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Large-scale stable isotope characterization of a Late Cretaceous dinosaur-dominated ecosystem

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- 1 Large-scale stable isotope characterization of a Late Cretaceous
- 2 dinosaur-dominated ecosystem
- 3
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23 ABSTRACT

24 In the Cretaceous of North America, environmental sensitivity and habitat specialization have 25 been hypothesized to explain the surprisingly restricted geographic ranges of many large-bodied 26 dinosaurs. Understanding the drivers behind this are key to predicting broader trends of dinosaur 27 species and community response to climate change under greenhouse conditions. However, 28 previous studies of this question have frequently examined only small sub-components of the 29 paleoecosystem or operated without comparison to similar modern systems from which to 30 constrain interpretations. Here we perform a high-resolution multi-taxic δ^{13} C and δ^{18} O study of a 31 Cretaceous coastal floodplain ecosystem, focusing on species interactions and paleotemperature 32 estimation, and compare with similar data from extant systems. Bioapatite δ^{13} C preserves 33 predator-prey offsets between tyrannosaurs and ornithischians, and between aquatic reptiles and 34 fish. Large ornithischians had broadly overlapping stable isotope ranges, contrary to 35 hypothesized niche partitioning driven by specialization on coastal or inland subhabitat use. 36 Comparisons to a modern analogue coastal floodplain show similar patterns of ecological guild 37 structure and aquatic-terrestrial resource interchange. Multi-taxic oxygen isotope temperature 38 estimations yield results for the Campanian of Alberta consistent with the few other 39 paleotemperature proxies available, and are validated when applied for extant species from a 40 modern coastal floodplain, suggesting this approach is a simple and effective avenue for 41 paleoenvironmental reconstruction. Together, these new data suggest dinosaur niche-partitioning 42 was more complex than previously hypothesized, and provide a framework for future research on 43 dinosaur-dominated Mesozoic floodplain communities.

45 INTRODUCTION

46 The Cretaceous is a key interval of Earth history for understanding paleocommunity structure 47 and ecological dynamics in greenhouse systems, as it contains a rich and well-studied fossil 48 record from multiple continents, with those from the Campanian of Alberta, Canada, being 49 among the richest and most intensively collected (Brown et al., 2013). To thoroughly 50 characterize paleoecological dynamics at this time, however, we must build upon the inferences 51 obtained from broad-scale biodiversity meta-analyses (e.g. Butler et al., 2011; Sakamoto et al., 52 2016) and examine species and communities at more ecologically sensitive spatial/temporal 53 scales (Price et al., 2018). Of particular interest is the testing of hypothesized subhabitat 54 partitioning, which has long been suggested to explain restricted distributions in coeval 55 Cretaceous dinosaurs (Fricke and Pearson, 2008; Horner et al., 1992; Lehman, 2001; Sampson et 56 al., 2010). These patterns are surprising given that large terrestrial tetrapods today commonly 57 have broad ranges and are relatively tolerant of changing environmental conditions (Lehman, 58 1987). The veracity of such patterns for dinosaur habitat zonation, however, remain debated (e.g. 59 Lucas et al., 2016), and are only variably recovered in studies of Cretaceous regional vertebrate 60 community dynamics (Brinkman et al., 1998; Cullen and Evans, 2016). Stable isotope analyses 61 provide a useful tool for testing these hypotheses, while simultaneously allowing assessment of 62 paleotemperature conditions in the terrestrial deposits of the Campanian of Alberta/Montana, 63 which remains relatively uncharacterized despite its densely-sampled and well-studied fossil 64 record (Barrick et al., 1999; Upchurch and Wolfe, 1987). Detailed assessments of 65 paleoecological and paleoenvironmental dynamics across both spatial and temporal gradients are 66 a key piece of data for patterns of biodiversity in greenhouse systems.

