

Alma Mater Studiorum Università di Bologna
Archivio istituzionale della ricerca

Factors influencing aerosol and precipitation ion chemistry in urban background of Moscow megacity

This is the final peer-reviewed author's accepted manuscript (postprint) of the following publication:

Published Version:

Zappi Alessandro, Popovicheva Olga, Tositti Laura, Chichaeva Marina, Eremina Irina, Kasper-Giebl Anne, et al. (2023). Factors influencing aerosol and precipitation ion chemistry in urban background of Moscow megacity. *ATMOSPHERIC ENVIRONMENT*, 294, 1-17 [10.1016/j.atmosenv.2022.119458].

Availability:

This version is available at: <https://hdl.handle.net/11585/905317> since: 2024-01-29

Published:

DOI: <http://doi.org/10.1016/j.atmosenv.2022.119458>

Terms of use:

Some rights reserved. The terms and conditions for the reuse of this version of the manuscript are specified in the publishing policy. For all terms of use and more information see the publisher's website.

This item was downloaded from IRIS Università di Bologna (<https://cris.unibo.it/>).
When citing, please refer to the published version.

(Article begins on next page)

This is the final peer-reviewed accepted manuscript of:

Alessandro Zappi, Olga Popovicheva, Laura Tositti, Marina Chichaeva, Irina Eremina, Anne Kasper-Giebl, Ying I. Tsai, Dmitry Vlasov, Nikolay Kasimov, Factors influencing aerosol and precipitation ion chemistry in urban background of Moscow megacity, Atmospheric Environment, Volume 294, 2023, 119458, ISSN 1352-2310,

<https://doi.org/10.1016/j.atmosenv.2022.119458>.

The final published version is available online at:
<https://www.sciencedirect.com/science/article/pii/S1352231022005234?via%3Dihub>

Terms of use:

Some rights reserved. The terms and conditions for the reuse of this version of the manuscript are specified in the publishing policy. For all terms of use and more information see the publisher's website.

This item was downloaded from IRIS Università di Bologna (<https://cris.unibo.it/>)

When citing, please refer to the published version.

Factors influencing aerosol and precipitation ion chemistry in urban background of Moscow megacity

Alessandro Zappi¹, Olga Popovicheva², Laura Tositti¹, Marina Chichaeva³, Irina Eremina³, Anne Kasper-Giebl⁴, Ying Tsai⁵, Dmitry Vlasov³, Nikolay Kasimov³

¹ Department of Chemistry “Giacomo Ciamician”, University of Bologna, Bologna, Italy

² SINP, Lomonosov Moscow State University, Moscow, Russia

³ Faculty of Geography, Lomonosov Moscow State University, Moscow, Russia

⁴ Institute of Chemical Technologies and Analytics, TU Wien, Vienna, Austria

⁵ Department of Environmental Engineering and Science, Chia Nan University of Pharmacy and Science, Tainan 71710, Taiwan

Corresponding authors:

Olga Popovicheva: olga.popovicheva@gmail.com

Lomonosov Moscow State University, Leninskie Gory, 1, 199991, Moscow, Russia

Tel: +7 939 49 54

Laura Tositti: laura.tositti@unibo.it

Via Selmi 2, 40126, Bologna, Italy

Tel: +39 051 2099488

Fax: +39 051 2099456

25 **ABSTRACT**

26 In-depth study of aerosol and precipitation ion chemistry was carried out in order to cover an existing
27 data gap in the seasonal-dependent ionic data in Moscow urban background atmosphere. Literature
28 about atmospheric pollution in this megacity is, indeed, still poor, despite its peculiar climatological
29 characteristics that make it a very interesting site for atmospheric research. Particulate matter (PM)
30 and precipitation were collected at Moscow urban background during spring of 2018, summer, and
31 autumn of 2019. Aerosol inorganic ions are characterized for the whole period, carboxylic acids and
32 sugars for summer. Inorganic ions constitute a major PM₁₀ fraction of 27% in autumn, it drops down
33 to 16.5% in summer. Dominance of secondary inorganic aerosol (SIA) over all the other inorganic
34 ion species is determined. Na⁺ and Ca²⁺ shows the largest variability between seasons, with maximum
35 in summer and spring, respectively. Excess of anions in autumn is indicated by ion balance while
36 adding carbonates, spring and summer aerosol is found electroneutral. Chloride depletion
37 phenomenon follows the snow melt and remobilization of salt applied to roads for deicing,
38 demonstrating the biggest impact in summer. Degree of neutralization of ammonia as sulfate and
39 nitrate is found by the correlation between the concentration of NH₄⁺ and calculated from SO₄²⁻ and
40 NO₃⁻. Inorganic fraction in PM₁₀ is higher (16.5%) than total carboxylic acids (3.5%), and total sugars
41 and anhydrosugars (0.8%). Concentration polar plots show ion variation jointly with wind speed and
42 wind direction and the source origin. Highest SIA concentrations relate a cluster of the prevalent air
43 mass transportation and indicate the source in the northwestern direction in spring and summer. K⁺
44 acts as a marker of fires in spring and summer as well as of the domestic biomass burning in the
45 region around a megacity in autumn. Source apportionment of inorganic ion and BC by Varimax
46 analysis reveals major sources namely SIA, soil resuspension, biomass burning, deicing salt, and
47 fossil fuel combustion. In summer, organic ions, anhydrosugars, and sugars shows additional impact
48 of photochemistry and biological activities. Rainfall is heavier in summer than in autumn, with pH of
49 acid rain during spring and summer. The highest concentrations for almost all ions were observed in
50 spring, except HCO₃⁻. Precipitation composition is described with aerosol wet removal by seasonal
51 variability of SIA scavenging factors, decreasing in summer and increasing in autumn and spring.

52

53 **Keywords:** Moscow; air pollution; rain pollution; Varimax; aerosol ion chemistry

54

55 **1. Introduction**

56 Airborne particulate matter (PM) is one of the atmospheric pollutants of major concern associated
57 not only to major air quality issues in connection with adverse effects on health (WHO, 2013) and
58 the environment, but also with climate change (Fuzzi et al., 2015). PM is a multiphase system
59 including a large number of components, of both primary and secondary origin, and in most cases
60 prone to further physicochemical modifications (Pöschl and Shiraiwa, 2015; Pöschl, 2005; Raes et
61 al., 2000). Owing to the recognized importance of this complex pollutant, great efforts have been
62 applied worldwide in order to monitor and characterize it as well as to fix thresholds and regulations,
63 allowing to prevent or limit adverse effects on human health and environment. The understanding of
64 the processes leading to PM formation and origin requires to associate them to source profiling and
65 quantification. For these reasons PM chemical speciation is of utmost importance since it allows to
66 identify harmful substances as well as sources, providing tools for mitigation and prevention of
67 potential hazards.

68 The major components of PM are elemental and organic carbon, sea salt, crustal materials from
69 soil/road dust, and water soluble species (ammonium, sulfate, and nitrate) known as Secondary
70 Inorganic Aerosol (SIA) (Seinfeld and Pandis, 2016). Among the PM components, ions play a major
71 role both because of their abundance and massive contribution to aerosol mass load. Each ion
72 determined is associated to specific sources of emission and formation pathways (Raes et al., 2000).
73 Due to their water-soluble nature, ions also affect the acidity of precipitation. They are a key in
74 precipitation formation processes, thus contributing to the hydrological cycle (Reiss et al., 2007;
75 Weber et al., 2016; Dadashazar et al., 2019).

76 In European urban areas, ammonium sulfate and nitrate concentrations are abundant (Putaud et al.,
77 2010; Tositti et al., 2014). Dusts (Brattich et al., 2015; Tositti et al., 2021) may contribute to calcium
78 and magnesium, usually in the form of relatively soluble carbonates capable, beside ammonia, to add
79 neutralization potential to acidic species (Morozzi et al., 2021; Tositti et al., 2018a). Coastal areas are
80 affected by sea salt compounds starting from the most abundant, i.e. NaCl, and extending to further
81 saline compounds (Koçak et al., 2007; Lung et al., 2004). In inland cities during the cold period the
82 sources of particulate sodium and chloride are de-icing agents (DIAs) (Kolesar et al., 2018; Kotowski
83 et al., 2020; Stieger et al., 2018; Tositti et al., 2018a).

84 Ambient organic ions are associated with secondary organic aerosol (SOA), mainly produced by
85 oxidation of volatile organic compounds (VOCs) (Kawamura and Bikkina, 2016; Pavuluri et al.,
86 2010; Yang and Yu, 2008). Though at significantly lower concentration level than SIA, they are
87 present in significant “chemodiversity” (Tositti et al., 2018a). The speciation and profiling of SOA is
88 a very hard and complex task (Gentner et al., 2017; Pio et al., 2005; Schauer et al., 2002; Wang et al.,

89 2006), owing to a number of precursors and complexity of oxidation patterns. Oxidation, in fact, is
90 carried out by a series of oxidizing species in the troposphere such as OH and NO₃ radicals, group of
91 peroxy-radicals and ozone (Seinfeld and Pandis, 2016).

92 High polarity of both SIA and SOA is a key in the so-called cloud processing, wherein aerosol
93 particles induce water vapor condensation upon cloud condensation nuclei until water droplets or
94 snow crystals are formed. Precipitation chemistry largely reflects the composition of precursor
95 aerosols (Pruppacher and Klett, 2010) whose degree of neutralization affects precipitation acidity
96 (pH). Acidity of rains constituted an extremely serious environmental problem in Europe which,
97 owing to effective policies of fuel management, led to a substantial solution of acid rain problems
98 from the late 70's to date (Tositti et al., 2018b; Vet et al., 2014; Vuorenmaa et al., 2018).

99 Factors influencing ion chemistry are more important in the urban environment where exposure and
100 hazard to air pollution are sensibly higher, owing to the significant population density and related
101 activities such as industry, transport, and heating etc. In this respect, megacities such as Moscow with
102 over ten million inhabitants require a special focus, since emissions and consequent air quality
103 standards might reveal even more unsafe for human health, material infrastructures as well as cultural
104 heritage (Baklanov et al., 2016).

105 Emission sources in Moscow as a high latitude megacity have several specific features. Centralized
106 heating supply is seasonally operated from the end of September until the beginning of May. Moscow
107 power and heating plants are almost totally supplied with natural gas. Biomass and coal are not used
108 either for industrial, domestic heating, or individual purposes, differently from many European and
109 Asian cities (Diapouli et al., 2017; Nava et al., 2015). Biomass burning (BB) is not recognized by the
110 local Environmental Agency as a significant source of aerosol pollution in Moscow (Kul'bachevskii,
111 2021). However, largely populated region in the city outskirts (Moscovskaya Oblast) applies biomass
112 for many scopes, such as wood burning for domestic stoves, agricultural cleaning work, and summer
113 recreational activities, that should be considered for air quality assessment. Seasonal light absorption
114 measurements supported the assessment for the BB contribution in the region around Moscow in
115 autumn and winter (Popovicheva et al., 2022). Moreover, combined organic and statistical analysis
116 differentiated the impact of agriculture fires in spring (Popovicheva et al., 2020b).

117 Differently from the rest of Europe (Finardi et al., 2014; Font et al., 2019; Tang et al., 2014), the
118 knowledge of Moscow megacity aerosol and its sources is limited. Black carbon (BC) pollution in
119 the Moscow urban background (Popovicheva et al., 2020a) as well as in the Moscow center (Golitsyn
120 et al., 2015) indicated a strong impact of transport. The first comprehensive aerosol characterization
121 in the urban background of Moscow megacity included ions in spring only (Popovicheva et al.,
122 2020c). BB markers (levoglucosan and K⁺ ions) proved the impact of wildfires during the extreme

123 smoke event in August of 2010 (Popovicheva et al., 2014) and agricultural fires (Popovicheva et al.,
124 2020c).

125 Annual records of precipitation in 2012 in Moscow showed the significant increase of its acidity due
126 to the presence of deicing salts (Eremina et al., 2015). The total ionic composition in rains did not
127 change in the last years, while chloride and sodium concentrations were increased. Such finding was
128 associated with DIAs, abundant due to the winter road management in Moscow (Eremina et al., 2015).
129 Despite long-term statistical data for rain acidity and composition (Chubarova et al., 2014), a limited
130 knowledge about factors influencing both particulate and precipitation ions chemistry is available to
131 date for Moscow environment. This situation is further deteriorated by the fact that the Environmental
132 Protection Agency “Mosecomonitoring” conducts the measurements for only PM₁₀ and PM_{2.5} mass
133 concentrations at some monitoring stations, without any PM speciation.

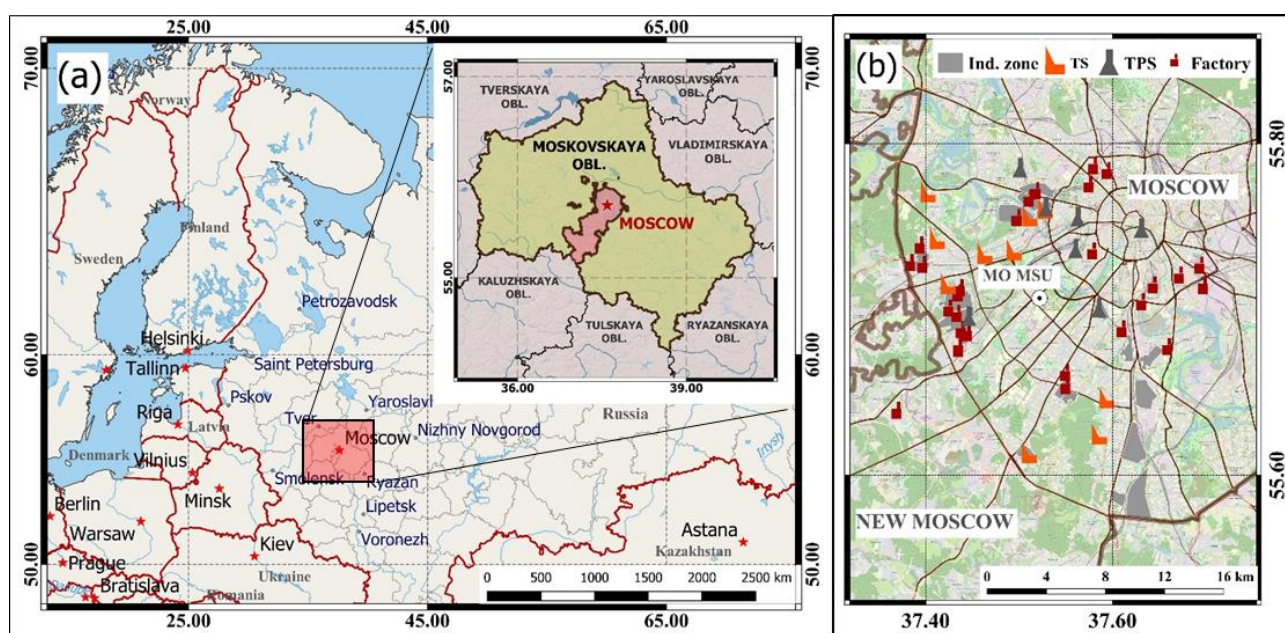
134 This paper therefore aims to cover an existing data gap for the aerosol and precipitation ion chemistry
135 in Moscow urban background. We present the results for the seasonal variation of ionic composition
136 obtained in a sampling campaign carried out in periods of spring 2018, summer and autumn 2019.
137 Descriptive statistics are used to provide PM₁₀ and ion seasonal characteristics as well as the
138 comparison between Moscow and some large cities. Constraints of aerosol electroneutrality
139 highlights the ion chemistry of neutralizing agents such as ammonia and carbonates. Varimax and
140 correlation analyses trace the main sources of aerosol and rain ionic species. Black carbon and
141 levoglucosan are considered together with ionic species for improved source apportionment in
142 summer. The relevance of the aerosol chemical speciation to meteorological parameters and air mass
143 transportation are investigated by using the both the conditional probability function (CPF) based on
144 the bivariate polar plot and the Lagrangian integrated trajectory (HYSPLIT) model with the cluster
145 analyses.

146

147 **2. Materials and methods**

148 **2.1 Study Area**

149 Moscow metropolitan area (55° 45' N; 37° 37' E at the city center; mean altitude 150 m a.sl.) is located
150 in the center of the European part of Russia (map in Fig.1). Moscow covers an area of 2561 km² and
151 has a registered population exceeding 12.5 million, which makes it one of the most densely populated
152 northernmost megacities in Europe. Moscow is characterized by a humid continental climate (Dfb
153 according Köppen climate classification) with ~685 mm precipitation per year (Chubarova et al.,
154 2014) which are represented by snow for about one third of year (mainly from late October to March),
155 and by rain for the other part of the year with a maximum rainfall amount in summer (on average ~61
156 mm in June). According to the record of the Meteorological Observatory (MO) MSU over the 60-
157 year period (1954–2013) (Chubarova et al., 2014), significant warming of regional climate with a
158 positive temperature trend of 0.07°C/year is observed. Southwestern wind (5-10 m s⁻¹) prevails during
159 all seasons (<https://weatherspark.com>). Highest wind speeds are typical of fall and winter. Clean
160 Arctic air masses transported from the North are recorded especially in winter while aged plumes
161 from agriculture and wildfires pollute the Moscow atmosphere in spring and summer. Moderate
162 average air temperatures, low solar UV radiation levels, and good ventilation make the accumulation
163 of primary emitted pollutants and the photochemical formation of secondary pollutants less intensive
164 in Moscow (Elansky et al., 2018). Operation of centralized heating system begins and ends in
165 Moscow when the temperature decreases below 8°C and increases above 8°C, respectively. Thus, the
166 heating period divides the whole year formally on cold and warm with respect to its operation.



167 **Figure 1.** a) Location of Moscow megacity in European part of Russia. Insert is Moscow and
168 Moskovskaya Oblast nearby. b) Southwest sector of Moscow area including New Moscow. Location
169 of sampling site at Meteorological Observatory of Moscow State University (MO MSU) (55° 42' N;
170

171 37° 31' E) is indicated. Industrial zones (Ind.zone), thermal power stations (TPS), thermal stations
172 (TS), and factories in the southwest sector of Moscow are indicated.

173

174 According to air quality characterization, Moscow is classified among the slightly polluted megacities
175 in Europe and North America (Elansky et al., 2018). Traffic exhaust constitutes 93% of the overall
176 gaseous megacity emissions (Bityukova and Mozgunov, 2019). While automobilization is growing
177 in Moscow, intensive implementation of higher ecological classes of engines, a declining share of
178 trucks and better quality of fuel promote the improvement of vehicle ecological parameters.
179 Assuming that currently in Moscow the vehicles are powered with desulphurized fuels, traffic is no
180 longer a significant source of SO₂. However, high concentrations of NO_x and CO are still an issue
181 (Elansky et al., 2014), with the potential contributing significantly to SIA by means of active
182 oxidation capacities of the atmosphere (Prinn, 2003; Seinfeld and Pandis, 2016). Around 50% of all
183 pollutant emissions from industrial sources and thermal power stations are emitted by companies
184 producing and redistributing energy, gas, and water. The currently fast urban expansion especially
185 in New Moscow area (Fig.1) additionally contributes to an aerosol buildup in Moscow.

186

187

188 **2.2. Aerosol and precipitation sampling**

189 The sampling campaign was conducted at the Aerosol Complex located at Meteorological
190 Observatory of Moscow State University (MO MSU) in the MSU campus, southwest of Moscow city
191 (Fig.1). MO MSU is not directly affected by local pollution sources such as large transport roads or
192 industrial facilities, the residential area and a highway takes place about 800 m south of the
193 observatory, and industrial area is situated at about 3 km. MO MSU site can be therefore classified
194 as an urban background (Chubarova et al., 2014).

