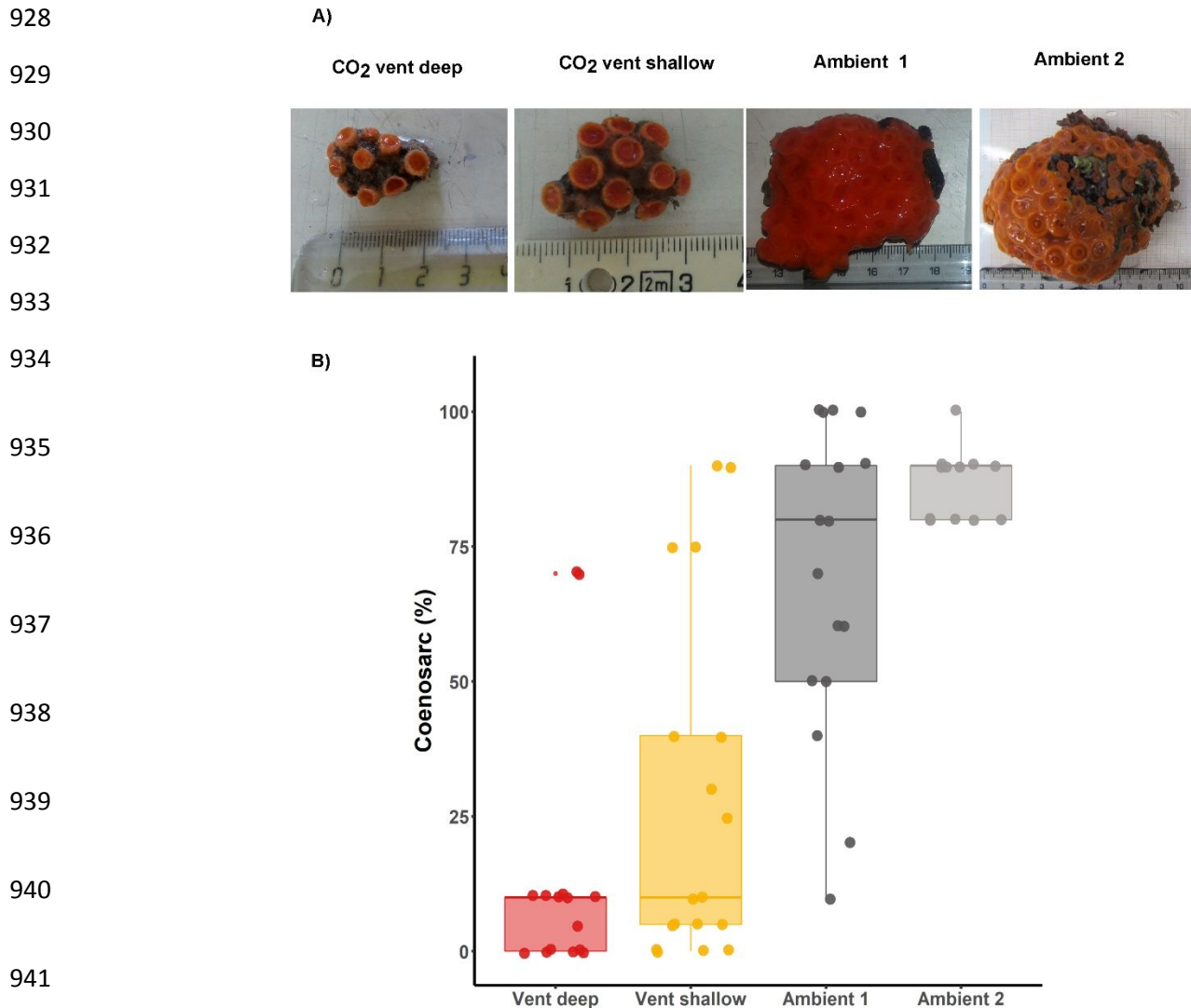
918 vent reference sites with ambient pH. B) The underwater volcanic vents occur in a semi-submerged

919 cave at 5 m depth, release continuously gaseous emissions (92-95% CO₂, no sulfur), and do not

920 elevate temperature (Supporting Note 1, Figure S1, Tables S1, S5). C, E) Time series and D, F)
921 and pH_T (total scale) variability at the CO_2 vent site at 2 m (orange), 3 m (red), and 4 m (violet)
922 depth and at reference sites at 2 m with ambient pH (blue). Dates of pH_T series: from September 8
923 to September 24, 2018 ($n= 1530$ for each depth at the CO_2 vent site and $n= 1331$ for the ambient
924 pH site 1), and from May 30 to June 18, 2019 ($n= 1840$ for each depth at the CO_2 vent site and $n=$
925 1691 for the ambient pH site 2). Measurements were taken every 15 minutes using SeaFETs
926 sensors.

