

Estimation of recharge in mountain hard-rock aquifers based on discrete spring discharge monitoring during base-flow recession

Stefano Segadelli^a, Maria Filippini^b*, Anna Monti^b, Fulvio Celico^c, Alessandro Gargini^b

^a Geological, Seismic and Soil Service, Emilia-Romagna Region Administration, Bologna, Italy

^b Department of Biological, Geological, and Environmental Sciences - BiGeA, Alma Mater Studiorum

University of Bologna, via Zamboni, 67, 40126; Bologna, Italy, maria.filippini3@unibo.it

^c Department of Chemistry, Life Sciences and Environmental Sustainability, University of Parma, Italy

* Corresponding author

HYDROGEOLOGY JOURNAL – ELECTRONIC SUPPLEMENTARY MATERIAL

Table of contents:

S1. Discharge, temperature and electrical conductivity measurements at the springs of Mt. Prinzera and Mt. Zirone

S2. Hydrographs of the springs of Mt. Prinzera and Mt. Zirone

S3. Estimation of annual rainfall along hydrogeologic years 2012-2013 (Mt. Prinzera) and 2016-2017 (Mt. Zirone)

S4. Hydrograph analysis by means of “recession plots”

S5. Estimation of recharge at Mt. Prinzera and Mt. Zirone from annual water budgeting

S6. Surface evidence for the infiltration potential at Mt. Prinzera and Mt. Zirone

		Z01			Z02			Z03			Z04			Z05			days after previous measure
		Q [l/s]	T [°C]	EC [µS/cm]	Q [l/s]	T [°C]	EC [µS/cm]	Q [l/s]	T [°C]	EC [µS/cm]	Q [l/s]	T [°C]	EC [µS/cm]	Q [l/s]	T [°C]	EC [µS/cm]	
Fall	17/10/2016	0.135	12.1	333	0.111	12.4	644	0.800	8.9	420	0.907	11.7	228	0.047	14.7	434	\
	27/10/2016	0.140	12.1	328	0.286	12.3	648	0.408	9.5	445	1.010	11.7	226	0.121	14.8	532	10
	31/10/2016	0.129	12.1	324	0.169	12.2	617	0.310	8.8	470	0.991	11.7	227	0.072	14.1	478	4
	08/11/2016	0.364	12.3	331	0.326	12.2	568	0.422	8.1	468	1.835	11.6	233	0.241	13.8	544	8
	15/11/2016	0.175	12.3	321	0.092	11.9	573	0.333	8.0	500	1.508	11.5	225	0.221	13.1	475	7
	22/11/2016	0.200	12.2	325	0.097	11.6	566	0.343	8.1	530	1.662	11.5	225	0.228	12.8	470	7
	29/11/2016	0.178	12.1	322	0.070	11.4	587	0.371	7.4	537	1.612	11.4	224	0.210	12.5	464	7
	09/12/2016	0.164	12.1	323	0.050	11.2	572	0.392	6.9	539	1.750	11.4	222	0.189	12.3	478	10
	14/12/2016	0.148	12.0	325	0.038	11.0	532	0.628	7.2	525	1.760	11.4	220	0.176	12.1	474	5
	22/12/2016	0.157	12.0	314	0.097	10.6	528	0.526	6.0	543	1.737	11.4	214	0.195	11.9	509	8
30/12/2016	0.149	11.9	326	0.063	10.4	529	0.353	4.7	540	1.615	11.2	221	0.152	11.4	470	8	
Fall average		0.176	12.1	325	0.127	11.6	579	0.444	7.6	502	1.490	11.5	224	0.168	13.0	484	
Winter	10/01/2017	0.130	11.5	321	0.042	9.8	508	0.427	4.3	528	1.458	11.1	219	0.130	10.8	466	11
	17/01/2017	0.121	11.4	319	0.031	9.4	498	0.410	4.5	500	1.372	11.1	216	0.120	10.5	464	7
	31/01/2017	0.111	11.4	304	0.020	9.2	468	0.405	5.1	480	1.229	11.1	207	0.112	10.1	443	14
	07/02/2017	0.380	11.5	324	1.054	8.5	480	3.720	5.1	435	2.030	11.1	223	0.531	8.7	522	7
	14/02/2017	0.190	11.5	321	0.380	8.9	470	0.913	6.1	478	1.608	11.0	213	0.236	10.1	453	7
	21/02/2017	0.153	11.4	310	0.113	8.7	466	0.843	6.8	489	1.501	11.0	215	0.223	10.2	442	7
	28/02/2017	0.133	11.3	303	0.093	8.8	466	0.750	7.3	472	1.431	11.0	215	0.236	10.4	438	7
	07/03/2017	0.163	11.2	313	0.117	8.8	506	1.080	7.3	499	1.537	10.9	218	0.257	10.3	477	7
	15/03/2017	0.168	11.2	319	0.134	8.8	502	0.793	7.5	505	1.502	10.9	221	0.209	10.7	446	8
	20/03/2017	0.147	11.2	317	0.125	9.0	508	0.653	8.1	504	1.466	11.0	218	0.203	10.8	444	5
30/03/2017	0.142	11.1	322	0.122	9.0	524	0.832	9.4	509	1.458	10.9	220	0.193	11.5	457	10	
Win. average		0.167	11.3	316	0.203	9.0	491	0.984	6.5	491	1.508	11.0	217	0.223	10.4	459	
Spring	06/04/2017	0.130	11.1	321	0.110	9.2	514	0.675	9.8	514	1.431	10.9	219	0.182	11.3	451	7
	11/04/2017	0.124	11.1	318	0.091	9.3	513	0.580	9.8	506	1.374	11.0	216	0.169	11.4	436	5
	18/04/2017	0.119	11.1	317	0.080	9.9	509	0.488	9.9	509	1.337	11.0	215	0.152	11.6	437	7
	28/04/2017	0.116	10.9	316	0.070	9.6	507	0.368	9.4	498	1.256	10.9	215	0.136	11.8	438	10
	08/05/2017	0.199	11.0	322	0.188	10.2	518	0.686	10.4	515	1.417	11.0	218	0.180	12.0	455	10
	12/05/2017	0.155	11.0	322	0.099	10.4	510	0.612	10.6	511	1.357	11.0	216	0.141	12.1	448	4
	19/05/2017	0.120	11.1	320	0.054	10.3	504	0.450	12.0	495	1.301	11.1	216	0.125	12.4	438	7
	26/05/2017	0.096	11.1	318	0.030	10.4	501	0.305	13.5	480	1.251	11.1	215	0.110	12.8	430	7
	31/05/2017	0.085	11.3	318	0.020	10.8	498	0.213	14.6	465	1.199	11.1	214	0.104	13.0	425	5
	12/06/2017	0.071	11.3	320	0.011	11.6	488	0.100	17.8	444	1.162	11.1	212	0.091	13.5	433	12
26/06/2017	0.058	11.3	321	0.008	12.2	476	0.056	16.6	428	1.109	11.2	209	0.088	14.0	444	14	
30/06/2017	0.160	11.4	327	0.143	12.5	520	0.383	17.0	448	1.801	11.2	223	0.118	14.4	450	4	
Spr. average		0.119	11.1	320	0.075	10.5	505	0.410	12.6	484	1.333	11.1	216	0.133	12.5	440	
Summer	07/07/2017	0.118	11.4	325	0.068	12.8	505	0.051	17.5	421	1.376	11.3	215	0.096	14.8	443	7
	15/07/2017	0.126	11.6	325	0.077	12.8	508	0.090	17.8	422	1.421	11.3	218	0.080	15.0	440	8
	23/07/2017	0.114	11.6	328	0.039	12.8	500	0.056	18.7	419	1.367	11.4	215	0.067	15.3	436	8
	25/07/2017	0.138	11.8	334	0.051	12.9	509	0.210	17.9	440	1.475	11.4	216	0.077	15.7	433	2
	10/08/2017	0.060	12.1	332	0.023	13.2	506	0.100	18.5	434	1.319	11.5	215	0.067	16.0	442	16
	07/09/2017	0.044	12.1	326	0.013	12.8	519	0.031	17.6	424	1.155	11.5	216	0.056	16.3	438	28
Sum. average		0.100	11.8	328	0.045	12.9	508	0.090	18.0	427	1.352	11.4	216	0.074	15.5	439	
Average		0.146	11.6	321	0.121	10.7	522	0.535	10.3	483	1.428	11.2	218	0.160	12.5	458	