195 Sampling of particulate matter and precipitation was performed during three seasons: spring 2018
196 (from April 3rd to May 22nd), summer 2019 (from May 29th to July 31st), and autumn 2019 (from
197 September 5th to December 31st). In this work the term “seasons” is used according to the astronomical
198 and not the meteorological convention. PM sampling was not available after December 31st due to
199 the harsh weather conditions which compromised the sampler, preventing the collection of further
200 winter samples. Samples collected in December (11) were merged to the autumn dataset according to
201 the astronomical convention, therefore relating December to autumn as an overall cold season

202 Particles with a diameter less than or equal to 10 µm (PM₁₀) were daily collected over 24-hour periods
203 from 5 p.m. of a given day to 5 p.m. of the next day. In spring and autumn sampling was performed
204 on 47 mm quartz fiber filters (Pall Science, New York, USA), previously heated at 600°C for 5 h.

205 Polyethylene terephthalate 47 mm filters (Merck, Germany) were used for summer sampling. An
206 aspirator using a PM₁₀ impactor was operated at an air flow of 1 m³ h⁻¹ by, and the pumped volume
207 was reported in standard atmospheric conditions. Totally, 35, 56 and 67 aerosol samples were
208 collected in spring, summer, and autumn, respectively.

209 Both types of precipitation (rainfall and snow) were collected at a height of 2 m from the ground
210 surface, using a PVC funnel with an exposed area of 1600 cm² and a polyethylene vessel as a bulk
211 collector. Totally, 16, 23 and 45 precipitation samples were collected in spring, summer, and autumn,
212 respectively.

213 Meteorological parameters (temperature, relative humidity, pressure, precipitation, wind speed, and
214 wind direction) were obtained at a 3 h time resolution by the MO MSU meteorological service. PM₁₀
215 mass concentrations were collected using the tapered element oscillating microbalance TEOM 1400a
216 (Thermo Environmental Instruments Inc., Waltham, MA, USA) by Mosecomonitoring
217 (<https://mosecom.mos.ru/home-page/>). Additionally, continuous aerosol equivalent BC (eBC)
218 concentrations were measured using the custom made portable aethalometer, for more details see
219 elsewhere (Popovicheva et al., 2020a). BC is mainly formed by the incomplete combustion of fossil
220 fuels (diesel, gasoline, gas, coal) in emissions of transport, energy production, residential heating and
221 biomass burning. Therefore, in this study BC is used as a marker of combustion for the source
222 apportionment together with ions. Further in the text we will call eBC as BC according to commonly
223 accepted terminology.

224

225 **2.3 Analytical Procedures**

226 Aerosol samples were analyzed for water-soluble ionic species, namely cations, including sodium
227 (Na⁺), ammonium (NH₄⁺), potassium (K⁺), magnesium (Mg²⁺) and calcium (Ca²⁺), and anions
228 including chloride (Cl⁻), nitrate (NO₃⁻), and sulphate (SO₄²⁻), using ion chromatography as described
229 elsewhere (Golobokova et al., 2020; Popovicheva et al., 2020c; Tsai and Kuo, 2013). For the
230 extraction phase, the polyethylene terephthalate filter was placed in 5.0 mL of ultrapure water, then
231 shaken for 90 min and filtered through an ester acetate syringe membrane of pore size 0.2 μm. Quartz
232 filter were extracted with 8 ml ultrapure water in an ultrasonic bath for 30 min, then the solution was
233 filtered through acetate cellulose filters (porosity 0.2 μm). Anions of spring samples were separated
234 by an ion-exchange column (Ion Pac AS17A, Thermo Fisher, USA) in isocratic conditions using 1.8
235 mM Na₂CO₃ + 1.7 mM NaHCO₃, an AMMS suppressor and conductivity detection. Cations were
236 separated on Ion Pac CS12A using 38 mM methane sulfonic acid (MSA), the CSRS suppressor and
237 a conductivity detector. The detection limits varied between 2 and 20 ng m⁻³, the lowest values were
238 obtained for sodium and chloride.

239 In summer and autumn samples, inorganic cations were determined using a Dionex ICS-1000
240 equipped with a gradient pump (ICS-3000 Model SP-1), Dionex CSRS Ultra 4mm self-regenerating
241 suppressor, conductivity detector, and an Ion Pac CS12A analytical column. Inorganic anions as well
242 those from carboxylic acids (formate, acetate, oxalate, carboxylic di- lactate, succinate, and glutarate)
243 were analyzed using a Dionex ICS-5000+ IC, which was equipped with a gradient pump (ICS-3000
244 Model SP-1) and AMMS 300 MicromembraneTH suppressor and conductivity detector. The guard
245 column and analytical column were Dionex IonPacTM AG11-HC and Dionex IonPacTM AS11-HC,
246 respectively. The MDLs for carboxylates varied between 0.88 and 4.05 ng m⁻³, were lowest for
247 formate and highest for succinate at a nominal sampling volume of 24 m³ for each typical 24 h PM
248 sampling period.

249 Additionally, in summer samples sugars (glucose and mannose) and anhydrosugars (levoglucosan
250 and mannosan) were determined using a Dionex ICS-5000+ IC equipped with a pulsed amperometric
251 detector, a CarboPac MA1 guard column and an anion-exchange analytical column, and a Dionex
252 ICS-5000 Electrochemical Detector. The MDLs for sugars and anhydrosugars varied between 1.59
253 and 3.35 ng m⁻³, were lowest for levoglucosan and highest for glucose at a nominal sampling volume
254 of 24 m³ for each typical 24 h PM sampling period.

255 In precipitation samples the concentrations of SO₄²⁻, Cl⁻, NO₃⁻ and H⁺, Ca²⁺, Mg²⁺, Na⁺, K⁺, NH₄⁺
256 ions as well as hydrogen carbonate (HCO₃⁻) and pH were measured. HCO₃⁻ and pH were analyzed
257 immediately after sampling. Precipitation samples were tested for pH using a pH-meter SP-701
258 (Suntex, Taipei, Taiwan) calibrated by standard pH solutions before measurements. H⁺ concentration
259 was calculated from pH values. Hydrogen carbonate was determined by acidimetric titration with
260 0.005N HCl by an ionmeter (Econics, RF). The concentrations of other cations and anions were
261 analyzed by using ion chromatography on a JetChrom instrument (Portlab, Moscow, Russia).
262 Detection limits varied between 0.02 mg/L for NO₃⁻, Mg²⁺, NH₄⁺, and K⁺, and 0.03 mg L⁻¹ for SO₄²⁺,
263 Na⁺ and 0.05 mg L⁻¹ for HCO₃⁻, Cl⁻, Ca²⁺.

264

265 **2.4. Statistical methods**

266 In order to compare PM₁₀ and ion data from different seasons and check for similarities between
267 seasons and variables, Mann-Whitney test (Mann and Whitney, 1947) and Spearman correlation
268 indexes (Spearman, 2010) were used. We decided to use such tests, instead of the classical Student's
269 *t*-test and Pearson correlation, because these are non-parametric. They allow the robust comparison
270 even in the absence of normally-distributed data, as typically observed for atmospheric variables (Ott,
271 1990). For Mann-Whitney test, compositional data were divided into three groups, based on the three
272 analyzed seasons, and for each variable the comparisons between each couple of seasons were

273 computed. Results are expressed in terms of p -values, used in this case as a “similarity index”: p -
274 values higher than the chosen significance level (0.05) indicates that the two compared variables are
275 not statistically different, while p -values lower than 0.05 indicate a statistically relevant difference.
276 Method based on the bivariate polar plot (Carslaw and Ropkins, 2012) shows how the concentration
277 of species varies jointly with wind speed and wind direction in polar coordinates. This approach
278 allows to trace the source location, preferably at scale ~ 1 -10 km where the wind direction is remaining
279 quasi-constant while the calculations of air mass transportation provides a limited support thanks to
280 the low spatial resolution of the model employed (Carslaw and Beevers, 2013). The highest values of
281 concentrations identify the directions of emission source.

282 Backward trajectories (BWT) were generated using NOAA HYbrid Single-Particle Lagrangian
283 Integrated Trajectory (HYSPLIT) model of the Air Resources Laboratory (Stein et al., 2015), with
284 coordinate resolution equal to $1^\circ \times 1^\circ$ of latitude and longitude. The potential source areas were
285 investigated using 3 days back for air masses arriving each one hour at 500 m height above sea level
286 (A.S.L.). BWTs were calculated in 1 hour for all periods of sampling.

287 Cluster analysis of trajectory data is an effective method of grouping the origin-like trajectories by
288 combining the geographically close data, it takes into account the features of seasonal atmospheric
289 circulation. The angular method is selected in which the matrix of angular distances shows how
290 similar are two points of the BWT in dependence on their angle with respect to the initial location
291 (Cui et al., 2021). Fire information was obtained from Resource Management System (FIRMS),
292 database MODIS (<https://firms.modaps.eosdis.nasa.gov/map>), operated by the NASA/GSFC Earth
293 Science Data Information System (ESDIS).

294 The scavenging ratio is a parameterization of the loss rate for gas or aerosol particles from the
295 atmosphere by their incorporation into larger drops. This factor can be calculated by taking into
296 account the volumes or masses of the chemical species being exchanged (Engelmann, 1971; Kasper-
297 Giebl et al., 1999). The scavenging ratio in our study is the quotient of the concentration of ions in
298 precipitation related to the concentration of those ions in aerosols. The scavenging ratios for each ion
299 (ω) were calculated using the following formula:

$$\omega = (C_p / \rho_p) / (C_a / \rho_a), \quad (1)$$

300 where C_p is the ion concentration in precipitation, $\mu\text{g L}^{-1}$, C_a is the concentration of the same ion in
301 aerosols, $\mu\text{g m}^{-3}$, ρ_p is the density of the precipitation, 1.0 kg L^{-1} , ρ_a is the density of the air. Air density
302 was assumed to be constant during all sampling periods (1.2 kg m^{-3}). With formula (1), it is possible
303 to estimate the impact of rainfall for the different chemical species. In this work we have limited our
304 calculation to season averages and to SIA ions, since their origin is more safely attributable to wet

305 deposition as the aerosol scavenging processes; the other ion species may come from dry deposition
306 which cannot be solved using a bulk precipitation collector.

307 Source apportionment was performed for both particulates and precipitation by factor analysis using
308 the Varimax algorithm (Van Espen and Adams, 1983), that is an extension of principal component
309 analyses (PCA) with a rotation of the principal components space to convert the PCs into factors. In
310 this way, each factor can be considered as the common source for the variables having higher loading
311 values for that factor. This method was preferred to other more efficient models such as Positive
312 Matrix Factorization owing to statistical constraints dictated by the limited number of observed
313 parameters (Hopke, 2016).

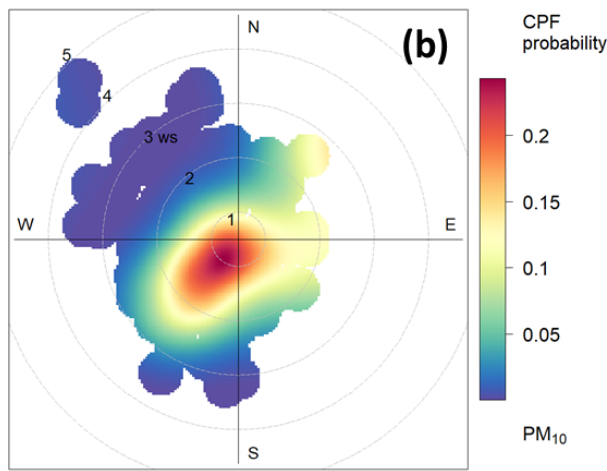
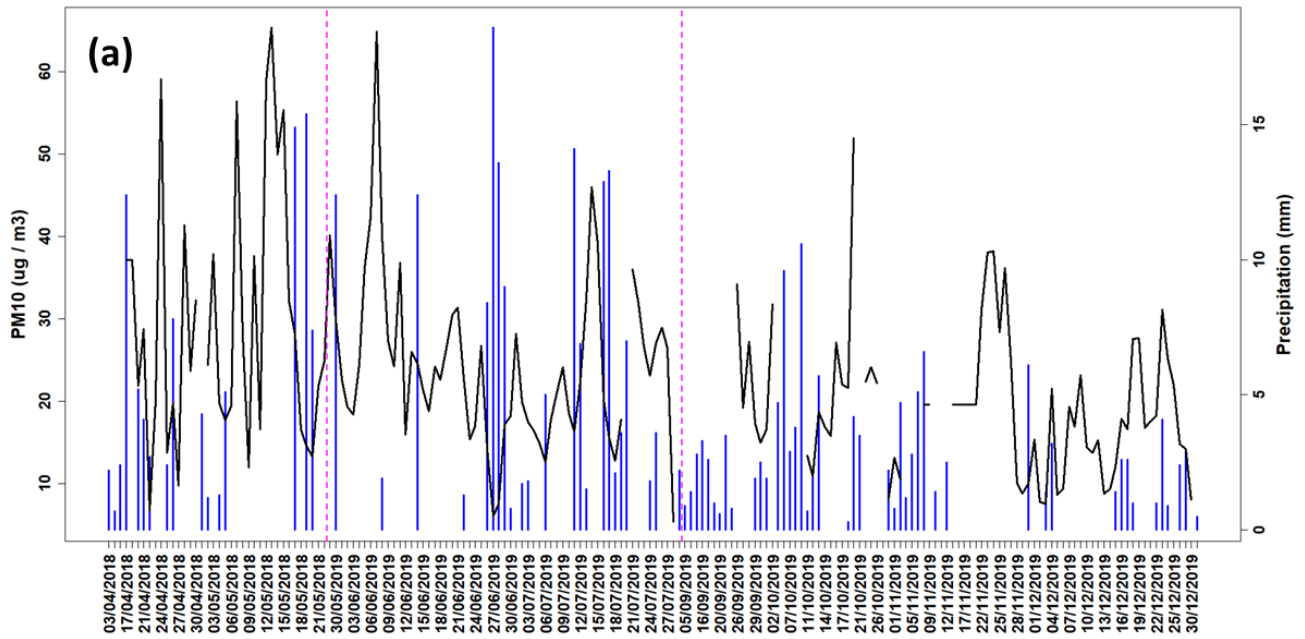
314

315 **3. Results and discussion**

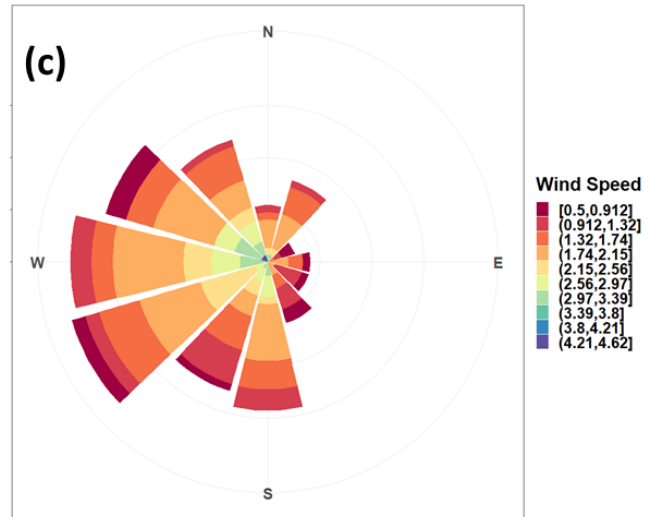
316 **3.1. Aerosol data overview**

317 **3.1.1. PM₁₀ and meteorological parameters overview**

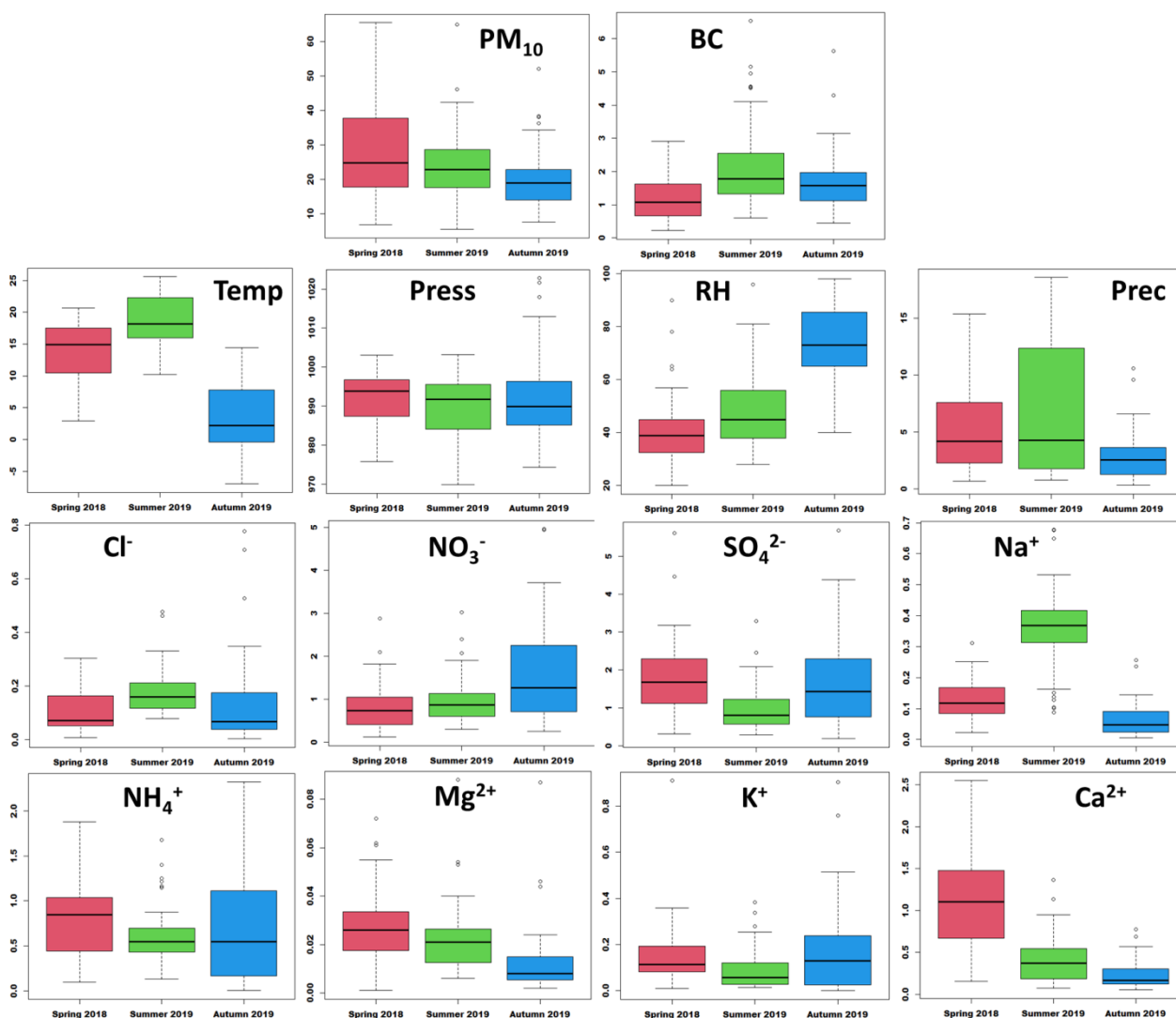
318 Table S1 presents the seasonal statistics of meteorological variables. Springtime of 2018 and
319 summertime of 2019 with mean temperature $14 \pm 4.7^\circ\text{C}$ and $18.7 \pm 3.8^\circ\text{C}$, respectively, corresponded
320 to the warm period while autumn with low mean temperature $3.2 \pm 5.4^\circ\text{C}$ belong to the cold period.
321 The season with the highest precipitation (6.53 ± 1.3 mm) was summer. Time series of PM₁₀ mass in
322 three seasons, i.e., spring 2018, summer 2019, and autumn 2019 was shown in Figure 2. PM₁₀ did not
323 exceed the EU threshold of $50 \mu\text{g m}^{-3}$ during any day if not in rare cases. The plot of PM₁₀ vs. the
324 main meteorological variables reveals that higher values were detected in association with periods of
325 low or no precipitation (Figure 2a), and low wind predominant from the southwest (Figure 2b). The
326 wind rose for the studied period is shown in Figure 2c, confirming westerly and south-westerly winds
327 with speeds between 0.5 and 2.5 m s^{-1} , similar to the predominant wind direction for the MO MSU
328 reported by (Chubarova et al., 2014). No particular correlation was observed between PM₁₀ and
329 temperature (Figure S1a). Lower PM₁₀ was encountered in connection with lower pressure (Figure
330 S1b), usually coinciding with higher precipitation (Figure 2a) and relative humidity (Figure S1c),
331 responsible for aerosol wet scavenging, in agreement with the PM₁₀ wet removal. The biggest
332 differences between meteorological parameters were found for temperature and relative humidity by
333 boxplots in Figure 3.