68 APPROACH

69 We perform the broadest multi-taxic stable isotope analysis to date of a vertebrate community in 70 the Mesozoic. We sample a phylogenetically, physiologically, and ecologically broad range of 71 vertebrate taxa from the Late Cretaceous of Alberta, so as to capture the isotopic range in this 72 ecosystem at a well-constrained spatial and temporal scale. Our analyses are used: (1) to 73 reconstruct paleocommunity structure from this ancient coastal floodplain and compare it with stable isotope data from modern analogues, (2) to test the hypothesis of subhabitat specialization 74 75 amongst co-occurring large ornithischians, and (3) to estimate temperature(s) for the sampled 76 assemblage, with comparisons to the available data from regional samples obtained via other 77 proxies, and confirmatory tests of oxygen isotope-derived temperatures from modern analogues . 78

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79 LOCALITY & GEOLOGICAL SETTING

80 We sampled a vertebrate microfossil bonebed, the 'Rainy Day Site', from the upper Oldman 81 Formation of Alberta (known as Judith River Fm. in Montana) (Figure 1) (Arbour and Evans, 82 2017; Cullen and Evans, 2016). Sedimentology in this locality is consistent with an active 83 channel belt with high aggradation and sediment support, grading into mud-dominated intervals 84 reflective of the wetter, more organic-rich settings of the uppermost Oldman Formation (Figure 85 1D). These deposits are constrained spatially, minimally time-averaged, represent local 86 deposition of wetland or similar materials in fluvial deposits in the broader coastal floodplain 87 environment, and, as the fossils were preserved together, were likely exposed to the same suite of 88 diagenetic/taphonomic effects (Rogers and Brady, 2010). This differs from many other isotopic 89 studies on Mesozoic dinosaur communities, which derive samples from multiple sites and 90 lithologies; that lack of control on diagenetic and other variation within a sample increases the

potential of significant analytical errors. Cretaceous data were compared to modern analogue
coastal floodplain stable isotope analyses of the Atchafalaya River Basin fauna of Louisiana
from Cullen et al. (2019). The Atchafalaya River Basin represents a useful analogue to the upper
Oldman Fm., as both are low-lying, seasonally flooded landscapes with warm temperate to
subtropical climates, contain relatively similar flora, and have numerous taxa that are
physiologically, ecologically, and phylogenetically similar (e.g. metatherian mammals,
crocodilians, lepisosteid fish, etc).

98

99 MATERIAL & METHODS

100 Stable isotope analyses performed on 83 samples (N_{taxa}=17), measured via laser-ablation gas-101 chromatography isotope-ratio-mass-spectrometry at the University of Western Ontario, reported 102 using standard isotopic practice (δ^{13} C and δ^{18} O), following the methods of Cullen et al (2019, 103 and references therein). Stable isotope compositions for each specimen and associated data were 104 recorded (Table DR1 & 2), with tests assessing sampling biases, XRD examining mineral 105 alteration, and relative grouping of isotopic compositions of sampled taxa (sensu Fricke et al., 106 2008 and references therein) finding no bias and a lack of pervasive alteration (Figure DR1). The 107 carbon isotope composition of diet ($\delta^{13}C_{diet}$) was calculated using trophic enrichment factors (or 108 TEF; $\Delta = \delta_{\text{bioapatite}} - \delta_{\text{diet}}$) obtained for extant relatives and/or physiological analogues of sampled 109 taxa (Table DR3). Given a lack of specific ecologically-relevant extant relatives, dinosaur 110 $\delta^{13}C_{diet}$ was reconstructed under three TEF scenarios, as birds, as reptiles, and as mammals, to 111 reveal potential effects of dietary-physiology differences. Stable isotope compositions and TEF-112 applied $\delta^{13}C_{diet}$ compositions for Cretaceous and modern taxa are included as species-level 113 means in Table DR4.