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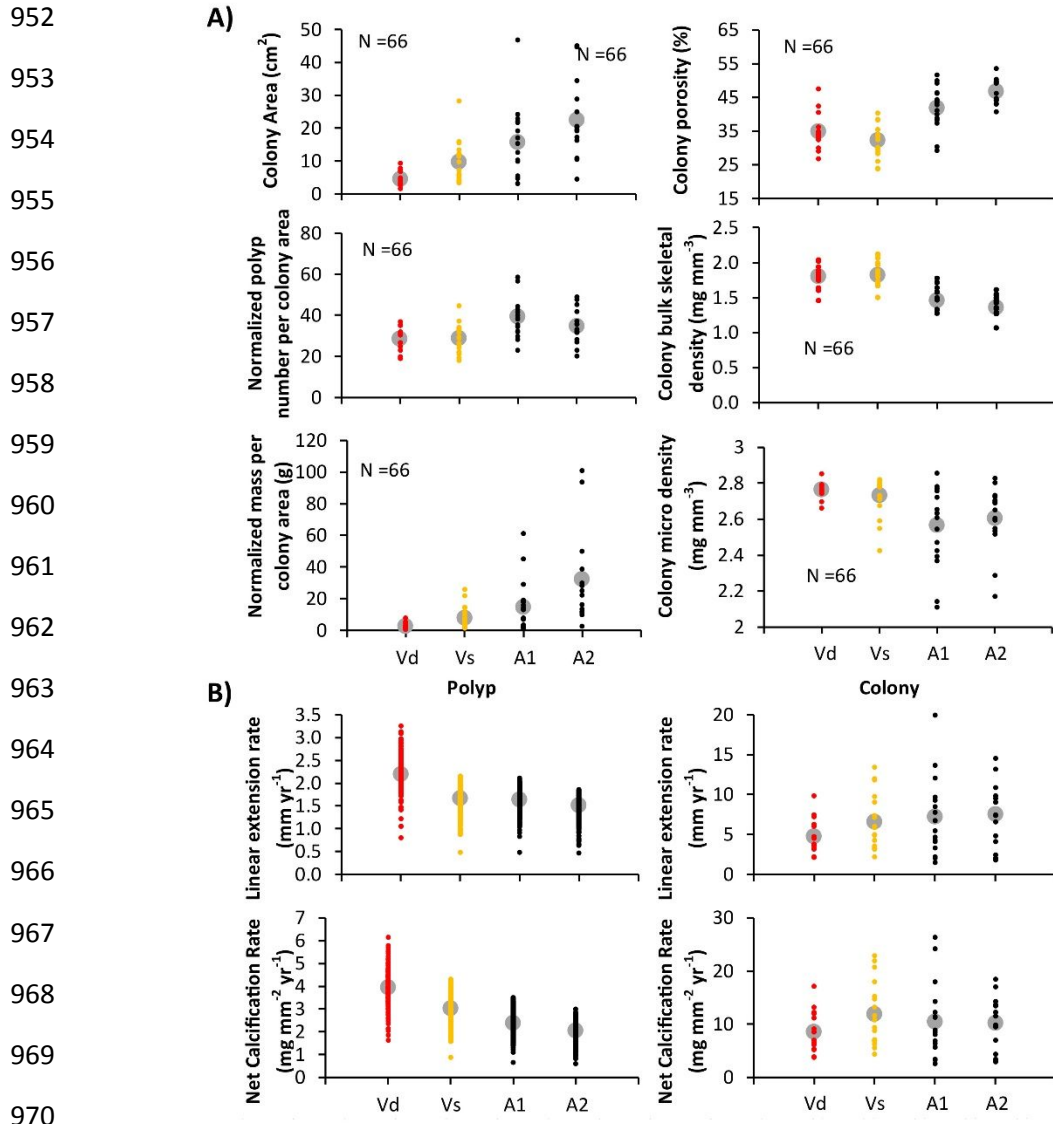
For Review Only