CPF at the 75th percentile (=28)



334
 335 **Figure 2.** a) Time series of PM₁₀ (black lines) and precipitation (blue bars), dashed lines indicate the
 336 separation between three seasons; b) CPF bivariate polar plots showing the wind speed (ws) and wind
 337 direction dependence of PM₁₀ concentration at 75th percentile, radial axis for wind speed in m s⁻¹; c)
 338 wind rose for the whole period of study.
 339



340

341 **Figure 3.** Boxplots of aerosol PM_{10} , BC, temperature (Temp, °C), pressure (Press, hPa), relative
 342 humidity (RH, %), and Precipitation (Prec, mm), and inorganic ions concentrations ($\mu\text{g m}^{-3}$) for spring
 343 2018, summer 2019, and autumn 2019. Units of measure are the same of Table S1. Red, green, and
 344 blue boxes indicate spring, summer, and autumn, respectively.

345

346 3.1.2. Descriptive statistics

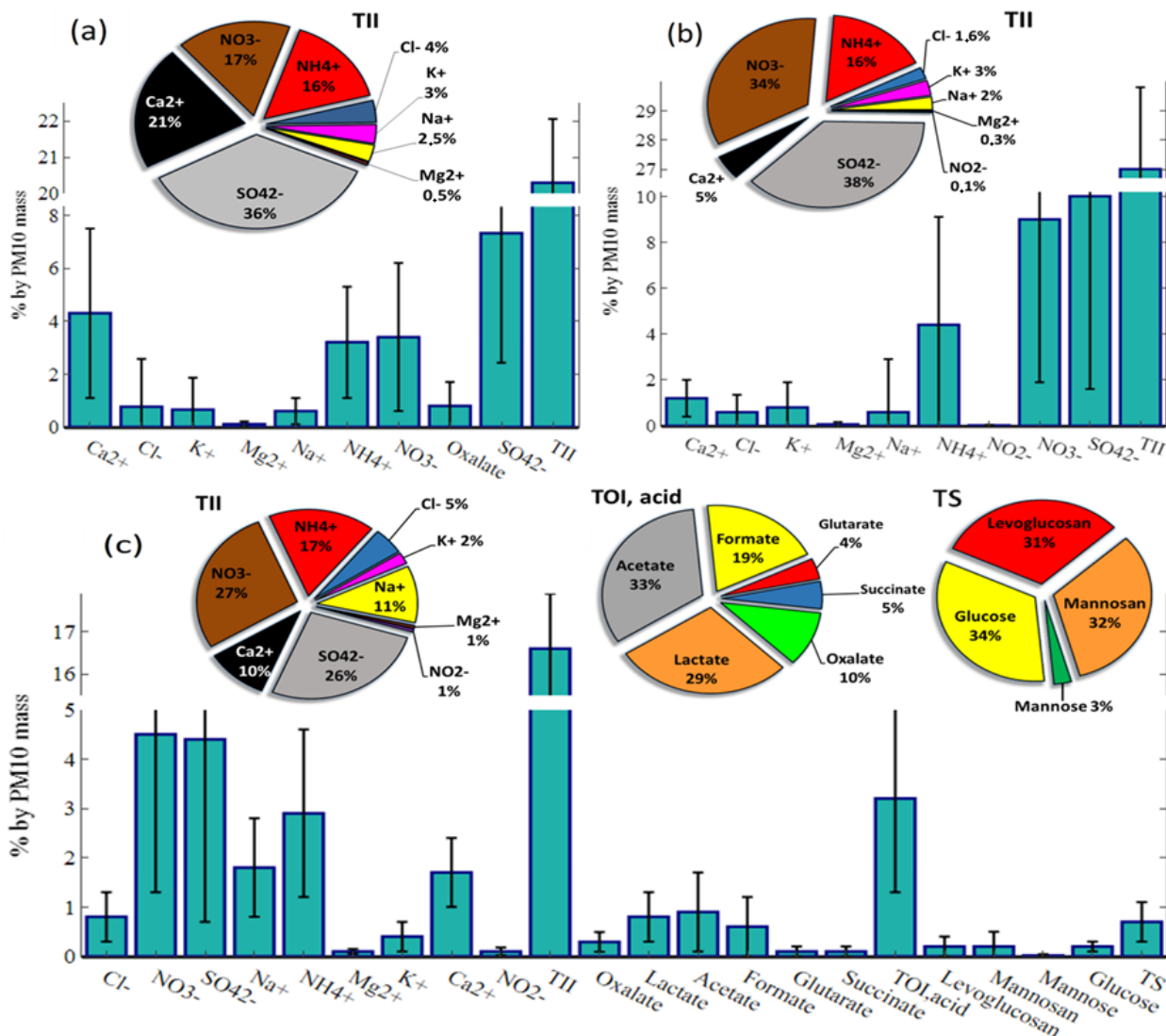
347 Basic statistics on seasonal concentrations of PM_{10} , inorganic ions (SO_4^{2-} , NO_3^- , Cl^- , Ca^{2+} , K^+ , Na^+ ,
 348 and Mg^{2+}), carboxylic acids, and sugars is shown in Table S1 for the whole period of study. Oxalate
 349 and nitrite (NO_2^-) could be measured above DL additionally to inorganic ions in spring, and in autumn
 350 and summer, respectively. Carboxylates (oxalate, acetate, formate, lactate, glutarate and succinate)
 351 as well as sugars (mannose, glucose) and anhydrosugars (levoglucosan and mannosan) were
 352 considered for summertime only. Also, total inorganic ions (TII), total organic ions (TOI) and total
 353 sugars (TS) are estimated in Table S1.

354 Maximum mean PM₁₀ (33.5±28.2 μg m⁻³) was found in spring and minimum (19.5±8.6 μg m⁻³) in
355 autumn. Statistically significant difference was estimated between autumn as compared to other two
356 seasons. In both spring and autumn mean concentrations for inorganic anions and cations was found
357 as following: SO₄²⁻>NO₃⁻>Cl⁻ and Ca²⁺>K⁺>Na⁺>Mg²⁺. In summer this tendency is the same, except
358 for Na⁺, higher than K⁺.

359 Boxplots for inorganic ions are shown in Figure 3. Na⁺ showed the largest variability between
360 seasons, with a median of 0.37 μg m⁻³ and maximum of 0.67 μg m⁻³ in summer and a median of 0.05
361 μg m⁻³ and minimum of 0.01 μg m⁻³ in autumn, respectively. Its variability is well-captured by Mann-
362 Whitney tests comparing chemical values between seasons: *p*-values are all statistically significant
363 (< 10⁻⁴), indicating significant differences between all couple of seasons. A similar behavior was
364 found also for Ca²⁺, although in this case a maximum (1.11 μg m⁻³) in spring and a minimum (0.16
365 μg m⁻³) in autumn were observed, all Mann-Whitney *p*-values are statistically significant. K⁺ is
366 different by the highest median of 0.11 μg m⁻³ and maximum in spring of 0.99 μg m⁻³, and in autumn
367 (0.124 and 0.76 μg m⁻³) with significant differences (*p*-value < 0.5) in summer. Median Cl⁻ (0.17 μg
368 m⁻³) and SO₄²⁻ (1.0 μg m⁻³) also demonstrate the significant differences in summer while similarities
369 are found for these ions in the other seasons. Median NO₃⁻ (1.6 μg m⁻³) and Mg²⁺ (0.12 μg m⁻³),
370 instead, showed significant differences in autumn as compared to other seasons, while a median NH₄⁺
371 (0.8 μg m⁻³) exhibited a slightly significant difference (*p*-value = 0.04) only when comparing spring
372 and summer values.

373 Massive contribution of measured species to PM₁₀ for each season is shown in Figure 4. Total
374 inorganic ions (TII) constitute the largest fraction of aerosol mass in the cold period reaching a
375 maximum of 27% in autumn, and a minimum of 16.5% in summer. Sulfates and nitrates dominate
376 among the major anion components, with maximum contribution of 38 and 34% to TII in autumn.
377 Cations are mainly represented by calcium and ammonium. Ca²⁺ had a maximum of 21% in spring
378 and minimum of 5% in autumn. Seasonality of NH₄⁺ was not prominent, ~15-17% in any season. A
379 lower though significant fraction of inorganic ions in TII was represented by sodium and chloride; in
380 summer the percentage of both Cl⁻ and Na⁺ approached their maxima respectively of 5 and 11%,
381 dropping to minima of 1.6 and 2% in autumn, respectively. Finally, maximum of K⁺ fraction of 3%
382 was observed in spring and autumn.

383



384

385 **Figure 4.** Percentage contribution in PM₁₀ mass of inorganic ions and oxalate in spring 2018 (a), of
 386 inorganic ions and NO₂⁻ in autumn 2019 (b), and of inorganic ions, organic ions, sugars and
 387 anhydrosugars in summer 2019 (c). Contribution of total inorganic ions (TII), total organic ions (TOI)
 388 and acids, and total sugars (TS) including anhydrosugars is shown. Pie diagrams show the relative
 389 abundance of individual measured species in total.

390

391 In European urban areas, ammonium sulfate and nitrate concentrations may reach up to 21-28% of
 392 PM₁₀, owing to the extensive use of combustion at all the scales from industry down to domestic and
 393 individual uses (Putaud et al., 2010; Tositti et al., 2014). In Moscow urban background we have
 394 obtained the maximum contribution of ammonium sulfate and nitrate to PM₁₀ around 15% in the cold
 395 period.

396 Summer species analyses showed that total inorganic fraction in PM₁₀ was higher (16.5%) than total
 397 carboxylates (3.5%), and total sugars and anhydrosugars (0.8%). Acetate and lactate dominated total
 398 organic ions (TOI), 33% and 29%, respectively. Acetate constitutes 1% to PM₁₀ mass in summer

399 samples. Levoglucosan and mannosan were the most prominent components of total sugars with
400 respectively 31% and 32%.

401

402 **3.1.3. Comparing Moscow PM data with other large cities**

403 Aerosol sources are seasonally modulated by a number of meteorological factors, e.g., precipitation,
404 atmospheric circulation as well as solar radiation, all influencing the planetary boundary layer height
405 and, therefore, air pollutant concentrations. Urban emissions may be also influenced by seasonal
406 variability, as in the case of domestic heating. In Table 1, data obtained for Moscow urban background
407 are compared with a few large cities such as Amsterdam, London, Athens, and Paris. The comparison
408 was carried out with urban background sites of big cities taking place also at the similar northern
409 latitudes as Moscow. Mean PM₁₀ in Amsterdam (Hama et al., 2018) is slightly lower than Moscow
410 but SIA concentrations are higher. SIA concentrations in London are also higher (Beccaceci et al.,
411 2015). In Athens NO₃⁻ concentrations are comparable with Moscow while other SIA are higher. The
412 average PM₁₀ in Paris suburbs is almost twice higher while the concentrations of SIA are 3-12 times
413 higher (Srivastava et al., 2018). Continental Toronto downtown, that is found at a similar latitude
414 than Moscow but is influenced by the Great Lakes, shows SIA (Jeong et al., 2020) comparable to
415 those observed in Moscow. Here we note that data for Toronto and Athens are presented for PM_{2.5}.
416 However, the comparison hold for SIA species may be reasonably considered because the SIA
417 populating the submicron range is common to both PM₁₀ and PM_{2.5}.

418 Despite Moscow has the largest population among European and Canadian cities of our comparison
419 (with exception of London closer with its 9 million of inhabitants) though their size is not comparable,
420 our data shows that Moscow urban background has the lowest SIA content with minima of SO₄²⁻ and
421 NH₄⁺, and comparable NO₃⁻. Such finding is well related to a fact that nitrogen oxides constitute a
422 large fraction of gaseous atmospheric pollution in Moscow, and their concentration is comparable to
423 most of largest cities in industrialized countries while the average concentration of sulfur dioxide,
424 ozone, and carbon monoxide is much lower than in the majority of world megacities (Elansky et al.,
425 2007). The reason for that is the dominant using of gas as a fuel, leading mainly to the emissions NO_x
426 and CO from heat and power plants (Elansky, 2014).

427 Na⁺ and Cl⁻ concentrations in all the cities mentioned above with marine climate are higher than in
428 Moscow a few times, indicating direct sea salt impact on their airsheds. Under ordinary circulation
429 conditions, sea salt cannot exceed a few kilometers from the coast-line inward. In Toronto (Srivastava
430 et al., 2018), Na⁺ and Cl⁻ are significantly less, revealing the absence either sea-salt or road
431 management impact on the ion composition. The similar trend of strongly higher Mg²⁺ concentrations
432 are found for coastal cities, probably the indication of the impact of sea salt aerosols on water-soluble

433 species as well. Road salt aerosol driving wintertime atmospheric chlorine chemistry have been
 434 observed for inland cities (McNamara et al., 2020) and related to pseudo-marine factor relating to
 435 winter street management (Tositti et al., 2014). Thus, Cl^- and Na^+ concentrations at the MO MSU site
 436 can be attributed to the impact of de-icing agents (DIAs) on aerosol composition in agreement with
 437 previous work by Eremina et al. (Eremina et al., 2015). Indeed, icy winters with long-lasting snow
 438 deposits require an intensive road management in Moscow to prevent traffic inconveniences. The
 439 composition of solid and liquid DIA used in Moscow megacity is dominated by NaCl salts, as
 440 reported in Supplementary Material. DIAs resuspension by traffic-induced turbulence significantly
 441 alter the snow composition along the roadsides (Vlasov et al., 2020). Near the MO MSU, soil
 442 salinization persists throughout the year and practically does not depend on the moisture regime
 443 affecting aerosol composition by soil resuspension in a continuous way and a modulation controlled
 444 by local circulation.

445 Mean PM_{10} in Moscow is comparable with continental large cities such as Berlin (Wieprecht et al.,
 446 2004) and Prague (Schwarz et al., 2019), and slightly lower than in Sofia (Hristova et al., 2020) and
 447 Bucharest (Perrone et al., 2018). All mentioned cities are characterized by higher SO_4^{2-} and NH_4^+
 448 concentrations than in Moscow. NO_3^- is higher for Prague and Berlin, and comparable for Sofia and
 449 Bucharest. Transport of marine aerosols and impact of road salt along with aged sea salt increases
 450 Na^+ and Cl^- concentrations in Sofia (Hristova et al., 2020). In Prague, road salt along with aged sea
 451 salt provide the similar concentrations as in Moscow (Schwarz et al., 2019). Moreover, in Prague, the
 452 concentrations of Mg^{2+} is comparable while Ca^{2+} is two times lower than in Moscow. In Budapest,
 453 Na^+ concentrations are lower, and Mg^{2+} concentrations are comparable to Moscow. K^+ is only one
 454 ion which concentration data for Moscow and other cities are very similar, an exception is Sofia
 455 where three times higher pollution by potassium related to residential heating is observed.

456 Although a longer sampling campaign would be required to draw conclusive considerations, our data
 457 indicate that the pollution in Moscow is comparable, if not lower, than that of also smaller cities. Paris
 458 and London, moreover, have significative higher pollution, but with certainly lower dimensions and
 459 populations.

460

461 **Table 1.** Mean PM_{10} and ions concentrations, in $\mu\text{g m}^{-3}$, in Moscow and large European and North
 462 America cities over periods indicted. Sites of comparison are described.

City, period	PM_{10}	Cl^-	NO_3^-	SO_4^{2-}	Na^+	NH_4^+	Mg^{2+}	K^+	Ca^{2+}	site
Moscow, Russia Apr – May 2018, Jun – Dec 2019	23.3	0.14	1.20	1.46	0.18	0.69	0.02	0.13	0.48	urban background
Amsterdam, Netherlands Apr 2013 – May 2014	20.7	0.98	4.79	2.2	0.86	1.75	0.13	0.10	0.27	urban background

London, UK Jan 2013 – Dec 2013	--	1.34	3.41	2.16	1.05	1.25	0.14	0.09	0.39	urban background
Paris, France 6-21 March 2015	49.0	0.58	14.4	4.3	0.41	5.66	0.06	0.14	0.86	suburban background
Athens, Greece June 2005-Sep 2006	24.2*	0.26*	1.42*	4.19*	0.73*	1.32*	0.12*	0.18*	1.2*	suburban background
Toronto, Canada Mar 2004 – May 2014	13.0*	0.03*	1.70*	1.88*	0.04*	0.91*	0.01*	--	--	Downtown
Sofia, Bulgaria Jan 2019 – Jan 2020	30.9	0.31	1.87	3.03	0.37	0.75	0.12	0.36	0.76	urban background
Budapest, Hungary Feb 2015 – May 2015	29.9	--	1.9*	2.0*	0.13*	1.2*	0.02*	--	--	urban background
Prague-Libuš, Czech Rep. Apr 2008 – Mar 2009	26.7	0.22	3.15	2.86	0.16	1.67	0.03	0.13	0.22	suburban background
Berlin, Germany Sep 2001 – Sep 2002	22.3	--	2.4	3.9	--	1.8	--	--	--	urban background

463 **Note.** *Data are given for PM_{2.5}

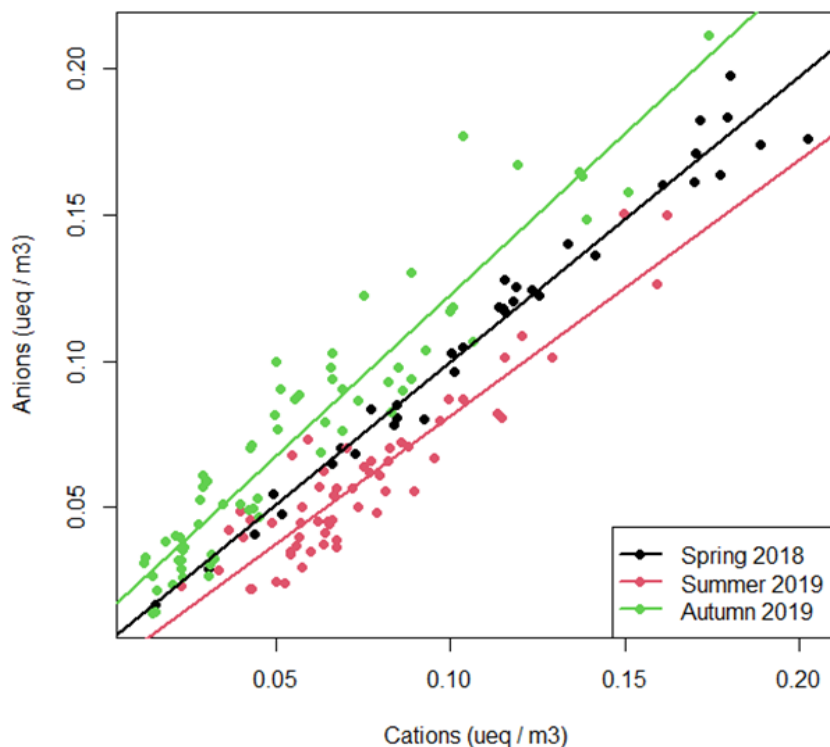
464

465 **3.2. Aerosol data analysis**

466 **3.2.1. Cations and anions balance**

467 For sake of electroneutrality constraints, cations and anions have been analyzed for their charge
468 balance. Ion data, converted into equivalents, show a deficit of anions in spring and summer, 33%
469 and 35%, respectively (Fig. S2). This fact may be ascribed to carbonate ions, that were not analyzed
470 directly. Their concentration was calculated from calcium and magnesium concentrations (Alastuey
471 et al., 2004; Popovicheva et al., 2020d; Tositti et al., 2014) attributed to the mineral components
472 associated to soil resuspension (Chow et al., 2015). Once added CO₃²⁻ to anions, linear regressions
473 are computed for three seasons (Fig. 5). Table 2 shows R², slopes, and intercepts, and their relative
474 standard deviations, σ , for the three seasons. Results indicate that spring aerosol is substantially
475 electroneutral with the slope and intercept not significantly different from their ideal values of 1 and
476 0, respectively. Summer data, instead, show a slope slightly, but significantly different from 1 (p -
477 value = 0.01). This indicates a slight excess of cations, but such a low difference, given that the
478 intercept is not significantly different from 0, is most likely due to the experimental uncertainties,
479 characterizing summer aerosol as almost neutralized. Autumn regression, instead, shows both slope
480 and intercept significantly different from 1 and 0 (p -values 0.005 and 0.0001). This suggests an excess
481 of anions, probably due to the presence of not neutralized H⁺ which can be a result of the gas to
482 particle conversion of SO₂ and NO_x to their respective acids and ammonium salts (not measured and
483 therefore unaccounted in the charge balance), due to a temporary reduction of alkaline substances
484 such as gaseous NH₃. It is to note that for Moscow atmosphere this component may include not only
485 actual soil and road components, but also the additional carbonate fraction and, what's more, CaCl₂

486 in DIA used for road management leading to a not predictable split of Ca and consequent bias in the
 487 charge balance (Eremina et al., 2015; Vlasov et al., 2021).