Temperature calculations derived from oxygen isotope compositions (δ^{18} O) were 114 115 performed using taxon-bioapatite-isotope geothermometers, and particularly the ectotherm-116 endotherm combined mean (EECM) equation of Cullen et al. (2019) (modified from dual taxon 117 approach of Fricke and Wing, 2004, including species-specific methods of others from 118 references therein)). This modified method calculates the combined mean of bioapatite $\delta^{18}O_{phosphate}$ (in this case, modified by 1‰ from laser-produced total $\delta^{18}O_{LPTO}$, sensu 119 Larson and Longstaffe, 2007) from each sampled ectotherm taxon, as well as the combined 120 121 mean of δ^{18} O_{water} determined for each sampled endotherm taxon, and uses them in the equation: T=111.4 – (4.3)(combined mean bioapatite $\delta^{18}O_{ecto phosphate}$ – combined mean $\delta^{18}O_{endo water}$) 122

123

124 PALEOCOMMUNITY STRUCTURE

Stable carbon compositions for Cretaceous taxa show offsets in δ^{13} C of ~1-3‰ present between 125 126 megacarnivorous tyrannosaurids and their large-bodied ornithischian prey, as well as between 127 aquatic reptiles and fish (Figure 2), consistent with expectations from predator-prey pairs known 128 for extant systems (Bocherens, 2000). There is considerable overlap in both δ^{13} C and δ^{18} O 129 among sampled bulk-feeding herbivores (hadrosaurids, ceratopsians, ankylosaurs) (Figures 2 and 130 3, Figure DR1C). While the full extent of plant isotopic heterogeneity in the Late Cretaceous of 131 this region of North America is not known, there is sufficient evidence to suggest that coastal vs. 132 inland signals should be detectable (Fricke and Pearson, 2008). This overlap thus challenges the 133 hypothesis that, at least in this model Campanian ecosystem, these taxa significantly partitioned 134 isotopic niches (a proxy for ecological niche, see Newsome et al., 2007) via subhabitat 135 specialization. Our results suggest instead that habitat use and dietary differences (if present) are 136 either not detectable using bioapatite δ^{13} C and δ^{18} O, or that they partitioned niches via other

137	means, such as temporal or feeding-height stratification, across a broad range of plant sources
138	(Byerly et al., 2017; Mallon et al., 2013). The lack of apparent canopy effects in these data
139	provide some evidence against the latter, but this requires further study (Kohn, 2010, and
140	references therein). As well, it should be noted that species-level niche-partitioning also may
141	exist in these systems (Farlow and Planka, 2002; Mallon, 2019), and these results primarily
142	inform hypotheses of subhabitat niche partitioning among broader groups like hadrosaurids and
143	ceratopsids. It is also conceivable that these greenhouse environments were ecologically
144	unsaturated, facilitating a greater degree of shared resource use and weaker niche partitioning
145	(Pinto-Sanchez et al., 2014).
146	The small theropod Troodon, hypothesized as omnivorous or herbivorous (Holtz Jr et al.,
147	1998; Sankey, 2008), is shown here to be isotopically similar to the faunivorous small theropods
148	Saurornitholestes and Dromaeosaurus, as well as indeterminate mammals, suggesting a mixed-
149	feeding omnivorous or faunivorous diet (Figures 2 and 3). The data for the enigmatic
150	Richardoestesia, known only from a pair of jaws and isolated teeth, are more difficult to
151	interpret; our results do not strongly support hypothesized piscivory (Sankey, 2008), as δ^{13} C
152	compositions could support feeding on fish and/or terrestrial resources (Figures 2 and 3). The
153	aquatic reptile <i>Champsosaurus</i> and the ray <i>Myledaphus</i> have more negative bioapatite δ^{13} C
154	compared to other aquatic taxa, similar to the that observed between otters and other aquatic taxa
155	in the Atchafalaya (Figure 3). This may suggest a larger dietary proportion from aquatic
156	invertebrates (e.g. crustaceans, molluscs) when compared to crocodilians and large predatory fish
157	(Cullen et al., 2019). Relative isospace structure is similar between Cretaceous and modern
158	systems, regardless of which TEF is used to reconstruct dinosaur $\delta^{13}C_{diet}$. This suggests that at
159	the guild-level both modern and Cretaceous coastal floodplains are ecologically similar in terms

of relative structure and composition, and that, despite some differences, the Atchafalaya
 provides a robust community and paleoenvironmental analogue for the Campanian of Alberta, and
 more generally, coastal plain ecosystems from the latest Cretaceous of North America.