Table S2 – Discharge (Q), temperature (T) and electrical conductivity (EC) measurements at the springs of Mt. Zirone along the hydrogeologic year 2016-2017 (the Q measures selected for summer recession analysis are highlighted in red).

S2. Hydrographs of the springs of Mt. Prinzera and Mt. Zirone

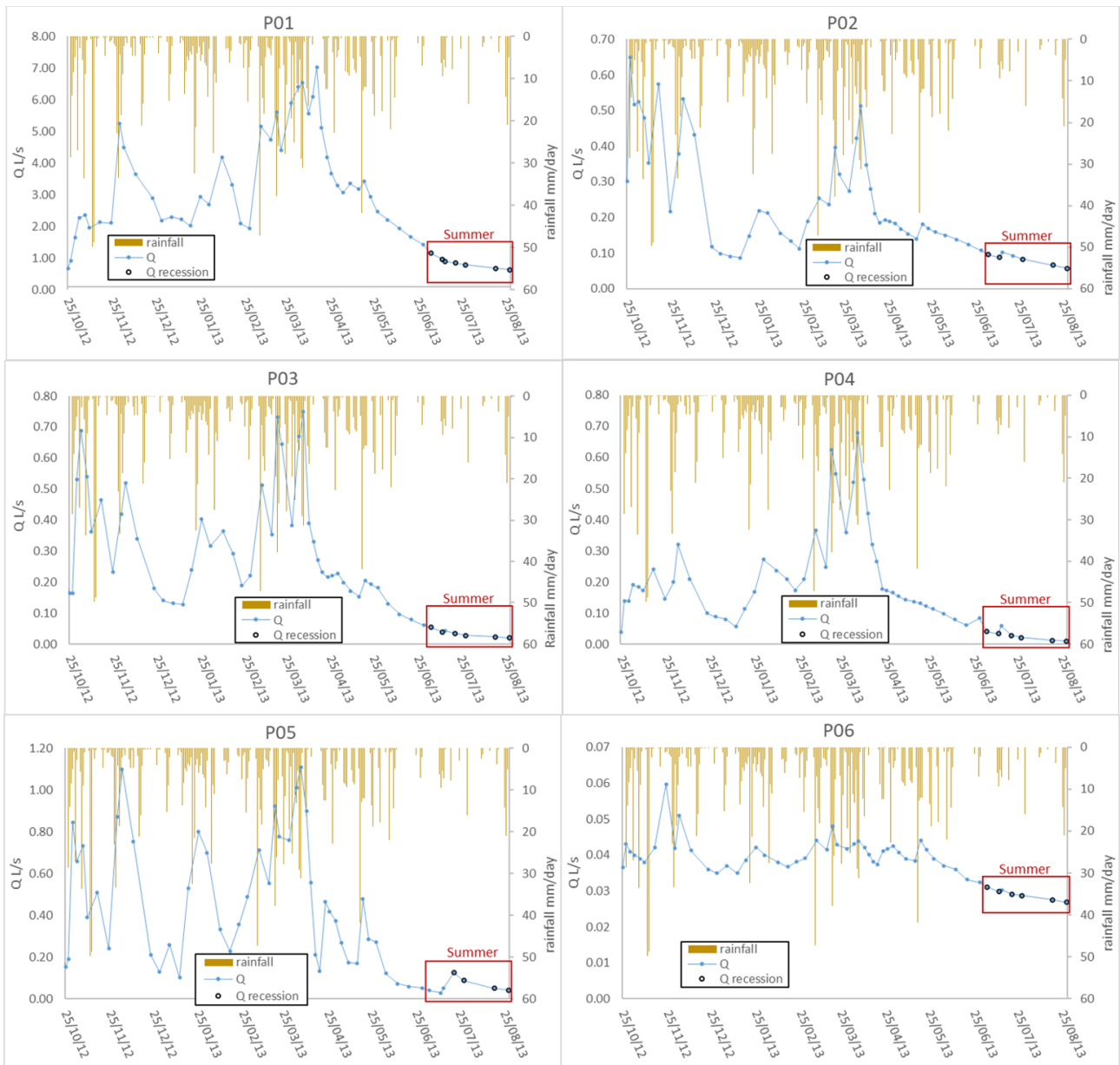


Figure S1 (I)- Hydrographs of the springs of Mt. Prinzera and Mt. Zirone. The summer part the hydrographs analyzed in Fig. 5 of the main article is highlighted in red. The maximum daily rainfall values among the ones registered at the selected monitoring stations (see Section S3 below) are reported on each hydrograph.