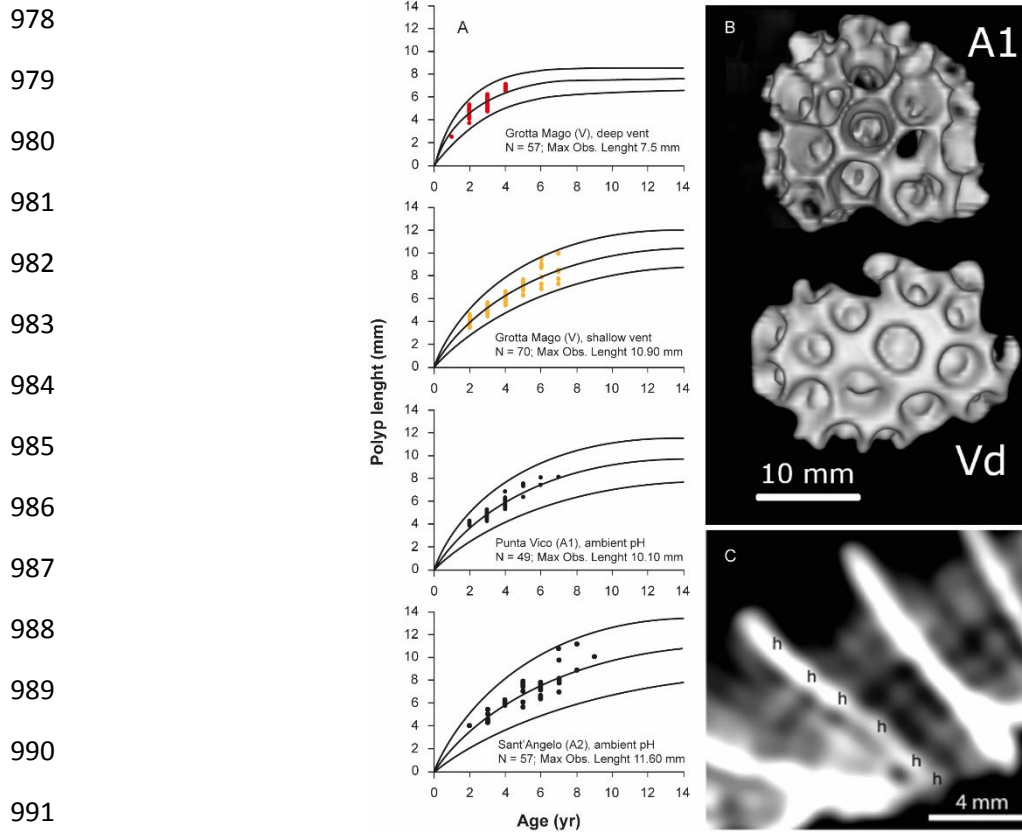
942 **FIGURE 2. A)** Photographs showing colonies sampled for skeletal characteristics and growth of
 943 the CO₂ vent system and off-vent reference sites with ambient pH. Vent system deep (3 m depth);
 944 Vent system shallow (1 – 2 m depth), Ambient 1= Punta Vico (1 – 2 m depth); Ambient 2 =
 945 Sant'Angelo (1 – 2 m depth). Colonies in the vent system exhibited encrusting form (flat growth
 946 form), whereas colonies in Ambient pH sites were a mixture of massive (extensive vertical and
 947 lateral growth) and encrusting forms (see also Figure S7). **B)** % Coenosarc (the living tissue
 948 connecting the polyps) of colonies among the study sites. Significant differences were found
 949 between vent (shallow and deep) and ambient pH populations ($F_{3,58} = 24.9$, $p < 0.0001$; pair-wise

950 comparisons between vent system deep -shallow and off-vent reference sites $p < 0.0001$). Number
951 of colonies = 66.