163 Overall, where overlap exists our results are in line with stable isotope compositions 164 recorded from other Late Cretaceous vertebrates (Amiot et al., 2015; Fricke and Pearson, 2008), 165 though our combination of broader taxon sampling and extant comparisons facilitate a greater degree of ecological resolution for reconstruction and hypothesis testing. However, we recover 166 167 distinct results from those of Ostrom et al (1993), the only other broad multi-taxic stable isotope 168 study to attempt to reconstruct these ecosystems. Their analyses of high-molecular-weight organic δ^{13} C and δ^{15} N extracted from a mix of fossil bones and teeth sampled from multiple 169 170 sites throughout the Dinosaur Park Formation (formerly Judith River Formation) of Alberta 171 produced somewhat equivocal results, with many compositions at odds with expectations (e.g. putatively herbivorous ornithomimids with higher δ^{15} N than faunivorous tyrannosaurs, 172 173 faunivorous lizards with lower δ^{15} N than herbivorous dinosaurs, etc). δ^{13} C compositions for their 174 samples were ~ -27 to -23%, rather than the ~ -10 to -3% compositions recorded here and in 175 other stable isotope studies of Cretaceous dinosaurs (Amiot et al., 2015; Fricke et al., 2008). The 176 differing sources of the material and distinct analytical methods may account for this disparity, 177 and it is also possible that their results may be reflective of a sediment bulk organic carbon 178 signal, as the latter has a nearly identical δ^{13} C range in the Late Cretaceous of western North 179 America (Fricke et al., 2008).

180

181 PALEOTEMPERATURE RECONSTRUCTION

182 Individual taxon oxygen isotope temperature estimates produce variable results, though the mean 183 paleotemperature estimate for the Late Campanian of Alberta derived from the multi-taxic 184 EECM method is 18°C (Figure 4, Table DR5). This is consistent with independent mean annual 185 paleotemperature proxies from the broader region, including leaf physiognomy (Upchurch and 186 Wolfe, 1987), clumped isotope analysis of paleosol carbonate nodules (Burgener et al., 2019), 187 and paired taxon oxygen isotope temperature estimates (Barrick et al., 1999), and produces 188 distinct results from the single taxon temperature estimate approaches (Figure 4, Amiot et al., 189 2004, and references therein; Kohn, 1996)). Applying EECM to extant taxa from the modern 190 system produced a temperature estimate within ~2 °C of mean annual temperature, providing an 191 important independent check that the method produces useful results and may moderate some of 192 the variation introduced by individual taxon physiologies when using the single-taxon or 'single 193 ectotherm, single endotherm' variations of the approach (Figure 4). Our modified EECM method 194 is therefore a robust and simple tool for estimating paleotemperature in conjunction with multi-195 taxic paleoecological stable isotope datasets.

196

197 CONCLUSIONS

Stable isotope analyses of a spatially and temporally constrained Mesozoic vertebrate ecosystem reveals evidence of predator-prey offsets in δ^{13} C compositions and ecological guild distributions that meet predictions derived from extant analogue ecosystems. The overlapping isotopic ranges of bioapatite from large ornithischians argues against a long-held hypothesis that these groups were strongly partitioning their niches based on subhabitat sensitivities. The bioapatite oxygenisotope EECM approach for paleotemperature estimation represents a and simple and robust technique for paleoclimate reconstruction. Our combined fossil + modern combined approach provides a framework for future paleocommunity analyses in these abundantly preserved coastalplain systems.

207

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224

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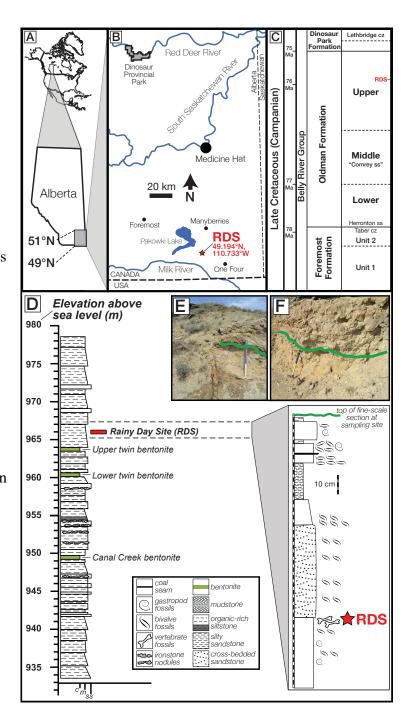
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333	
334	