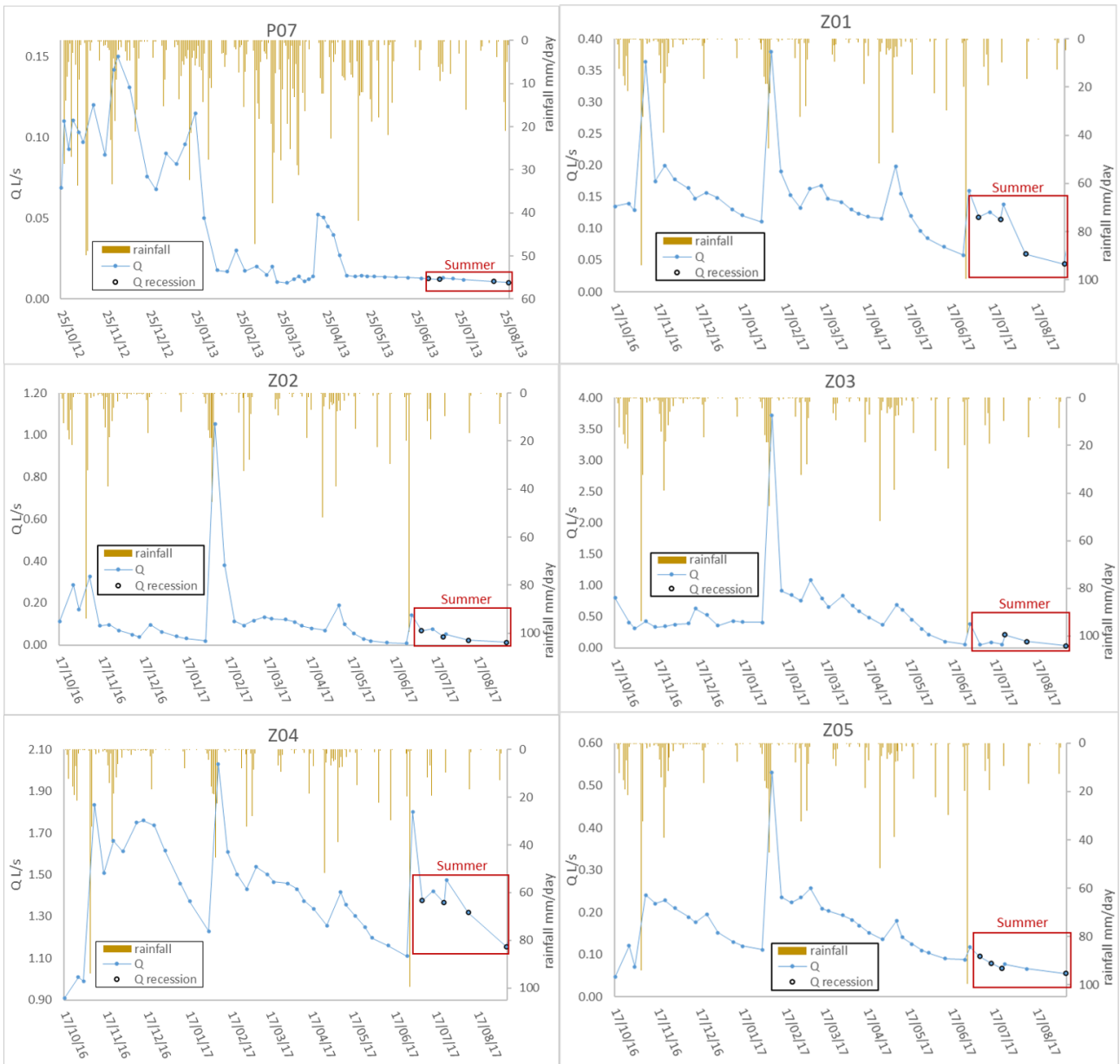


Figure S1 (II)- Hydrographs of the springs of Mt. Prinzera and Mt. Zirone. The summer part the hydrographs analyzed in Fig. 5 of the main article is highlighted in red. The maximum daily rainfall values among the ones registered at the selected monitoring stations (see Section S3 below) are reported on each hydrograph.

S3. Estimation of annual precipitation along hydrogeologic years 2012-2013 (Mt. Prinzerza) and 2016-2017 (Mt. Zirone)

Daily precipitation and air temperature data were acquired from several meteorological stations (rain and temperature gauges) of the Hydro-meteorological Service of the Environmental Protection Agency for the Emilia Romagna Region (ARPAE) (Fig. S2).

In order to estimate the annual precipitation P over the catchments of Mt. Prinzerza and Mt. Zirone, a linear relationship was identified in the two areas between P and the elevation of selected rain gauges (Fig. S3). Three and four rain gauges were selected for Mt. Prinzerza and Mt. Zirone, respectively, among the ones active in the hydrological year of the survey (2012-2013 or 2016-2017). The selected gauges are located at different elevations between 169 and 808 m a.s.l. (Tab. S3).

The areas of Mt. Prinzerza and Mt. Zirone were split into four altimetric belts with a 100 m altitude spacing. The rainfall volume corresponding to each belt was determined by multiplying the belt surface for the value of P corresponding to the averaged belt elevation. The sum of P volumes from the different belts was divided by the total area to obtain a value of annual P representative for the whole massif (" P_{TOT} "; Tab. S4).

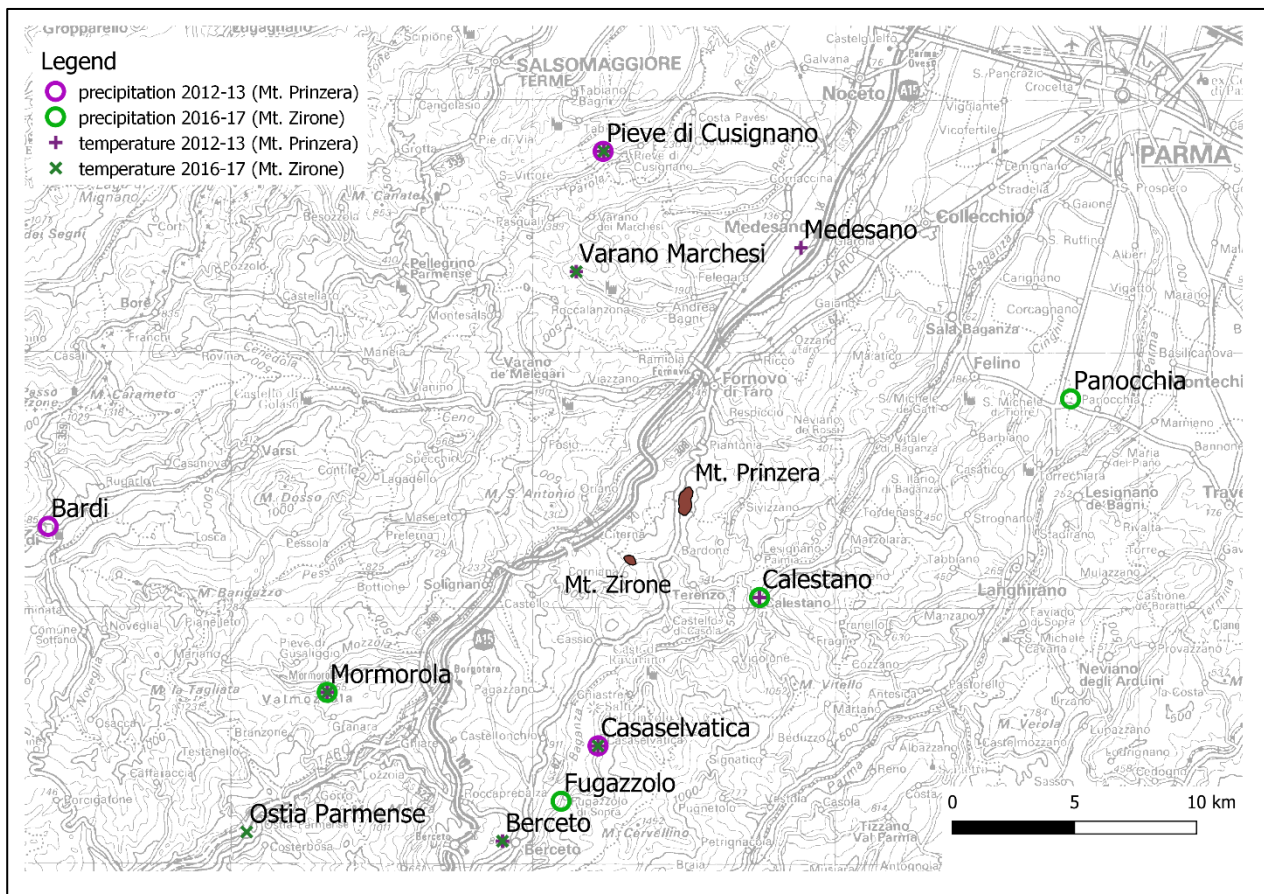


Figure S2 – Location of the meteorological stations selected among the ones active in the hydrogeologic years 2012-2013 and 2016-2017 for daily precipitation and temperature.