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971 **FIGURE 3. Skeletal and growth parameters measured in *A. calycularis*.** A) Skeletal
 972 parameters of colonies and **B)** Growth parameters at polyp and colony levels at the CO_2 vent site
 973 (Vd and Vs) and ambient pH conditions (A1 and A2), respectively. Colony mass and polyp number
 974 were normalized to colony area. Grey circles in the plot represent the mean. Vd= Vent system
 975 deep, number of colonies =16; Vs= Vent system shallow, number of colonies = 18; A1= Ambient
 976 1, number of colonies =17; A2 = Ambient 2, number of colonies = 15. Total number of colonies
 977 analyzed = 66.



992 **FIGURE 4. Relationships between age-length growth curves of *A. calycularis*.** A) Age-length

993 von Bertalanffy growth curves at the polyp (corallite) level. Dots indicate the age determined by

994 counting the growth bands on computerized tomography. Lines indicate the mean von Bertalanffy

995 growth curve and the 95% confidence interval). Values of L_{∞} (maximum expected length in the

996 population) and K (a growth constant, larger for fast growth) were: 7.6 mm and 0.5 for Vent deep;

997 11.0 mm and 0.2 for Vent shallow; 10.2 mm and 0.2 for Ambient 1; and 11.8 mm and 0.17 for

998 Ambient 2, respectively. B) Surface 3D rendering of the CT scans performed on a colony of *A.*

999 *calycularis* from Ambient 1 (A1) and on another colony of similar surface area from Vent deep

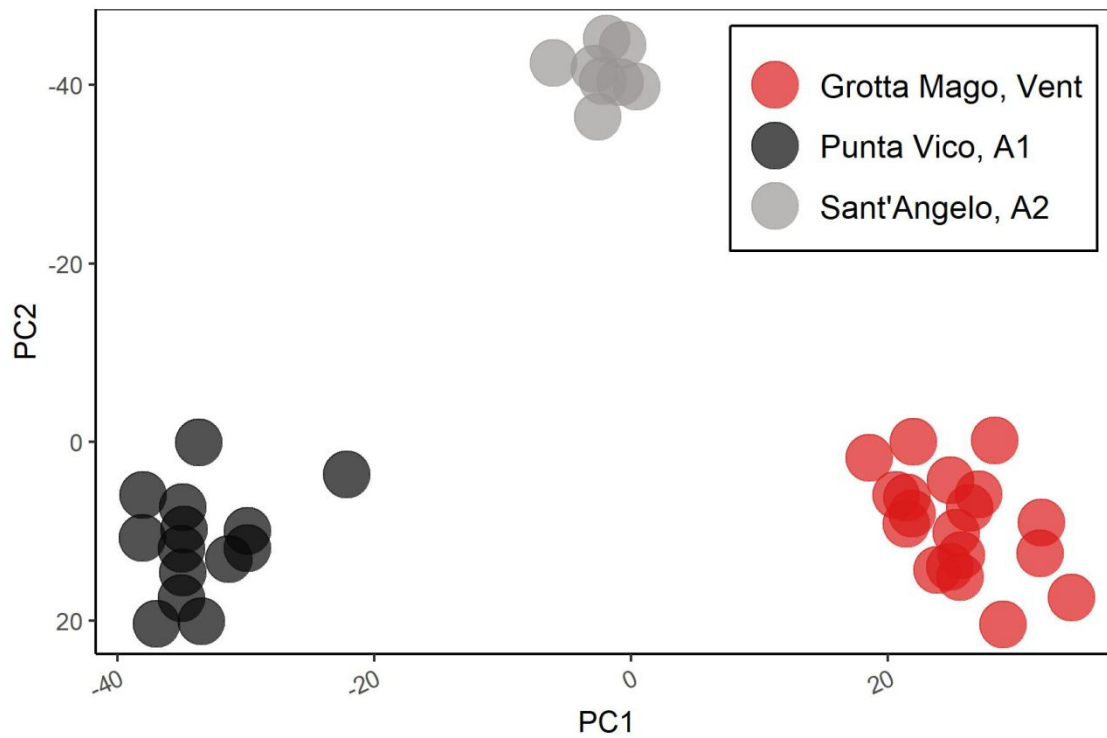
1000 (Vd). Photograph of the deep vent (Vd) colony shows that same calcification is allocated to a minor

1001 number of polyps, and these few polyps result in having a more dense skeleton. C) Computerized

1002 tomography scan to count the growth bands on a single corallite. In this photograph, the corallite

1003 of *A. calycularis* is 6 years old; h indicates high density annual bands.