FIGURES & FIGURE CAPTIONS

336

337 Figure 1. Geographic and 338 stratigraphic position of 339 Cretaceous locality ('Rainy Day 340 Site', or RDS). A, position of 341 study area in North America; B, 342 position of RDS in southern 343 Alberta (specific GPS coordinates 344 on-file at Royal Tyrrell Museum 345 and available upon request); C, 346 Campanian stratigraphy of southern Alberta with 347 stratigraphic position of RDS 348 349 indicated; D, Stratigraphic section 350 at RDS locality indicating 351 location of RDS in respect to 352 other nearby sites and marker 353 beds; E, photograph of locality 354 showing position of RDS and 355 nearby sites and marker beds; F, 356 photograph of broader area stratigraphy. Green line in E & F indicates position relative to



357 stratigraphic column in D inset. Abbreviations: c-clay; cz-coal zone; m-mud; ss-sandstone.

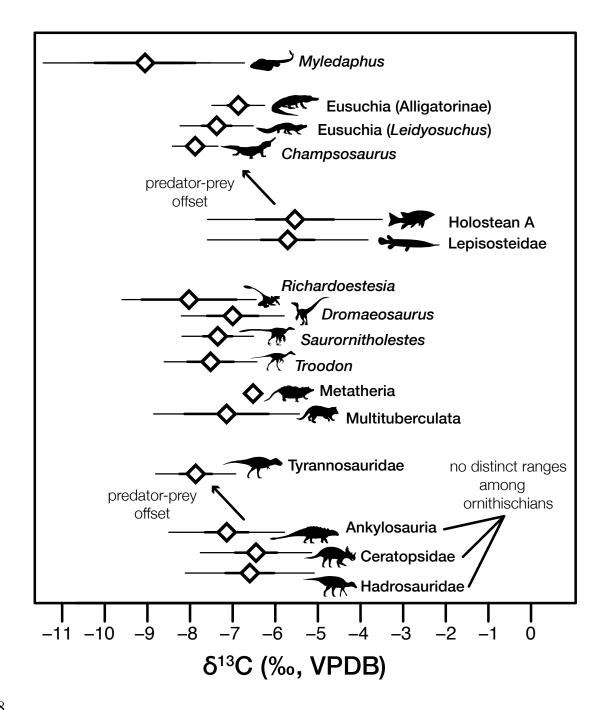
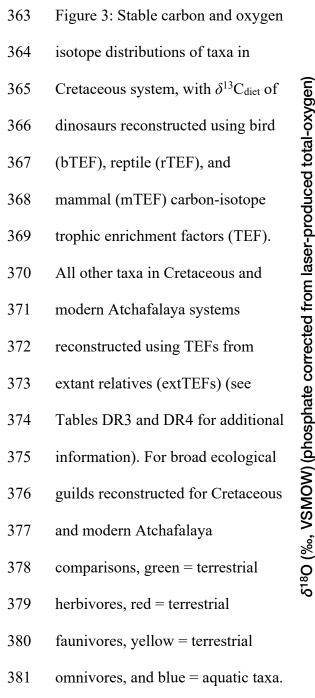
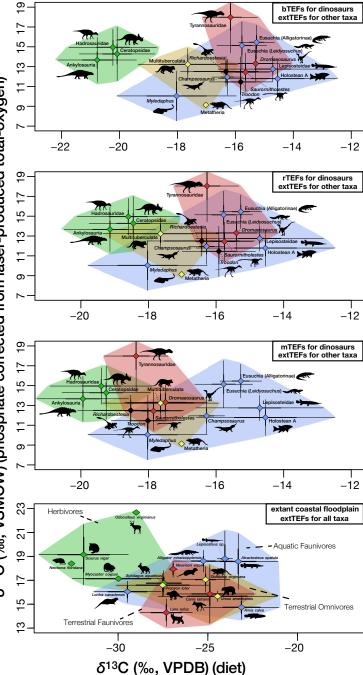
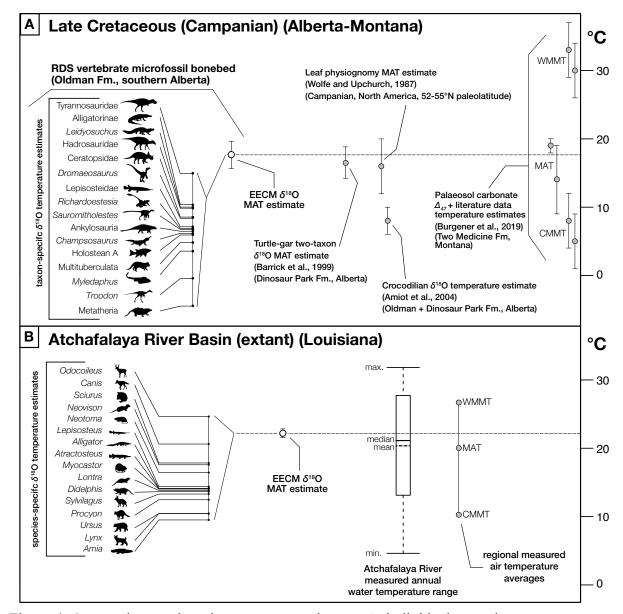