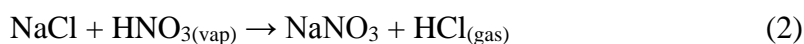
488
 489 **Figure 5.** Linear regressions of anions vs cations for spring 2018, summer 2019, and autumn 2019
 490 with the contribution of CO_3^{2-} added to anions.

491
 492 **Table 2.** Regression lines parameters for ion charge balance for three seasons.

Season	R^2	Slope	σ_{slope}	Intercept	$\sigma_{\text{intercept}}$
Spring 2018	0.97	0.97	0.03	0.002	0.0035
Summer 2019	0.84	0.87	0.05	-0.006	0.0039
Autumn 2019	0.89	1.10	0.04	0.012	0.0032

493
 494 **3.2.2. Chloride depletion**

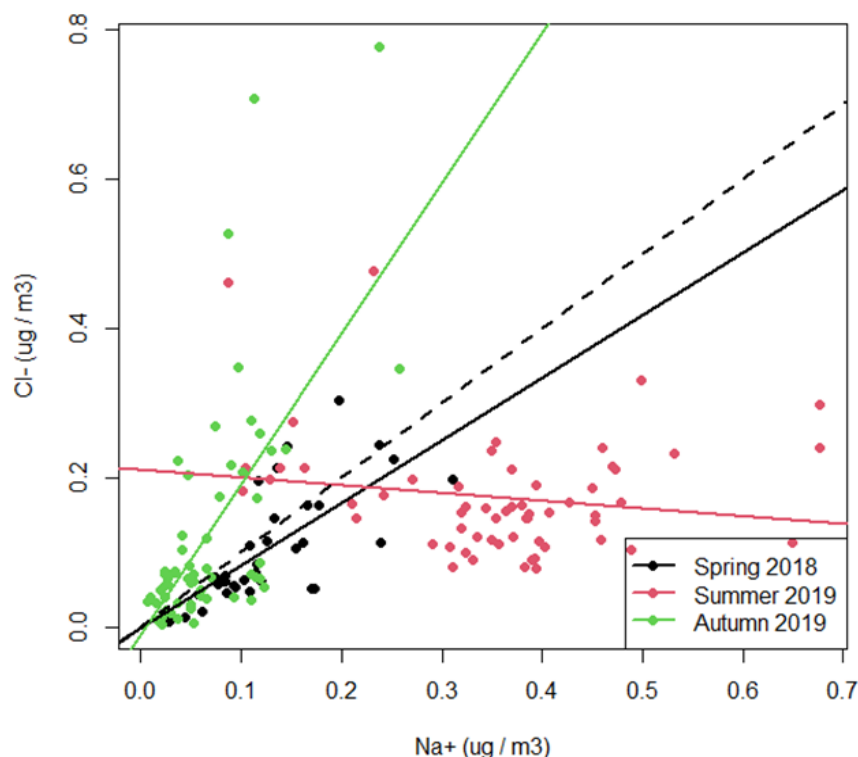
495 We check for the occurrence of chloride depletion in Moscow urban background, deriving from the
 496 likely interactions of HNO_3 with NaCl (Hoffman et al., 2004; McNamara et al., 2020). The reaction
 497 involved is the following:



499 It was previously shown that the appearance of hydrogen chloride in the atmosphere of Moscow is
 500 the result of heterogeneous multiphase chemical reactions, including the above one, with the
 501 participation of DIAs (Eremina et al., 2015) .

502 In order to evaluate the chloride depletion in three seasons, regressions were computed using Na^+ as
 503 independent variable, owing to its remarkable chemical stability in the environment, and Cl^- as the

504 dependent variable. The results are reported in Figure 6 and Table 3. Ion concentrations reveal a poor
 505 but positive ($R^2 > 0.4$) correlation in spring and autumn with slopes of 0.87 and 2.02, respectively.
 506 Both slopes are far from 1.54, the value expected for NaCl composition. In spring, Cl^- concentration
 507 is less than expected in sodium chloride, demonstrating a feature of chlorine depletion following the
 508 snow melt and the remobilization of DIA in road dust.
 509 In summer, the correlation between Cl^- and Na^+ is negative, suggesting a significant influence of
 510 reaction (2) in displacing chloride from sodium chloride. Probably, DIAs are not washed off
 511 efficiently by snow thawing and runoff as soon as the weather becomes milder, allowing for its
 512 aerosolization at the inception of warmer and drier conditions in summer. NaCl resuspension is
 513 promoted by both a significant low-level turbulence and dryness and the required photochemistry
 514 leading to HNO_3 accumulation, less easily neutralized by ammonia which reacts preferably to H_2SO_4
 515 (Hewitt, 2001; Seinfeld and Pandis, 2016).



516
 517 **Figure 6.** Correlation of Na^+ and Cl^- for the three seasons. Dashed line corresponds to concentration
 518 ratio of Na^+ and $\text{Cl}^- = 1:1$.

519
 520 **Table 3.** Regression lines parameters for chloride depletion in three seasons.

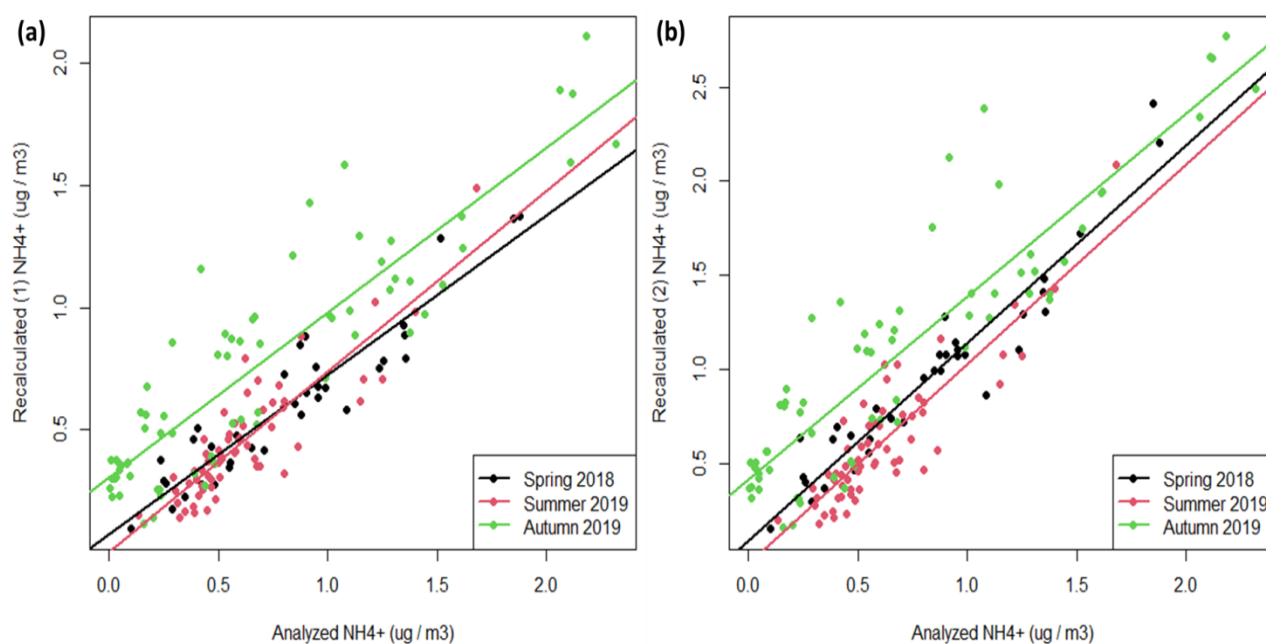
Season	R^2	Slope	σ_{slope}	Intercept	$\sigma_{\text{intercept}}$
Spring 2018	0.49	0.87	0.14	-0.003	0.02
Summer 2019	0.01	-0.10	0.08	0.21	0.03
Autumn 2019	0.41	2.02	0.31	-0.01	0.025

521

522 3.2.3. Secondary inorganic aerosol (SIA)

523 SIA derives from the oxidation of SO_2 to H_2SO_4 and of NO_x to HNO_3 , respectively, under the
524 influence of photochemically driven oxidation processes (Hewitt, 2001). These reactions are followed
525 by variable degrees of neutralization by two alkaline substances, i.e. gaseous NH_3 , the most abundant
526 atmospheric alkaline substance, and to a lesser extent CaCO_3 of mineral origin (Fountoukis and
527 Nenes, 2007; Heal et al., 2012; Karydis et al., 2020; Tositti, 2018). It is to note that the neutralization
528 of H_2SO_4 by NH_3 is thermodynamically favored as compared to the analogous reaction with HNO_3 ,
529 while NH_4NO_3 is less thermally stable, both conditions promoting the availability of HNO_3 for
530 reaction (2) and the subsequent chloride depletion.

531 Lai et al. (Lai et al., 2007) proposed two methods to evaluate the degree of neutralization of ammonia
532 as sulfate or nitrate by using the correlation between the concentration of NH_4^+ and that calculated
533 from the concentrations of SO_4^{2-} and NO_3^- . Both methods are based on stoichiometric considerations
534 and are based respectively on the neutralization of H_2SO_4 and HNO_3 by ammonia as NH_4NO_3 and
535 NH_4HSO_4 (method 1), or as NH_4NO_3 and $(\text{NH}_4)_2\text{SO}_4$ (method 2). In our work, we used nss-SO_4^{2-}
536 fraction instead of total SO_4^{2-} . The results obtained are depicted in Figure 7, Table 4 shows the
537 regression results. According to Lai et al., the good correlations ($R^2 > 0.7$) and intercepts not
538 significantly different from zero indicate that sulfates and nitrates are fully neutralized by ammonia
539 in spring and summer. In autumn, the intercept is significantly higher than 0, sulfate and nitrate are
540 not fully neutralized by ammonia, but other cations (as Mg^{2+} and Ca^{2+}) contribute to the equilibrium.
541 Ammonia, indeed, is supposedly less abundant in autumn owing to a decrease in biogenic activities
542 and to soil capping due to snow cover as soon as the cold season proceeds. These seasonal features
543 of ammonium participation in the formation of SIA and their further washing out by precipitation
544 lead to the fact that the share of ammonium in the total amount of ions in the snow cover is on average
545 about 6% (Eremina and Vasil'chuk, 2019), while in precipitation during the warm period it averages
546 11-12% (Eremina, 2019). The slope closer to 1, found by method 2, suggests that nitrate and sulfate
547 exist mostly as NH_4NO_3 and $(\text{NH}_4)_2\text{SO}_4$ in all seasons.



548

549 **Figure 7.** Regression lines of recalculated vs analyzed NH_4^+ concentrations according to a) method
 550 1 and b) method 2 (Lai et al., 2007).

551

552 **Table 4.** Parameters of the regression lines for neutralization of sulfates, nitrates, and ammonia as
 553 NH_4NO_3 and NH_4HSO_4 (method 1), and as NH_4NO_3 and $(\text{NH}_4)_2\text{SO}_4$ (method 2).

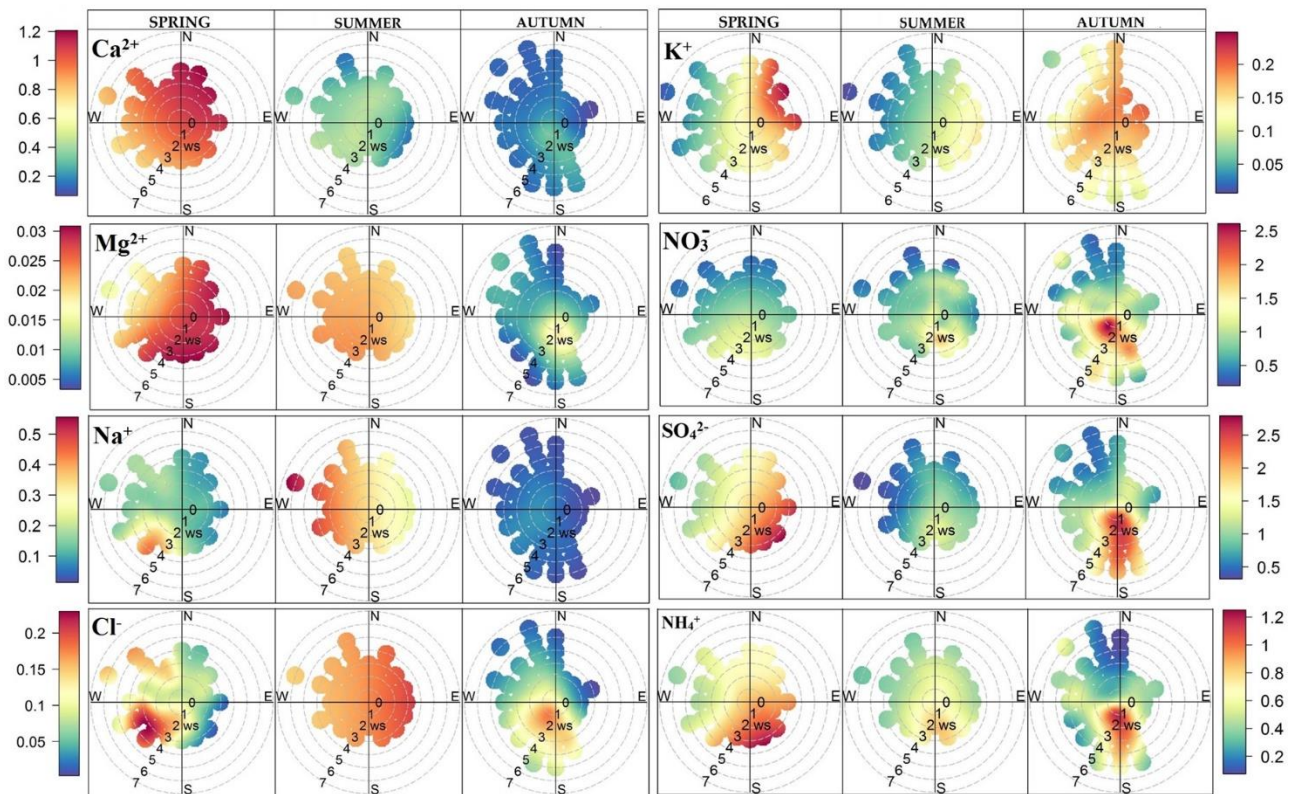
Method 1	R²	Slope	σ_{slope}	Intercept	$\sigma_{\text{intercept}}$
Spring 2018	0.87	0.65	0.043	0.072	0.040
Summer 2019	0.74	0.74	0.056	-0.002	0.038
Autumn 2019	0.79	0.67	0.042	0.305	0.039
Method 2					
Spring 2018	0.90	1.05	0.057	0.087	0.053
Summer 2019	0.76	1.06	0.076	-0.036	0.052
Autumn 2019	0.80	0.97	0.058	0.413	0.054

554

555 3.3. Potential source analyses

556 3.3.1. Anemological analysis

557 Meteorological factors contribute significantly to variability of the ionic content in aerosols
 558 (Dadashazar et al., 2019). Relationships among the source factor and wind direction and speed, i.e.
 559 anemological analysis, are derived from polar plots (Saraga et al., 2021). Figure 8 shows the bivariate
 560 concentration polar plots during spring 2018, summer and autumn 2019, providing a graphical
 561 representation of the source origin for the MO MSU site. Map of Figure 1 depicts the MO MSU
 562 sampling site located in the southwest sector of Moscow, where industrial facilities are placed within
 563 a distance up to 10 km, including thermal power stations, thermal stations, mechanical engineering,
 564 metallurgical, food, reinforced concrete, chemical and pharmaceutical factories.



565

566 **Figure 8.** Concentration bivariate polar plots showing the wind speed and wind direction dependence
 567 during spring 2018, summer and autumn 2019. The radial axis is wind speed in m s⁻¹ and the color
 568 scale is the concentrations in μg m⁻³.

569

570 Polar plots of SO₄²⁻ and NH₄⁺ show a very similar pattern indicating the same origin for ammonium
 571 sulfates and reflects the high correlation between these ions as a result of the significant
 572 thermodynamic stability of ammonium sulfates originated from SO₂ gas-to-particle conversion
 573 (Hewitt, 2001). In spring and autumn, they indicate the pollution sources in the southeast where the
 574 most industrially developed area takes place in Moscow (Bityukova and Saulskaya, 2017). Prevailing
 575 winds were from north and northwest in summer, therefore the MO MSU site is not downwind the
 576 heavily impacting sources as in spring and autumn.

577 High concentrations of NO₃⁻ (about 1.5 μg m⁻³) in spring and summer are observed at weak wind
 578 conditions when non-buoyant ground-level sources such traffic are important. Increasing wind speed
 579 generally results in lower concentrations due to dispersion and mechanical turbulence. In autumn,
 580 NO₃⁻ increases probably due to the centralized heating system operation when the MO MSU site is
 581 downwind from southwest stacks of the large industrial zone (in particular, from the thermal power
 582 stations (Fig. 1)) as well as from the southeast industrial region.

583 Cl⁻ and Na⁺ polar plots in spring indicate high concentrations, above 0.2 and 0.32 μg m⁻³, respectively,
 584 from southwest at wind speeds higher than 2 m s⁻¹. In fact, it coincides with heavily trafficked roads

585 southwest MO NSU where efficient road management aiming at reducing snow/ice deposits in the
586 cold period is carried out systematically. In summer, the amount of salts in soil interstitial water drops
587 by 2-7 times (Nikiforova et al., 2014). Also, Cl^- , Na^+ and NO_3^- decrease due to salt leaching, runoff,
588 and absorption of nitrates by biota, but Ca^{2+} , Mg^{2+} and HCO_3^- increase because dust and crushed
589 marble chips accumulation in the upper horizon (Kosheleva et al., 2018). Complete leaching of salts
590 from the soil profile does not occur during summer, that leads together with soil solonchization to a
591 progressive salinization of Moscow surface soils and to its substantial and everlasting availability to
592 air-resuspension in the periods without snow cover (Nikiforova et al., 2017). Moreover, due to the
593 greater amount of precipitation in summer compared to other seasons, Cl^- can migrate to the middle
594 part of the soil profile while Na^+ fixed in surface soil horizon can be resuspended into the atmosphere.
595 Na^+ polar plot in summer shows the resuspension of Na^+ -enriched soil particles at any wind directions
596 and speeds. Later, in autumn road salt is likely trapped in snow and ice with consequent scarce
597 resuspension and no pronounced features of sources by polar plot.