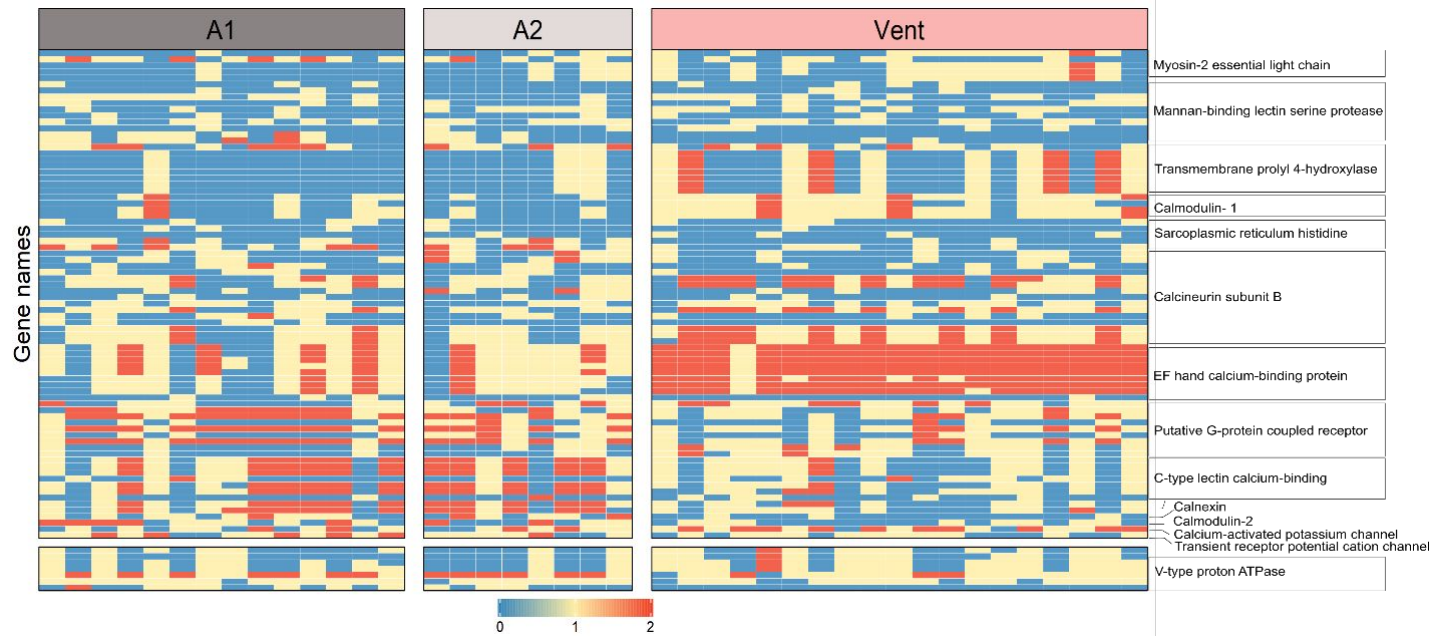
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1006 **FIGURE 5. Population genetic structure of *A. calycularis* based on 46,784 single nucleotide**
1007 **polymorphisms (SNPs).** Number of individuals: CO₂ vent site (Vent, Grotta Mago) = 19; ambient
1008 pH sites: Ambient 1 (A1, Punta Vico) = 14, Ambient 2 (A2, Sant' Angelo) = 8.

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1012 **FIGURE 6.** Upper: SNP genotypes for 13 calcium ion related loci showing high levels of
 1013 divergence in the Grotta Mago population. Lower: Genotypes for 7 SNPs within a highly
 1014 differentiated V-type proton ATPase potentially involved in pH regulation in the calcicoblastic
 1015 layer where calcification occurs. Vent: CO₂ vent (Grotta Mago), A1: Ambient 1 (Punta Vico); A2:
 1016 Ambient 2 (Sant'Angelo). Legend: 0 = homozygous major allele, 1 = heterozygous; 2 =
 1017 homozygous minor allele.