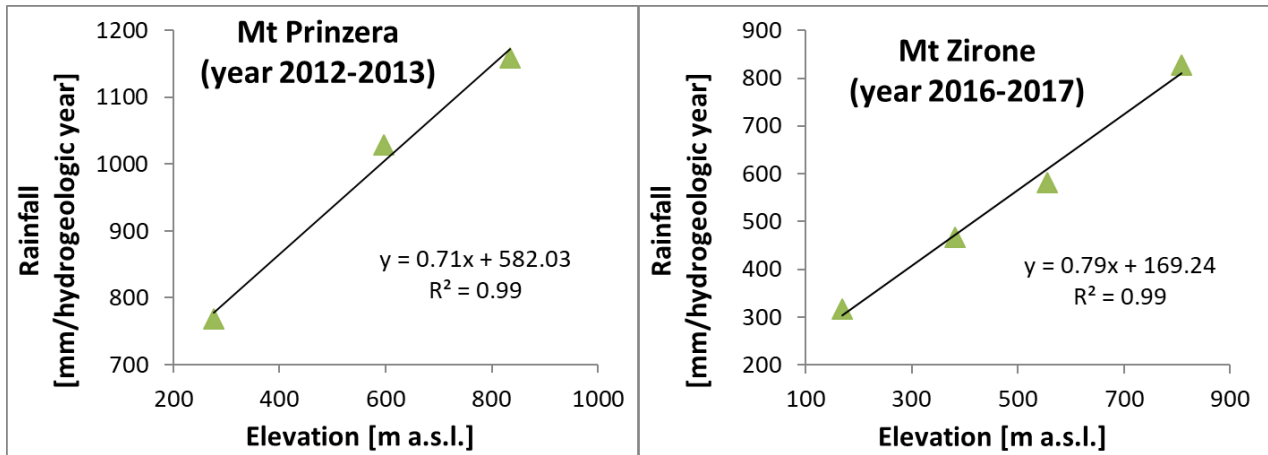


Figure S3 – Linear relationship between the elevation of the meteorological stations and the annual rainfall in the hydrogeologic year of interest.

	Meteorological station (ARPAE)	Elevation	Precipitation [mm]*
Mt. Prinzero	Pieve di Cusignano	277	768
	Bardi	597	1028
	Casaselvatica	834	1159
Mt. Zirone 2016-17	Pannocchia	169	317
	Calestano	381	468
	Mormorola	556	582
	Fugazzolo	808	828

* total precipitation over the monitored hydrologic year

Table S3 – Ground elevation at the selected meteorological stations and annual rainfall during the hydrogeologic year of interest.

	Altimetric belt	Precipitation [mm]	Belt area [m ²]	Precipitation volume [m ³]	P _{TOT} [mm]
Mt. Prinzero 2012-13	A1 (400-500)	900.45	27721.00	24961.37	1010.8
	A2 (500-600)	971.21	324291.00	314954.66	
	A3 (600-700)	1041.97	372578.00	388215.10	
	A4 (700-735)	1102.12	46888.00	51676.02	
Mt. Zirone 2016-17	A1 (400-500)	525.82	32000.00	16826.24	631.94
	A2 (500-600)	605.06	290000.00	175467.40	
	A3 (600-700)	684.30	210000.00	143703.00	
	A4 (700-707)	729.47	2000.00	1458.93	

Table S4 – Altimetric belt areas and estimation of R_{TOT} for Mt. Prinzero and Mt. Zirone. Calculations are described in Section 3.3 of the main text.

S4. Hydrograph analysis by means of “recession plots”

A well-known hydrological method to examine recession hydrographs of streams is to plot the rate of change in discharge (dQ/dt) versus the mean discharge over the dt interval (Q). This kind of plot is also known as “recession plot”. The method was first proposed by Brutsaert and Nieber (1977) to avoid picking the exact time at which recession begins, and further investigated e.g. by Mendoza et al. (2003), Shaw and Riha (2012), Troch et al. (2013).

The recession plot method has been here applied to the spring hydrographs of Mt. Prinzerera and Mt. Zirone as an alternative to the Maillet model for the analysis of recession.

We built a recession plot for each spring considering the whole depletion hydrograph, i.e. from the maximum peak of discharge (April and February in the cases in Mt. Prinzerera and Mt. Zirone, respectively) down to the end of the hydrologic recession (end of August in both aquifers) (Fig. S4). We removed secondary discharge peaks along the falling limb of the hydrograph by taking into account only decreasing discharge values (i.e. negative values of dQ/dt). The correlation coefficient of the individual recession plots is much higher for the Mt. Prinzerera springs (R^2 of 0.71 on average) compared to Mt. Zirone springs (R^2 of 0.54 on average).

The averaged value of the slopes (S) of the recession plots is 2.0 for the springs of both aquifers, consistently with the values obtained by Shaw and Riha (2012) that analyzed individual recession events in streams.

A subdivision in “recession classes” is proposed in the main text that is based on the value of the recession coefficient α of Maillet. The value of S of each spring was chosen as the parameter to be compared with the α of the same spring to assess if the subdivision in recession classes proposed in the main text was maintained when analyzing spring hydrographs with the recession plot method instead of the Maillet model. The comparison between α and S is shown and discussed in the main text.

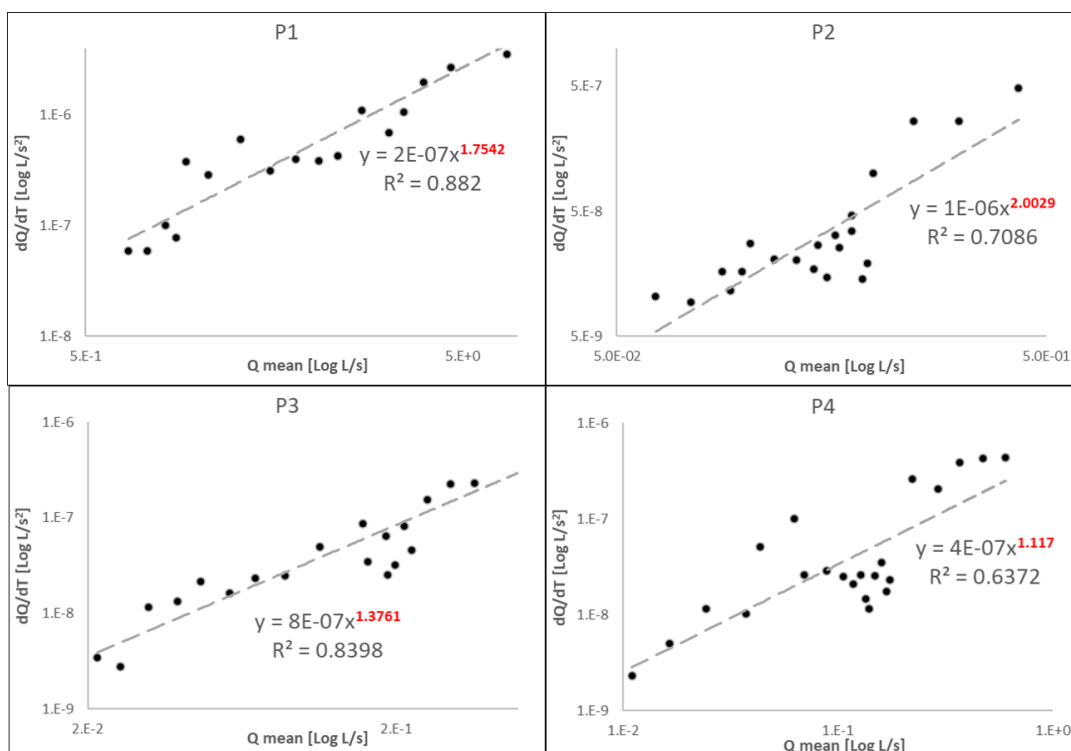


Figure S4 (I) – Recession plots from the depletion hydrographs of the springs of Mt. Prinzerera and Mt. Zirone. The value of S is highlighted in red.