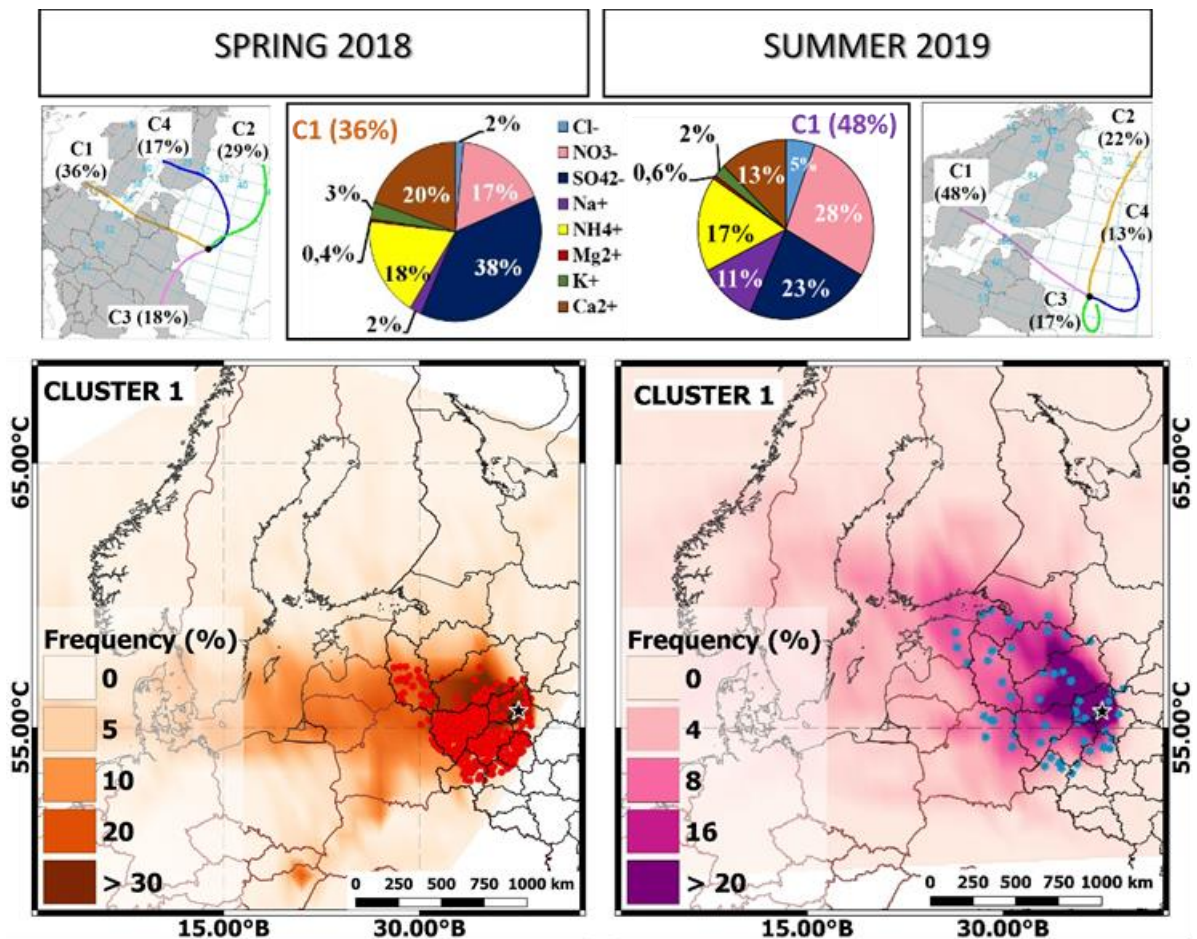
598 For Ca^{2+} and Mg^{2+} , the high concentrations above 0.025 and 0.63 $\mu\text{g m}^{-3}$, respectively, in spring
599 occurs from all the wind directions, relating to local source of soil resuspension after snow thaw. Ca^{2+}
600 polar plot in summer shows additional source areas southwest and southeast, associated with intensive
601 building activities in the New Moscow area. In the other seasons these ion sources are not clearly
602 emerging.

603 K^+ can be originated from soil (mineral potassium) and biomass burning. It is to note that mineral
604 potassium is less soluble than the corresponding fraction from biomass burning, owing to strong
605 crystal lattice bonds in the rocks and minerals. K^+ in our experimental data, obtained from ion
606 chromatography, represents most likely soluble potassium source. Thus, K^+ polar plot indicates
607 biomass burning in spring from the east and in summer from the southeast at wind speeds higher than
608 2 m s^{-1} , with a likely dominating influence of remote sources in Moskovskaya Oblast as well as
609 seasonal wildfires. Autumn plot, instead, indicates the presence of a major source of K^+ from all the
610 directions and low wind speeds, though more intense from west and southeast. Since in autumn no
611 wildfires are observed in the northern European part of Russia, the likely origin for K^+ is biomass
612 burning in the region surrounding Moscow. The spectral dependence of the aerosol light attenuation
613 observed in Popovicheva et al. (Popovicheva et al., 2022) supported the assessment for a contribution
614 of biomass burning in the region around Moscow. We note that autumn is a cold season requiring
615 domestic heating since September, it is recognized that the residential sector near Moscow employs
616 wood for heating.

617

618 **3.3.2. Backward trajectory analysis**

619 Air pollutant transport relates to the atmospheric circulation and spatial distribution of regional
 620 sources. In spring and summer agriculture fires and wildfires may occur in the region around Moscow,
 621 thus providing the strong impact on city air quality. Additionally, the residential sector in
 622 Moscovskaya Oblact can impact with using of biomass for house heating and cleaning the gardens.
 623 Major directions of air mass transportations by BWT cluster analysis are shown in Figure 9. In spring,
 624 the main cluster C1 includes 36% of analyzed BWTs and indicates the northwestern and western
 625 direction as dominant. Clusters C2 (29%) and C3 (18%) of northwestern and southern directions
 626 support the prevalent air mass transportation, respectively. 17% of BWTs compose the cluster C4
 627 from the north. In summer, the largest number of BWTs came from the northwestern direction (48%),
 628 they compose the cluster C1. The northern and southern directions include the cluster C2 and C3,
 629 characterized by 22% and 17%, respectively, while the trajectories coming from the southwest
 630 directions make a less significant percentage contribution to the cluster C4 (13%).



631
 632 **Figure 9.** BTW cluster analysis and pie diagrams for ion concentrations which correspond to the
 633 trajectories within clusters C1 for spring 2018 (left) and summer 2019 (right) (top panel). 72-hour
 634 BWT frequencies of C1 cluster for spring 2018 (left) and summer 2019 (right) (bottom panel). Fire
 635 locations related to the time of BWT pass are indicated by stars, MO MSU site is marked by black
 636 star.

637

638 Pie diagrams in Figure 9 show ions concentrations related to trajectories within the clusters C1 of the
639 prevalent air mass transportation for each season. SO_4^{2-} , NO_3^- , and NH_4^+ of 38%, 17%, and 18% and
640 23%, 28%, and 17% of total ion concentrations, compose the highest SIA source in the northwestern
641 direction in spring and summer, respectively. Highest frequencies of BWTs in this direction relate to
642 the area of large fires occurred in the European part of Russia and around Moscow in both seasons.
643 High K^+ concentrations observed in this direction act as a marker of BB source impacted the air
644 pollution of a megacity.

645

646 3.3.3. Varimax analysis: ions and BC sources

647 In order to obtain the aerosol source profile, the receptor modelling and factor analysis was applied
648 (Hopke, 2016). In our study, Varimax analysis was computed using ion and BC data for the whole
649 dataset, with the aim of finding the potential common sources. Five factors were chosen as the best
650 solution, being the one explaining more than 80% of the total variance. Table 5 shows the Varimax
651 loadings obtained by the model.

652

653 **Table 5.** Varimax loadings for ion and BC concentrations during the whole study period. Bold values
654 correspond to the most representative variables for each factor. EV stands for the explained variance,
655 F for a factor.

	F1	F2	F3	F4	F5
EV(%)	27.6	20.8	13.3	13.1	12.1
Cl ⁻	-0.050	0.040	0.089	-0.898	-0.123
NO ₃ ⁻	-0.541	0.187	-0.149	-0.236	0.048
SO ₄ ²⁻	-0.550	-0.102	0.044	0.160	-0.128
Na ⁺	0.252	-0.340	-0.342	-0.287	0.359
NH ₄ ⁺	-0.578	-0.142	-0.015	0.014	0.210
Mg ²⁺	-0.051	-0.622	-0.096	-0.070	-0.061
K ⁺	0.043	-0.067	-0.048	-0.056	-0.873
Ca ²⁺	-0.031	-0.654	0.138	0.079	-0.098
BC	-0.030	0.070	-0.905	0.094	-0.131

656

657 The five factors reported in Table 5 were considered as corresponding to as many emission sources.
658 Factor 1 (F1) shows high loadings for NO_3^- , SO_4^{2-} , and NH_4^+ , indicating SIA as the dominant PM_{10}
659 component, corresponding to a non-resolvable cumulative contribution from all high temperature
660 emission sources encompassing traffic, heat and power generation, and industry (Seinfeld and Pandis,
661 2016). The absence of BC in this factor indicates that there is no direct link to combustion sources
662 like those using fossil fuel. SIA is originated in secondary processes which are not directly relate to

663 primary sources as BC but rather reflect the regional emissions affecting a megacity due to the
 664 relatively low chemical kinetics of formation and to non-local air mass transportation (Fig. 9).
 665 Moreover, comparable patterns of concentration polar plots for SO_4^{2-} and NH_4^+ (Fig. 8) confirm their
 666 common source as well. Factor 2 (F2) associates Ca^{2+} and Mg^{2+} , indicating the mineral component
 667 of which calcium is a proxy. It represents the more abundant components in crustal materials so-
 668 called soil/road dust resuspension, derived from rock reworking, soil development, and the built
 669 environment (Lai et al., 2007). Factor 3 (F3) indicates the BC sources associated with fossil fuel
 670 combustion i.e. traffic, heating system, and industry in a megacity (Popovicheva et al., 2022). High
 671 loading only for BC in this factor suggests the absence of any strong ion relevance to combustion
 672 sources. The decoupling of BC from K^+ likely indicates that the main source of BC in the Moscow
 673 atmosphere is fossil fuel combustion while K^+ is a marker of biomass burning presenting the highest
 674 loading in Factor 5 (F5). Cl^- derived from road salt, is found alone in Factor 4 (F4). It is to note that
 675 the decoupling from Na^+ does not mean that Cl^- is not due to road salt, their association has been
 676 shown above. Relatively significant loadings of Na^+ are found both in F2 and F4. This reasonably
 677 means that, differently from chloride, which is a highly reactive mobile substance and can be degassed
 678 by atmospheric processing, the same does not hold for Na^+ . Sodium is a highly conservative ion,
 679 which accumulates in the course of years in soil, where it is stabilized by ion exchange chemistry in
 680 the soil matrix to be recirculated in the lower troposphere following soil resuspension. Note that
 681 previously cited research work on the effects of DIA on soil water chemistry in Moscow revealed a
 682 decoupling of chloride ion from its cation partner Na^+ , circumstance which may support the
 683 occurrence of Cl^- in factor 4 without other significant correlation.

684

685 3.3.4 Varimax analysis in summer case study: organic ions and sugars sources

686 For summer ion chemistry and BC are integrated by a series of chemical species including carboxylic
 687 and di-carboxylic acids, anhydrosugars (levoglucosan with its isomer mannosan), and sugars. These
 688 species were analyzed only in summer since this is the only season of high photochemistry as well as
 689 biogenic activities and biomass burning in Russia. Although it is not possible to evaluate the seasonal
 690 variation of such species, their presence in summer samples can be used to evaluate their source, also
 691 in correlation with inorganic ones. The whole dataset is subjected to an independent elaboration for
 692 receptor modelling using Varimax factor analysis (Table 6).

693

694 **Table 6.** Varimax loadings of the summer case-study. Bold values correspond to the most
 695 representative variables for each factor. EV stands for the explained variance, F for a factor.

	F1	F2	F3	F4	F5	F6
--	----	----	----	----	----	----

EV(%)	18.8	18.1	13.7	13.5	12.0	7.80
Cl ⁻	-0.121	0.533	-0.116	-0.007	-0.019	-0.080
NO ₃ ⁻	0.453	-0.075	-0.143	-0.006	-0.023	0.100
SO ₄ ²⁻	0.497	-0.024	-0.021	0.017	0.188	-0.047
Na ⁺	-0.085	-0.098	-0.260	0.399	0.256	0.228
NH ₄ ⁺	0.424	-0.095	-0.047	0.196	0.154	-0.055
Mg ²⁺	0.048	-0.020	0.059	0.570	-0.021	-0.039
K ⁺	0.030	0.518	-0.014	-0.053	-0.055	0.058
Ca ²⁺	0.047	0.047	0.049	0.464	-0.249	-0.066
BC	0.140	0.321	-0.064	0.125	-0.082	0.013
Oxalate	0.454	0.116	-0.035	-0.067	0.040	-0.096
Lactate	-0.141	0.178	-0.544	0.205	0.209	-0.222
Acetate	0.050	-0.035	-0.484	-0.218	-0.290	0.001
Formate	0.005	-0.068	-0.548	-0.092	-0.026	0.181
Glutarate	0.140	0.027	0.011	0.020	-0.567	0.074
Succinate	0.105	0.019	0.009	0.116	-0.530	-0.035
Levogluconan	0.064	0.425	0.073	-0.122	0.090	0.039
Mannosan	0.015	-0.095	-0.085	-0.112	-0.053	0.707
Mannose	0.238	0.121	0.119	-0.207	0.230	0.074
Glucose	-0.024	0.246	0.164	0.238	0.077	0.559

696

697 Factor 1 (F1) presents high loadings of SIA and oxalates, confirming the role of photochemistry in
698 the formation of both inorganic and organic secondary aerosols (Tsai et al., 2013). In particular,
699 oxalates and sulfates undergo a similar chemical processing mainly associated with cloud processing
700 (Laongsri and Harrison, 2013). Factor 2 (F2) includes Cl⁻, BC, K⁺, and levogluconan. Time series of
701 BC, nss-K, and levogluconan in Figure S3 show a remarkably good agreement throughout the summer
702 period. It indicates the variable influence of biomass burning which relates to large wildfires observed
703 in the European part of Russia and particular in the region around Moscow, as shown by BWT cluster
704 analyses (Fig. 9). It is to note that the high loading of Cl⁻ in F2, similar to BB markers, shows its
705 seasonal distinct source together with a complex relationship to local origin. It means that Cl⁻ is
706 contributed by two sources, i.e. road salts from which it is partially removed to the gaseous phase,
707 and biomass burning, as was also observed by Brattich et al. (Brattich et al., 2021).

708 Factor 3 (F3) includes three carboxylic acids sharing similar sources. In particular, acetate is
709 correlated with biomass burning (Tsai et al., 2013). Acetate comes mostly from primary emissions,
710 while formate come from secondary transformations. The mass ratio of acetate and formate (A/F) is
711 used to distinguish primary and secondary sources of these acids (Tsai et al., 2013). An A/F ratio
712 higher than 1 indicates a prevalence of primary sources, while A/F lower than 1 indicates the contrary.
713 In our case, the mean A/F mass ratio is 2.19, therefore the main source of these acids can be
714 considered the urban primary sources, probably vehicular emissions (Tsai et al., 2013; Wang et al.,
715 2007).

716 Factor 4 (F4), including Ca^{2+} , Mg^{2+} , and Na^+ , represents the influence of road dust and soil
717 resuspension contribution thanks to the low-level turbulence and relatively drier conditions of
718 summer, capable to recirculate these components. Succinate and glutarate in Factor 5 (F5) relates to
719 oxalate and SIA as products of secondary formation from organic precursors by photochemical and
720 heterogeneous reactions (Guo et al., 2021). Factor 6 (F6) with a little influence on the overall variance
721 has a high loading of mannosan and glucose. Since mannosan is acting as biomass burning marker,
722 and glucose is the product of biogenic activity, they both reflect the co-existence of biomass burning
723 and bioaerosols which was observed in other studies (Bauer et al., 2008; Popovicheva et al., 2020).
724 High temperature in summer (Fig. S1a) definitely relates with intensive biological activities.
725 Moreover, the co-existence of sugars with anhydrosugars may be explained by suspension of soil and
726 plant debris in the heat wave of wildfires (Medeiros et al., 2006).

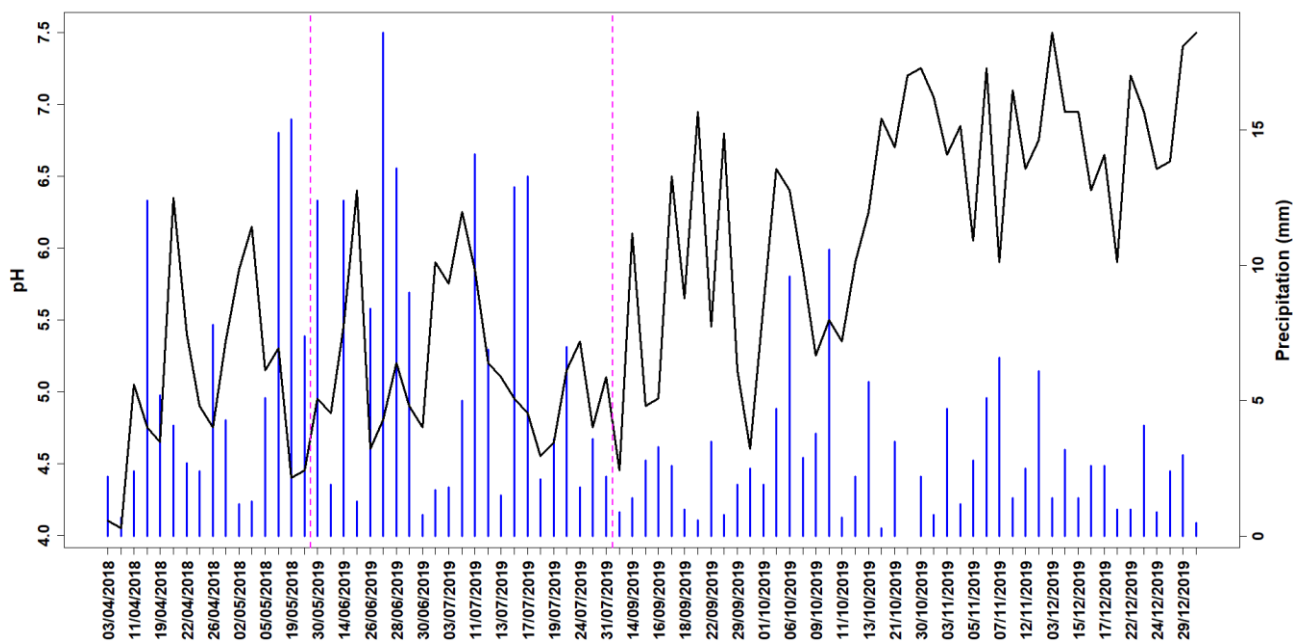
727

728 **3.4. Precipitation chemistry**

729 Precipitation in the form of rain and snow reflects the aerosol composition through complex in-cloud
730 and below clouds removal. The ion composition and pH are a function of the amount of condensed
731 water and the relative degree of neutralization of SIA, while pH may affect solubility of removed
732 aerosols and of those trapped in rain/snow bulk collectors (Pruppacher and Klett, 2010; Tositti et al.,
733 2018b; Vlasov et al., 2021).