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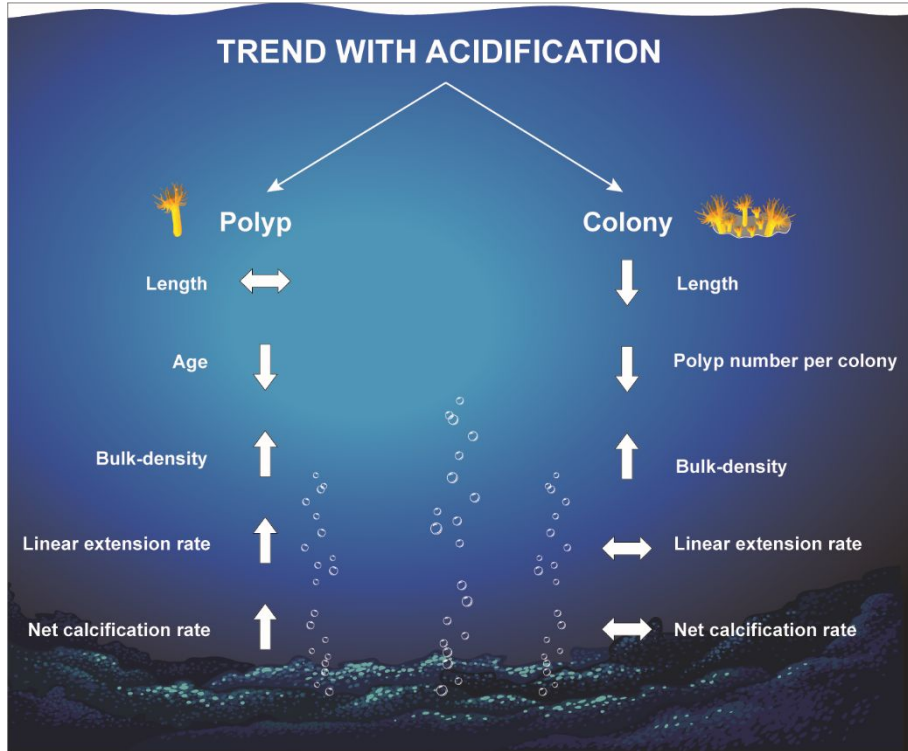
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1035 **FIGURE 7. Schematic summary of the responses on skeletal and growth parameters to ocean**1036 **acidification measured in the coral *A. calycularis* at the polyp and colony levels.**

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1038 **TABLE 1.** Measured and estimated seawater physiochemical parameters at the CO₂ vent site and
 1039 reference areas with ambient pH for salinity (S), temperature (T), total alkalinity (A_T), dissolved
 1040 inorganic carbon (C_T), pH_T, pCO₂, calcite (Ω_c) and aragonite(Ω_a) saturation. Values are means, ±
 1041 SD with 25th and 75th percentiles. Calculated concentrations of C_T, pCO₂, Ω_c and Ω_a are shown.
 1042 1: Parameters measured from discrete water samples; 2: parameters measured with *in situ* sensors.
 1043 For detailed SeaFET pH sensor statistics and the carbonate system parameters, see Figure S5 and
 1044 Table S5, respectively.