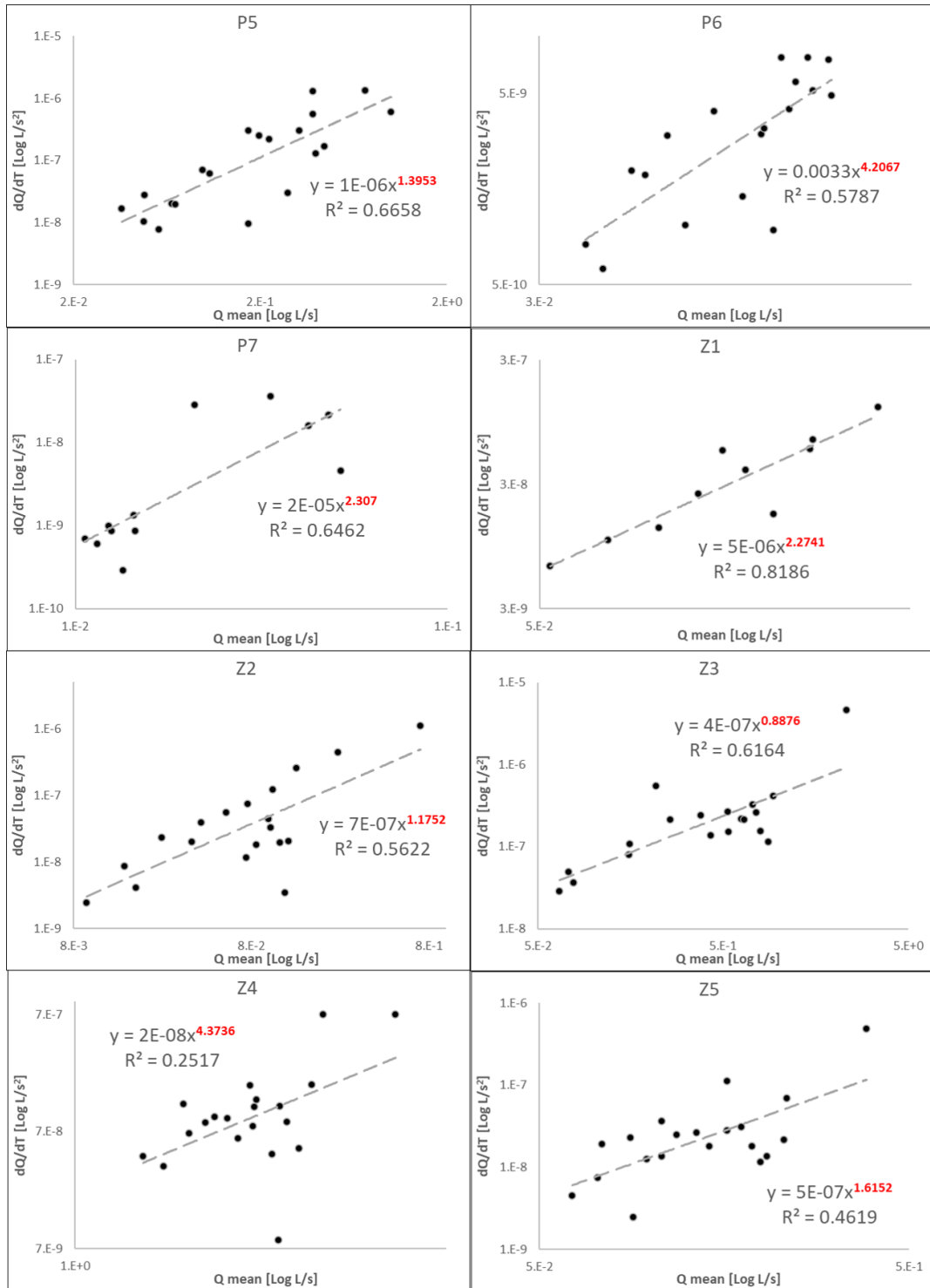


Figure S4 (II) – Recession plots from the depletion hydrographs of the springs of Mt. Prinzera and Mt. Zirone. The value of S is highlighted in red.

S5. Estimation of recharge at Mt. Prinzerza and Mt. Zirone from annual water budgeting

The aquifer recharge R_{wb} of Mt. Prinzerza and Mt. Zirone was estimated through a water budget equation (S1) for the two monitored years (2012-2013 and 2016-2017, respectively):

$$(S1) \quad R_{wb} = (P_{TOT} - ET) \times CPI$$

where CPI is the Coefficient of Potential Infiltration that accounts for loss of recharge due to runoff and other minor processes (Civita, 2005) and ET is the annual evapotranspiration estimated using the Turc equation (Turc, 1951) (S2):

$$(S2) \quad ET = P_{TOT} / \sqrt{0.9 + P_{TOT}^2/L^2}$$

where L is a “thermal indicator” that depends on mean annual air temperature (T) and is defined by (S3)

$$(S3) \quad L = 300 + 25 \times T + 0.05 \times T^3$$

In order to estimate T over the catchments of Mt. Prinzerza and Mt. Zirone, a linear relationship was identified in the two areas between the mean annual air T at a meteorological station and the elevation of the station (Fig. S5). Seven and six stations were selected for Mt. Prinzerza and Mt. Zirone, respectively, among the ones active in the hydrological year of the survey (2012-13 or 2016-2017) (Fig. S2). The selected stations are located at different elevations between 104 and 834 m a.s.l. (Tab. S5). A mean annual air temperature was assigned to each altimetric belt (Tab. S6). The mean annual air T for the entire catchment was estimated as the average of mean annual temperatures assigned to each belt weighted on the belt areas (“averaged air T” in Tab. S6).

ET was estimated equal to 543 and 486 mm at Mt. Prinzerza and Mt. Zirone, respectively, corresponding to the 54 and 77% of P_{TOT} .

Typical ranges of CPI are suggested by (Civita (2005)) for different lithologies. In the case of Mt. Prinzerza, a CPI in the mid range of fissured plutonites (25%) allowed estimating a R_{wb} of 117 mm that fits well the R estimated in the main text (see Fig. S6). At Mt. Zirone, a much higher CPI (87%) had to be considered to obtain a good fit between R_{wb} and R, with R_{wb} of 127 mm. Such high value of the coefficient is uncommonly observed for the investigated lithologies. However, several reasons are discussed in the next Section 6 that would justify a much higher infiltration potential at Mt. Zirone compared to Mt. Prinzerza.