734 Long-term records of precipitation chemistry at MO MSU site show the significant increase of the
735 acid precipitation number from almost absence until the level of 30% in last years (Eremina et al.,
736 2015). The percentage of acid precipitation with $\text{pH} < 5$ was 20.7% for a year on average, 31.0% for
737 a warm season, and only 5.7% for winter. The total ionic mineral composition does not change in
738 average, while chloride and sodium concentration were increasing.

739 Figure 2a shows precipitation events scattered throughout three seasons, they appear more abundant
740 in summer and autumn (Table S1). In the end of autumn, in December, rainfall was subsequently
741 substituted by snowfall. The pH recorded in precipitation reveals a certain tendency towards lower
742 values in spring and summer (Fig. 10), such a fact is typically recognized as pertaining to acid rains
743 (Bricker and Rice, 2003). Observed pH in spring and summer oscillate around 5-5.6.



744

745 **Fig. 10.** pH (black line) and precipitation intensity (blue bars) for the whole period of study.

746

747 Since during some time in autumn soil is frozen and covered with snow at the Moscow latitude, the
 748 agricultural contribution is likely negligible while traffic and industrial sources (with the support of
 749 ventilation) appear to be the most important ones. These leads to scarce availability of ammonia and
 750 therefore to a reduced neutralizing effect on sulfuric and nitric acids formed by SO₂ and NO_x
 751 oxidation rather than to an excess of acidifying species. Increase of pH in colder seasons, above 5.6,
 752 suggests both the neutralizing influence of ammonia and the contribution from carbonates in
 753 resuspended alkaline urban soils (Nikiforova et al., 2017), road dust (Kasimov et al., 2019), and
 754 undissolved DIAs (Korolev and Gornyakov, 2018).

755 Precipitation ion chemistry and data variability is associated with the seasonal precipitation rate as
 756 well as with the type of precipitation event in terms of total volume of the deposited hydrometeor as
 757 well as of the progression of the single event. Table 7 reports basic statistics for precipitation during
 758 three seasons, together with pH and ion concentrations. The highest mean precipitation is observed
 759 in summer while in autumn it is the lowest. The highest mean and median concentrations for almost
 760 all ions were observed in spring, except HCO₃⁻. Its maximum is in autumn, that is generally consistent
 761 with observations in previous years according to which the average concentrations of HCO₃⁻ in
 762 September-December are higher than in April-May (Eremina et al., 2017).

763

764 **Table 7.** Statistics of precipitation, pH, and ions concentrations (in µg L⁻¹) in three seasons.

Spring 2018				Summer 2019				Autumn 2019			
mean	median	min	max	mean	median	min	max	mean	median	min	max

Precipitation (mm)	5.59	4.20	0.70	15.4	6.53	4.30	0.80	18.6	2.91	2.55	0.30	10.6
pH	5.04	4.98	4.05	6.35	5.16	5.03	4.45	6.40	6.37	6.55	4.60	7.50
H ⁺	21.7	10.8	0.45	89.1	11.4	9.55	0.40	35.5	2.03	0.28	0.03	25.1
HCO ₃ ⁻	1.24	0.00	0.00	7.40	0.77	0.00	0.00	9.39	3.93	2.80	0.00	13.5
SO ₄ ²⁻	2.54	2.60	0.60	4.90	1.66	1.06	0.31	6.10	2.32	1.70	0.40	8.70
Cl ⁻	13.2	10.9	3.00	39.5	10.4	7.09	2.13	34.9	7.09	6.80	1.30	20.90
NO ₃ ²⁻	2.90	3.20	0.50	6.50	2.97	1.50	0.45	16.2	2.23	1.90	0.70	5.00
Ca ²⁺	5.96	4.50	1.10	21.0	5.03	3.45	0.68	18.9	4.29	3.70	0.80	11.0
Mg ²⁺	0.34	0.20	0.10	1.60	0.19	0.09	0.02	1.12	0.21	0.20	0.00	0.50
Na ⁺	1.71	1.40	0.30	4.90	0.50	0.35	0.07	1.80	0.96	0.80	0.10	2.70
K ⁺	0.43	0.30	0.10	2.90	0.40	0.27	0.02	2.22	0.51	0.40	0.10	3.30
NH ₄ ⁺	1.73	1.50	0.30	4.30	1.39	0.97	0.27	4.90	1.36	1.20	0.30	3.20

765

766 Spearman correlation analysis has been carried out. Based on Spearman indexes, it is found that ions
767 pertaining to SIA appear highly correlated (> 0.70) in connection with their origin and source. Other
768 correlations are in general lower than those observed for SIA (< 0.7), suggesting multiple sources for
769 the other ions. An exception is represented by the ions already cited as coming from road management
770 practices (Ca²⁺, Mg²⁺, and Na⁺, correlations between 0.65 and 0.75). This likely reflects the influence
771 of soil resuspension in dry periods on the precipitation chemistry, since precipitation is collected as
772 bulk deposition, i.e. the sum of wet and dry deposition, and not with a wet only approach (Tositti et
773 al., 2018b). Finally, H⁺ and HCO₃⁻ appear uncorrelated in respect to all the other ions (< 0.3) and they
774 are reciprocally anticorrelated (-0.936). This is reasonably due to the neutralizing effects of calcium
775 and magnesium carbonates from soil resuspension acting as neutralizing agents toward H⁺ which is
776 decreased through the formation of HCO₃⁻. In Moscow, during years with a deficit of acid
777 precipitation, the predominant anion in precipitation is HCO₃⁻, the source of which is carbonates of
778 alkali and alkaline earth metals. This is confirmed by the results of thermodynamic calculations; in
779 episodes with an excess of acids in precipitation it is associated with the absorption of ammonia from
780 the atmosphere (Eremina et al., 2017).

781 Efficiency of wet removal can be appreciated analyzing simultaneously the trend of PM₁₀ and
782 precipitation, as reported in Figure 2a. Seasonal variability of scavenging ratios (ω) of SIA species
783 (Table 9) have been calculated by taking into account the ion concentrations in aerosols, their
784 concentrations in precipitation, and the air density, using eq.1. Scavenging ratios are determined on
785 a seasonal basis because they have less variability compared to ones calculated using precipitation
786 event concentrations and consider the average ion concentrations in PM₁₀ during both precipitation
787 and dry periods (Cheng and Zhang, 2017).

788

789 **Table 9.** Scavenging ratios calculated for SIA compounds during three seasons.

	ωNO_3^-	ωSO_4^{2-}	ωNH_4^+
Spring 2018	$7.70 \cdot 10^3$	$2.12 \cdot 10^3$	$3.17 \cdot 10^3$
Summer 2019	$3.45 \cdot 10^3$	$2.40 \cdot 10^3$	$2.70 \cdot 10^3$
Autumn 2019	$5.29 \cdot 10^3$	$3.31 \cdot 10^3$	$2.74 \cdot 10^4$

790

791 The lowest ωNO_3^- , ωNH_4^+ , and ωSO_4^{2-} are determined in summer and spring, respectively. The
792 maximum ωNO_3^- is found in spring, in autumn we observe the maximum ωSO_4^{2-} and ωNH_4^+ .
793 Gaseous species such as HNO_3 , SO_2 , and NH_3 can influence the SIA scavenging ratios since the
794 composition of the precipitation is determined by the scavenging process of both gases and particles.
795 That might be the reason of the overestimated levels of ω especially in spring and summer, when the
796 formation of SIA is favored by the photochemical reactions (Oduber et al., 2021). For instance, 88-
797 96% of NO_3^- content in rainwater is formed due to gas-phase scavenging of HNO_3 , whereas 89-96%
798 of SO_4^{2-} concentration in rainwater may be due to particulate sulfate (Kasper-Giebl et al., 1999).
799 However, our findings showing ω decreasing in summer and increasing in autumn and spring are
800 consistent with estimates obtained for atmospheric air and precipitation at Canadian rural locations
801 (Cheng and Zhang, 2017).

802 Comparing the obtained values with those found in other works (Cheng and Zhang, 2017; Encinas et
803 al., 2010) the secondary ωSO_4^{2-} , ωNO_3^- and ωNH_4^+ resulted to be comparable with those calculated
804 in our study.

805 Though the scavenging ratio is widely used, it must bear in mind that its meaning and use have some
806 limitation. In fact, owing to the complexity of aerosol wet removal processes (Andronache, 2003),
807 many uncertainties in the physical meaning of the scavenging factors remain. The wet removal
808 process may indeed be split into in-cloud processes, characterized by hygroscopicity-induced
809 nucleation mechanisms, and below-cloud processes, a highly stochastic process based on the
810 impaction of droplets (snowflakes) with aerosol particles with a huge dependency on particle size.
811 Variability of ω is also influenced by a large number of factors, such as wind speed, temperature,
812 precipitation amount and type, the time elapsed between two precipitation events, nucleation
813 efficiency, vertical concentration differences, cloud type, etc. (Cheng and Zhang, 2017; Duce et al.,
814 1991; Luan et al., 2019; Oduber et al., 2021).

815

816 4. Conclusions

817 Characterization of ion composition together with anemological analyses sheds a light into the high
818 seasonal variability and various sources of origin, therefore add the understanding of factors
819 influencing aerosol and precipitation ion chemistry in Moscow urban background. Total inorganic
820 ions compose the biggest fraction of aerosol PM_{10} mass, up to 27% in autumn. Degree of

821 neutralization in summer and anions in autumn, and substantially electroneutrality in spring
822 constraints by the season-related neutralized agents such as ammonia and carbonates. Dominance of
823 secondary inorganic aerosols (SIA) over all the other ion species exists mostly as NH_4NO_3 and
824 $(\text{NH}_4)_2\text{SO}_4$ in all seasons. Backward trajectories within a cluster of the prevalent air mass
825 transportation indicate the highest regional SIA source in the northwestern direction in spring and
826 summer. Sulfates and nitrates dominate among the SIA components, with maximum contribution of
827 38 and 34%, respectively, in autumn. Similarity of SO_4^{2-} and NH_4^+ relationships among the wind
828 direction and speed reflects the significant thermodynamic stability of ammonium sulfates originated
829 from SO_2 gas-to-particle conversion preferably in the southeast industrially developed Moscow area.
830 High concentrations of NO_3^- are due to ground-level traffic as well as the centralized heating system
831 operation. Salt components (Na^+ and Cl^-) are attributed to a high amount of de-icing agents used in
832 road management. Na^+ shows a maximum in summer and minimum in autumn, with the largest
833 variability between seasons, differently from Cl^- . While chloride is a highly reactive mobile substance
834 and degassed by atmospheric processing, conservative potassium accumulates in soil and stabilizes
835 by ion exchange chemistry then recirculates in the atmosphere due to soil resuspension. Chlorine
836 depletion phenomenon follows the snow melt and demonstrates the biggest Cl^- displacing from NaCl
837 promoted in summer by both a significant low-level turbulence and dryness, and the required HNO_3
838 photochemistry. Ca^{2+} and Mg^{2+} relate to local soil resuspension after snow thaw, Ca^{2+} shows an
839 additional source from intensive building activities in the New Moscow area. Biomass burning is
840 indicated from remote sources in the region surrounding Moscow by high K^+ concentrations in all
841 seasons. High frequencies of backward trajectories from the European part of Russia in warm seasons
842 confirm the large agriculture and wildfires impact to the air pollution of a megacity. Main sources of
843 PM_{10} inorganic ions are secondary sulfates and nitrates, soil resuspension, road salt, and biomass
844 burning in spring and autumn. Organic acid anions, K^+ , and anhydrosugars detect biomass burning
845 associated to summer wildfires. Concerning PM_{10} and ionic data there is no particularly critical issues
846 about atmospheric pollution in Moscow urban background, in comparison to evaluated large
847 European and Canadian cities. Lower content of SO_2^- , probably due to centralizing gas operated
848 heating system, and comparable NO_3^- concentrations impacted by intensive traffic and industry, are
849 found. Impact of deicing agents on Cl^- and Na^+ concentration is similar to inland cities with cold
850 climate.

851 Precipitation collected in the same period shows the link between aerosol and precipitation
852 scavenging, with increasing acidity in spring and summer. The lowest values of scavenging factor
853 for NO_3^- and NH_4^+ was determined in summer and for SO_4^{2-} in spring, while the maximum for NO_3^-

854 in spring. The finding obtained can support the mitigation efforts contributing to the reduction of
855 aerosol mass load and consequently of aerosol and precipitation acidity.

856

857 **Acknowledgements**

858 Data collection and initial treatment was funded by Russian Science Foundation, grant number 19-
859 77-30004. Authors thank to a grant of the Ministry of Science and Higher Education of Russian
860 Federation under the Agreement (075-15-2021-574) for data analyses support. This research was
861 performed according to the Development program of the Interdisciplinary Scientific and Educational
862 School of Lomonosov Moscow State University «Future Planet and Global Environmental Change»,
863 it was carried out using the equipment of MSU Shared Research Equipment Center “Technologies
864 for obtaining new nanostructured materials and their complex study” which were purchased by MSU
865 in the frame of the Equipment Renovation Program (National Project “Science”). Dr. L. Golobokova
866 and Dr. T. Khodzher (Limnological Institute RAS) for partial performance of ion composition
867 analyses.

868

869

870 **References**

- 871 Alastuey, A., Querol, X., Rodríguez, S., Plana, F., Lopez-Soler, A., Ruiz, C., Mantilla, E., 2004.
872 Monitoring of atmospheric particulate matter around sources of secondary inorganic aerosol.
873 *Atmos. Environ.* 38, 4979–4992. <https://doi.org/10.1016/j.atmosenv.2004.06.026>
- 874 Andronache, C., 2003. Estimated variability of below-cloud aerosol removal by rainfall for observed
875 aerosol size distributions. *Atmos. Chem. Phys.* 3, 131–143. [https://doi.org/10.5194/ACP-3-131-](https://doi.org/10.5194/ACP-3-131-2003)
876 2003
- 877 Baklanov, A., Molina, L.T., Gauss, M., 2016. Megacities, air quality and climate. *Atmos. Environ.*
878 126, 235–249. <https://doi.org/10.1016/J.ATMOSENV.2015.11.059>
- 879 Bauer, H., Claeys, M., Vermeylen, R., Schueller, E., Weinke, G., Berger, A., Puxbaum, H., 2008.
880 Arabitol and mannitol as tracers for the quantification of airborne fungal spores. *Atmos. Environ.*
881 42, 588–593. <https://doi.org/10.1016/J.ATMOSENV.2007.10.013>
- 882 Beccaceci, S., McGhee, E.A., Brown, R.J.C., Green, D.C., 2015. A Comparison Between a Semi-
883 Continuous Analyzer and Filter-Based Method for Measuring Anion and Cation Concentrations
884 in PM10 at an Urban Background Site in London. *Aerosol Sci. Technol.* 49, 793–801.
885 <https://doi.org/10.1080/02786826.2015.1073848>
- 886 Bityukova, V., Saulskaya, T., 2017. Changes of the anthropogenic impact of Moscow industrial
887 zones during the recent decades. *Vestn. Mosk. Unviersiteta, Seriya Geogr.* 24–33.
- 888 Bityukova, V.R., Mozgunov, N.A., 2019. Spatial Features Transformation of Emission from Motor
889 Vehicles in Moscow. *Geogr. Environ. Sustain.* 12, 57–73. [https://doi.org/10.24057/2071-9388-](https://doi.org/10.24057/2071-9388-2019-75)
890 2019-75
- 891 Brattich, E., Liu, H., Zhang, B., Hernández-Ceballos, M.Á., Paatero, J., Sarvan, D., Djurdjevic, V.,
892 Tositti, L., Ajtić, J., 2021. Observation and modeling of high-⁷Be concentration events at the
893 surface in northern Europe associated with the instability of the Arctic polar vortex in early 2003.
894 *Atmos. Chem. Phys.* 21, 17927–17951. <https://doi.org/10.5194/ACP-21-17927-2021>
- 895 Brattich, E., Riccio, A., Tositti, L., Cristofanelli, P., Bonasoni, P., 2015. An outstanding Saharan dust
896 event at Mt. Cimone (2165 m a.s.l., Italy) in March 2004. *Atmos. Environ.* 113, 223–235.
897 <https://doi.org/10.1016/j.atmosenv.2015.05.017>
- 898 Bricker, O.P., Rice, K.C., 2003. Acid Rain. *Annu. Rev. Earth Planet. Sci.* 21, 151–174.
899 <https://doi.org/http://dx.doi.org/10.1146/annurev.ea.21.050193.001055>
- 900 Carslaw, D.C., Beevers, S.D., 2013. Characterising and understanding emission sources using

901 bivariate polar plots and k-means clustering. *Environ. Model. Softw.* 40, 325–329.
902 <https://doi.org/10.1016/j.envsoft.2012.09.005>

903 Carslaw, D.C., Ropkins, K., 2012. Openair - An r package for air quality data analysis. *Environ.*
904 *Model. Softw.* 27–28, 52–61. <https://doi.org/10.1016/j.envsoft.2011.09.008>

905 Cheng, I., Zhang, L., 2017. Long-term air concentrations, wet deposition, and scavenging ratios of
906 inorganic ions, HNO₃, and SO₂ and assessment of aerosol and precipitation acidity at Canadian
907 rural locations. *Atmos. Chem. Phys.* 17, 4711–4730. [https://doi.org/10.5194/ACP-17-4711-](https://doi.org/10.5194/ACP-17-4711-2017)
908 2017

909 Chow, J., Lowenthal, D.H., Antony Chen, L., Wang, X., Watson, J.G., 2015. Mass reconstruction
910 methods for PM_{2.5}: a review. *Air Qual. Atmos. Heal.* 8, 243–263.
911 <https://doi.org/10.1007/s11869-015-0338-3>

912 Chubarova, N.E., Nezval', E.I., Belikov, I.B., Gorbarenko, E. V., Eremina, I.D., Zhdanova, E.Y.,
913 Korneva, I.A., Konstantinov, P.I., Lokoshchenko, M.A., Skorokhod, A.I., Shilovtseva, O.A.,
914 2014. Climatic and environmental characteristics of Moscow megalopolis according to the data
915 of the Moscow State University Meteorological Observatory over 60 years. *Russ. Meteorol.*
916 *Hydrol.* 39, 602–613. <https://doi.org/10.3103/S1068373914090052>

917 Cui, L., Song, X., Zhong, G., 2021. Comparative analysis of three methods for hysplit atmospheric
918 trajectories clustering. *Atmosphere (Basel)*. 12, 698. <https://doi.org/10.3390/ATMOS12060698>

919 Dadashazar, H., Ma, L., Sorooshian, A., 2019. Sources of pollution and interrelationships between
920 aerosol and precipitation chemistry at a central California site. *Sci. Total Environ.* 651, 1776–
921 1787. <https://doi.org/10.1016/j.scitotenv.2018.10.086>

922 Diapouli, E., Kalogridis, A.C., Markantonaki, C., Vratolis, S., Fetfatzis, P., Colombi, C.,
923 Eleftheriadis, K., 2017. Annual Variability of Black Carbon Concentrations Originating from
924 Biomass and Fossil Fuel Combustion for the Suburban Aerosol in Athens, Greece. *Atmosphere*
925 (Basel). 8, 234. <https://doi.org/10.3390/ATMOS8120234>

926 Duce, R.A., Liss, P.S., Merrill, J.T., Atlas, E.L., Buat- Menard, P., Hicks, B.B., Miller, J.M.,
927 Prospero, J.M., Arimoto, R., Church, T.M., Ellis, W., Galloway, J.N., Hansen, L., Jickells, T.D.,
928 Knap, A.H., Reinhardt, K.H., Schneider, B., Soudine, A., Tokos, J.J., Tsunogai, S., Wollast, R.,
929 Zhou, M., 1991. The atmospheric input of trace species to the world ocean. *Global Biogeochem.*
930 *Cycles* 5, 193–259. <https://doi.org/10.1029/91GB01778>