Figure 2. Stable carbon isotope compositions of analyzed fossil taxa. Expected predator-prey
isotopic offsets exist between tyrannosaurids and large ornithischians, and between aquatic
reptiles and fish. No distinct separation present in the isotopic ranges of sampled ornithischians.





Stable oxygen isotope compositions are 'phosphate-corrected' (i.e. modified to display only $\delta^{18}O_{phosphate}$, and remove the small fraction of $\delta^{18}O_{carbonate}$ and $\delta^{18}O_{hydroxyl}$ also measured in laserproduced total-oxygen) in order to facilitate comparisons with modern Atchafalaya data from Cullen et al. (2019). Full description of method contained in that reference.



386 Figure 4: Oxygen isotope-based temperature estimates. A, individual taxon isotope-temperature 387 estimates from Cretaceous paleoecosystem compared to endotherm-ectotherm-combined-mean 388 (EECM) estimate from same data, alongside temperature estimates from other proxies (MAT= 389 mean annual temperature, WMMT= warm month mean temperature, and CMMT= cold month 390 mean temperature). B, individual taxon isotope-temperature estimates from the modern 391 Atchafalaya River Basin ecosystem compared to EECM estimate from same data, to measured 392 temperatures from Atchafalaya River (taken multiple times per day during one full year), and to measured air temperatures for that region. All temperatures and $\delta^{18}O_{water}$ calculated from 393 394 $\delta^{18}O_{phosphate}$ compositions. See Table DR5 for detailed information and original data sources.

396 SUPPLEMENTARY INFORMATION

- ³⁹⁷ ¹GSA Data Repository item 2020155. This contains the data, including Figure DR1 (tests of
- 398 sampling bias and geochemical alteration), Table DR1 (mean δ^{13} C, δ^{18} O and collection data for
- all specimens), Table DR2 (δ^{13} C and δ^{18} O measurements for all samples), Table DR3 (Trophic
- 400 Enrichment Factors applied for $\delta^{13}C_{diet}$ comparisons between Cretaceous and modern faunal
- 401 communities), Table DR4 (species mean δ^{13} C and δ^{18} O for Cretaceous and modern taxa), Table
- 402 DR5 (Calculated oxygen isotope-based temperature estimates for Cretaceous and Atchafalaya,
- 403 other Campanian paleotemperature proxies, and modern Louisiana temperature data), and
- 404 associated supplementary references mentioned within the included DR figures and tables. Data
- 405 Repository available online at http://www.geosociety.org/datarepository/2020/, or on request
- 406 from editing@geosociety.org