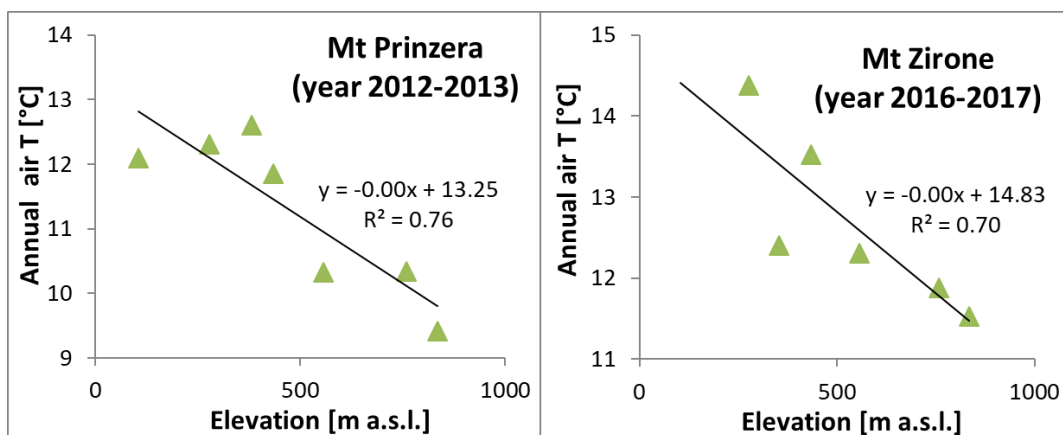


Figure S5 – Linear relationship between the elevation of the meteorological stations and the mean annual air temperature in the hydrogeologic year of interest.

	Meteorological station (ARPAE)	Elevation	T [°C]
Mt. Prinzera 2012-13	Pieve di Cusignano	277	12.30887
	Varano Marchesi	434	11.84751
	Mormorola	556	10.33055
	Berceto	758	10.3336
	Calestano	381	12.59756
	Medesano	104	12.09259
	Casaselvatica	834	9.409191
Mt. Zirone 2016-17	Pieve di Cusignano	277	14.37936
	Ostia Parmense	354	12.40505
	Varano Marchesi	434	13.52554
	Mormorola	556	12.3085
	Berceto	758	11.88294
	Casaselvatica	834	11.52532

* mean annual air temperature at the gauge

Table S5 – Ground elevation at the selected meteorological stations and mean annual air temperature during the hydrogeologic year of interest.

	Altimetric belt	annual air T [°C]	Belt area [m ²]	averaged air T [°C]
Mt. Prinzera 2012-13	A1 (400-500)	11.41	27721.00	10.76
	A2 (500-600)	11.00	324291.00	
	A3 (600-700)	10.59	372578.00	
	A4 (700-735)	10.18	46888.00	
Mt. Zirone 2016-17	A1 (400-500)	13.03	32000.00	12.49
	A2 (500-600)	12.63	290000.00	
	A3 (600-700)	12.23	210000.00	
	A4 (700-707)	11.83	2000.00	

Table S6 – Estimation of averaged air T at Mt. Prinzera and Mt. Zirone as the average of the mean annual temperature at different elevations weighted on belt areas of the same elevation.

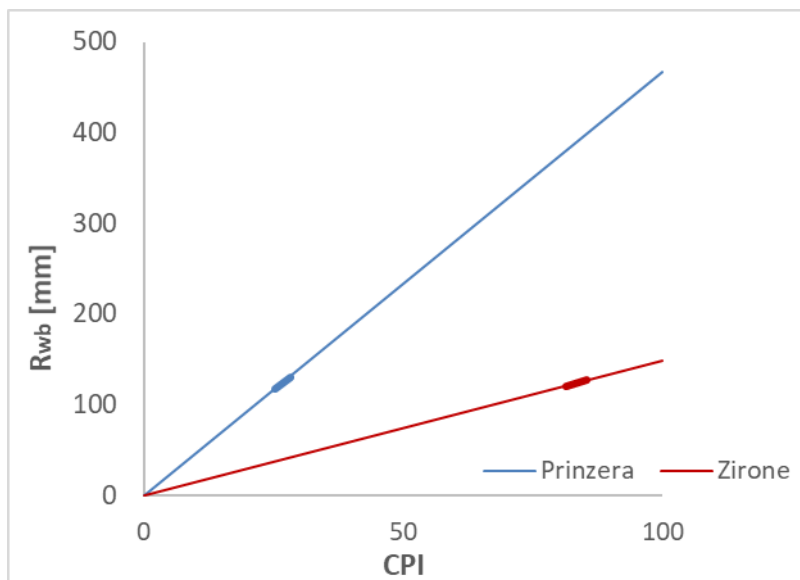


Figure S6 – R_{wb} variation at varying CPI. The thicker segments identify the R values estimated in the main text.

S6. Surface evidences for the infiltration potential at Mt Prinzera and Mt. Zirone

Field observations on surface morphology, Quaternary cover and fracturing were conducted at Mt. Prinzera and Mt. Zirone providing insights into the infiltration potential of the two aquifers.

The surface of Mt. Zirone is interested by a large number of rock slope deformations (*sensu* Hungr et al., 2014) likely set on the preexisting tectonic structures (Fig. S7). Such generalized stress-release condition is expected to enhance infiltration of recharging water from the topographic surface into the aquifer. Infiltration is likely exacerbated due to the low thickness of the aquifer (up to 150 m, which is half the maximum thickness of Mt Prinzera aquifer). In contrast, the olistolithe of Mt. Prinzera appears rather intact, with only two rock slope deformations in its northernmost and southernmost edges (Fig. S7). The different degrees of structural relaxation characterizing the two areas are also well discernible in the field, where the fractures at Mt. Zirone appear much wider than that of Mt. Prinzera (Fig. S8). The above observations suggest a lower infiltration potential at Mt Prinzera compared to Mt Zirone.

Preliminary structural surveys were performed at Mt. Prinzera and Mt. Zirone in summer 2014 and 2015, respectively, along two 20 m long scan lines. Whereas the total number of fractures along the scan line was similar in the two cases (1961 at Mt. Prinzera and 2443 at Mt. Zirone), the aperture and persistence of fractures was much lower at Mt. Prinzera, corroborating the hypothesis of higher infiltration potential and overall higher permeability at Mt. Zirone (Fig. S9).

The Quaternary covers in the areas of Mt. Prinzera and Mt. Zirone were analyzed during a field survey from March 2011 to October 2011, using the Technical Regional Map as a base map (1:5000), integrated with aerial photographs (Tab. S7). The results allowed refining the available geological map (Di Dio et al., 2005) based on CARG (Italian Geological Cartography Project) data. The Quaternary deposits consist of eluvial and colluvial deposit, residual cover, active and dormant landslide and mass movements. The woodland cover mostly consists of scattered oak trees and juniper shrubs (Corticelli et al., 2011). The areas not intersected by woodland or Quaternary covers were classified as bedrock outcrop and assumed as the areas contributing most actively to groundwater recharge due to easier recharge infiltration. The overall lower percent of bedrock outcrop over the total aquifer surface at Mt. Prinzera (60%) compared to Mt. Zirone (69%), may contribute to the hypothesized higher infiltration at Mt. Zirone.