931 Elansky, N., 2014. Air quality and CO emissions in the Moscow megacity. *Urban Clim.* 8, 42–56.

- 932 <https://doi.org/10.1016/j.uclim.2014.01.007>
- 933 Elansky, N.F., Lavrova, O.V., Rakin, A.A., Skorokhod, A.I., 2014. Anthropogenic disturbances of
934 the atmosphere in Moscow region. *Dokl. Earth Sci.* 454, 158–162.
935 <https://doi.org/10.1134/S1028334X14020020>
- 936 Elansky, N.F., Lokoshchenko, M.A., Belikov, I.B., Skorokhod, A.I., Shumskii, R.A., 2007.
937 Variability of trace gases in the atmospheric surface layer from observations in the city of
938 Moscow. *Izv. - Atmos. Ocean Phys.* 43, 219–231. <https://doi.org/10.1134/S0001433807020089>
- 939 Elansky, N.F., Ponomarev, N.A., Verevkin, Y.M., 2018. Air quality and pollutant emissions in the
940 Moscow megacity in 2005–2014. *Atmos. Environ.* 175, 54–64.
941 <https://doi.org/10.1016/j.atmosenv.2017.11.057>
- 942 Encinas, D., Calzada, I., Casado, H., 2010. Scavenging Ratios in an Urban Area in the Spanish Basque
943 Country. *Aerosol Sci. Technol.* 38, 685–691.
944 <https://doi.org/http://dx.doi.org/10.1080/02786820490460716>
- 945 Engelmann, R.J., 1971. Scavenging Prediction Using Ratios of Concentrations in Air and
946 Precipitation in: *Journal of Applied Meteorology and Climatology* Volume 10 Issue 3 (1971). *J.*
947 *Appl. Meteorol. Climatol.* 10, 493–497. [https://doi.org/https://doi.org/10.1175/1520-0450\(1971\)010<0493:SPUROC>2.0.CO;2](https://doi.org/https://doi.org/10.1175/1520-0450(1971)010<0493:SPUROC>2.0.CO;2)
- 949 Eremina, I.D., 2019. Chemical composition of atmospheric precipitation in Moscow and the trends
950 of its long-term changes | Eremina | *Vestnik Moskovskogo universiteta. Seriya 5, Geografiya.*
951 *Vestn. Mosk. Unviersiteta, Seriya Geogr.* 3–10.
- 952 Eremina, I.D., Aloyan, A.E., Arutyunyan, V.O., Larin, I.K., Chubarova, N.E., Yermakov, A.N., 2017.
953 Hydrocarbonates in atmospheric precipitation of Moscow: Monitoring data and analysis. *Izv.*
954 *Atmos. Ocean. Phys.* 53, 334–342. <https://doi.org/10.1134/S0001433817030069>
- 955 Eremina, I.D., Aloyan, A.E., Arutyunyan, V.O., Larin, I.K., Chubarova, N.E., Yermakov, A.N., 2015.
956 Acidity and mineral composition of precipitation in Moscow: Influence of deicing salts. *Izv. -*
957 *Atmos. Ocean Phys.* 51, 624–632. <https://doi.org/10.1134/S0001433815050047>
- 958 Eremina, I.D., Vasil'chuk, J.Y., 2019. Temporal Variations in Chemical Composition of Snow Cover
959 in Moscow. *Geogr. Environ. Sustain.* 12, 148–158. <https://doi.org/10.24057/2071-9388-2019-79>
960 79
- 961 Finardi, S., Silibello, C., D'Allura, A., Radice, P., 2014. Analysis of pollutants exchange between the
962 Po Valley and the surrounding European region. *Urban Clim.* 10, 682–702.

963 <https://doi.org/10.1016/J.UCLIM.2014.02.002>

964 Finlayson-Pitts, B.J., Pitts, J.N., 2000. Chemistry of the Upper and Lower Atmosphere: Theory,
965 Experiments, and Applications, First Ed. ed. Academic Press, San Diego, CA.
966 <https://doi.org/https://doi.org/10.1016/B978-0-12-257060-5.X5000-X>

967 Font, A., Guiseppin, L., Blangiardo, M., Ghersi, V., Fuller, G.W., 2019. A tale of two cities: is air
968 pollution improving in Paris and London? *Environ. Pollut.* 249, 1–12.
969 <https://doi.org/10.1016/J.ENVPOL.2019.01.040>

970 Fountoukis, C., Nenes, A., 2007. ISORROPIAII: A computationally efficient thermodynamic
971 equilibrium model for K⁺-Ca²⁺-Mg²⁺-NH₄⁺-Na⁺-SO₄²⁻-NO₃⁻-Cl⁻-H₂O aerosols. *Atmos.*
972 *Chem. Phys.* 7, 4639–4659. <https://doi.org/10.5194/acp-7-4639-2007>

973 Fuzzi, S., Baltensperger, U., Carslaw, K., Decesari, S., Denier Van Der Gon, H., Facchini, M.C.,
974 Fowler, D., Koren, I., Langford, B., Lohmann, U., Nemitz, E., Pandis, S., Riipinen, I., Rudich,
975 Y., Schaap, M., Slowik, J.G., Spracklen, D. V., Vignati, E., Wild, M., Williams, M., Gilardoni,
976 S., 2015. Particulate matter, air quality and climate: Lessons learned and future needs. *Atmos.*
977 *Chem. Phys.* 15, 8217–8299. <https://doi.org/10.5194/ACP-15-8217-2015>

978 Gentner, D.R., Jathar, S.H., Gordon, T.D., Bahreini, R., Day, D.A., El Haddad, I., Hayes, P.L., Pieber,
979 S.M., Platt, S.M., De Gouw, J., Goldstein, A.H., Harley, R.A., Jimenez, J.L., Prévôt, A.S.H.,
980 Robinson, A.L., 2017. Review of Urban Secondary Organic Aerosol Formation from Gasoline
981 and Diesel Motor Vehicle Emissions. *Environ. Sci. Technol.* 51, 1074–1093.
982 https://doi.org/10.1021/ACS.EST.6B04509/SUPPL_FILE/ES6B04509_SI_001.PDF

983 Golitsyn, G.S., Grechko, E.I., Wang, G., Wang, P., Dzhola, A. V., Emilenko, A.S., Kopeikin, V.M.,
984 Rakitin, V.S., Safronov, A.N., Fokeeva, E. V., 2015. Studying the pollution of Moscow and
985 Beijing atmospheres with carbon monoxide and aerosol. *Izv. - Atmos. Ocean Phys.* 51, 1–11.
986 <https://doi.org/10.1134/S0001433815010041>

987 Golobokova, L., Khodzher, T., Khuriganova, O., Marinayte, I., Onishchuk, N., Rusanova, P.,
988 Potemkin, V., 2020. Variability of chemical properties of the atmospheric aerosol above lake
989 baikal during large wildfires in siberia. *Atmosphere (Basel)*. 11, 1–21.
990 <https://doi.org/10.3390/atmos11111230>

991 Guo, W., Zhang, X., Zhang, Z., Zheng, N., Xiao, Hongwei, Xiao, Huayun, 2021. Low-molecular-
992 weight carboxylates in urban southwestern China: Source identification and effects on aerosol
993 acidity. *Atmos. Pollut. Res.* 12, 101141. <https://doi.org/10.1016/J.APR.2021.101141>

- 994 Hama, S.M.L., Cordell, R.L., Staelens, J., Mooibroek, D., Monks, P.S., 2018. Chemical composition
995 and source identification of PM₁₀ in five North Western European cities. *Atmos. Res.* 214, 135–
996 149. <https://doi.org/10.1016/J.ATMOSRES.2018.07.014>
- 997 Heal, M.R., Kumar, P., Harrison, R.M., 2012. Particles, air quality, policy and health. *Chem. Soc.*
998 *Rev.* 41, 6606–6630. <https://doi.org/10.1039/c2cs35076a>
- 999 Hewitt, C.N., 2001. The atmospheric chemistry of sulphur and nitrogen in power station plumes.
1000 *Atmos. Environ.* [https://doi.org/10.1016/S1352-2310\(00\)00463-5](https://doi.org/10.1016/S1352-2310(00)00463-5)
- 1001 Hoffman, R.C., Laskin, A., Finlayson-Pitts, B.J., 2004. Sodium nitrate particles: physical and
1002 chemical properties during hydration and dehydration, and implications for aged sea salt
1003 aerosols. *J. Aerosol Sci.* 35, 869–887. <https://doi.org/10.1016/J.JAEROSCI.2004.02.003>
- 1004 Hopke, P.K., 2016. Review of receptor modeling methods for source apportionment. *J. Air Waste*
1005 *Manag. Assoc.* <https://doi.org/10.1080/10962247.2016.1140693>
- 1006 Hristova, E., Veleva, B., Georgieva, E., Branzov, H., 2020. Application of Positive Matrix
1007 Factorization Receptor Model for Source Identification of PM₁₀ in the City of Sofia, Bulgaria.
1008 *Atmos.* 2020, Vol. 11, Page 890 11, 890. <https://doi.org/10.3390/ATMOS11090890>
- 1009 Jeong, C.H., Traub, A., Huang, A., Hilker, N., Wang, J.M., Herod, D., Dabek-Zlotorzynska, E., Celio,
1010 V., Evans, G.J., 2020. Long-term analysis of PM_{2.5} from 2004 to 2017 in Toronto: Composition,
1011 sources, and oxidative potential. *Environ. Pollut.* 263, 114652.
1012 <https://doi.org/10.1016/J.ENVPOL.2020.114652>
- 1013 Karydis, V., Tsimpidi, A., Pozzer, A., Lelieveld, J., 2020. How alkaline compounds control
1014 atmospheric aerosol acidity. *Atmos. Chem. Phys.* 2018, 1–22. <https://doi.org/10.5194/acp-2020-1222>
- 1016 Kasimov, N.S., Kosheleva, N.E., Vlasov, D. V., Nabelkina, K.S., Ryzhov, A. V., 2019.
1017 Physicochemical Properties of Road Dust in Moscow. *Geogr. Environ. Sustain.* 12, 96–113.
1018 <https://doi.org/10.24057/2071-9388-2019-55>
- 1019 Kasper-Giebl, A., Kalina, M.F., Puxbaum, H., 1999. Scavenging ratios for sulfate, ammonium and
1020 nitrate determined at Mt. Sonnblick (3106 m a.s.l.). *Atmos. Environ.* 33, 895–906.
1021 [https://doi.org/10.1016/S1352-2310\(98\)00279-9](https://doi.org/10.1016/S1352-2310(98)00279-9)
- 1022 Kawamura, K., Bikkina, S., 2016. A review of dicarboxylic acids and related compounds in
1023 atmospheric aerosols: Molecular distributions, sources and transformation. *Atmos. Res.*
1024 <https://doi.org/10.1016/j.atmosres.2015.11.018>

- 1025 Koçak, M., Mihalopoulos, N., Kubilay, N., 2007. Chemical composition of the fine and coarse
1026 fraction of aerosols in the northeastern Mediterranean. *Atmos. Environ.* 41, 7351–7368.
1027 <https://doi.org/10.1016/j.atmosenv.2007.05.011>
- 1028 Kolesar, K.R., Mattson, C.N., Peterson, P.K., May, N.W., Prendergast, R.K., Pratt, K.A., 2018.
1029 Increases in wintertime PM_{2.5} sodium and chloride linked to snowfall and road salt application.
1030 *Atmos. Environ.* 177, 195–202. <https://doi.org/10.1016/J.ATMOSENV.2018.01.008>
- 1031 Korolev, V.A., Gorniyakov, A.K., 2018. Impact assessment of the anti-icing reagents application
1032 when conducting engineering-ecological surveys in cities. *Eng. Surv.* 1, 66–78.
1033 <https://doi.org/https://doi.org/10.25296/1997-8650-2018-1-2-66-78>
- 1034 Kosheleva, N.E., Dorokhova, M.F., Kuzminskaya, N.Y., Ryzhov, A.V., Kasimov, N.S., 2018. Impact
1035 of motor vehicles on the ecological state of soils in the western district of Moscow. *Moscow*
1036 *Univ. Bull. Ser. 5. Geogr.* 2, 16–27.
- 1037 Kotowski, T., Motyka, J., Knap, W., Bielewski, J., 2020. 17-Year study on the chemical composition
1038 of rain, snow and sleet in very dusty air (Krakow, Poland). *J. Hydrol.* 582, 124543.
1039 <https://doi.org/10.1016/J.JHYDROL.2020.124543>
- 1040 Kul'bachevskii, A.O., 2021. Report on the state of the environment in Moscow in 2020 [WWW
1041 Document]. *Dep. Nat. Environ. Prot. Moscow Gov.* URL [https://www.akm.ru/eng/press/state-](https://www.akm.ru/eng/press/state-report-on-the-state-and-environmental-protection-in-russia-in-2020/)
1042 [report-on-the-state-and-environmental-protection-in-russia-in-2020/](https://www.akm.ru/eng/press/state-report-on-the-state-and-environmental-protection-in-russia-in-2020/) (accessed 1.27.22).
- 1043 Lai, S. chao, Zou, S. chun, Cao, J. ji, Lee, S. cheng, Ho, K. fai, 2007. Characterizing ionic species in
1044 PM_{2.5} and PM₁₀ in four Pearl River Delta cities, South China. *J. Environ. Sci.* 19, 939–947.
1045 [https://doi.org/10.1016/S1001-0742\(07\)60155-7](https://doi.org/10.1016/S1001-0742(07)60155-7)
- 1046 Laongsri, B., Harrison, R.M., 2013. Atmospheric behaviour of particulate oxalate at UK urban
1047 background and rural sites. *Atmos. Environ.* 71, 319–326.
1048 <https://doi.org/10.1016/j.atmosenv.2013.02.015>
- 1049 Luan, T., Guo, X., Zhang, T., Guo, L., 2019. Below-Cloud Aerosol Scavenging by Different-Intensity
1050 Rains in Beijing City. *J. Meteorol. Res.* 2019 331 33, 126–137. [https://doi.org/10.1007/S13351-](https://doi.org/10.1007/S13351-019-8079-0)
1051 [019-8079-0](https://doi.org/10.1007/S13351-019-8079-0)
- 1052 Mann, H.B., Whitney, D.R., 1947. On a Test of Whether one of Two Random Variables is
1053 Stochastically Larger than the Other. *Ann. Math. Stat.* 18, 50–60.
1054 <https://doi.org/https://doi.org/10.1214/aoms/1177730491>
- 1055 McNamara, S.M., Kolesar, K.R., Wang, S., Kirpes, R.M., May, N.W., Gunsch, M.J., Cook, R.D.,

- 1056 Fuentes, J.D., Hornbrook, R.S., Apel, E.C., China, S., Laskin, A., Pratt, K.A., 2020. Observation
1057 of Road Salt Aerosol Driving Inland Wintertime Atmospheric Chlorine Chemistry. *ACS Cent.*
1058 *Sci.* 6, 684–694. <https://doi.org/10.1021/ACSCENTSCI.9B00994>
- 1059 Medeiros, P.M., Conte, M.H., Weber, J.C., Simoneit, B.R.T., 2006. Sugars as source indicators of
1060 biogenic organic carbon in aerosols collected above the Howland Experimental Forest, Maine.
1061 *Atmos. Environ.* 40, 1694–1705. <https://doi.org/10.1016/J.ATMOSENV.2005.11.001>
- 1062 Morozzi, P., Bolelli, L., Brattich, E., Ferri, E.N., Girotti, S., Sangiorgi, S., Orza, J.A.G., Piñero-
1063 García, F., Tositti, L., 2021. Chemiluminescent fingerprints from airborne particulate matter: A
1064 luminol-based assay for the characterization of oxidative potential with kinetical implications.
1065 *Sci. Total Environ.* 789, 148005. <https://doi.org/10.1016/J.SCITOTENV.2021.148005>
- 1066 Nava, S., Lucarelli, F., Amato, F., Becagli, S., Calzolari, G., Chiari, M., Giannoni, M., Traversi, R.,
1067 Udusti, R., 2015. Biomass burning contributions estimated by synergistic coupling of daily and
1068 hourly aerosol composition records. *Sci. Total Environ.* 511, 11–20.
1069 <https://doi.org/10.1016/J.SCITOTENV.2014.11.034>
- 1070 Nikiforova, E.M., Kasimov, N.S., Kosheleva, N.E., 2017. Long-term dynamics of anthropogenic
1071 solonchicity in soils of the Eastern okrug of Moscow under the impact of deicing salts. *Eurasian*
1072 *Soil Sci.* 50, 84–94. <https://doi.org/10.1134/S1064229317010100>
- 1073 Nikiforova, E.M., Kasimov, N.S., Kosheleva, N.E., 2014. Long-term dynamics of the anthropogenic
1074 salinization of soils in Moscow (by the example of the Eastern district). *Eurasian Soil Sci.* 47,
1075 203–215. <https://doi.org/10.1134/S1064229314030041>
- 1076 Oduber, F., Calvo, A.I., Blanco-Alegre, C., Castro, A., Alves, C., Cerqueira, M., Lucarelli, F., Nava,
1077 S., Calzolari, G., Martin-Villacorta, J., Esteves, V., Fraile, R., 2021. Towards a model for aerosol
1078 removal by rain scavenging: The role of physical-chemical characteristics of raindrops. *Water*
1079 *Res.* 190, 116758. <https://doi.org/10.1016/J.WATRES.2020.116758>
- 1080 Ott, W.R., 1990. A physical explanation of the lognormality of pollutant concentrations. *J. Air Waste*
1081 *Manage. Assoc.* 40, 1378–1383. <https://doi.org/10.1080/10473289.1990.10466789>
- 1082 Pavuluri, C.M., Kawamura, K., Swaminathan, T., 2010. Water-soluble organic carbon, dicarboxylic
1083 acids, ketoacids, and α -dicarbonyls in the tropical Indian aerosols. *J. Geophys. Res. Atmos.* 115,
1084 D11302. <https://doi.org/10.1029/2009JD012661>
- 1085 Perrone, M.G., Vratolis, S., Georgieva, E., Török, S., Šega, K., Veleva, B., Osán, J., Bešlić, I.,
1086 Kertész, Z., Pernigotti, D., Eleftheriadis, K., Belis, C.A., 2018. Sources and geographic origin

1087 of particulate matter in urban areas of the Danube macro-region: The cases of Zagreb (Croatia),
1088 Budapest (Hungary) and Sofia (Bulgaria). *Sci. Total Environ.* 619–620, 1515–1529.
1089 <https://doi.org/10.1016/J.SCITOTENV.2017.11.092>

1090 Pio, C.A., Silva, P.A., Cerqueira, M.A., Nunes, T. V., 2005. Diurnal and seasonal emissions of
1091 volatile organic compounds from cork oak (*Quercus suber*) trees. *Atmos. Environ.* 39, 1817–
1092 1827. <https://doi.org/10.1016/j.atmosenv.2004.11.018>

1093 Popovicheva, O., Chichaeva, M., Kobelev, V., Sinitskiy, A., Hansen, A., 2020a. Black Carbon in
1094 urban emissions on the Polar Circle, in: 26th International Symposium on Atmospheric and
1095 Ocean Optics, Atmospheric Physics. SPIE-Intl Soc Optical Eng, p. 344.
1096 <https://doi.org/10.1117/12.2577550>

1097 Popovicheva, O., Chichaeva, M., Kovach, R., Zhdanova, E., Kasimov, N., 2022. Seasonal, Weekly,
1098 and Diurnal Black Carbon in Moscow Megacity Background under Impact of Urban and
1099 Regional Sources. *Atmosphere (Basel)*. 13, 563. <https://doi.org/10.3390/ATMOS13040563>

1100 Popovicheva, O., Ivanov, A., Vojtisek, M., 2020b. Functional Factors of Biomass Burning
1101 Contribution to Spring Aerosol Composition in a Megacity: Combined FTIR-PCA Analyses.
1102 *Atmosphere (Basel)*. 11, 319. <https://doi.org/https://doi.org/10.3390/atmos11040319>

1103 Popovicheva, O., Kistler, M., Kireeva, E., Persiantseva, N., Timofeev, M., Kopeikin, V., Kasper-
1104 Giebl, A., 2014. Physicochemical characterization of smoke aerosol during large-scale wildfires:
1105 Extreme event of August 2010 in Moscow. *Atmos. Environ.* 96, 405–414.