Month/year	Vent site (GM)			A1-PV	A2-SA
	2m	3 m	4 m	2 m	2 m
September 2018					
S	37.3 ¹ ± 0.2 (37.2, 37.5), n=9	37.3 ¹ ± 0.2 (37.2, 37.5), n=9	37.3 ¹ ± 0.2 (37.2, 37.5), n=9	37.3 ¹ ± 0 (37.3, 37.3), n=5	37.4 ¹ ± 0 (37.4, 37.4), n=3
T (°C)	25.9 ² ± 0.2 (25.8, 26.0), n=1530	26.0 ² ± 0.2 (25.8, 26.0), n=1530	26.0 ² ± 0.2 (25.9, 26.1), n=1530	17.3 ² ± 0.4 (17.0, 17.6), n=1331 ²	25.9 ² ± 0.2 (59.9, 26.1), n=408
A _T (μmol kg ⁻¹)	2564 ¹ ± 7 (2561, 2566), n=9	2562 ± 8 (2557, 2565), n=9	2562 ¹ ± 8 (2556, 2566), n=9	2618 ¹ ± 15 (2607, 2633), n=5	2610 ¹ ± 1 (2609, 2611), n=3
C _T (μmol kg ⁻¹)	2542 ± 79 (2477, 2585), n=1530	2552 ± 84 (2485, 2593), n=1530	2555 ± 80 (2495, 2598), n=1530	2262 ± 24 (2246, 2276), n=1331	2275 ± 1 (2275, 2276), n=3
pH _T	7.65 ² (7.58, 7.90), n=1530	7.62 ² (7.55, 7.87), n=1530	7.60 ² (7.53, 7.85), n=1530	8.05 ² (8.03, 8.07), n=1331	8.02 ¹ (8.02, 8.03), n=3
pCO ₂ (μatm)	2905 ± 1664 (1724, 3438), n=1530	3146 ± 1928 (1837, 3668), n=1530	3192 ± 1806 (1958, 3799), n=1530	322 ± 34 (298, 341), n=1331	375 ± 1 (374, 375), n=3
Ω _c	1.68 ± 0.59 (1.21, 2.21), n=1530	1.58 ± 0.56 (1.14, 2.09), n=1530	1.54 ± 0.54 (1.10, 1.98), n=1530	5.96 ± 0.36 (5.74, 6.21), n=1331	5.70 ± 0.01 (5.69, 5.70), n=3
Ω _a	1.11 ± 0.39 (0.80, 1.47), n=1530	1.05 ± 0.37 (0.75, 1.39), n=1530	1.02 ± 0.36 (0.73, 1.32), n=1530	3.86 ± 0.23 (3.72, 4.02), n=1331	3.71 ± 0.01 (3.70, 3.71), n=3
June 2019					
S	37.8 ¹ ± 0 (37.8, 37.8), n=7	37.8 ¹ ± 0 (37.8, 37.8), n=7	37.8 ¹ ± 0 (37.8, 37.8), n=7	38.0 ¹ ± 0 (38.0, 38.0), n=3	37.9 ¹ ± 0 (37.9, 37.9), n=7
T (°C)	21.9 ² ± 2.1 (19.9, 24.1), n=1840	21.8 ² ± 2.1 (19.8, 23.8), n=1840 ²	21.7 ² ± 2.0 (19.6, 23.4), n=1840 ²	26.2 ² ± 0.2 (26.1, 26.3), n=408	26.2 ² ± 1.1 (25.8, 27.0), n=1691 ²
A _T (μmol kg ⁻¹)	2539 ¹ ± 22 (2593, 2552), n=7	2541 ¹ ± 20 (2532, 2550), n=7	2551 ¹ ± 22 (2538, 2568), n=7	2630 ¹ ± 1 (2630.1, 2630.9), n=3	2642 ¹ ± 17 (2629, 2659), n= 7
C _T (μmol kg ⁻¹)	2450 ± 42 (2424, 2464), n=1840	2489 ± 75 (2443, 2509), n=1840	2535 ± 104 (2461, 2574), n=1840	2336 ± 3 (2320, 2353), n=3	2336 ± 23 (2320, 2353), n=1691
pH _T	7.88 ² (7.86, 7.98), n=1840	7.74 ² (7.74, 7.93), n=1840	7.60 ² (7.56, 7.90), n=1840	8.04 ¹ (8.04, 8.04), n=3	7.97 ² (7.94, 7.99), n=1691
pCO ₂ (μatm)	1531 ± 627 (1167, 1653), n=1840	2082 ± 1502 (1352, 2127), n=1840	2812 ± 2310 (1495, 3090), n=1840	372 ± 1 (372, 373), n=3	586 ± 76 (532, 635), n=1691
Ω _c	2.20 ± 0.41 (2.01, 2.46), n=1840	1.87 ± 0.54 (1.57, 2.26), n=1840	1.60 ± 0.65 (1.10, 2.14), n=1840	5.92 ± 0.01 (5.92, 5.93), n=3	5.34 ± 0.35 (5.08, 5.58), n=1691
Ω _a	1.44 ± 0.27 (1.32, 1.61), n=1840	1.23 ± 0.35 (1.03, 1.48), n=1840	1.05 ± 0.42 (0.72, 1.40), n=1840	3.89 ± 0.01 (3.86, 3.87), n=3	3.54 ± 0.23 (3.38, 3.71), n=1691

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