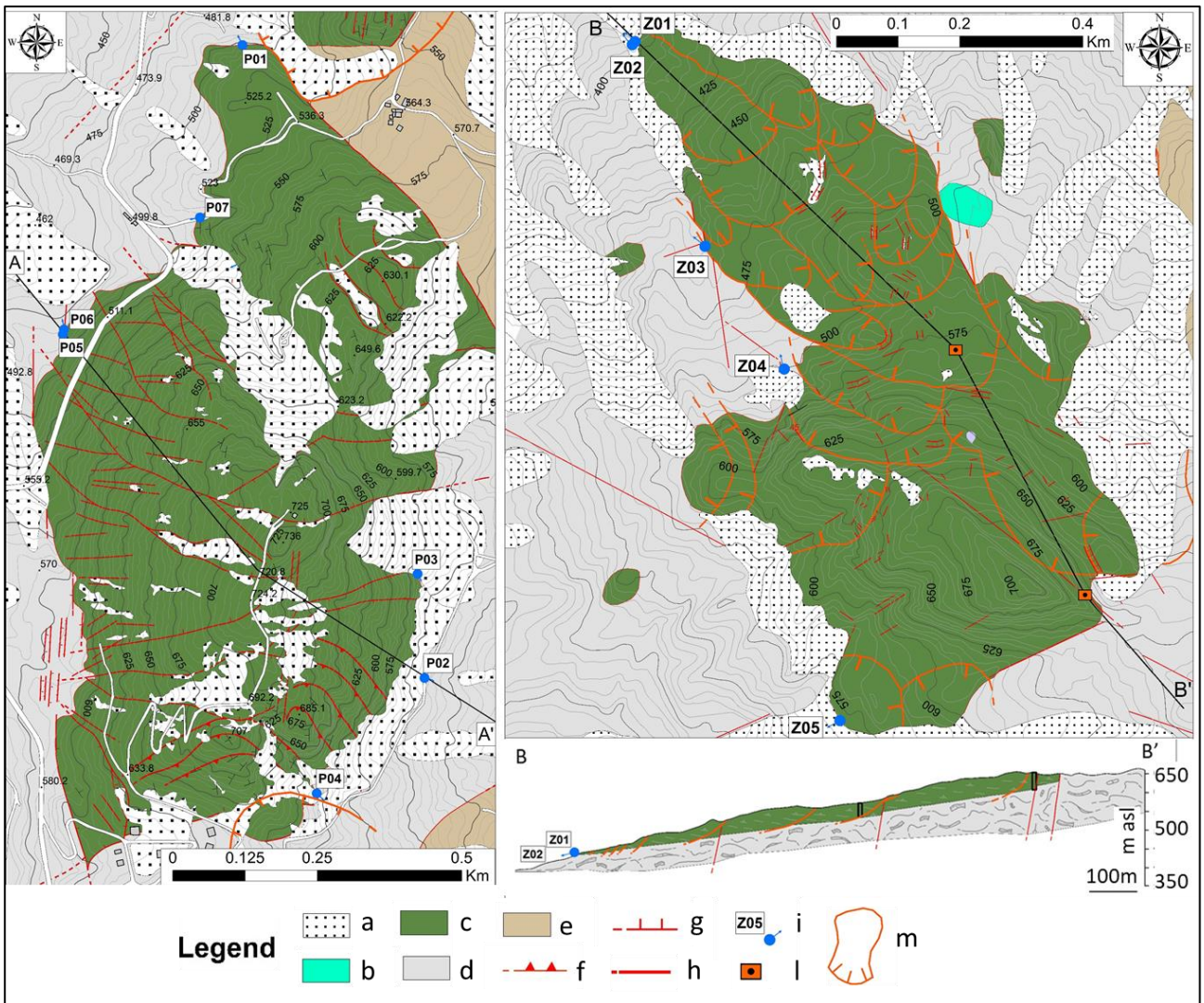


Figure S7 - Geological sketch maps of Mt. Prinzer (left) and Mt. Zirone (right) a: Quaternary deposits; b: Calpionella limestones; c: ophiolite hard-rock aquifers; d: polygenic breccias in clay matrix (aquitard); e: Helminthoid flysch; f: thrust; g: fault (the teeth indicates the downwards moved side); h: tectonic contact; i: perennial spring; l: borehole; m: rock slope deformation boundary (the teeth indicates the downwards moved side).

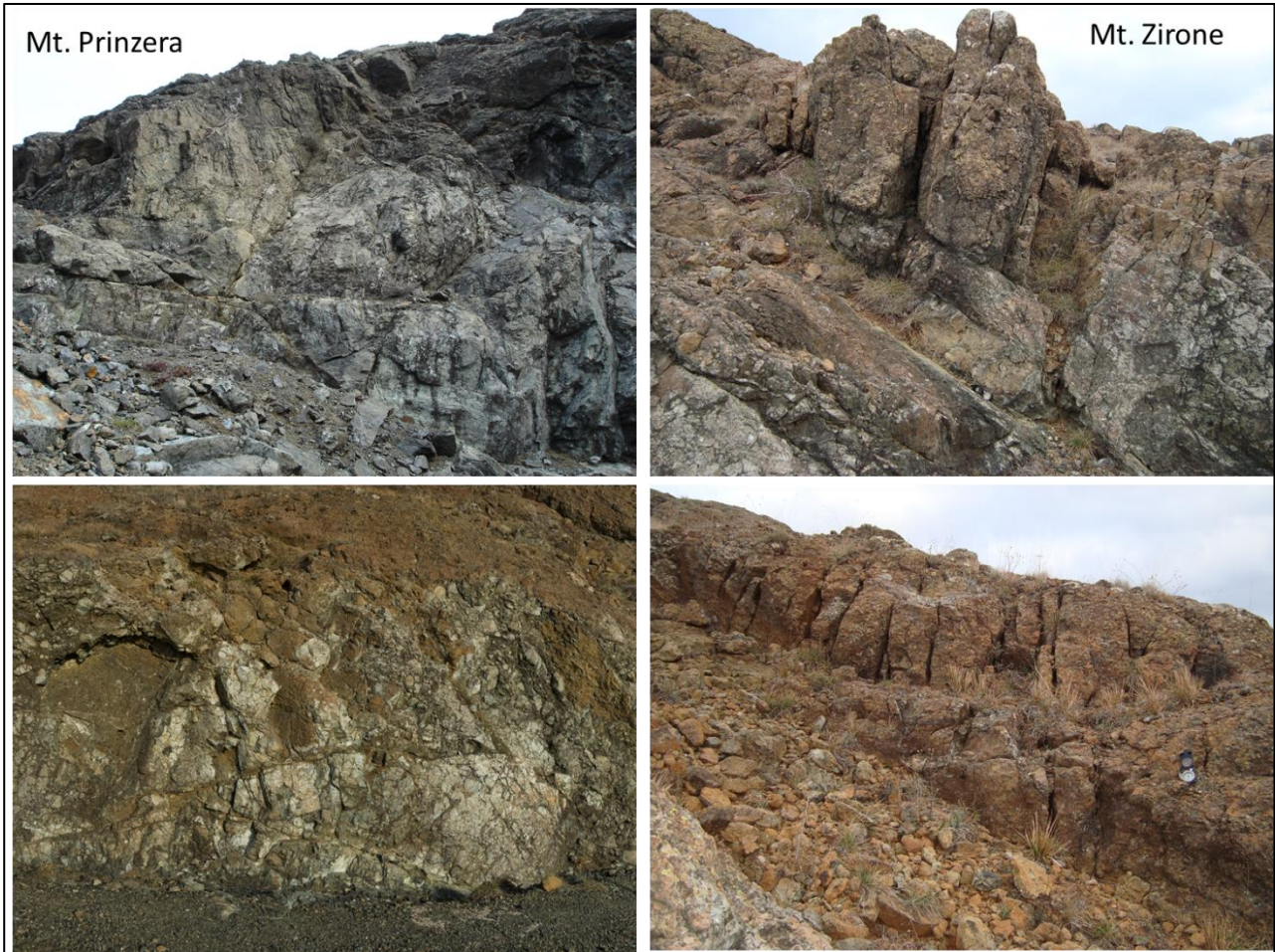


Figure S8 – Pictures from Mt. Prinzer (left) and Mt. Zirone (right) highlighting different stress-release conditions.