1106 Popovicheva, O., Padoan, S., Schnelle-Kreis, J., Nguyen, D.-L., Adam, T., Kistler, M., Steinkogler,
1107 T., Kasper-Giebl, A., Zimmerman, R., Chubarova, N., 2020c. Spring aerosol in urban
1108 atmosphere of megacity: analytical and statistical assessment for source impacts. *Aerosol Air*
1109 *Qual. Res.* 20. <https://doi.org/10.4209/aaqr.2019.08.0412>

1110 Popovicheva, O., Volpert, E., Sitnikov, N.M., Chichaeva, M.A., Padoan, S., 2020d. Black carbon in
1111 spring aerosols of Moscow urban background. *Geogr. Environ. Sustain.* 13, 233–243.
1112 <https://doi.org/10.24057/2071-9388-2019-90>

1113 Pöschl, U., 2005. Atmospheric aerosols: Composition, transformation, climate and health effects.
1114 *Angew. Chemie - Int. Ed.* 44, 7520–7540. <https://doi.org/10.1002/anie.200501122>

1115 Pöschl, U., Shiraiwa, M., 2015. Multiphase Chemistry at the Atmosphere-Biosphere Interface
1116 Influencing Climate and Public Health in the Anthropocene. *Chem. Rev.* 115, 4440–4475.
1117 <https://doi.org/10.1021/cr500487s>

- 1118 Prinn, R.G., 2003. The Cleansing Capacity of the Atmosphere. *Annu. Rev. Environ. Resour.* 28, 29–
1119 57. <https://doi.org/https://doi.org/10.1146/annurev.energy.28.011503.163425>
- 1120 Pruppacher, H.R., Klett, J.D., 2010. *Microphysics of Clouds and Precipitation*, Atmospheric and
1121 Oceanographic Sciences Library. Springer Netherlands, Dordrecht. [https://doi.org/10.1007/978-](https://doi.org/10.1007/978-0-306-48100-0)
1122 0-306-48100-0
- 1123 Putaud, J.P., Van Dingenen, R., Alastuey, A., Bauer, H., Birmili, W., Cyrys, J., Flentje, H., Fuzzi, S.,
1124 Gehrig, R., Hansson, H.C., Harrison, R.M., Herrmann, H., Hitzenberger, R., Hüglin, C., Jones,
1125 A.M., Kasper-Giebl, A., Kiss, G., Koussa, A., Kuhlbusch, T.A.J., Löschau, G., Maenhaut, W.,
1126 Molnar, A., Moreno, T., Pekkanen, J., Perrino, C., Pitz, M., Puxbaum, H., Querol, X., Rodriguez,
1127 S., Salma, I., Schwarz, J., Smolik, J., Schneider, J., Spindler, G., ten Brink, H., Tursic, J., Viana,
1128 M., Wiedensohler, A., Raes, F., 2010. A European aerosol phenomenology - 3: Physical and
1129 chemical characteristics of particulate matter from 60 rural, urban, and kerbside sites across
1130 Europe. *Atmos. Environ.* 44, 1308–1320. <https://doi.org/10.1016/j.atmosenv.2009.12.011>
- 1131 Raes, F., Van Dingenen, R., Vignati, E., Wilson, J., Putaud, J.-P., Seinfeld, J.H., Adams, P., 2000.
1132 Formation and cycling of aerosols in the global troposphere. *Atmos. Environ.*
1133 [https://doi.org/10.1016/S1352-2310\(00\)00239-9](https://doi.org/10.1016/S1352-2310(00)00239-9)
- 1134 Saraga, D., Maggos, T., Degrendele, C., Klánová, J., Horvat, M., Kocman, D., Kanduč, T., Garcia
1135 Dos Santos, S., Franco, R., Gómez, P.M., Manousakas, M., Bairachtari, K., Eleftheriadis, K.,
1136 Kermenidou, M., Karakitsios, S., Gotti, A., Sarigiannis, D., 2021. Multi-city comparative PM_{2.5}
1137 source apportionment for fifteen sites in Europe: The ICARUS project. *Sci. Total Environ.* 751,
1138 141855. <https://doi.org/10.1016/j.scitotenv.2020.141855>
- 1139 Schauer, J.J., Kleeman, M.J., Cass, G.R., Simoneit, B.R.T., 2002. Measurement of emissions from
1140 air pollution sources. 5. C₁ - C₃₂ organic compounds from gasoline-powered motor vehicles.
1141 *Environ. Sci. Technol.* 36, 1169–1180. <https://doi.org/10.1021/es0108077>
- 1142 Schwarz, J., Pokorná, P., Rychlík, Š., Škáčová, H., Vlček, O., Smolík, J., Ždímal, V., Hůnová, I.,
1143 2019. Assessment of air pollution origin based on year-long parallel measurement of PM_{2.5} and
1144 PM₁₀ at two suburban sites in Prague, Czech Republic. *Sci. Total Environ.* 664, 1107–1116.
1145 <https://doi.org/10.1016/J.SCITOTENV.2019.01.426>
- 1146 Seinfeld, J.H., Pandis, S.N., 2016. *Atmospheric Chemistry and Physics: From Air Pollution to*
1147 *Climate Change*, Environment: Science and Policy for Sustainable Development.
1148 <https://doi.org/10.1080/00139157.1999.10544295>
- 1149 Seinfeld, J.H., Pandis, S.N., 1998. Chemistry of the atmospheric aqueous phase. *Atmos. Chem. Phys.*

1150 From Air Pollut. to Clim. Chang. 337407.

1151 Spearman, C., 2010. The Proof and Measurement of Association between Two Things. *Int. J.*
1152 *Epidemiol.* 39, 1137–1150. <https://doi.org/https://doi.org/10.1093/ije/dyq191>

1153 Srivastava, D., Favez, O., Bonnaire, N., Lucarelli, F., Haeffelin, M., Perraudin, E., Gros, V.,
1154 Villenave, E., Albinet, A., 2018. Speciation of organic fractions does matter for aerosol source
1155 apportionment. Part 2: Intensive short-term campaign in the Paris area (France). *Sci. Total*
1156 *Environ.* 634, 267–278. <https://doi.org/10.1016/J.SCITOTENV.2018.03.296>

1157 Stein, A.F., Draxler, R.R., Rolph, G.D., Stunder, B.J.B., Cohen, M.D., Ngan, F., 2015. Noaa’s hysplit
1158 atmospheric transport and dispersion modeling system. *Bull. Am. Meteorol. Soc.* 96, 2059–
1159 2077. <https://doi.org/10.1175/BAMS-D-14-00110.1>

1160 Stieger, B., Spindler, G., Fahlbusch, B., Müller, K., Grüner, A., Poulain, L., Thöni, L., Seitler, E.,
1161 Wallasch, M., Herrmann, H., 2018. Measurements of PM10 ions and trace gases with the online
1162 system MARGA at the research station Melpitz in Germany – A five-year study. *J. Atmos.*
1163 *Chem.* 75, 33–70. <https://doi.org/10.1007/S10874-017-9361-0>

1164 Tang, L., Haeger-Eugensson, M., Sjöberg, K., Wichmann, J., Molnár, P., Sallsten, G., 2014.
1165 Estimation of the long-range transport contribution from secondary inorganic components to
1166 urban background PM10 concentrations in south-western Sweden during 1986-2010. *Atmos.*
1167 *Environ.* 89, 93–101. <https://doi.org/10.1016/J.ATMOSENV.2014.02.018>

1168 Tositti, L., 2018. The relationship between health effects and airborne particulate constituents, in:
1169 *Clinical Handbook of Air Pollution-Related Diseases*. Springer International Publishing, pp. 33–
1170 54. https://doi.org/10.1007/978-3-319-62731-1_3

1171 Tositti, L., Brattich, E., Cassardo, C., Morozzi, P., Bracci, A., Marinoni, A., Di Sabatino, S., Porcù,
1172 F., Zappi, A., 2021. Development and evolution of an anomalous Asian dust event across Europe
1173 in March 2020. *Atmos. Chem. Phys. Discuss.* 1–39. <https://doi.org/10.5194/ACP-2021-429>

1174 Tositti, L., Brattich, E., Masiol, M., Baldacci, D., Ceccato, D., Parmeggiani, S., Stracquadanio, M.,
1175 Zappoli, S., 2014. Source apportionment of particulate matter in a large city of southeastern Po
1176 Valley (Bologna, Italy). *Environ. Sci. Pollut. Res.* 21, 872–890. <https://doi.org/10.1007/s11356-013-1911-7>

1178 Tositti, L., Brattich, E., Parmeggiani, S., Bolelli, L., Ferri, E., Girotti, S., 2018a. Airborne particulate
1179 matter biotoxicity estimated by chemometric analysis on bacterial luminescence data. *Sci. Total*
1180 *Environ.* 640–641, 1512–1520. <https://doi.org/10.1016/J.SCITOTENV.2018.06.024>

- 1181 Tositti, L., Pieri, L., Brattich, E., Parmeggiani, S., Ventura, F., 2018b. Chemical characteristics of
1182 atmospheric bulk deposition in a semi-rural area of the Po Valley (Italy). *J. Atmos. Chem.* 75,
1183 97–121. <https://doi.org/10.1007/s10874-017-9365-9>
- 1184 Tsai, Y.I., Kuo, S.C., 2013. Contributions of low molecular weight carboxylic acids to aerosols and
1185 wet deposition in a natural subtropical broad-leaved forest environment. *Atmos. Environ.* 81,
1186 270–279. <https://doi.org/10.1016/j.atmosenv.2013.08.061>
- 1187 Tsai, Y.I., Sopajaree, K., Chotrukha, A., Wu, H.C., Kuo, S.C., 2013. Source indicators of biomass
1188 burning associated with inorganic salts and carboxylates in dry season ambient aerosol in Chiang
1189 Mai Basin, Thailand. *Atmos. Environ.* 78, 93–104.
1190 <https://doi.org/10.1016/J.ATMOSENV.2012.09.040>
- 1191 Van Espen, P., Adams, F., 1983. The application of principal component and factor analysis
1192 procedures to data for element concentrations in aerosols from a remote region. *Anal. Chim.*
1193 *Acta* 150, 153–161. [https://doi.org/10.1016/S0003-2670\(00\)85467-7](https://doi.org/10.1016/S0003-2670(00)85467-7)
- 1194 Vet, R., Artz, R.S., Carou, S., Shaw, M., Ro, C.U., Aas, W., Baker, A., Bowersox, V.C., Dentener,
1195 F., Galy-Lacaux, C., Hou, A., Pienaar, J.J., Gillett, R., Forti, M.C., Gromov, S., Hara, H.,
1196 Khodzher, T., Mahowald, N.M., Nickovic, S., Rao, P.S.P., Reid, N.W., 2014. A global
1197 assessment of precipitation chemistry and deposition of sulfur, nitrogen, sea salt, base cations,
1198 organic acids, acidity and pH, and phosphorus. *Atmos. Environ.* 93, 3–100.
1199 <https://doi.org/10.1016/j.atmosenv.2013.10.060>
- 1200 Vlasov, D., Kosheleva, N., Kasimov, N., 2021. Spatial distribution and sources of potentially toxic
1201 elements in road dust and its PM10 fraction of Moscow megacity. *Sci. Total Environ.* 761.
1202 <https://doi.org/10.1016/j.scitotenv.2020.143267>
- 1203 Vuorenmaa, J., Augustaitis, A., Beudert, B., Bochenek, W., Clarke, N., de Wit, H.A., Dirnböck, T.,
1204 Frey, J., Hakola, H., Kleemola, S., Kobler, J., Krám, P., Lindroos, A.J., Lundin, L., Löfgren, S.,
1205 Marchetto, A., Pecka, T., Schulte-Bisping, H., Skotak, K., Srybny, A., Szpikowski, J.,
1206 Ukonmaanaho, L., Váňa, M., Åkerblom, S., Forsius, M., 2018. Long-term changes (1990–2015)
1207 in the atmospheric deposition and runoff water chemistry of sulphate, inorganic nitrogen and
1208 acidity for forested catchments in Europe in relation to changes in emissions and
1209 hydrometeorological conditions. *Sci. Total Environ.* 625, 1129–1145.
1210 <https://doi.org/10.1016/j.scitotenv.2017.12.245>
- 1211 Wang, G., Kawamura, K., Lee, S., Ho, K., Cao, J., 2006. Molecular, seasonal and spatial distributions
1212 of organic aerosols from fourteen Chinese cities. *Environ. Sci. Technol.* 40, 4619–4625.

1213 <https://doi.org/10.1021/es060291x>

1214 Wang, Y., Zhuang, G., Chen, S., An, Z., Zheng, A., 2007. Characteristics and sources of formic,
1215 acetic and oxalic acids in PM_{2.5} and PM₁₀ aerosols in Beijing, China. *Atmos. Res.* 84, 169–
1216 181. <https://doi.org/10.1016/J.ATMOSRES.2006.07.001>

1217 WHO, 2013. Health effects of particulate matter. Policy implications for countries in eastern Europe,
1218 Caucasus and central Asia (2013).

1219 Wieprecht, W., Lutz, M., John, A., Lüdtke, C., Möller, D., Acker, K., 2004. PM₁₀ Aerosol mass and
1220 composition in and around Berlin (Germany). *J. Aerosol Sci.* 35, 453–464.
1221 <https://doi.org/10.1016/J.JAEROSCI.2004.06.028>

1222 Yang, L., Yu, L.E., 2008. Measurements of oxalic acid, oxalates, malonic acid, and malonates in
1223 atmospheric particulates. *Environ. Sci. Technol.* 42, 9268–9275.
1224 <https://doi.org/10.1021/es801820z>

1225

1226



Click here to access/download

Highlighted Manuscript

Zappi_et_al_2022_manuscript_revised_round2_trackON.
docx

Reviewer #1: Comments on the manuscript "Factors influencing aerosol and precipitation ion chemistry in urban background of northern megacity"

The revised manuscript has made some responses to my previous concerns. It is clear that the authors have taken our comments seriously and have provided new information, clarifications, and new lines of evidence that help support their claim. Also, it seems to me too that the conclusions of the paper might well be supported by the observations. However, there are still some minor comments that need to be addressed.

Line 354: The high PM₁₀ concentration in spring is mainly caused by the deposition of terrestrial crustal weathering materials. So, where do these natural particulate materials mainly come from? Whether there are arid areas around the study area? And the natural environment around the study area needs to be introduced in this manuscript.

As for the crustal component the spring contribution seems to be significantly contributed by the road salt which includes a mineral insoluble fraction up to 20 % by weight beside saline compounds, added to promote friction over ice. Towards the end of spring following ice melt, it may become easily available for resuspension, together with the soil component from the surrounding area around the sampling station.

Line 401: In this part, the authors compared and analyzed the physicochemical composition of aerosols between the study area and other European cities. I suggest that the author add a few sentences to summarize the level of atmospheric environment in Moscow compared with these cities, for example, serious, moderate or light pollution.

The data presented in this work have pros and cons. Pros refer to its novelty, since it concerns a location so far limitedly investigated for its air quality, despite its utmost importance.

Cons refer to the relatively low number of samples analyzed, limiting deductions and environmental diagnosis to the period concerned.

Although a longer sampling campaign would be required to draw conclusive considerations, our data indicate that the pollution in Moscow is comparable, if not lower, than that of also smaller cities. Paris and London, moreover, have significative higher pollution, but with certainly lower dimensions and populations.

Lines 541-542: Why do these two substances exist in all seasons?

While nitrates are well known combustion byproducts of secondary origin, the origin of ammonia, gaseous precursor of nitrate counter-cation as a result of HNO₃ neutralization in the troposphere, are known to be agricultural activities, industrial chemosynthesis, especially for fertilizer production and vehicle catalytic converters due to excess NO reduction. The latter source is a potentially significant source across Moscow megacity throughout the year given the importance of vehicular traffic as a pollution source which can therefore compare to the expected seasonality of the former contribution, which is in turn affected by mesoscale to regional transportation

Line 551: The authors determined the seasonal source of aerosol in the study area based on wind direction and backward air mass trajectory during the traceability analysis of atmospheric materials. It would be nice to add a description of potential source areas.

As already stated, we suppose that most of the observed pollution is due to urban activities, and in this sense, we already described the surroundings of the sampling site (e.g.: Figure 1) with the industrial facilities location and the highway network near the Moscow State University where the sampling station is operated. The two most important and to some extent peculiar pollution sources are those emitted by wildfires in the area surrounding Moscow, already described at the end of paragraph 3.3.1, and by the resuspension of road salt used for traffic management, which is clearly described along the text. Therefore, we decided not to further modify the text.

Line 723: In this study, the authors analyzed the seasonal variation in the concentrations of various ions and BC of aerosol and precipitation, and traced their provenance by using wind direction distribution and backward air mass trajectory. However, the seasonal differences of target substances in aerosols and precipitation have not been systematically expounded. These two environmental media retain information on dry and wet deposition, and it is important to understand the differences in their chemical compositions for further understanding the different atmospheric deposition mechanisms.

In the absence of a wet and dry sampler, it is hard to discriminate and characterize accurately deposition processes, which can be further biased by the influence of local soil resuspension. All these factors have been explained accurately, including some comments concerning the difficulties in associating the origin of aerosols from which precipitation may form with local aerosols. In fact, one of the main unresolved problems in the understanding of wet scavenging mechanisms and their influence on the properties of deposition, is the split between in-cloud processes (often involving aerosol emitted/produced at the regional scales or farther) and below-cloud processes which are likely to involve local aerosol. As a result, we find hazardous to speculate further over the available data.

Figure 4: The decimal point is wrong.

We modified the Figure with the correct decimal point

Some new references about urban and remote region pollutants study should be cited in this work, e.g. Dong Z, et al., 2019. Spatial variability, mixing states and composition of various haze particles in atmosphere during winter and summertime in northwest China. *Environmental Pollution*, 246, (2019) 79-88; 21. Dong Z, et al. Individual particles of cryoconite deposited on the mountain glaciers of the Tibetan Plateau: Insights into chemical composition and sources. *Atmospheric Environment*, 2016, 138: 114~124. doi: 10.1016/j. atmosenv.2016.05.020

As previously commented, the emission sources and climatological conditions of Moscow megacities have been widely described and motivated by the authors both in the submitted paper and in this letter, with emphasis on soil and road salt resuspension which make up a relevant local contribution to the megacity aerosol.

The atmospheric dynamics behind the processes mentioned in the papers suggested by the estimated reviewer, though appealing, do not apply to our case. Moscow is not a remote area at all, as compared to Tibetan plateau, therefore the relative amount of potentially similar aerosol component is widely different. Moreover, snow cover is not permanent throughout the year; therefore, ground sources of particles become abundantly and promptly available for resuspension and recirculation at the local scale even in weakly energetic conditions. As explained in Lee et al. (2007) Lee, H. N., L. Tositti, X. Zheng, and P. Bonasoni (2007), the Tibetan plateau may be affected by China pollution clouds, but the history and fate of air masses from ground region up to the glacier resume a series of different sources and losses during transport. Analyses and comparisons of variations of ^{7}Be , ^{210}Pb , and $^{7}\text{Be}/^{210}\text{Pb}$ with ozone observations at two Global Atmosphere Watch stations from high mountains, *J. Geophys. Res.*, 112, D05303, doi:10.1029/2006JD007421.1.

Reviewer #2: The authors have revised the manuscript according to the reviewers' comments, and I recommend it to be accepted for publication.

We thank the reviewer for His/Her positive comment