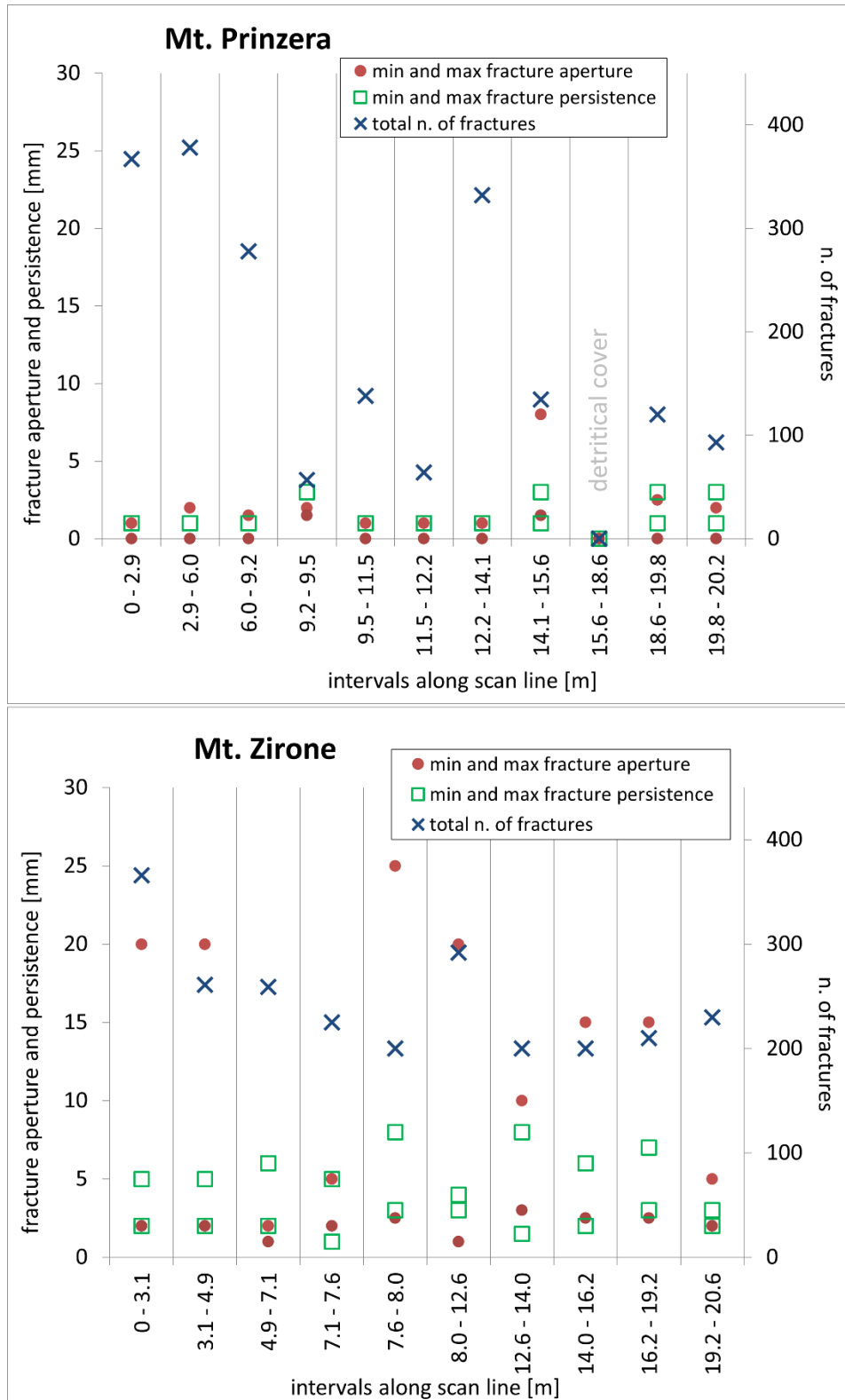


Figure S9 – Results of a preliminary structural survey along 20 m long scan lines at Mt. Prinzera and Mt. Zirone.

			% of total area		% of covered area	
	Mt. Prinzerà	Mt. Zirone	Mt. Prinzerà	Mt. Zirone	Mt. Prinzerà	Mt. Zirone
Aquifer area (m²)	771478	534000				
Bed rock outcrop or scattered trees (m²)	462427	367726	59.94	68.86		
Total cover (m²)	309051	166274	40.06	31.14		
Quaternary deposit cover (m²)	148691	25730	19.27	4.82	48.11	15.47
Woodland cover (m²)	160360	140544	20.79	26.32	51.89	84.53

Table S7- Woodland and Quaternary covers and bedrock outcrop at Mt. Prinzerà and Mt. Zirone.

ESM References:

- Brutsaert, W. and Nieber, J.L., 1977. Regionalized drought flow hydrographs from a mature glaciated plateau. *Water Resources Research*, 13(3): 637-643.
- Civita, M., 2005. *Idrogeologia applicata e ambientale (Applied and Environmental Hydrogeology)*. In Italian). Casa editrice ambrosiana, 794 pp.
- Corticelli, S., Garberi, M.C., Mariani, M.C. and Masi, S., 2011. *Usò del suolo 2008 (Land use in 2008)*, Community Network of the Emilia-Romagna Region.
- Di Dio, G., Martini, A., Lasagna, S. and Zanzucchi, G., 2005. Note illustrative della Carta Geologica d'Italia alla scala 1:50.000, Foglio n° 199 Parma Sud-Ovest (Explanatory notes of the Geologic Map of Italy at the scale 1:50.000, Sheet n° 199 Parma Sud-Ovest. In Italian).
- Hungr, O., Leroueil, S. and Picarelli, L., 2014. The Varnes classification of landslide types, an update. *Landslides*, 11(2): 167-194.
- Mendoza, G.F., Steenhuis, T.S., Walter, M.T. and Parlange, J.Y., 2003. Estimating basin-wide hydraulic parameters of a semi-arid mountainous watershed by recession-flow analysis. *Journal of Hydrology*, 279(1): 57-69.
- Shaw, S.B. and Riha, S.J., 2012. Examining individual recession events instead of a data cloud: Using a modified interpretation of $dQ/dt-Q$ streamflow recession in glaciated watersheds to better inform models of low flow. *Journal of hydrology*, 434: 46-54.
- Troch, P.A. et al., 2013. The importance of hydraulic groundwater theory in catchment hydrology: The legacy of Wilfried Brutsaert and Jean-Yves Parlange. *Water Resources Research*, 49(9): 5099-5116.
- Turc, L., 1951. Nouvelles formule pour le bilan de Peau en fonction des valeurs moyennes annuelles des précipitations et de la temperature: *Comptes Rendus de l'Academie Sciences (New Formulas for the Estimation of Runoff, Using the Average Annual Values of Temperature and Precipitation)*. *Proc. Natl. Acad. Sci. USA*, 233: 633-635.